

# Nature

## An overview

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# Main article

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## Nature

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*"Natural" and "Natural World" redirect here. For other uses, see [Nature \(disambiguation\)](#) and [Natural \(disambiguation\)](#).*

**Nature**, in the broadest sense, is equivalent to the **natural world**, **physical world**, or **material world**. "Nature" refers to the phenomena of the physical world, and also to life in general. It ranges in scale from the subatomic to the cosmic.

The word *nature* is derived from the Latin word *natura*, or "essential qualities, innate disposition", and in ancient times, literally meant "birth".<sup>[1]</sup> *Natura* was a Latin translation of the Greek word *physis* (φύσις), which originally related to the intrinsic characteristics that plants, animals, and other features of the world develop of their own accord.<sup>[2][3]</sup> The concept of nature as a whole, the physical universe, is one of several expansions of the original notion; it began with certain core applications of the word φύσις by pre-Socratic philosophers, and has steadily gained currency ever since. This usage was confirmed during the advent of modern scientific method in the last several centuries.<sup>[4][5]</sup>

Within the various uses of the word today, "nature" often refers to geology and wildlife. Nature may refer to the general realm of various types of living plants and animals, and in some cases to the processes associated with inanimate objects – the way

that particular types of things exist and change of their own accord, such as the weather and geology of the Earth, and the matter and energy of which all these things are composed. It is often taken to mean the "natural environment" or wilderness—wild



Hopetoun Falls, Australia



Bachalpsee in the Swiss Alps

animals, rocks, forest, beaches, and in general those things that have not been substantially altered by human intervention, or which persist despite human intervention. For example, manufactured objects and human interaction generally are not considered part of nature, unless qualified as, for example, "human nature" or "the whole of nature". This more traditional concept of natural things which can still be found today implies a distinction between the natural and the artificial, with the artificial being understood as that which has

been brought into being by a human consciousness or a human mind. Depending on the particular context, the term "natural" might also be distinguished from the unnatural, the supernatural, or synthetic.



Lightning strikes during the eruption of the huge Galunggung volcano, West Java, in 1982.

## Earth

Earth (or, "the earth") is the only planet presently known to support life, and its natural features are the subject of many fields of scientific research. Within the solar system, it is third closest to the sun; it is the largest terrestrial planet and the fifth largest overall. Its most prominent climatic features are its two large polar regions, two relatively narrow temperate zones, and a wide equatorial tropical to subtropical region.<sup>[6]</sup> Precipitation varies widely with location, from several metres of water per year to less than a millimetre. 71 percent of the Earth's surface is covered by salt-water oceans. The remainder consists of continents and islands, with most of the inhabited land in the Northern Hemisphere.

Earth has evolved through geological and biological processes that have left traces of the original conditions. The outer surface is divided into several gradually migrating tectonic plates. The interior remains active, with a thick layer of plastic mantle and an iron-filled core that generates a magnetic field.

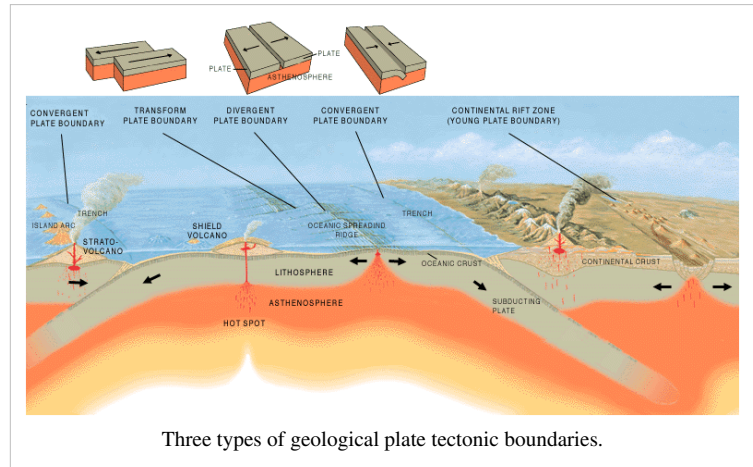
The atmospheric conditions have been significantly altered from the original conditions by the presence of life-forms,<sup>[7]</sup> which create an ecological balance that stabilizes the surface conditions. Despite the wide regional variations in climate by latitude and other geographic factors, the long-term average global climate is quite stable during interglacial periods,<sup>[8]</sup> and variations of a degree or two of average global temperature have historically had major effects on the ecological balance, and on the actual geography of the Earth.<sup>[9][10]</sup>



View of the Earth, taken in 1972 by the Apollo 17 astronaut crew. This image is the only photograph of its kind to date, showing a fully sunlit hemisphere of the Earth.

## Geology

Geology is the science and study of the solid and liquid matter that constitutes the Earth. The field of geology encompasses the study of the composition, structure, physical properties, dynamics, and history of Earth materials, and the processes by which they are formed, moved, and changed. The field is a major academic discipline, and is also important for mineral and hydrocarbon extraction, knowledge about and mitigation of natural hazards, some Geotechnical engineering fields, and understanding past climates and environments.



## Geological evolution

The geology of an area evolves through time as rock units are deposited and inserted and deformational processes change their shapes and locations.

Rock units are first emplaced either by deposition onto the surface or intrude into the overlying rock. Deposition can occur when sediments settle onto the surface of the Earth and later lithify into sedimentary rock, or when as volcanic material such as volcanic ash or lava flows, blanket the surface. Igneous intrusions such as batholiths, laccoliths, dikes, and sills, push upwards into the overlying rock, and crystallize as they intrude.

After the initial sequence of rocks has been deposited, the rock units can be deformed and/or metamorphosed. Deformation typically occurs as a result of horizontal shortening, horizontal extension, or side-to-side (strike-slip) motion. These structural regimes broadly relate to convergent boundaries, divergent boundaries, and transform boundaries, respectively, between tectonic plates.

## Historical perspective



Plankton inhabit oceans, seas and lakes, and have existed in various forms for at least 2 billion years.<sup>[11]</sup>

Earth is estimated to have formed 4.54 billion years ago from the solar nebula, along with the Sun and other planets.<sup>[12]</sup> The moon formed roughly 20 million years later. Initially molten, the outer layer of the planet cooled, resulting in the solid crust. Outgassing and volcanic activity produced the primordial atmosphere. Condensing water vapor, most or all of which came from ice delivered by comets, produced the oceans and other water sources.<sup>[13]</sup> The highly energetic chemistry is believed to have produced a self-replicating molecule around 4 billion years ago.<sup>[14]</sup>

Continents formed, then broke up and reformed as the surface of Earth reshaped over hundreds of millions of years, occasionally combining to make a supercontinent. Roughly 750 million years ago, the earliest known supercontinent Rodinia, began to break apart. The continents later recombined



to form Pannotia which broke apart about 540 million years ago, then finally Pangaea, which broke apart about 180 million years ago.<sup>[15]</sup>

There is significant evidence that a severe glacial action during the Neoproterozoic era covered much of the planet in a sheet of ice. This hypothesis has been termed the "Snowball Earth", and it is of particular interest as it precedes the Cambrian explosion in which multicellular life forms began to proliferate about 530–540 million years ago.<sup>[16]</sup>

Since the Cambrian explosion there have been five distinctly identifiable mass extinctions.<sup>[17]</sup> The last mass extinction occurred some 65 million years ago, when a meteorite collision probably triggered the extinction of the non-avian dinosaurs and other large reptiles, but spared small animals such as mammals, which then resembled shrews. Over the past 65 million years, mammalian life diversified.<sup>[18]</sup>

Several million years ago, a species of small African ape gained the ability to stand upright.<sup>[19]</sup> The subsequent advent of human life, and the development of agriculture and further civilization allowed humans to affect the Earth more rapidly than any previous life form, affecting both the nature and quantity of other organisms as well as global climate. By comparison, the Great Oxygenation Event, produced by the proliferation of algae during the Siderian period, required about 300 million years to culminate.

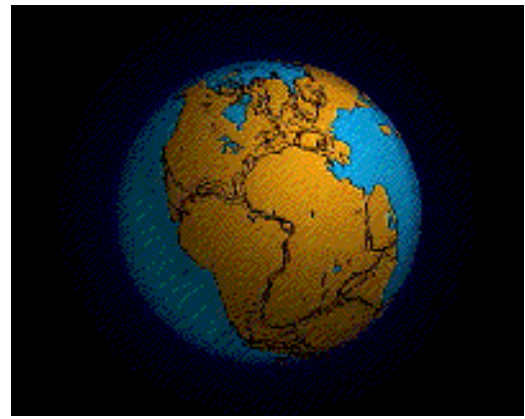
The present era is classified as part of a mass extinction event, the Holocene extinction event, the fastest ever to have occurred.<sup>[20][21]</sup> Some, such as E. O. Wilson of Harvard University, predict that human destruction of the biosphere could cause the extinction of one-half of all species in the next 100 years.<sup>[22]</sup> The extent of the current extinction event is still being researched, debated and calculated by biologists.<sup>[23]</sup>

## Atmosphere, climate, and weather



Lightning

serves to protect life at the surface. The atmosphere also retains heat during the night, thereby reducing the daily temperature extremes.



An animation showing the movement of the continents from the separation of Pangaea until the present day.

The atmosphere of the Earth serves as a key factor in sustaining the planetary ecosystem. The thin layer of gases that envelops the Earth is held in place by the planet's gravity. Dry air consists of 78% nitrogen, 21% oxygen, 1% argon and other inert gases, carbon dioxide, etc.; but air also contains a variable amount of water vapor. The atmospheric pressure declines steadily with altitude, and has a scale height of about 8 kilometres at the Earth's surface: the height at which the atmospheric pressure has declined by a factor of  $e$  (a mathematical constant equal to 2.71...)<sup>[24][25]</sup> The ozone layer of the Earth's atmosphere plays an important role in depleting the amount of ultraviolet (UV) radiation that reaches the surface. As DNA is readily damaged by UV light, this

Terrestrial weather occurs almost exclusively in the lower part of the atmosphere, and serves as a convective system for redistributing heat. Ocean currents are another important factor in determining climate, particularly the major underwater thermohaline circulation which distributes heat energy from the equatorial oceans to the polar regions. These currents help to moderate the differences in temperature between winter and summer in the temperate zones. Also, without the redistributions of heat energy by the ocean currents and atmosphere, the tropics would be much hotter, and the polar regions much colder.

Weather can have both beneficial and harmful effects. Extremes in weather, such as tornadoes or hurricanes and cyclones, can expend large amounts of energy along their paths, and produce devastation. Surface vegetation has evolved a dependence on the seasonal variation of the weather, and sudden changes lasting only a few years can have a dramatic effect, both on the vegetation and on the animals which depend on its growth for their food.

The planetary climate is a measure of the long-term trends in the weather. Various factors are known to influence the climate, including ocean currents, surface albedo, greenhouse gases, variations in the solar luminosity, and changes to the planet's orbit. Based on historical records, the Earth is known to have undergone drastic climate changes in the past, including ice ages.

The climate of a region depends on a number of factors, especially latitude. A latitudinal band of the surface with similar climatic attributes forms a climate region. There are a number of such regions, ranging from the tropical climate at the equator to the polar climate in the northern and southern extremes. Weather is also influenced by the seasons, which result from the Earth's axis being tilted relative to its orbital plane. Thus, at any given time during the summer or winter, one part of the planet is more directly exposed to the rays of the sun. This exposure alternates as the Earth revolves in its orbit. At any given time, regardless of season, the northern and southern hemispheres experience opposite seasons.

Weather is a chaotic system that is readily modified by small changes to the environment, so accurate weather forecasting is currently limited to only a few days. Overall, two things are currently happening worldwide: (1) temperature is increasing on the average; and (2) regional climates have been undergoing noticeable changes.<sup>[26]</sup>



Blue light is scattered more than other wavelengths by the gases in the atmosphere, giving the Earth a blue halo when seen from space



A tornado in central Oklahoma

## Water on Earth

**Water** is a chemical substance that is composed of hydrogen and oxygen and is vital for all known forms of life.<sup>[27]</sup> In typical usage, *water* refers only to its liquid form or state, but the substance also has a solid state, ice, and a gaseous state, water vapor or steam. Water covers 71% of the Earth's surface.<sup>[28]</sup> On Earth, it is found mostly in oceans and other large water

bodies, with 1.6% of water below ground in aquifers and 0.001% in the air as vapor, clouds (formed of solid and liquid water particles suspended in air), and precipitation.<sup>[29]</sup> Oceans hold 97% of surface water, glaciers and polar ice caps 2.4%, and other land surface water such as rivers, lakes and ponds 0.6%. Additionally, a minute amount of the Earth's water is contained within biological bodies and manufactured products.



The Iguazu Falls on the border between Brazil and Argentina

## Oceans



A view of the Atlantic Ocean from Leblon, Rio de Janeiro.

An ocean is a major body of saline water, and a principal component of the hydrosphere. Approximately 71% of the Earth's surface (an area of some 361 million square kilometers) is covered by ocean, a continuous body of water that is customarily divided into several principal oceans and smaller seas. More than half of this area is over 3,000 meters (9,800 ft) deep. Average oceanic salinity is around 35 parts per thousand (ppt) (3.5%), and nearly all seawater has a salinity in the range of 30 to 38 ppt. Though generally recognized as several 'separate' oceans, these waters comprise one global, interconnected body of salt water often referred to as the World Ocean or global ocean.<sup>[30][31]</sup> This concept of a global ocean as a continuous body of water with relatively free interchange among its parts is of fundamental importance to

oceanography.<sup>[32]</sup>

The major oceanic divisions are defined in part by the continents, various archipelagos, and other criteria: these divisions are (in descending order of size) the Pacific Ocean, the Atlantic Ocean, the Indian Ocean, the Southern Ocean and the Arctic Ocean. Smaller regions of the oceans are called seas, gulfs, bays and other names. There are also salt lakes, which are smaller bodies of landlocked saltwater that are not interconnected with the World Ocean. Two notable examples of salt lakes are the Aral Sea and the Great Salt Lake.



## Lakes

A lake (from Latin *lacus*) is a terrain feature (or physical feature), a body of liquid on the surface of a world that is localized to the bottom of basin (another type of landform or terrain feature; that is, it is not global) and moves slowly if it moves at all. On Earth, a body of water is considered a lake when it is inland, not part of the ocean, is larger and deeper than a pond, and is fed by a river.<sup>[33][34]</sup> The only world other than Earth known to harbor lakes is Titan, Saturn's largest moon, which has lakes of ethane, most likely mixed with methane. It is not known if Titan's lakes are fed by rivers, though Titan's surface is carved by numerous river beds. Natural lakes on Earth are generally found in mountainous areas, rift zones, and areas with ongoing or recent glaciation. Other lakes are found in endorheic basins or along the courses of mature rivers. In some parts of the world, there are many lakes because of chaotic drainage patterns left over from the last Ice Age. All lakes are temporary over geologic time scales, as they will slowly fill in with sediments or spill out of the basin containing them.



Lake Mapourika, New Zealand

## Ponds

A **pond** is a body of standing water, either natural or man-made, that is usually smaller than a lake. A wide variety of man-made bodies of water are classified as ponds, including water gardens designed for aesthetic ornamentation, fish ponds designed for commercial fish breeding, and solar ponds designed to store thermal energy. Ponds and lakes are distinguished from streams via current speed. While currents in streams are easily observed, ponds and lakes possess thermally driven microcurrents and moderate wind driven currents. These features distinguish a pond from many other aquatic terrain features, such as stream pools and tide pools.



The Westborough Reservoir (Mill Pond) in Westborough, Massachusetts.

## Rivers



The Nile river in Cairo, Egypt's capital city

A river is a natural watercourse,<sup>[35]</sup> usually freshwater, flowing toward an ocean, a lake, a sea or another river. In a few cases, a river simply flows into the ground or dries up completely before reaching another body of water. Small rivers may also be called by several other names, including stream, creek, brook, rivulet, and rill; there is no general rule that defines what can be called a river. Many names for small rivers are specific to geographic location; one example is *Burn* in Scotland and North-east England. Sometimes a river is said to be larger than a creek,<sup>[36]</sup> but this is not always the case, due to vagueness in the language.<sup>[37]</sup> A river is part of the hydrological cycle. Water within a river is generally collected from precipitation through surface runoff, groundwater recharge, springs, and the release of stored water in natural ice and snowpacks (i.e., from glaciers).



## Streams

A stream is a flowing body of water with a current, confined within a bed and stream banks. In the United States a stream is classified as a watercourse less than 60 feet (18 metres) wide. Streams are important as conduits in the water cycle, instruments in groundwater recharge, and they serve as corridors for fish and wildlife migration. The biological habitat in the immediate vicinity of a stream is called a riparian zone. Given the status of the ongoing Holocene extinction, streams play an important corridor role in connecting fragmented habitats and thus in conserving biodiversity. The study of streams and waterways in general involves many branches of inter-disciplinary natural science and engineering, including hydrology, fluvial geomorphology, aquatic ecology, fish biology, riparian ecology and others.

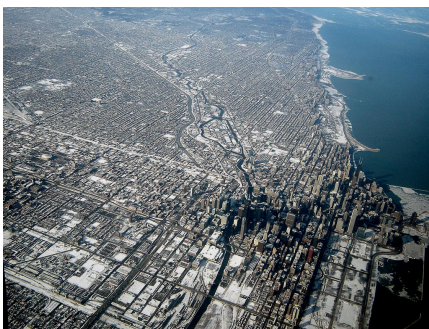


A rocky stream in Hawaii

## Ecosystems



Loch Lomond in Scotland forms a relatively isolated ecosystem. The fish community of this lake has remained unchanged over a very long period of time.<sup>[38]</sup>



An aerial view of a human ecosystem. Pictured is the city of Chicago

Ecosystems are composed of a variety of abiotic and biotic components that function in an interrelated way.<sup>[39]</sup> The structure and composition is determined by various environmental factors that are interrelated. Variations of these factors will initiate dynamic modifications to the ecosystem. Some of the more important components are: soil, atmosphere, radiation from the sun, water, and living organisms.

Central to the ecosystem concept is the idea that living organisms interact with every other element in their local environment. Eugene Odum, a founder of ecology, stated: "Any unit that includes all of the organisms (ie: the "community") in a given area interacting with the physical environment so that a flow of energy leads to clearly defined trophic structure, biotic diversity, and material cycles (i.e.: exchange of materials between living and nonliving parts) within the system is an ecosystem."<sup>[40]</sup> Within the ecosystem, species are connected and dependent upon one another in the food chain, and exchange energy and matter between themselves as well as with their environment.<sup>[41]</sup> The human ecosystem concept is grounded in the deconstruction of the human/nature dichotomy and the premise that all species are ecologically integrated with each other, as well as with the abiotic constituents of their biotope.

A smaller unit of size is called a microecosystem. For example, a microsystem can be a stone and all the life under it. A *macroecosystem* might involve a whole ecoregion, with its

drainage basin.<sup>[42]</sup>

## Wilderness

**Wilderness** is generally defined as areas that have not been significantly modified by human activity. The WILD Foundation <sup>[43]</sup> goes into more detail, defining wilderness as: "The most intact, undisturbed wild natural areas left on our planet – those last truly wild places that humans do not control and have not developed with roads, pipelines or other industrial infrastructure." Wilderness areas can be found in preserves, estates, farms, conservation preserves, ranches, national forests, national parks and even in urban areas along rivers, gulches or otherwise undeveloped areas. Wilderness areas <sup>[43]</sup> and protected parks are considered important for the survival of certain species, ecological studies, conservation, solitude, and recreation.

Some nature writers believe wilderness areas are vital for the human spirit and creativity,<sup>[44]</sup> and some Ecologists consider wilderness areas to be an integral part of the planet's self-sustaining natural ecosystem (the biosphere). They may also preserve historic genetic traits and that they provide habitat for wild flora and fauna that may be difficult to recreate in zoos, arboretums or laboratories.



Old growth European Beech forest in Biogradska Gora National Park, Montenegro.

## Life



Female mallard and ducklings – reproduction is essential for continuing life

Although there is no universal agreement on the definition of life, scientists generally accept that the biological manifestation of life is characterized by organization, metabolism, growth, adaptation, response to stimuli and reproduction.<sup>[45]</sup> Life may also be said to be simply the characteristic state of organisms.

Properties common to terrestrial organisms (plants, animals, fungi, protists, archaea and bacteria) are that they are cellular, carbon-and-water-based with complex organization, having a metabolism, a capacity to grow, respond to stimuli, and reproduce. An entity with these properties is generally considered life. However, not every definition of life considers all of these properties to be essential. Human-made analogs of life may also be considered to be life.

The biosphere is the part of Earth's outer shell – including land, surface rocks, water, air and the atmosphere – within which life occurs, and which biotic processes in turn alter or transform. From the broadest geophysiological point of view, the biosphere is the global ecological system integrating all living beings and their relationships, including their interaction with the elements of the lithosphere (rocks), hydrosphere (water), and atmosphere (air). Currently the entire Earth contains over 75 billion tons (150 *trillion* pounds or about  $6.8 \times 10^{13}$  kilograms) of biomass (life), which lives within various environments within the biosphere.<sup>[46]</sup>

Over nine-tenths of the total biomass on Earth is plant life, on which animal life depends very heavily for its existence.<sup>[47]</sup> More than 2 million species of plant and animal life have been identified to date,<sup>[48]</sup> and estimates of the actual number of existing species range from several million to well over 50 million.<sup>[49][50][51]</sup> The number of individual species of life is constantly in some degree of flux, with new species appearing and others ceasing to exist on a continual basis.<sup>[52][53]</sup> The total number of species is presently in rapid decline.<sup>[54][55][56]</sup>

## Evolution

Life is only known to exist on the planet Earth.(cf Astrobiology) The origin of life is still a poorly understood process, but it is thought to have occurred about 3.9 to 3.5 billion years ago during the hadean or archean eons on a primordial earth that had a substantially different environment than is found at present.<sup>[59]</sup> These life forms possessed the basic traits of self-replication and inheritable traits. Once life had appeared, the process of evolution by natural selection resulted in the development of ever-more diverse life forms.



An area of the Amazon Rainforest in Brazil. The tropical rainforests of South America contain the largest diversity of species on Earth.<sup>[57][58]</sup>

Species that were unable to adapt to the changing environment and competition from other life forms became extinct. However, the fossil record retains evidence of many of these older species. Current fossil and DNA evidence shows that all existing species can trace a continual ancestry back to the first primitive life forms.<sup>[59]</sup>

The advent of photosynthesis in very basic forms of plant life worldwide allowed the sun's energy to be harvested to create conditions allowing for more complex life. The resultant oxygen accumulated in the atmosphere and gave rise to the ozone layer. The incorporation of smaller cells within larger ones resulted in the development of yet more complex cells called eukaryotes.<sup>[60]</sup> Cells within colonies became increasingly specialized, resulting in true multicellular organisms. With the ozone layer absorbing harmful ultraviolet radiation, life colonized the surface of Earth.

## Microbes

The first form of life to develop on the Earth were microbes, and they remained the only form of life on the planet until about a billion years ago when multi-cellular organisms began to appear.<sup>[61]</sup> Microorganisms are single-celled organisms that are generally microscopic, and smaller than the human eye can see. They include Bacteria, Fungi, Archaea and Protista.

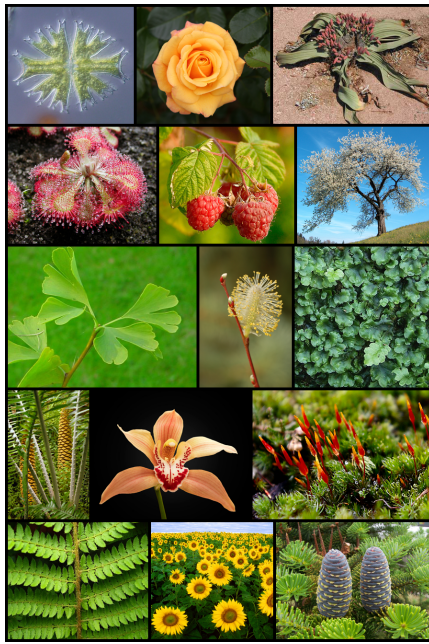
These life forms are found in almost every location on the Earth where there is liquid water, including the interior of rocks within the planet.<sup>[62]</sup> Their reproduction is both rapid and profuse. The combination of a high mutation rate and a horizontal gene transfer<sup>[63]</sup> ability makes them highly adaptable, and able to survive in new environments, including outer space.<sup>[64]</sup> They form an essential part of the planetary ecosystem. However some microorganisms are pathogenic and can post health risk to other organisms.



A microscopic mite *Lorryia formosa*.



## Plants and animals



A selection of diverse plant species



A selection of diverse animal species

Originally Aristotle divided all living things between plants, which generally do not move fast enough for humans to notice, and animals. In Linnaeus' system, these became the kingdoms Vegetabilia (later Plantae) and Animalia. Since then, it has become clear that the Plantae as originally defined included several unrelated groups, and the fungi and several groups of algae were removed to new kingdoms. However, these are still often considered plants in many contexts. Bacterial life is sometimes included in flora,<sup>[65][66]</sup> and some classifications use the term *bacterial flora* separately from *plant flora*.

Among the many ways of classifying plants are by regional floras, which, depending on the purpose of study, can also include *fossil flora*, remnants

of plant life from a previous era. People in many regions and countries take great pride in their individual arrays of characteristic flora, which can vary widely across the globe due to differences in climate and terrain.

Regional floras commonly are divided into categories such as *native flora* and *agricultural and garden flora*, the lastly mentioned of which are intentionally grown and cultivated. Some types of "native flora" actually have been introduced centuries ago by people migrating from one region or continent to another, and become an integral part of the native, or natural flora of the place to which they were introduced. This is an example of how human interaction with nature can blur the boundary of what is considered nature.

Another category of plant has historically been carved out for *weeds*. Though the term has fallen into disfavor among botanists as a formal way to categorize "useless" plants, the informal use of the word "weeds" to describe those plants that are deemed worthy of elimination is illustrative of the general tendency of people and societies to seek to alter or shape the course of nature. Similarly, animals are often categorized in ways such as *domestic*, *farm animals*, *wild animals*, *pests*, etc. according to their relationship to human life.

Animals as a category have several characteristics that generally set them apart from other living things. Animals are eukaryotic and usually multicellular (although see Myxozoa), which separates them from bacteria, archaea and most protists. They are heterotrophic,

generally digesting food in an internal chamber, which separates them from plants and algae. They are also distinguished from plants, algae, and fungi by lacking cell walls.

With a few exceptions, most notably the sponges (Phylum Porifera), animals have bodies differentiated into separate tissues. These include muscles, which are able to contract and control locomotion, and a nervous system, which sends and processes signals. There is also typically an internal digestive chamber. The eukaryotic cells possessed by all animals are surrounded by a characteristic extracellular matrix composed of collagen and elastic glycoproteins.

This may be calcified to form structures like shells, bones, and spicules, a framework upon which cells can move about and be reorganized during development and maturation, and which supports the complex anatomy required for mobility.

## Human interrelationship

Although humans currently comprise only a minuscule proportion of the total living biomass on Earth, the human effect on nature is disproportionately large. Because of the extent of human influence, the boundaries between what humans regard as nature and "made environments" is not clear cut except at the extremes. Even at the extremes, the amount of natural environment that is free of discernible human influence is presently diminishing at an increasingly rapid pace.

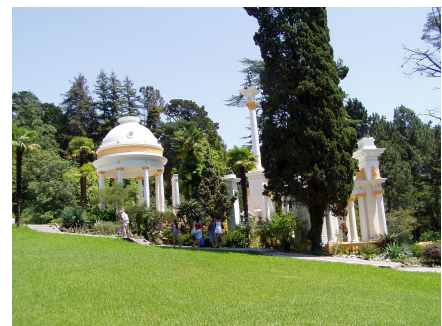
The development of technology by the human race has allowed the greater exploitation of natural resources and has helped to alleviate some of the risk from natural hazards. In spite of this progress, however, the fate of human civilization remains closely linked to changes in the environment. There exists a highly complex feedback loop between the use of advanced technology and changes to the environment that are only slowly becoming understood.<sup>[67]</sup> Man-made threats to the Earth's natural environment include pollution, deforestation, and disasters such as oil spills. Humans have contributed to the extinction of many plants and animals.

Humans employ nature for both leisure and economic activities. The acquisition of natural resources for industrial use remains the primary component of the world's economic system. Some activities, such as hunting and fishing, are used for both sustenance and leisure, often by different people. Agriculture was first adopted around the 9th millennium BCE. Ranging from food production to energy, nature influences economic wealth.

Although early humans gathered uncultivated plant materials for food and employed the medicinal properties of vegetation for healing,<sup>[68]</sup> most modern human use of plants is through agriculture. The clearance of large tracts of land for crop growth has led to a significant reduction in the amount available of forestation and wetlands, resulting in the loss of habitat for many plant and animal species as well as increased erosion.<sup>[69]</sup>



Despite their natural beauty, the secluded valleys along the Na Pali Coast in Hawaii are heavily modified by introduced invasive species such as She-oak.



Sochi dendrarium is an example of confluence of "natural" and a "made" environment

## Aesthetics and beauty



*Pinguicula grandiflora*, commonly known as a Butterwort

Beauty in nature has historically been a prevalent theme in art and books, filling large sections of libraries and bookstores. That nature has been depicted and celebrated by so much art, photography, poetry and other literature shows the strength with which many people associate nature and beauty. Reasons why this association exists, and what the association consists of, is studied by the branch of philosophy called aesthetics. Beyond certain basic characteristics that many philosophers agree about to explain what is seen as beautiful, the opinions are virtually endless.<sup>[70]</sup> Nature and wildness have been important subjects in various eras of world history. An early tradition of landscape art began in China during the Tang Dynasty (618–907). The tradition of representing nature *as it is* became one of the aims of Chinese painting and was a significant influence in Asian art.

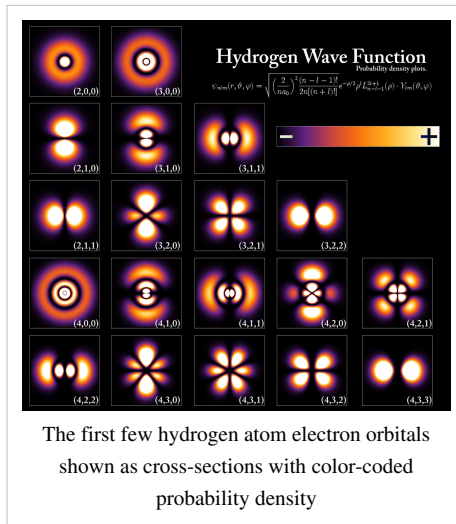
Although natural wonders are celebrated in the Psalms and the Book of Job, wilderness portrayals in art became more prevalent in the 1800s, especially in the works of the Romantic movement. British artists John Constable and J. M. W. Turner turned their attention to capturing the beauty of the natural world in their paintings. Before that, paintings had been primarily of religious scenes or of human beings. William Wordsworth's poetry described the wonder of the natural world, which had formerly been viewed as a threatening place. Increasingly the valuing of nature became an aspect of Western culture.<sup>[71]</sup> This artistic movement also coincided with the Transcendentalist movement in the Western world. A common classical idea of beautiful art involves the word mimesis, the imitation of nature. Also in the realm of ideas about beauty in nature is that the perfect is implied through perfect mathematical forms and more generally by patterns in nature. As David Rothenburg writes, "The beautiful is the root of science and the goal of art, the highest possibility that humanity can ever hope to see".<sup>[72]:281</sup>

## Matter and energy

Some fields of science see nature as matter in motion, obeying certain laws of nature which science seeks to understand. For this reason the most fundamental science is generally understood to be "physics" – the name for which is still recognizable as meaning that it is the study of nature.

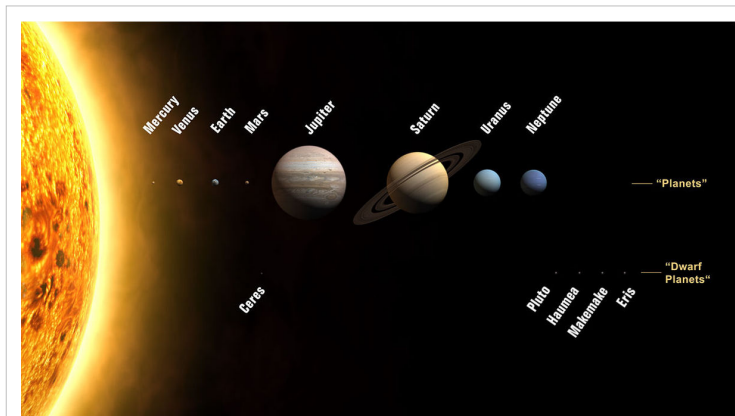
Matter is commonly defined as the substance of which physical objects are composed. It constitutes the observable universe. The visible components of the universe are now believed to compose only 4 percent of the total mass. The remainder is believed to consist of 23 percent cold dark matter and 73 percent dark energy.<sup>[73]</sup> The exact nature of these components is still unknown and is currently under intensive investigation by physicists.

The behavior of matter and energy throughout the observable universe appears to follow well-defined physical laws. These laws have been employed to produce cosmological models that successfully explain the structure and the evolution of the universe we can observe. The mathematical expressions of the laws of physics employ a set of twenty physical constants<sup>[74]</sup> that appear to be static across the observable universe.<sup>[75]</sup> The values of these constants have been carefully measured, but the reason for their specific values remains a mystery.





## Beyond Earth



Planets and dwarf planets of the Solar System (*Sizes to scale, distances not to scale*)



**NGC 4414** is a spiral galaxy in the constellation Coma Berenices about 56,000 light years in diameter and approximately 60 million light years from Earth

Outer space, also simply called *space*, refers to the relatively empty regions of the universe outside the atmospheres of celestial bodies. *Outer* space is used to distinguish it from airspace (and terrestrial locations). There is no discrete boundary between the Earth's atmosphere and space, as the atmosphere gradually attenuates with increasing altitude. Outer space within the Solar System is called interplanetary space, which passes over into interstellar space at what is known as the heliopause.

Outer space is certainly spacious, but it is far from empty. Outer space is sparsely filled with several dozen types of organic molecules discovered to date by microwave spectroscopy, blackbody radiation left over from the big bang and the origin of the universe, and cosmic rays, which include ionized atomic nuclei and various subatomic particles. There is also some gas, plasma and dust, and small meteors. Additionally, there are signs of human life in outer space today, such as material left over from previous manned and unmanned launches which are a potential hazard to spacecraft. Some of this debris re-enters the atmosphere

periodically.

Although the planet Earth is currently the only known body within the solar system to support life, current evidence suggests that in the distant past the planet Mars possessed bodies of liquid water on the surface.<sup>[76]</sup> For a brief period in Mars' history, it may have also been capable of forming life. At present though, most of the water remaining on Mars is frozen. If life exists at all on Mars, it is most likely to be located underground where liquid water can still exist.<sup>[77]</sup>

Conditions on the other terrestrial planets, Mercury and Venus, appear to be too harsh to support life as we know it. But it has been conjectured that Europa, the fourth-largest moon of Jupiter, may possess a sub-surface ocean of liquid water and could potentially host life.<sup>[78]</sup>

Recently, the team of Stéphane Udry have discovered a new planet named Gliese 581 g, which is an extrasolar planet orbiting the red dwarf star Gliese 581. Gliese 581 g appears to lie in the habitable zone of space surrounding the star, and therefore could possibly host life as we know it.

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## External links

- Nature Journal – Weekly journal of science (<http://www.nature.com/nature/index.html>).
  - The WILD Foundation (<http://www.wild.org/>).
  - BBC – Science and Nature (<http://www.bbc.co.uk/sn/>).
  - European Commission – Nature and Biodiversity homepage ([http://ec.europa.eu/environment/nature/index\\_en.htm](http://ec.europa.eu/environment/nature/index_en.htm)).
  - The Nature Conservancy (<http://www.nature.org/>)
  - Science.gov – Environment & Environmental Quality ([http://www.science.gov/browse/w\\_123.htm](http://www.science.gov/browse/w_123.htm)).
  - NatureWatch – Wiki for documenting biodiversity (german) ([http://www.naturewatch.wikia.com/wiki/NatureWatch\\_Wiki](http://www.naturewatch.wikia.com/wiki/NatureWatch_Wiki))
  - Ministry of Environment Republic of Korea – Nature ([http://eng.me.go.kr/docs/sub2/policy\\_view.html?topmenu=C&cat=240&class=13](http://eng.me.go.kr/docs/sub2/policy_view.html?topmenu=C&cat=240&class=13))
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# Etymology

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## Nature (philosophy)

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Nature is a concept with two major sets of inter-related meanings, referring on the one hand to the things which are natural, or subject to the normal working of "laws of nature", or on the other hand to the essential properties and causes of those things to be what they naturally are, or in other words the laws of nature themselves.

How to understand the meaning and significance of nature has been a consistent theme of discussion within the history of Western Civilization, in the philosophical fields of metaphysics and epistemology, as well as in theology and science. The study of natural things and the regular laws which seem to govern them, as opposed to discussion about what it means to be natural, is the area of natural science.

The word "nature" derives from Latin *nātūra*, a philosophical term derived from the verb for birth, which was used as a translation for the earlier Ancient Greek term *phusis* which was derived from the verb for natural growth, for example that of a plant. Already in classical times, philosophical use of these words combined two related meanings which have in common that they refer to the way in which things happen by themselves, "naturally", without "interference" from human deliberation, divine intervention, or anything outside of what is considered normal for the natural things being considered.

Understandings of nature depend on the subject and age of the work where they appear. For example Aristotle's explanation of natural properties differs from what is meant by natural properties in modern philosophical and scientific works, which can also differ from other scientific and conventional usage.

### Classical nature and Aristotelian metaphysics

The *Physics* (from *physis*, Greek for "nature") is Aristotle's principal work on nature. In *Physics* II.1, Aristotle defines a nature as "a source or cause of being moved and of being at rest in that to which it belongs primarily".<sup>[1]</sup> In other words, a nature is the principle within a natural raw material that is the source of tendencies to change or rest in a particular way unless stopped. For example a rock would fall unless stopped. Natural things stand in contrast to artifacts, which are formed by human artifice, not because of an innate tendency. (The raw materials of a bed have no tendency to become a bed.) In terms of Aristotle's theory of four causes, the word natural is applied both to the innate potential of matter cause and the forms which the matter tends to become naturally.<sup>[2]</sup>

According to Leo Strauss,<sup>[3]</sup> the beginning of Western philosophy involved the "discovery or invention of nature" and the "pre-philosophical equivalent of nature" was supplied by "such notions as 'custom' or 'ways'". In ancient Greek philosophy on the other hand, Nature or natures are ways that are "really universal" "in all times and places". What makes nature different is that it presupposes not only that not all customs and ways are equal, but also that one can "find one's bearings in the cosmos" "on the basis of inquiry" (not for example on the basis of traditions or religion). To put this "discovery or invention" into the traditional terminology, what is "by nature" is contrasted to what is "by convention". The concept of nature taken this far remains a strong tradition in modern western thinking. Science, according to Strauss' commentary of Western history is the contemplation of nature, while technology was or is an attempt to imitate it.<sup>[4]</sup>

Going further, the philosophical concept of nature or natures as a special type of causation - for example that the way particular humans are is partly caused by something called "human nature" is an essential step towards Aristotle's teaching concerning causation, which became standard in all Western philosophy until the arrival of modern science.

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Whether it was intended or not, Aristotle's inquiries into this subject were long felt to have resolved the discussion about nature in favor of one solution. In this account, there are four different types of cause:

- The material cause is the "raw material" - the matter which undergoes change. One of the causes of a statue being what it is might be that it is bronze. All meanings of the word nature encompass this simple meaning.
- The efficient cause is the motion of another thing, which makes a thing change, for example a chisel hitting a rock causes a chip to break off. This is the way which the matter is forming into a form so that it become substance like what Aristotle said that a substance must have a form and matter in order to call it substance. This is the motion of changing a single being into two. This is the most obvious way in which cause and effect works, as in the descriptions of modern science. But according to Aristotle, this does not yet explain that of which the motion is, and we must "apply ourselves to the question whether there is any other cause per se besides matter".<sup>[5]</sup>
- The formal cause is the form or idea which serves as a template towards which things develop - for example following an approach based upon Aristotle we could say that a child develops in a way partly determined by a thing called "human nature". Here, nature is a cause.
- The final cause is the aim towards which something is directed. For example a human aims at something perceived to be good, as Aristotle says in the opening lines of the Nicomachean Ethics.

The formal and final cause are an essential part of Aristotle's "Metaphysics" - his attempt to go beyond nature and explain nature itself. In practice they imply a human-like consciousness involved in the causation of all things, even things which are not man-made. Nature itself is attributed with having aims.<sup>[6]</sup>

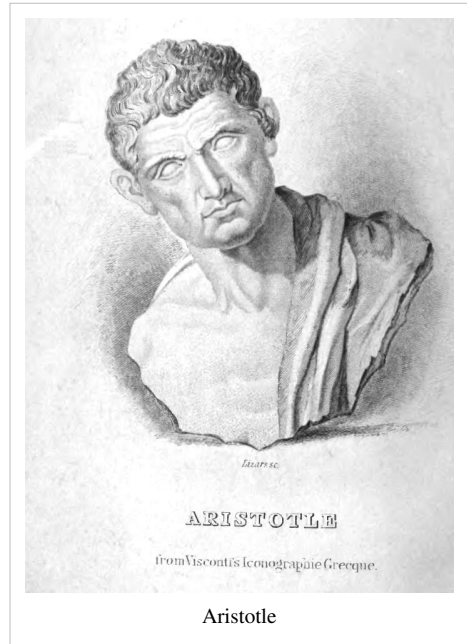
The artificial, like the conventional therefore, is within this branch of Western thought, traditionally contrasted with the natural. Technology was contrasted with science, as mentioned above. And another essential aspect to this understanding of causation was the distinction between the accidental properties of a thing and the substance - another distinction which has lost favor in the modern era, after having long been widely accepted in medieval Europe.

To describe it another way, Aristotle treated organisms and other natural wholes as existing at a higher level than mere matter in motion. Aristotle's argument for formal and final causes is related to a doctrine about how it is possible that people know things: "If nothing exists apart from individual things, nothing will be intelligible; everything will be sensible, and there will be no knowledge of anything—unless it be maintained that sense-perception is knowledge".<sup>[7]</sup> Those philosophers who disagree with this reasoning therefore also see knowledge differently than Aristotle.

Aristotle then, described nature or natures as follows, in a way quite differently to modern science...

"**Nature**" means:

- (a) in one sense, the genesis of growing things — as would be suggested by pronouncing the  $\nu$  of φύσις<sup>[8]</sup> long—and
- (b) in another, that immanent thing from which a growing thing first begins to grow.
- (c) The source from which the primary motion in every natural object is induced in that object as such. All things are said to grow which gain increase through something else by contact and organic unity (or adhesion, as in the case of embryos). Organic unity differs from contact; for in the latter case there need be nothing



except contact, but in both the things which form an organic unity there is some one and the same thing which produces, instead of mere contact, a unity which is organic, continuous and quantitative (but not qualitative). Again, "nature" means

(d) the primary stuff, shapeless and unchangeable from its own potency, of which any natural object consists or from which it is produced; e.g., bronze is called the "nature" of a statue and of bronze articles, and wood that of wooden ones, and similarly in all other cases. For each article consists of these "natures," the primary material persisting. It is in this sense that men call the elements of natural objects the "nature," some calling it fire, others earth or air or water, others something else similar, others some of these, and others all of them. Again in another sense "nature" means

(e) the substance of natural objects; as in the case of those who say that the "nature" is the primary composition of a thing, or as Empedocles says: *Of nothing that exists is there nature, but only mixture and separation of what has been mixed*; nature is but a name given to these by men. Hence as regards those things which exist or are produced by nature, although that from which they naturally are produced or exist is already present, we say that they have not their nature yet unless they have their form and shape. That which comprises both of these exists by nature; e.g. animals and their parts. And **nature is both the primary matter** (and this in two senses: either primary in relation to the thing, or primary in general; e.g., in bronze articles the primary matter in relation to those articles is bronze, but in general it is perhaps water—that is if all things which can be melted are water) **and the form or essence**, i.e. the end of the process, of generation. Indeed from this sense of "nature," by an extension of meaning, every essence in general is called "nature," because the nature of anything is a kind of essence. From what has been said, then, the primary and proper sense of "nature" is the essence of those things which contain in themselves as such a source of motion; for the matter is called "nature" because it is capable of receiving the nature, and the processes of generation and growth are called "nature" because they are motions derived from it. And nature in this sense is the source of motion in natural objects, which is somehow inherent in them, either potentially or actually.

— Metaphysics 1014b-1015a, translated by Hugh Tredennick, emphasis added.<sup>[9]</sup>

It might be argued, as indeed it has been, that this type of theory represented an oversimplifying diversion from the debates within Classical philosophy, possibly even that Aristotle saw it as a simplification or summary of the debates himself. But in any case the theory of the four causes became a standard part of any advanced education in the Middle Ages.

## Modern science and laws of nature: trying to avoid metaphysics

In contrast, Modern Science took its distinctive turn with Francis Bacon, who rejected the four distinct causes, and saw Aristotle as someone who "did proceed in such a spirit of difference and contradiction towards all antiquity: undertaking not only to frame new words of science at pleasure, but to confound and extinguish all ancient wisdom". He felt that lesser known Greek philosophers such as Democritus "who did not suppose a mind or reason in the frame of things", have been arrogantly dismissed because of Aristotelianism leading to a situation in his time wherein "the search of the physical causes hath been neglected, and passed in silence".<sup>[10]</sup>

And so Bacon advised...

Physic doth make inquiry, and take consideration of the same natures : but how? Only as to the material and efficient causes of them, and not as to the forms. For example; if the cause of whiteness in snow or froth be inquired, and it be rendered thus, that the subtile intermixture of air and water is the cause, it is well rendered ; but, nevertheless, is this the form of whiteness? No; but it is the efficient, which is ever but *vehiculum formæ*. This part of metaphysique I do not find laboured and performed...

— Francis Bacon, *Advancement of Learning* II.VII.6<sup>[11]</sup>

In his *Novum Organum* Bacon argued that the only forms or natures we should hypothesize are the "simple" (as opposed to compound) ones such as the ways in which heat, movement, etc. work. For example in aphorism 51 he writes:

51. The human understanding is, by its own nature, prone to abstraction, and supposes that which is fluctuating to be fixed. But it is better to dissect than abstract nature; such was the method employed by the school of Democritus, which made greater progress in penetrating nature than the rest. It is best to consider matter, its conformation, and the changes of that conformation, its own action, and the law of this action or motion, for forms are a mere fiction of the human mind, unless you will call the laws of action by that name.

Following Bacon's advice, the scientific search for the formal cause of things is now replaced by the search for "laws of nature" or "laws of physics" in all scientific thinking. To use Aristotle's well-known terminology these are descriptions of efficient cause, and not formal cause or final cause. It means modern science limits its hypothesizing about non-physical things to the assumption that there are regularities to the ways of all things which do not change.

These general laws, in other words, replace thinking about *specific* "laws", for example "*human nature*". In modern science, human nature is part of the same general scheme of cause and effect, obeying the same general laws, as all other things. The above-mentioned difference between accidental and substantial properties, and indeed knowledge and opinion, also disappear within this new approach that aimed to avoid metaphysics.



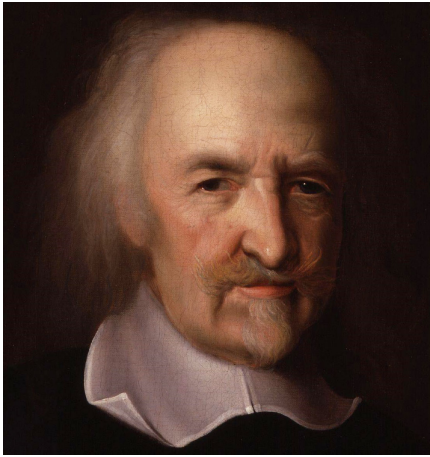
A Renaissance representation of Democritus the laughing philosopher, by Agostino Carracci



Francis Bacon



As Bacon knew, the term "laws of nature" was one taken from medieval Aristotelianism. St Thomas of Aquinas for example, defined law so that nature really was legislated to consciously achieve aims, like human law: "an ordinance of reason for the common good, made by him who has care of the community and promulgated".<sup>[12]</sup> In contrast, roughly contemporary with Bacon, Hugo Grotius described the law of nature as "a rule that [can] be deduced from fixed principles by a sure process of reasoning".<sup>[13]</sup> And later still, Montesquieu was even further from the original legal metaphor, describing laws vaguely as "the necessary relations deriving from the nature of things".<sup>[14]</sup>



Thomas Hobbes

One of the most important implementors of Bacon's proposal was Thomas Hobbes, whose remarks concerning nature are particularly well-known. His most famous work, *Leviathan*, opens with the word "Nature" and then parenthetically defines it as "the art whereby God hath made and governes the world". Despite this pious description, he follows a Baconian approach. Following his contemporary, Descartes, Hobbes describes life itself as mechanical, caused in the same way as clockwork:

For seeing life is but a motion of Limbs, the beginning whereof is in some principall part within; why may we not say, that all Automata (Engines that move themselves by springs and wheelles as doth a watch) have an artificiall life?

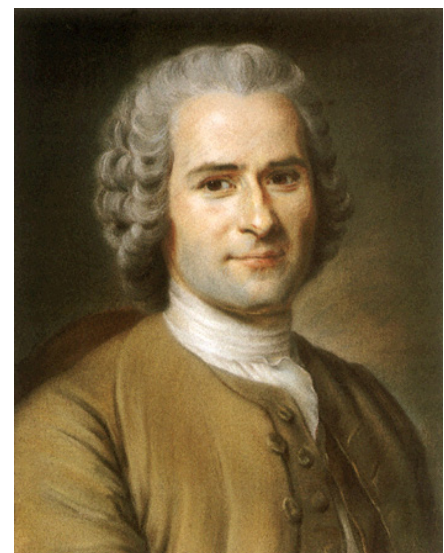
On this basis, already being established in natural science in his lifetime, Hobbes sought to discuss politics and human life in terms of "laws of nature". But in the new modern approach of Bacon and Hobbes, and before them Machiavelli (who however never clothed his criticism of the Aristotelian approach in medieval terms like "laws of nature"),<sup>[15]</sup> such laws of nature are quite different to human laws: they no longer imply any sense of better or worse, but simply how things really are, and, when in reference to laws of *human* nature, what sorts of human behavior can be most relied upon.

## "Late modern" nature

Having disconnected the term "law of nature" from the original medieval metaphor of human-made law, the term "law of nature" is now used less than in early modern times.

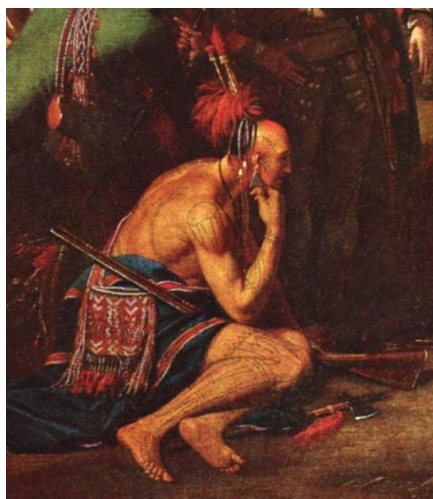
To take the critical example of human nature, as discussed in ethics and politics, once early modern philosophers such as Hobbes had described human nature as whatever you could expect from a mechanism called a human, the point of speaking of human nature became problematic in some contexts.

In the late 18th century, Rousseau took a critical step in his *Second Discourse*, reasoning that human nature as we know it, rational, and with language, and so on, is a result of historical accidents, and the specific up-bringing of an individual. The consequences of this line of reasoning were to be enormous. It was all about the question of nature. In effect it was being claimed that human nature, one of the most important types of nature in Aristotelian thinking, did not exist as it had been understood to exist.



Jean-Jacques Rousseau: a civilized man, but a person who questioned whether civilization was according to human nature.





Benjamin West's "The Death of General Wolfe". The section shows the Native American. West's portrayal of the Native American has been cited as an example of the "noble savage", a concept associated with Rousseau's Second Discourse. The Discourse itself used stories of great apes to help explain what man would be like in the state of nature.

## The survival of metaphysics

The approach of modern science, like the approach of Aristotelianism, is apparently not universally accepted by all people who accept the concept of nature as a reality which we can pursue with reason.

Bacon and other opponents of Metaphysics claim that all attempts to go beyond nature are bound to fall into the same errors, but Metaphysicians themselves see differences between different approaches.

Immanuel Kant for example, expressed the need for a Metaphysics in quite similar terms to Aristotle.

...though we cannot know these objects as things in themselves, we must yet be in a position at least to think them as things in themselves; otherwise we should be landed in the absurd conclusion that there can be appearance without anything that appears.

— *Critique of Pure Reason* pp. Bxxvi-xxvii

As in Aristotelianism then, Kantianism claims that the human mind must itself have characteristics which are beyond nature, metaphysical, in some way. Specifically Kant argued that the human mind comes ready-made with *a priori* programming, so to speak, which allows it to make sense of nature.

## The study of nature without metaphysics

Authors from Nietzsche to Richard Rorty have claimed that science, the study of nature, can and should exist without metaphysics. But this claim has always been controversial. Authors like Bacon and Hume never denied that their use of the word "nature" implied a metaphysics, but tried to follow Machiavelli's approach of talking about what works, instead of claiming to understand what seems impossible to understand.

## Eastern civilization and the philosophical question of nature

The discussion so far above focuses upon the Western philosophical tradition, where the word "nature" has a very specific history. But despite claims mentioned above to the contrary, it is not universally accepted that Greek philosophy was the one occasion upon which the concept of nature was discovered and emphasized in this way.

In Chinese, the term "nature" may be rendered as either *ziran* (自然), or *xing* (性). The same terms appear in the philosophical literature of nations that adopted the Chinese writing such as Japan and Korea. In the early Chinese literature, nature appears in what might be called, a "pre-Socratic" sense akin to *Dao* (道), or "the Way", in antiquity, similar to *fa* (法) or "Law". Indeed, in ancient Daoism, *the Way* is above all, *the way of nature* (自然之道 *ziran zhi dao*). The term "Dao" is sometimes compared to the enigmatic way Heraclitus used "Logos". In older extant Chinese texts (e.g. 黃帝四經 *Huangdi Sijing*, or *Scripture of the Yellow Emperor*), Dao (as the Dao of nature) has at once a metaphysical and legal character, strongly suggesting that the source of legislation is to be found in the nature of things. While at first, the nature of things was intended as an impulse (志 *zhi* or 心 *xin*), in later Confucianism the distinction would be stressed between mind and will, or between life and the "principle" or "mind" of life (性 *xing*). In Mencius, for instance, life and its principle are juxtaposed in a way that later scholars establish to be *mind*, as a principle, independent of human will (thus, for example, the *mind of nature*). Confucius articulates, a question of natural principle, or the standard of interpretation of *names*. When Confucius seeks beyond the plane of convention

or custom—when he reaches out to the roots of names—he does not find the will of gods and spirits. What he did find remains the subject of interpretation for the scholarship of thousands of years. That subject is usually called nature or the mind thereof.

The philosophical tradition of Legalism, generally may be understood as a quest for the *mind of nature*, and as a struggle to preserve that quest against "heretical" (邪道 *xiadao*) tendencies to seek nature (or the mind thereof) outside the law. Accordingly, throughout ancient China, scholarship of the period generally remained tied to political problems, or problems of legal interpretation. Metaphysical problems were understood as eminently legal problems (and vice versa), so that the interpretation or study (學 *xue*) of Justice or Right (義 *yi*) emerged as the philosophical activity par excellence: to ask "what is Justice?", a favorite question of Confucius, is both to probe the essential (interior) nature of names, or "to know speech" (知言 *zhiyan*), in principle, by virtue of the constituent names.

The rise of Buddhism in ancient China, stimulated the debate on nature once again. Now nature was unequivocally regarded as the *mind of all things*, or as "Buddha-nature" (佛性 *foxing*). This was also the mind of "the Empire" (天下 *tianxia*), or the monarchic principle common to all nations, (hence an identification—notably in Japan—of Buddha as the essence of the Emperor). nature, regarded as utterly beyond both imagination and speech, was that which "cannot be imagined or deliberated", (不可思議 *bukeshiyi*) In the act of being revealed universally, that which is neither externally, nor internally, accessible to the delusional sensory capabilities of one's misappropriating *ego* disappears. Facing of the threat of a solar eclipse, or the depths of the problem of nature, and the consequent decay of civil life into "chaos" or (亂 *luan*), the Chan or Zen (禪) revival of "the Buddha Way" (佛道 *fodao*) emerged. This emphasized the original coincidence of the *Buddha mind* (the metaphysical) and the *everyday mind* (the political). The name Buddha, refers neither to something outside the ego (我 *wo*), nor to ego as a self-appropriating poetic faculty. From this approach, the Buddha was understood as "original" nature, or original mind, yet "very ordinary" because Buddha is not the constitutive principle of an order beyond the civil order, or public morality. Of "this very order", not the ego's deluded physical motion or nominal forms, the ordering principle of both speech and sensory experience "gathers" common experience under a global scope with universal names declared as "direct pointers", for example, (直指 *zhi*) to "the moon" (月 *yue*) or "the original mind". The foremost task of a *student of the way* was thus to recover the constitutive principle of the common experience, their original mind. When understood in this way, "original mind" was thought to conform with normative public morality. Ultimately, Chan was no less a return to "piety" (孝 *xiao*), than it was a return to nature as the common principle of the constitution of civil life.

## References

- [1] Aristotle *Physics* 192b21
- [2] Aristotle *Physics* 193b21
- [3] "Progress or Return" in *An Introduction to Political Philosophy: Ten Essays* by Leo Strauss. (Expanded version of *Political Philosophy: Six Essays* by Leo Strauss, 1975.) Ed. Hilail Gilden. Detroit: Wayne State UP, 1989.
- [4] Strauss and Cropsey eds. *History of Political Philosophy*, Third edition, p.209.
- [5] *Metaphysics* 995b, translated by Hugh Tredennick. Greek: μάλιστα δὲ ζητητέον καὶ πραγματευτέον πότερον ἔστι τι παρὰ τὴν ὕλην αἴτιον καθ' αὐτὸ ἢ οὐ
- [6] As for example Aristotle *Politics* 1252b.1: "Thus the female and the slave are by nature distinct (for nature makes nothing as the cutlers make the Delphic knife, in a niggardly way, but one thing for one purpose; for so each tool will be turned out in the finest perfection, if it serves not many uses but one"
- [7] *Metaphysics* 999b, translated by Hugh Tredennick. Greek: εἰ μὲν οὖν μηδὲν ἔστι παρὰ τὰ καθ' ἑκάστα, οὐθὲν ἂν εἴη νοητὸν ἀλλὰ πάντα αἰσθητὰ καὶ ἐπιστήμη οὐδενός, εἰ μὴ τις εἶναι λέγει τὴν αἰσθησιν ἐπιστήμην.
- [8] *Phusis* is the Greek word for Nature, and Aristotle is drawing attention to the similarity it has to the verb used to describe natural growth in a plant, *phusei*. Indeed the first use of the word involves a plant: ὧς ἄρα φωνήσας πόρε φάρμακον ἀργεῖφόντης ἐκ γαίης ἐρύσας, καὶ μοι φύσιν αὐτοῦ ἔδειξε. "So saying, Argeiphontes [=Hermes] gave me the herb, drawing it from the ground, and showed me its nature." *Odyssey* 10.302-3 (ed. A.T. Murray).
- [9] Greek, with emphasis added as a guide: φύσις λέγεται ἕνα μὲν τρόπον ἢ τῶν φyuμένων γένεσις, οἷον εἴ τις ἐπεκτείνας λέγοι τὸ υ, ἕνα δὲ ἔξ οὗ φύεται πρῶτον τὸ φυόμενον ἐνυπάρχοντος: ἔτι ὅθεν ἢ κίνησις ἢ πρώτη ἐν ἐκάστῳ τῶν φύσει ὄντων ἐν αὐτῷ ἢ αὐτὸ [20]

ὑπάρχει: **φύεσθαι** δὲ λέγεται ὅσα αὐξήσιν ἔχει δι' ἐτέρου τῷ ἄπτεσθαι καὶ συμπεφυκέναι ἢ προσπεφυκέναι ὥσπερ τὰ ἔμβρυα: διαφέρει δὲ σύμφυσις ἀφῆς, ἔνθα μὲν γὰρ οὐδὲν παρὰ τὴν ἀφήν ἕτερον ἀνάγκη εἶναι, ἐν δὲ τοῖς συμπεφυκόσιν ἔστι τι ἐν τῷ αὐτῷ ἐν ἀμφοῖν ὃ ποιεῖ ἀντὶ τοῦ [25] ἄπτεσθαι συμπεφυκέναι καὶ εἶναι ἐν κατὰ τὸ συνεχὲς καὶ ποσόν, ἀλλὰ μὴ κατὰ τὸ ποιόν. ἔτι δὲ φύσις λέγεται ἐξ οὗ πρῶτου ἢ ἔστιν ἢ γίγνεται τὴν φύσει ὄντων, ἀρρυθμίστου ὄντος καὶ ἀμεταβλήτου ἐκ τῆς δυνάμεως τῆς αὐτοῦ, οἷον ἀνδριάντος καὶ τῶν σκευῶν τῶν χαλκῶν ὁ χαλκός ἢ [30] φύσις λέγεται, τῶν δὲ ξυλίνων ξύλον: ὁμοίως δὲ καὶ ἐπὶ τῶν ἄλλων: ἐκ τούτων γὰρ ἔστιν ἕκαστον διασωζομένης τῆς πρώτης ὕλης: τοῦτον γὰρ τὸν τρόπον καὶ τῶν φύσει ὄντων τὰ στοιχεῖα φασι εἶναι φύσιν, οἱ μὲν πῦρ οἱ δὲ γῆν οἱ δ' ἀέρα οἱ δ' ὕδωρ οἱ δ' ἄλλο τι τοιοῦτον λέγοντες, οἱ δ' [35] ἔνια τούτων οἱ δὲ πάντα ταῦτα. ἔτι δ' ἄλλον τρόπον λέγεται ἢ φύσις ἢ τῶν φύσει ὄντων οὐσία, οἷον οἱ λέγοντες τὴν φύσιν εἶναι τὴν πρώτην σύνθεσιν, ἢ ὥσπερ Ἐμπεδοκλῆς λέγει ὅτι "φύσις οὐδενός ἔστιν ἐόντων, ἀλλὰ μόνον μῖξις τε διάλλαξις τε μιγέντων ἔστι, φύσις δ' ἐπὶ τοῖς ὀνομάζεται ἀνθρώποισιν. "Empedocles Fr. 8 διὸ καὶ ὅσα φύσει ἔστιν ἢ γίγνεται, ἥδη ὑπάρχοντος ἐξ οὗ πέφυκε γίγνεσθαι ἢ εἶναι, οὕτω φαμέν [5] τὴν φύσιν ἔχειν ἐὰν μὴ ἔχῃ τὸ εἶδος καὶ τὴν μορφήν. φύσει μὲν οὖν τὸ ἐξ ἀμφοτέρων τούτων ἐστίν, οἷον τὰ ζῶα καὶ τὰ μόρια αὐτῶν: **φύσις δὲ ἢ τε πρώτη ὕλη** (καὶ αὕτη διχῶς, ἢ ἢ πρὸς αὐτὸ πρώτη ἢ ἢ ὅλως πρώτη, οἷον τῶν χαλκῶν ἔργων πρὸς αὐτὰ μὲν πρῶτος ὁ χαλκός, ὅλως δ' [10] ἴσως ὕδωρ, εἰ πάντα τὰ τηκτὰ ὕδωρ) **καὶ τὸ εἶδος καὶ ἡ οὐσία**: τοῦτο δ' ἐστὶ τὸ τέλος τῆς γενέσεως. μεταφορᾷ δ' ἥδη καὶ ὅλως πᾶσα οὐσία φύσις λέγεται διὰ ταύτην, ὅτι καὶ ἡ φύσις οὐσία τίς ἐστιν. ἐκ δὲ τῶν εἰρημένων ἢ πρώτη φύσις καὶ κυρίως λεγομένη ἐστὶν ἡ οὐσία ἢ τῶν ἐχόντων [15] ἀρχὴν κινήσεως ἐν αὐτοῖς ἢ αὐτὰ: ἢ γὰρ ὕλη τῷ ταύτης δεκτικῇ εἶναι λέγεται φύσις, καὶ αἱ γενέσεις καὶ τὸ φύεσθαι τῷ ἀπὸ ταύτης εἶναι κινήσεις. καὶ ἢ ἀρχὴ τῆς κινήσεως τῶν φύσει ὄντων αὕτη ἐστίν, ἐνυπάρχουσα πῶς ἢ δυνάμει ἢ ἐντελεχείᾳ.

[10] Bacon *Advancement of Learning* II.VII.7 (<http://www.archive.org/stream/advancementlea00bacouoft#page/90/mode/2up/search/democritus>)

[11] <http://www.archive.org/stream/advancementlea00bacouoft#page/88/mode/2up/search/whiteness>

[12] Summa Theologiae I-II Q90, A4

[13] *On the Law of War and Peace*, Proleg. 40

[14] The Spirit of the Laws, opening lines

[15] The Prince 15:- "...since my intent is to write something useful to whoever understands it, it has appeared to me more fitting to go directly to the effectual truth of the thing than to the imagination of it. And many have imagined republics and principalities that have never been seen or known to exist in truth; for it is so far from how one lives to how one should live that he who lets go of what is done for what should be done learns his ruin rather than his preservation. For a man who wants to make a profession of good in all regards must come to ruin among so many who are not good. Hence it is necessary to a prince, if he wants to maintain himself, to learn to be able not to be good, and to use this and not use it according to necessity."

# Earth

## Earth

Earth ⊕



"The Blue Marble" photograph of Earth, taken from *Apollo 17*

Designations	
Alternative names	Terra, Gaia
Orbital characteristics	
Epoch J2000.0 <sup>[1]</sup>	
Aphelion	152,098,232 km 1.01671388 AU <sup>[2]</sup>
Perihelion	147,098,290 km 0.98329134 AU <sup>[2]</sup>
Semi-major axis	149,598,261 km 1.00000261 AU <sup>[3]</sup>
Eccentricity	0.01671123 <sup>[3]</sup>
Orbital period	365.256363004 days <sup>[4]</sup> 1.000017421 yr
Average orbital speed	29.78 km/s <sup>[5]</sup> 107,200 km/h
Mean anomaly	357.51716° <sup>[5]</sup>
Inclination	7.155° to Sun's equator 1.57869° <sup>[6]</sup> to invariable plane
Longitude of ascending node	348.73936° <sup>[5]</sup> <sup>[7]</sup>
Argument of perihelion	114.20783° <sup>[5]</sup> <sup>[8]</sup>

Satellites	1 natural (the Moon) 8,300+ artificial (as of 1 March 2001) <sup>[9]</sup>											
Physical characteristics												
Mean radius	6,371.0 km <sup>[10]</sup>											
Equatorial radius	6,378.1 km <sup>[11][12]</sup>											
Polar radius	6,356.8 km <sup>[13]</sup>											
Flattening	0.0033528 <sup>[14]</sup>											
Circumference	40,075.017 km (equatorial) <sup>[12]</sup> 40,007.86 km (meridional) <sup>[15]</sup>											
Surface area	510,072,000 km <sup>2</sup> <sup>[16][17][18]</sup> 148,940,000 km <sup>2</sup> land (29.2 %) 361,132,000 km <sup>2</sup> water (70.8 %)											
Volume	1.08321×10 <sup>12</sup> km <sup>3</sup> <sup>[5]</sup>											
Mass	5.9736×10 <sup>24</sup> kg <sup>[5]</sup>											
Mean density	5.515 g/cm <sup>3</sup> <sup>[5]</sup>											
Equatorial surface gravity	9.780327 m/s <sup>2</sup> <sup>[19]</sup> 0.99732 g											
Escape velocity	11.186 km/s <sup>[5]</sup>											
Sidereal rotation period	0.99726968 d <sup>[20]</sup> 23 <sup>h</sup> 56 <sup>m</sup> 4.100 <sup>s</sup>											
Equatorial rotation velocity	1674.4 km/h ( <b>unknown operator: u'strong'</b> m/s) <sup>[21]</sup>											
Axial tilt	23°26′21″.4119 <sup>[4]</sup>											
Albedo	0.367 (geometric) <sup>[5]</sup> 0.306 (Bond) <sup>[5]</sup>											
Surface temp. Kelvin Celsius	<table><tr><th>min</th><th>mean</th><th>max</th></tr><tr><td>184 K<sup>[22]</sup></td><td>287.2 K<sup>[23]</sup></td><td>331 K<sup>[24]</sup></td></tr><tr><td>−89.2 °C</td><td>14 °C</td><td>57.8 °C</td></tr></table>			min	mean	max	184 K <sup>[22]</sup>	287.2 K <sup>[23]</sup>	331 K <sup>[24]</sup>	−89.2 °C	14 °C	57.8 °C
min	mean	max										
184 K <sup>[22]</sup>	287.2 K <sup>[23]</sup>	331 K <sup>[24]</sup>										
−89.2 °C	14 °C	57.8 °C										
Atmosphere												
Surface pressure	101.325 kPa (MSL)											
Composition	78.08% nitrogen (N <sub>2</sub> ) <sup>[5]</sup> (dry air) 20.95% oxygen (O <sub>2</sub> ) 0.93% argon 0.038% carbon dioxide About 1% water vapor (varies with climate)											

**Earth** (or **the Earth**) is the third planet from the Sun, and the densest and fifth-largest of the eight planets in the Solar System. It is also the largest of the Solar System's four terrestrial planets. It is sometimes referred to as the world, the Blue Planet,<sup>[25]</sup> or by its Latin name, *Terra*.<sup>[26]</sup>

Earth formed 4.54 billion years ago, and life appeared on its surface within one billion years.<sup>[27]</sup> The planet is home to millions of species, including humans.<sup>[28]</sup> Earth's biosphere has significantly altered the atmosphere and other abiotic conditions on the planet, enabling the proliferation of aerobic organisms as well as the formation of the ozone layer which, together with Earth's magnetic field, blocks harmful solar radiation, permitting life on land.<sup>[29]</sup> The physical properties of the Earth, as well as its geological history and orbit, have allowed life to persist during this period. Estimates on how much longer the planet will be able to continue to support life range from a mere 500 million years, to as long as 2.3 billion years.<sup>[30][31][32]</sup>

Earth's outer surface is divided into several rigid segments, or tectonic plates, that migrate across the surface over periods of many millions of years. About 71% of the surface is covered by salt water oceans, with the remainder consisting of continents and islands which together have many lakes and other sources of water that contribute to the hydrosphere. Earth's poles are mostly covered with solid ice (Antarctic ice sheet) or sea ice (Arctic ice cap). The planet's interior remains active, with a thick layer of relatively solid mantle, a liquid outer core that generates a magnetic field, and a solid iron inner core.

Earth interacts with other objects in space, especially the Sun and the Moon. At present, Earth orbits the Sun once every 366.26 times it rotates about its own axis, which is equal to 365.26 solar days, or one sidereal year.<sup>[33]</sup> The Earth's axis of rotation is tilted 23.4° away from the perpendicular of its orbital plane, producing seasonal variations on the planet's surface with a period of one tropical year (365.24 solar days).<sup>[34]</sup> Earth's only known natural satellite, the Moon, which began orbiting it about 4.53 billion years ago, provides ocean tides, stabilizes the axial tilt, and gradually slows the planet's rotation. Between approximately 3.8 billion and 4.1 billion years ago, numerous asteroid impacts during the Late Heavy Bombardment caused significant changes to the greater surface environment.

Both the mineral resources of the planet and the products of the biosphere contribute resources that are used to support a global human population.<sup>[35]</sup> These inhabitants are grouped into about 200 independent sovereign states, which interact through diplomacy, travel, trade, and military action. Human cultures have developed many views of the planet, including personification as a deity, a belief in a flat Earth or in the Earth as the center of the universe, and a modern perspective of the world as an integrated environment that requires stewardship.

## Chronology

The earliest dated Solar System material was formed  $4.5672 \pm 0.0006$  billion years ago,<sup>[36]</sup> and by 4.54 billion years ago (within an uncertainty of 1%)<sup>[27]</sup> the Earth and the other planets in the Solar System had formed out of the solar nebula—a disk-shaped mass of dust and gas left over from the formation of the Sun. This assembly of the Earth through accretion was thus largely completed within 10–20 million years.<sup>[37]</sup> Initially molten, the outer layer of the planet Earth cooled to form a solid crust when water began accumulating in the atmosphere. The Moon formed shortly thereafter, 4.53 billion years ago.<sup>[38]</sup>

The current consensus model<sup>[39]</sup> for the formation of the Moon is the giant impact hypothesis, in which the Moon was created when a Mars-sized object (sometimes called Theia) with about 10% of the Earth's mass<sup>[40]</sup> impacted the Earth in a glancing blow.<sup>[41]</sup> In this model, some of this object's mass would have merged with the Earth and a portion would have been ejected into space, but enough material would have been sent into orbit to coalesce into the Moon.

Outgassing and volcanic activity produced the primordial atmosphere of the Earth. Condensing water vapor, augmented by ice and liquid water delivered by asteroids and the larger proto-planets, comets, and trans-Neptunian objects produced the oceans.<sup>[42]</sup> The newly formed Sun was only 70% of its present luminosity, yet evidence shows that the early oceans remained liquid—a contradiction dubbed the faint young Sun paradox. A combination of greenhouse gases and higher levels of solar activity served to raise the Earth's surface temperature, preventing the oceans from freezing over.<sup>[43]</sup> By 3.5 billion years ago, the Earth's magnetic field was established, which helped prevent the atmosphere from being stripped away by the solar wind.<sup>[44]</sup>

Two major models have been proposed for the rate of continental growth:<sup>[45]</sup> steady growth to the present-day<sup>[46]</sup> and rapid growth early in Earth history.<sup>[47]</sup> Current research shows that the second option is most likely, with rapid initial growth of continental crust<sup>[48]</sup> followed by a long-term steady continental area.<sup>[49][50][51]</sup> On time scales lasting hundreds of millions of years, the surface continually reshaped as continents formed and broke up. The continents migrated across the surface, occasionally combining to form a supercontinent. Roughly 750 million years ago (Ma), one of the earliest known supercontinents, Rodinia, began to break apart. The continents later recombined to form Pannotia, 600–540 Ma, then finally Pangaea, which broke apart 180 Ma.<sup>[52]</sup>

## Evolution of life

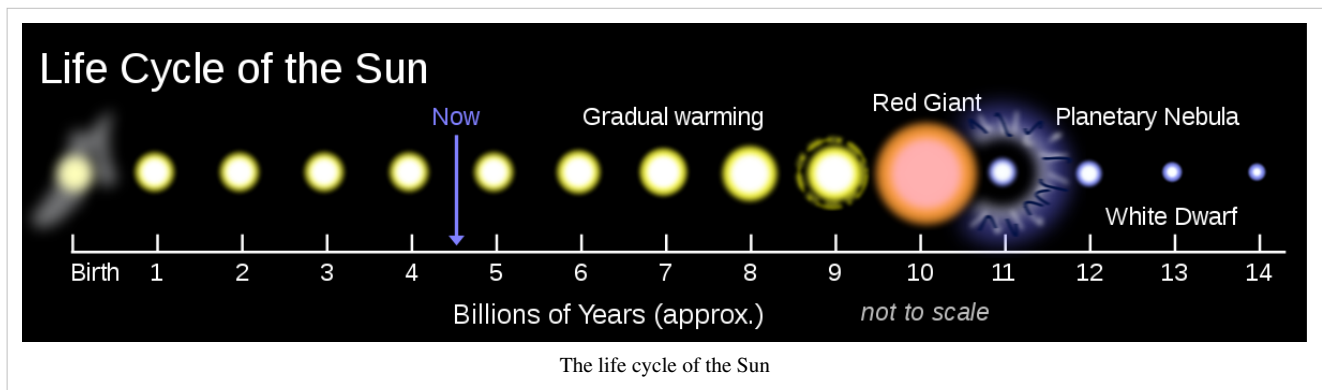
Highly energetic chemistry is believed to have produced a self-replicating molecule around 4 billion years ago and half a billion years later the last common ancestor of all life existed.<sup>[53]</sup> The development of photosynthesis allowed the Sun's energy to be harvested directly by life forms; the resultant oxygen accumulated in the atmosphere and formed a layer of ozone (a form of molecular oxygen [O<sub>3</sub>]) in the upper atmosphere. The incorporation of smaller cells within larger ones resulted in the development of complex cells called eukaryotes.<sup>[54]</sup> True multicellular organisms formed as cells within colonies became increasingly specialized. Aided by the absorption of harmful ultraviolet radiation by the ozone layer, life colonized the surface of Earth.<sup>[55]</sup>

Since the 1960s, it has been hypothesized that severe glacial action between 750 and 580 Ma, during the Neoproterozoic, covered much of the planet in a sheet of ice. This hypothesis has been termed "Snowball Earth", and is of particular interest because it preceded the Cambrian explosion, when multicellular life forms began to proliferate.<sup>[56]</sup>

Following the Cambrian explosion, about 535 Ma, there have been five major mass extinctions.<sup>[57]</sup> The most recent such event was 65 Ma, when an asteroid impact triggered the extinction of the (non-avian) dinosaurs and other large reptiles, but spared some small animals such as mammals, which then resembled shrews. Over the past 65 million years, mammalian life has diversified, and several million years ago an African ape-like animal such as *Ororin tugenensis* gained the ability to stand upright.<sup>[58]</sup> This enabled tool use and encouraged communication that provided the nutrition and stimulation needed for a larger brain, which allowed the evolution of the human race. The development of agriculture, and then civilization, allowed humans to influence the Earth in a short time span as no other life form had,<sup>[59]</sup> affecting both the nature and quantity of other life forms.

The present pattern of ice ages began about 40 Ma and then intensified during the Pleistocene about 3 Ma. High-latitude regions have since undergone repeated cycles of glaciation and thaw, repeating every 40–100,000 years. The last continental glaciation ended 10,000 years ago.<sup>[60]</sup>

## Future



The future of the planet is closely tied to that of the Sun. As a result of the steady accumulation of helium at the Sun's core, the star's total luminosity will slowly increase. The luminosity of the Sun will grow by 10% over the next

1.1 Gyr (1.1 billion years) and by 40% over the next 3.5 Gyr.<sup>[61]</sup> Climate models indicate that the rise in radiation reaching the Earth is likely to have dire consequences, including the loss of the planet's oceans.<sup>[62]</sup>

The Earth's increasing surface temperature will accelerate the inorganic CO<sub>2</sub> cycle, reducing its concentration to levels lethally low for plants (10 ppm for C4 photosynthesis) in approximately 500 million<sup>[30]</sup> to 900 million years. The lack of vegetation will result in the loss of oxygen in the atmosphere, so animal life will become extinct within several million more years.<sup>[63]</sup> After another billion years all surface water will have disappeared<sup>[31]</sup> and the mean global temperature will reach 70 °C<sup>[63]</sup> (158 °F). The Earth is expected to be effectively habitable for about another 500 million years from that point,<sup>[30]</sup> although this may be extended up to 2.3 billion years if the nitrogen is removed from the atmosphere.<sup>[32]</sup> Even if the Sun were eternal and stable, the continued internal cooling of the Earth would result in a loss of much of its CO<sub>2</sub> due to reduced volcanism,<sup>[64]</sup> and 35% of the water in the oceans would descend to the mantle due to reduced steam venting from mid-ocean ridges.<sup>[65]</sup>

The Sun, as part of its evolution, will become a red giant in about 5 Gyr. Models predict that the Sun will expand out to about 250 times its present radius, roughly 1 AU (**unknown operator: u'strong'** km).<sup>[61][66]</sup> Earth's fate is less clear. As a red giant, the Sun will lose roughly 30% of its mass, so, without tidal effects, the Earth will move to an orbit 1.7 AU (**unknown operator: u'strong'** km) from the Sun when the star reaches its maximum radius. The planet was therefore initially expected to escape envelopment by the expanded Sun's sparse outer atmosphere, though most, if not all, remaining life would have been destroyed by the Sun's increased luminosity (peaking at about 5000 times its present level).<sup>[61]</sup> A 2008 simulation indicates that Earth's orbit will decay due to tidal effects and drag, causing it to enter the red giant Sun's atmosphere and be vaporized.<sup>[66]</sup>

## Composition and structure

Further information: Earth physical characteristics tables

Earth is a terrestrial planet, meaning that it is a rocky body, rather than a gas giant like Jupiter. It is the largest of the four solar terrestrial planets in size and mass. Of these four planets, Earth also has the highest density, the highest surface gravity, the strongest magnetic field, and fastest rotation,<sup>[67]</sup> and is probably the only one with active plate tectonics.<sup>[68]</sup>



Size comparison of inner planets (left to right):  
Mercury, Venus, Earth and Mars

## Shape

The shape of the Earth approximates an oblate spheroid, a sphere flattened along the axis from pole to pole such that there is a bulge around the equator.<sup>[70]</sup> This bulge results from the rotation of the Earth, and causes the diameter at the equator to be 43 km larger than the pole-to-pole diameter.<sup>[71]</sup> For this reason the furthest point on the surface from the Earth's center of mass is the Chimborazo volcano in Ecuador.<sup>[72]</sup> The average diameter of the reference spheroid is about 12,742 km, which is approximately 40,000 km/π, as the meter was originally defined as 1/10,000,000 of the distance from the equator to the North Pole through Paris, France.<sup>[73]</sup>



Chimborazo, Ecuador. The furthestmost point on  
the Earth's surface from its center.<sup>[69]</sup>

Local topography deviates from this idealized spheroid, although on a global scale, these deviations are small: Earth has a tolerance of about one part in about 584, or 0.17%, from the reference spheroid, which is less than the 0.22%



tolerance allowed in billiard balls.<sup>[74]</sup> The largest local deviations in the rocky surface of the Earth are Mount Everest (8848 m above local sea level) and the Mariana Trench (10,911 m below local sea level). Because of the equatorial bulge, the surface locations farthest from the center of the Earth are the summits of Mount Chimborazo in Ecuador and Huascarán in Peru.<sup>[75][76][77]</sup>

### Chemical composition of the crust<sup>[78]</sup>

Compound	Formula	Composition	
		Continental	Oceanic
silica	SiO <sub>2</sub>	60.2%	48.6%
alumina	Al <sub>2</sub> O <sub>3</sub>	15.2%	16.5%
lime	CaO	5.5%	12.3%
magnesia	MgO	3.1%	6.8%
iron(II) oxide	FeO	3.8%	6.2%
sodium oxide	Na <sub>2</sub> O	3.0%	2.6%
potassium oxide	K <sub>2</sub> O	2.8%	0.4%
iron(III) oxide	Fe <sub>2</sub> O <sub>3</sub>	2.5%	2.3%
water	H <sub>2</sub> O	1.4%	1.1%
carbon dioxide	CO <sub>2</sub>	1.2%	1.4%
titanium dioxide	TiO <sub>2</sub>	0.7%	1.4%
phosphorus pentoxide	P <sub>2</sub> O <sub>5</sub>	0.2%	0.3%
<b>Total</b>		<b>99.6%</b>	<b>99.9%</b>

### Chemical composition

The mass of the Earth is approximately  $5.98 \times 10^{24}$  kg. It is composed mostly of iron (32.1%), oxygen (30.1%), silicon (15.1%), magnesium (13.9%), sulfur (2.9%), nickel (1.8%), calcium (1.5%), and aluminium (1.4%); with the remaining 1.2% consisting of trace amounts of other elements. Due to mass segregation, the core region is believed to be primarily composed of iron (88.8%), with smaller amounts of nickel (5.8%), sulfur (4.5%), and less than 1% trace elements.<sup>[79]</sup>

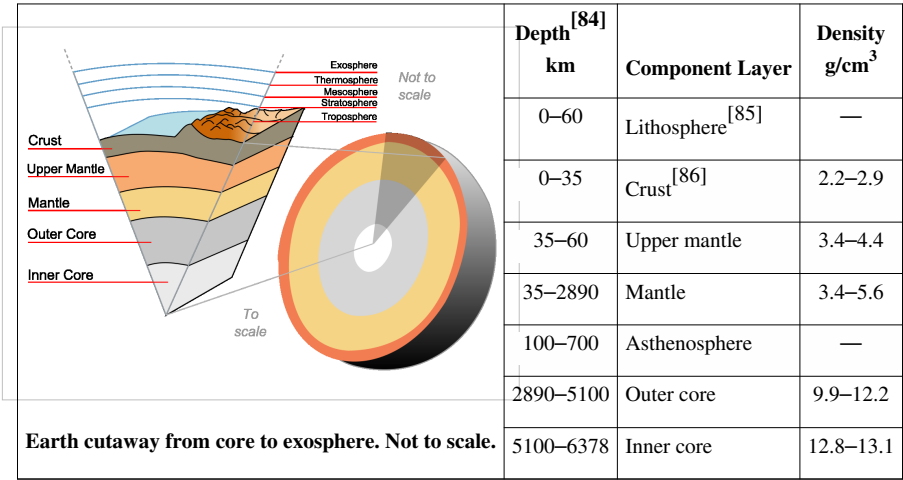
The geochemist F. W. Clarke calculated that a little more than 47% of the Earth's crust consists of oxygen. The more common rock constituents of the Earth's crust are nearly all oxides; chlorine, sulfur and fluorine are the only important exceptions to this and their total amount in any rock is usually much less than 1%. The principal oxides are silica, alumina, iron oxides, lime, magnesia, potash and soda. The silica functions principally as an acid, forming silicates, and all the commonest minerals of igneous rocks are of this nature. From a computation based on 1,672 analyses of all kinds of rocks, Clarke deduced that 99.22% were composed of 11 oxides (see the table at right), with the other constituents occurring in minute quantities.<sup>[80]</sup>

### Internal structure

The interior of the Earth, like that of the other terrestrial planets, is divided into layers by their chemical or physical (rheological) properties, but unlike the other terrestrial planets, it has a distinct outer and inner core. The outer layer of the Earth is a chemically distinct silicate solid crust, which is underlain by a highly viscous solid mantle. The crust is separated from the mantle by the Mohorovičić discontinuity, and the thickness of the crust varies: averaging 6 km under the oceans and 30–50 km on the continents. The crust and the cold, rigid, top of the upper mantle are collectively known as the lithosphere, and it is of the lithosphere that the tectonic plates are comprised. Beneath the

lithosphere is the asthenosphere, a relatively low-viscosity layer on which the lithosphere rides. Important changes in crystal structure within the mantle occur at 410 and 660 kilometers below the surface, spanning a transition zone that separates the upper and lower mantle. Beneath the mantle, an extremely low viscosity liquid outer core lies above a solid inner core.<sup>[81]</sup> The inner core may rotate at a slightly higher angular velocity than the remainder of the planet, advancing by 0.1–0.5° per year.<sup>[82]</sup>

Geologic layers of the Earth<sup>[83]</sup>



Heat

Earth's internal heat comes from a combination of residual heat from planetary accretion (about 20%) and heat produced through radioactive decay (80%).<sup>[87]</sup> The major heat-producing isotopes in the Earth are potassium-40, uranium-238, uranium-235, and thorium-232.<sup>[88]</sup> At the center of the planet, the temperature may be up to 7,000 K and the pressure could reach 360 GPa.<sup>[89]</sup> Because much of the heat is provided by radioactive decay, scientists believe that early in Earth history, before isotopes with short half-lives had been depleted, Earth's heat production would have been much higher. This extra heat production, twice present-day at approximately 3 billion years ago,<sup>[87]</sup> would have increased temperature gradients within the Earth, increasing the rates of mantle convection and plate tectonics, and allowing the production of igneous rocks such as komatiites that are not formed today.<sup>[90]</sup>

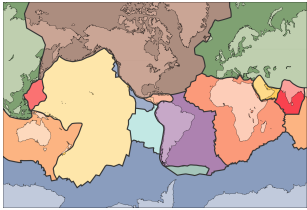
Present-day major heat-producing isotopes<sup>[91]</sup>

Isotope	Heat release W/kg isotope	Half-life years	Mean mantle concentration kg isotope/kg mantle	Heat release W/kg mantle
<sup>238</sup> U	9.46 × 10 <sup>-5</sup>	4.47 × 10 <sup>9</sup>	30.8 × 10 <sup>-9</sup>	2.91 × 10 <sup>-12</sup>
<sup>235</sup> U	5.69 × 10 <sup>-4</sup>	7.04 × 10 <sup>8</sup>	0.22 × 10 <sup>-9</sup>	1.25 × 10 <sup>-13</sup>
<sup>232</sup> Th	2.64 × 10 <sup>-5</sup>	1.40 × 10 <sup>10</sup>	124 × 10 <sup>-9</sup>	3.27 × 10 <sup>-12</sup>
<sup>40</sup> K	2.92 × 10 <sup>-5</sup>	1.25 × 10 <sup>9</sup>	36.9 × 10 <sup>-9</sup>	1.08 × 10 <sup>-12</sup>

The mean heat loss from the Earth is 87 mW m<sup>-2</sup>, for a global heat loss of 4.42 × 10<sup>13</sup> W.<sup>[92]</sup> A portion of the core's thermal energy is transported toward the crust by mantle plumes; a form of convection consisting of upwellings of higher-temperature rock. These plumes can produce hotspots and flood basalts.<sup>[93]</sup> More of the heat in the Earth is lost through plate tectonics, by mantle upwelling associated with mid-ocean ridges. The final major mode of heat loss is through conduction through the lithosphere, the majority of which occurs in the oceans because the crust there is much thinner than that of the continents.<sup>[94]</sup>

Tectonic plates

Earth's main plates<sup>[95]</sup>

	
Plate name	Area 10 <sup>6</sup> km <sup>2</sup>
Pacific Plate	103.3
African Plate <sup>[96]</sup>	78.0
North American Plate	75.9
Eurasian Plate	67.8
Antarctic Plate	60.9
Indo-Australian Plate	47.2
South American Plate	43.6

The mechanically rigid outer layer of the Earth, the lithosphere, is broken into pieces called tectonic plates. These plates are rigid segments that move in relation to one another at one of three types of plate boundaries: Convergent boundaries, at which two plates come together, Divergent boundaries, at which two plates are pulled apart, and Transform boundaries, in which two plates slide past one another laterally. Earthquakes, volcanic activity, mountain-building, and oceanic trench formation can occur along these plate boundaries.<sup>[97]</sup> The tectonic plates ride on top of the asthenosphere, the solid but less-viscous part of the upper mantle that can flow and move along with the plates,<sup>[98]</sup> and their motion is strongly coupled with convection patterns inside the Earth's mantle.

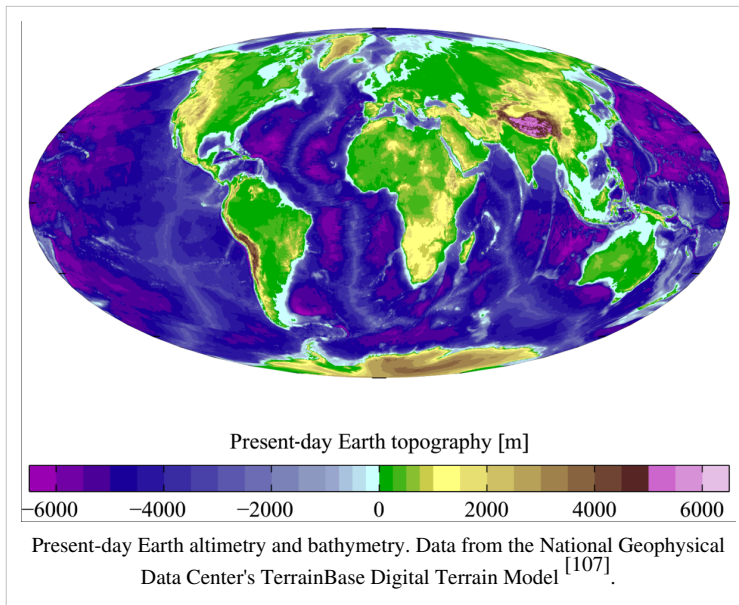
As the tectonic plates migrate across the planet, the ocean floor is subducted under the leading edges of the plates at convergent boundaries. At the same time, the upwelling of mantle material at divergent boundaries creates mid-ocean ridges. The combination of these processes continually recycles the oceanic crust back into the mantle. Because of this recycling, most of the ocean floor is less than 100 million years in age. The oldest oceanic crust is located in the Western Pacific, and has an estimated age of about 200 million years.<sup>[99][100]</sup> By comparison, the oldest dated continental crust is 4030 million years old.<sup>[101]</sup>

The seven major plates are the Pacific, North American, Eurasian, African, Antarctic, Indo-Australian, and South American. Other notable plates include the Arabian Plate, the Caribbean Plate, the Nazca Plate off the west coast of South America and the Scotia Plate in the southern Atlantic Ocean. The Australian Plate fused with the Indian Plate between 50 and 55 million years ago. The fastest-moving plates are the oceanic plates, with the Cocos Plate advancing at a rate of 75 mm/year<sup>[102]</sup> and the Pacific Plate moving 52–69 mm/year. At the other extreme, the slowest-moving plate is the Eurasian Plate, progressing at a typical rate of about 21 mm/year.<sup>[103]</sup>

Surface

The Earth's terrain varies greatly from place to place. About 70.8%<sup>[104]</sup> of the surface is covered by water, with much of the continental shelf below sea level. This equates to 148.94 million sq km (57.51 million sq mi)<sup>[105]</sup>. The submerged surface has mountainous features, including a globe-spanning mid-ocean ridge system, as well as undersea volcanoes,<sup>[71]</sup> oceanic trenches, submarine canyons, oceanic plateaus and abyssal plains. The remaining 29.2% not covered by water consists of mountains, deserts, plains, plateaus, and other geomorphologies.

The planetary surface undergoes reshaping over geological time periods because of tectonics and erosion. The surface features built up or deformed through plate tectonics are subject to steady weathering from precipitation, thermal cycles, and chemical effects. Glaciation, coastal erosion, the build-up of coral reefs, and large meteorite impacts<sup>[106]</sup> also act to reshape the landscape.



The continental crust consists of lower density material such as the igneous rocks granite and andesite. Less common is basalt, a denser volcanic rock that is the primary constituent of the ocean floors.<sup>[108]</sup> Sedimentary rock is formed from the accumulation of sediment that becomes compacted together. Nearly 75% of the continental surfaces are covered by sedimentary rocks, although they form only about 5% of the crust.<sup>[109]</sup> The third form of rock material found on Earth is metamorphic rock, which is created from the transformation of pre-existing rock types through high pressures, high temperatures, or both. The most abundant silicate minerals

on the Earth's surface include quartz, the feldspars, amphibole, mica, pyroxene and olivine.<sup>[110]</sup> Common carbonate minerals include calcite (found in limestone) and dolomite.<sup>[111]</sup>

The pedosphere is the outermost layer of the Earth that is composed of soil and subject to soil formation processes. It exists at the interface of the lithosphere, atmosphere, hydrosphere and biosphere. Currently the total arable land is 13.31% of the land surface, with only 4.71% supporting permanent crops.<sup>[17]</sup> Close to 40% of the Earth's land surface is presently used for cropland and pasture, or an estimated  $1.3 \times 10^7$  km<sup>2</sup> of cropland and  $3.4 \times 10^7$  km<sup>2</sup> of pastureland.<sup>[112]</sup>

The elevation of the land surface of the Earth varies from the low point of -418 m at the Dead Sea, to a 2005-estimated maximum altitude of 8,848 m at the top of Mount Everest. The mean height of land above sea level is 840 m.<sup>[113]</sup>

## Hydrosphere

The abundance of water on Earth's surface is a unique feature that distinguishes the "Blue Planet" from others in the Solar System. The Earth's hydrosphere consists chiefly of the oceans, but technically includes all water surfaces in the world, including inland seas, lakes, rivers, and underground waters down to a depth of 2,000 m. The deepest underwater location is Challenger Deep of the Mariana Trench in the Pacific Ocean with a depth of  $-10,911.4$  m.<sup>[114][115]</sup>

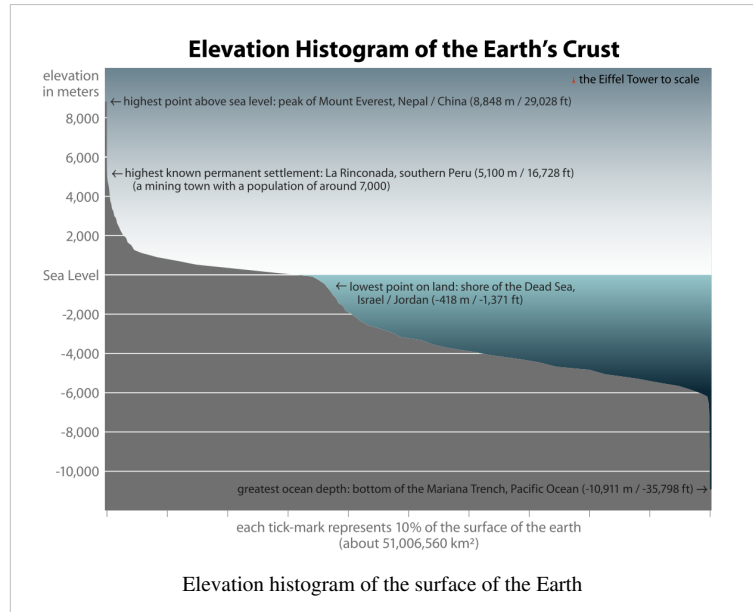
The mass of the oceans is approximately  $1.35 \times 10^{18}$  metric tons, or about 1/4400 of the total mass of the Earth. The oceans cover an area of  $3.618 \times 10^8$  km<sup>2</sup> with a mean depth of 3,682 m, resulting in an estimated volume of  $1.332 \times 10^9$  km<sup>3</sup>.<sup>[116]</sup> If all the land on Earth were spread evenly, water would rise to an altitude of more than 2.7 km.<sup>[117]</sup> About 97.5% of the water is saline, while the remaining 2.5% is fresh water. Most fresh water, about 68.7%, is currently ice.<sup>[118]</sup>

The average salinity of the Earth's oceans is about 35 grams of salt per kilogram of sea water (35 ‰).<sup>[119]</sup> Most of this salt was released from volcanic activity or extracted from cool, igneous rocks.<sup>[120]</sup> The oceans are also a reservoir of dissolved atmospheric gases, which are essential for the survival of many aquatic life forms.<sup>[121]</sup> Sea water has an important influence on the world's climate, with the oceans acting as a large heat reservoir.<sup>[122]</sup> Shifts in the oceanic temperature distribution can cause significant weather shifts, such as the El Niño-Southern Oscillation.<sup>[123]</sup>

## Atmosphere

The atmospheric pressure on the surface of the Earth averages 101.325 kPa, with a scale height of about 8.5 km.<sup>[5]</sup> It is 78% nitrogen and 21% oxygen, with trace amounts of water vapor, carbon dioxide and other gaseous molecules. The height of the troposphere varies with latitude, ranging between 8 km at the poles to 17 km at the equator, with some variation resulting from weather and seasonal factors.<sup>[124]</sup>

Earth's biosphere has significantly altered its atmosphere. Oxygenic photosynthesis evolved 2.7 billion years ago, forming the primarily nitrogen-oxygen atmosphere of today. This change enabled the proliferation of aerobic organisms as well as the formation of the ozone layer which blocks ultraviolet solar radiation, permitting life on land. Other atmospheric functions important to life on Earth include transporting water vapor, providing useful gases, causing small meteors to burn up before they strike the surface, and moderating temperature.<sup>[125]</sup> This last phenomenon is known as the greenhouse effect: trace molecules within the atmosphere serve to capture thermal energy emitted from the ground, thereby raising the average temperature. Water vapor, carbon dioxide, methane and ozone are the primary greenhouse gases in the Earth's atmosphere. Without this heat-retention effect, the average surface would be  $-18$  °C, in contrast to the current  $+15$  °C, and life would likely not exist.<sup>[104]</sup>



## Weather and climate

The Earth's atmosphere has no definite boundary, slowly becoming thinner and fading into outer space. Three-quarters of the atmosphere's mass is contained within the first 11 km of the planet's surface. This lowest layer is called the troposphere. Energy from the Sun heats this layer, and the surface below, causing expansion of the air. This lower density air then rises, and is replaced by cooler, higher density air. The result is atmospheric circulation that drives the weather and climate through redistribution of heat energy.<sup>[126]</sup>

The primary atmospheric circulation bands consist of the trade winds in the equatorial region below 30° latitude and the westerlies in the mid-latitudes between 30° and 60°.<sup>[127]</sup> Ocean currents are also important factors in determining climate, particularly the thermohaline circulation that distributes heat energy from the equatorial oceans to the polar regions.<sup>[128]</sup>

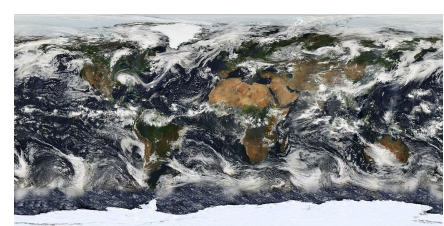
Water vapor generated through surface evaporation is transported by circulatory patterns in the atmosphere. When atmospheric conditions permit an uplift of warm, humid air, this water condenses and settles to the surface as precipitation.<sup>[126]</sup> Most of the water is then transported to lower elevations by river systems and usually returned to the oceans or deposited into lakes. This water cycle is a vital mechanism for supporting life on land, and is a primary factor in the erosion of surface features over geological periods. Precipitation patterns vary widely, ranging from several meters of water per year to less than a millimeter. Atmospheric circulation, topological features and temperature differences determine the average precipitation that falls in each region.<sup>[129]</sup>

The amount of solar energy reaching the Earth's decreases with increasing latitude. At higher latitudes the sunlight reaches the surface at a lower angles and it must pass through thicker columns of the atmosphere. As a result, the mean annual air temperature at sea level decreases by about 0.4 °C per per degree of latitude away from the equator.<sup>[130]</sup> The Earth can be sub-divided into specific latitudinal belts of approximately homogeneous climate. Ranging from the equator to the polar regions, these are the tropical (or equatorial), subtropical, temperate and polar climates.<sup>[131]</sup> Climate can also be classified based on the temperature and precipitation, with the climate regions characterized by fairly uniform air masses. The commonly used Köppen climate classification system (as modified by Wladimir Köppen's student Rudolph Geiger) has five broad groups (humid tropics, arid, humid middle latitudes, continental and cold polar), which are further divided into more specific subtypes.<sup>[127]</sup>

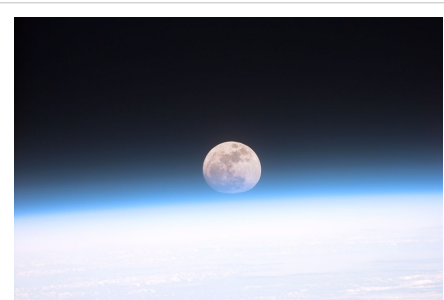
## Upper atmosphere

Above the troposphere, the atmosphere is usually divided into the stratosphere, mesosphere, and thermosphere.<sup>[125]</sup> Each layer has a different lapse rate, defining the rate of change in temperature with height. Beyond these, the exosphere thins out into the magnetosphere, where the Earth's magnetic fields interact with the solar wind.<sup>[132]</sup> Within the stratosphere is the ozone layer, a component that partially shields the surface from ultraviolet light and thus is important for life on Earth. The Kármán line, defined as 100 km above the Earth's surface, is a working definition for the boundary between atmosphere and space.<sup>[133]</sup>

Thermal energy causes some of the molecules at the outer edge of the Earth's atmosphere have their velocity increased to the point where they can escape from the planet's gravity. This results in a slow but steady leakage of the atmosphere into space. Because unfixed hydrogen has a low molecular weight, it can achieve escape velocity more readily and it leaks into outer space at a greater rate than other



Satellite cloud cover image of Earth using NASA's Moderate-Resolution Imaging Spectroradiometer



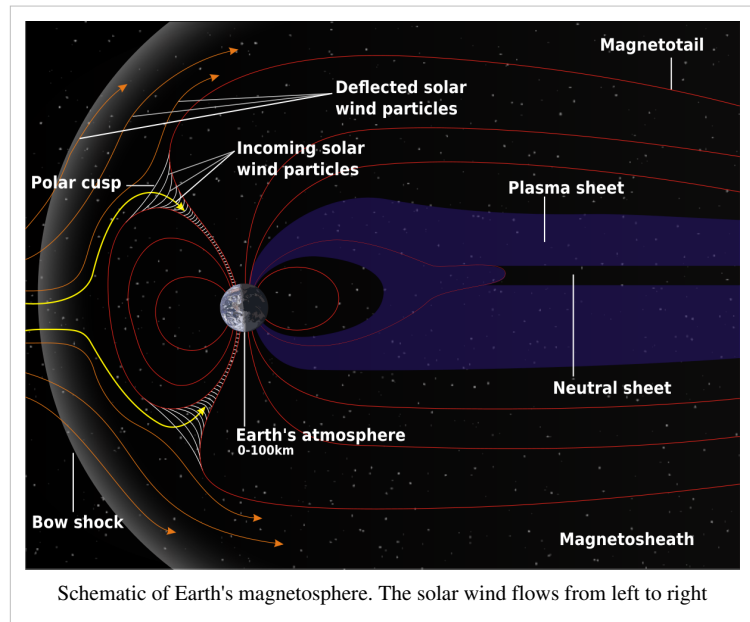
This view from orbit shows the full Moon partially obscured and deformed by the Earth's atmosphere. NASA image



gasses.<sup>[134]</sup> The leakage of hydrogen into space contributes to the pushing of the Earth from an initially reducing state to its current oxidizing one. Photosynthesis provided a source of free oxygen, but the loss of reducing agents such as hydrogen is believed to have been a necessary precondition for the widespread accumulation of oxygen in the atmosphere.<sup>[135]</sup> Hence the ability of hydrogen to escape from the Earth's atmosphere may have influenced the nature of life that developed on the planet.<sup>[136]</sup> In the current, oxygen-rich atmosphere most hydrogen is converted into water before it has an opportunity to escape. Instead, most of the hydrogen loss comes from the destruction of methane in the upper atmosphere.<sup>[137]</sup>

## Magnetic field

The Earth's magnetic field is shaped roughly as a magnetic dipole, with the poles currently located proximate to the planet's geographic poles. At the equator of the magnetic field, the magnetic field strength at the planet's surface is  $3.05 \times 10^{-5}$  T, with global magnetic dipole moment of  $7.91 \times 10^{15}$  T m<sup>3</sup>.<sup>[138]</sup> According to dynamo theory, the field is generated within the molten outer core region where heat creates convection motions of conducting materials, generating electric currents. These in turn produce the Earth's magnetic field. The convection movements in the core are chaotic; the magnetic poles drift and periodically change alignment. This results in field reversals at irregular intervals averaging a few times every million years. The most recent reversal occurred approximately 700,000 years ago.<sup>[139][140]</sup>



The field forms the magnetosphere, which deflects particles in the solar wind. The sunward edge of the bow shock is located at about 13 times the radius of the Earth. The collision between the magnetic field and the solar wind forms the Van Allen radiation belts, a pair of concentric, torus-shaped regions of energetic charged particles. When the plasma enters the Earth's atmosphere at the magnetic poles, it forms the aurora.<sup>[141]</sup>

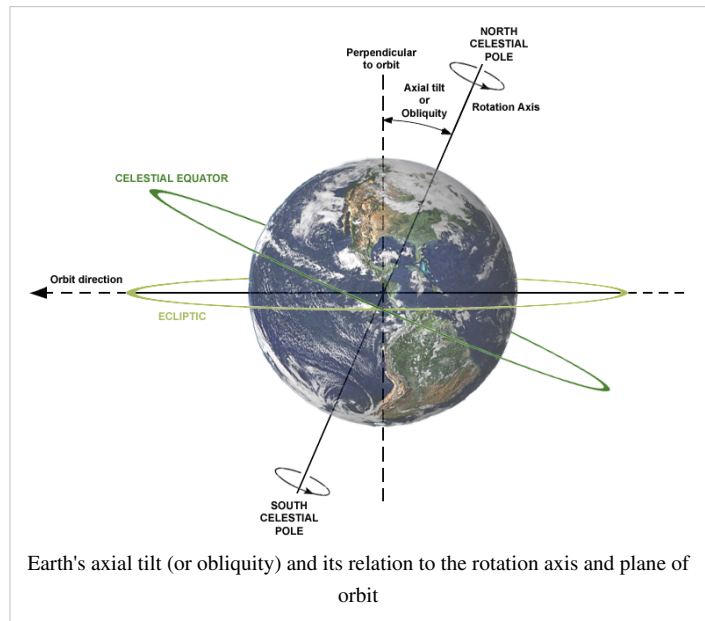
## Orbit and rotation

### Rotation

Earth's rotation period relative to the Sun—its mean solar day—is 86,400 seconds of mean solar time (86,400.0025 SI seconds).<sup>[142]</sup> As the Earth's solar day is now slightly longer than it was during the 19th century because of tidal acceleration, each day varies between 0 and 2 SI ms longer.<sup>[143][144]</sup>

Earth's rotation period relative to the fixed stars, called its *stellar day* by the International Earth Rotation and Reference Systems Service (IERS), is 86164.098903691 seconds of mean solar time (UT1), or 23<sup>h</sup> 56<sup>m</sup> 4.098903691<sup>s</sup>.<sup>[4][145]</sup> Earth's rotation period relative to the precessing or moving mean vernal equinox, misnamed its *sidereal day*, is 86164.09053083288 seconds of mean solar time (UT1) (23<sup>h</sup> 56<sup>m</sup> 4.09053083288<sup>s</sup>).<sup>[4]</sup> Thus the sidereal day is shorter than the stellar day by about 8.4 ms.<sup>[146]</sup>

The length of the mean solar day in SI seconds is available from the IERS for the periods 1623–2005<sup>[147]</sup> and 1962–2005.<sup>[148]</sup>



Apart from meteors within the atmosphere and low-orbiting satellites, the main apparent motion of celestial bodies in the Earth's sky is to the west at a rate of 15°/h = 15'/min. For bodies near the celestial equator, this is equivalent to an apparent diameter of the Sun or Moon every two minutes; from the planet's surface, the apparent sizes of the Sun and the Moon are approximately the same.<sup>[149][150]</sup>

### Orbit

Earth orbits the Sun at an average distance of about 150 million kilometers every 365.2564 mean solar days, or one sidereal year. From Earth, this gives an apparent movement of the Sun eastward with respect to the stars at a rate of about 1°/day, or a Sun or Moon diameter, every 12 hours. Because of this motion, on average it takes 24 hours—a solar day—for Earth to complete a full rotation about its axis so that the Sun returns to the meridian. The orbital speed of the Earth averages about 29.8 km/s (107,000 km/h), which is fast enough to cover the planet's diameter (about 12,600 km) in seven minutes, and the distance to the Moon (384,000 km) in four hours.<sup>[5]</sup>

The Moon revolves with the Earth around a common barycenter every 27.32 days relative to the background stars. When combined with the Earth–Moon system's common revolution around the Sun, the period of the synodic month, from new moon to new moon, is 29.53 days. Viewed from the celestial north pole, the motion of Earth, the Moon and their axial rotations are all counter-clockwise. Viewed from a vantage point above the north poles of both the Sun and the Earth, the Earth appears to revolve in a counterclockwise direction about the Sun. The orbital and axial planes are not precisely aligned: Earth's axis is tilted some 23.4 degrees from the perpendicular to the Earth–Sun plane, and the Earth–Moon plane is tilted about 5 degrees against the Earth–Sun plane. Without this tilt, there would be an eclipse every two weeks, alternating between lunar eclipses and solar eclipses.<sup>[5][151]</sup>

The Hill sphere, or gravitational sphere of influence, of the Earth is about 1.5 Gm (or 1,500,000 kilometers) in radius.<sup>[152][153]</sup> This is maximum distance at which the Earth's gravitational influence is stronger than the more distant Sun and planets. Objects must orbit the Earth within this radius, or they can become unbound by the gravitational perturbation of the Sun.

Earth, along with the Solar System, is situated in the Milky Way galaxy, orbiting about 28,000 light years from the center of the galaxy. It is currently about 20 light years above the galaxy's equatorial plane in the Orion spiral arm.<sup>[154]</sup>

## Axial tilt and seasons

Because of the axial tilt of the Earth, the amount of sunlight reaching any given point on the surface varies over the course of the year. This results in seasonal change in climate, with summer in the northern hemisphere occurring when the North Pole is pointing toward the Sun, and winter taking place when the pole is pointed away. During the summer, the day lasts longer and the Sun climbs higher in the sky. In winter, the climate becomes generally cooler and the days shorter. Above the Arctic Circle, an extreme case is reached where there is no daylight at all for part of the year—a polar night. In the southern hemisphere the situation is exactly reversed, with the South Pole oriented opposite the direction of the North Pole.

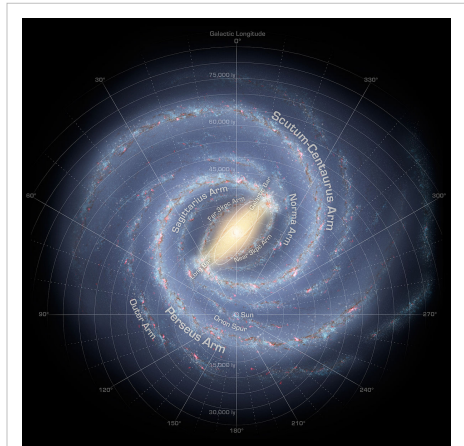
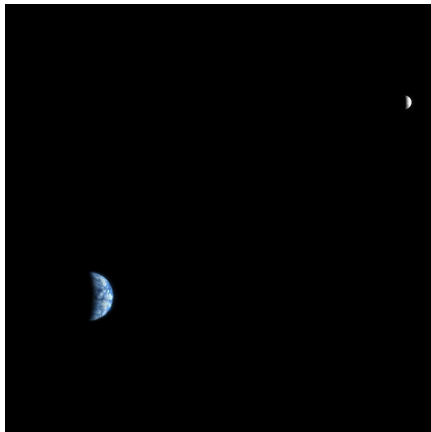


Illustration of the Milky Way Galaxy, showing the location of the Sun



Earth and Moon from Mars, imaged by Mars Reconnaissance Orbiter. From space, the Earth can be seen to go through phases similar to the phases of the Moon.

By astronomical convention, the four seasons are determined by the solstices—the point in the orbit of maximum axial tilt toward or away from the Sun—and the equinoxes, when the direction of the tilt and the direction to the Sun are perpendicular. In the northern hemisphere, Winter Solstice occurs on about December 21, Summer Solstice is near June 21, Spring Equinox is around March 20 and Autumnal Equinox is about September 23. In the Southern hemisphere, the situation is reversed, with the Summer and Winter Solstices exchanged and the Spring and Autumnal Equinox dates switched.<sup>[155]</sup>

The angle of the Earth's tilt is relatively stable over long periods of time. The tilt does undergo nutation; a slight, irregular motion with a main period of 18.6 years.<sup>[156]</sup> The orientation (rather than the angle) of the Earth's axis also changes over time, precessing around in a complete circle over each 25,800 year cycle; this precession is the reason for the difference between a sidereal year and a tropical year.

Both of these motions are caused by the varying attraction of the Sun and Moon on the Earth's equatorial bulge. From the perspective of the Earth, the poles also migrate a few meters across the surface. This polar motion has multiple, cyclical components, which collectively are termed quasiperiodic motion. In addition to an annual component to this motion, there is a 14-month cycle called the Chandler wobble. The rotational velocity of the Earth also varies in a phenomenon known as length of day variation.<sup>[157]</sup>

In modern times, Earth's perihelion occurs around January 3, and the aphelion around July 4. These dates change over time due to precession and other orbital factors, which follow cyclical patterns known as Milankovitch cycles. The changing Earth-Sun distance results in an increase of about 6.9%<sup>[158]</sup> in solar energy reaching the Earth at perihelion relative to aphelion. Since the southern hemisphere is tilted toward the Sun at about the same time that the Earth reaches the closest approach to the Sun, the southern hemisphere receives slightly more energy from the Sun than does the northern over the course of a year. This effect is much less significant than the total energy change due to the axial tilt, and most of the excess energy is absorbed by the higher proportion of water in the southern hemisphere.<sup>[159]</sup>

Moon

Characteristics

Diameter	3,474.8 km
Mass	$7.349 \times 10^{22}$ kg
Semi-major axis	384,400 km
Orbital period	27 d 7 h 43.7 m

The Moon is a relatively large, terrestrial, planet-like satellite, with a diameter about one-quarter of the Earth's. It is the largest moon in the Solar System relative to the size of its planet, although Charon is larger relative to the dwarf planet Pluto. The natural satellites orbiting other planets are called "moons" after Earth's Moon.

The gravitational attraction between the Earth and Moon causes tides on Earth. The same effect on the Moon has led to its tidal locking: its rotation period is the same as the time it takes to orbit the Earth. As a result, it always presents the same face to the planet. As the Moon orbits Earth, different parts of its face are illuminated by the Sun, leading to the lunar phases; the dark part of the face is separated from the light part by the solar terminator.

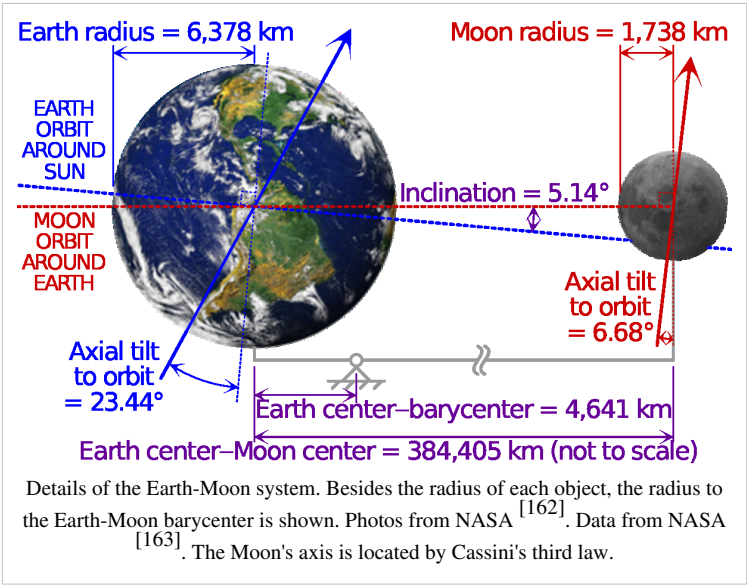
Because of their tidal interaction, the Moon recedes from Earth at the rate of approximately 38 mm a year. Over millions of years, these tiny modifications—and the lengthening of Earth's day by about 23  $\mu$ s a year—add up to significant changes.<sup>[160]</sup> During the Devonian period, for example, (approximately 410 million years ago) there were 400 days in a year, with each day lasting 21.8 hours.<sup>[161]</sup>

The Moon may have dramatically affected the development of life by moderating the planet's climate. Paleontological evidence and computer simulations show that Earth's axial tilt is stabilized by tidal interactions with the Moon.<sup>[164]</sup> Some theorists believe that without this stabilization against the torques applied by the Sun and planets to the Earth's equatorial bulge, the rotational axis might be chaotically unstable, exhibiting chaotic changes over millions of years, as appears to be the case for Mars.<sup>[165]</sup>

Viewed from Earth, the Moon is just far enough away to have very nearly the same apparent-sized disk as the Sun. The angular size (or solid angle) of these two bodies match because, although the Sun's diameter is about 400 times as large as the Moon's, it is also 400 times more distant.<sup>[150]</sup> This allows total and annular solar eclipses to occur on Earth.

The most widely accepted theory of the Moon's origin, the giant impact theory, states that it formed from the collision of a Mars-size protoplanet called Theia with the early Earth. This hypothesis explains (among other things) the Moon's relative lack of iron and volatile elements, and the fact that its composition is nearly identical to that of the Earth's crust.<sup>[166]</sup>

Earth has at least five co-orbital asteroids, including 3753 Cruithne and 2002 AA<sub>29</sub>.<sup>[167][168]</sup> As of 2011, there are 931 operational, man-made satellites orbiting the Earth.<sup>[169]</sup> On July 27, 2011, astronomers reported a trojan asteroid companion, 2010 TK<sub>7</sub>, librating around the leading Lagrange triangular point, L4, of Earth in Earth's orbit around the Sun.<sup>[170][171]</sup>





A scale representation of the relative sizes of, and average distance between, Earth and Moon

**Habitability**

A planet that can sustain life is termed habitable, even if life did not originate there. The Earth provides liquid water—an environment where complex organic molecules can assemble and interact, and sufficient energy to sustain metabolism.<sup>[172]</sup> The distance of the Earth from the Sun, as well as its orbital eccentricity, rate of rotation, axial tilt, geological history, sustaining atmosphere and protective magnetic field all contribute to the current climactic conditions at the surface.<sup>[173]</sup>

**Biosphere**

The planet's life forms are sometimes said to form a "biosphere". This biosphere is generally believed to have begun evolving about 3.5 billion years ago. The biosphere is divided into a number of biomes, inhabited by broadly similar plants and animals. On land, biomes are separated primarily by differences in latitude, height above sea level and humidity. Terrestrial biomes lying within the Arctic or Antarctic Circles, at high altitudes or in extremely arid areas are relatively barren of plant and animal life; species diversity reaches a peak in humid lowlands at equatorial latitudes.<sup>[174]</sup>

**Natural resources and land use**

The Earth provides resources that are exploitable by humans for useful purposes. Some of these are non-renewable resources, such as mineral fuels, that are difficult to replenish on a short time scale.

Large deposits of fossil fuels are obtained from the Earth's crust, consisting of coal, petroleum, natural gas and methane clathrate. These deposits are used by humans both for energy production and as feedstock for chemical production. Mineral ore bodies have also been formed in Earth's crust through a process of Ore genesis, resulting from actions of erosion and plate tectonics.<sup>[175]</sup> These bodies form concentrated sources for many metals and other useful elements.

The Earth's biosphere produces many useful biological products for humans, including (but far from limited to) food, wood, pharmaceuticals, oxygen, and the recycling of many organic wastes. The land-based ecosystem depends upon topsoil and fresh water, and the oceanic ecosystem depends upon dissolved nutrients washed down from the land.<sup>[176]</sup> Humans also live on the land by using building materials to construct shelters. In 1993, human use of land is approximately:

Land use	Arable land	Permanent crops	Permanent pastures	Forests and woodland	Urban areas	Other
Percentage	13.13% <sup>[17]</sup>	4.71% <sup>[17]</sup>	26%	32%	1.5%	30%

The estimated amount of irrigated land in 1993 was 2,481,250 km<sup>2</sup>.<sup>[17]</sup>

## Natural and environmental hazards

Large areas of the Earth's surface are subject to extreme weather such as tropical cyclones, hurricanes, or typhoons that dominate life in those areas. From 1980–2000, these events caused an average of 11,800 deaths per year.<sup>[177]</sup> Many places are subject to earthquakes, landslides, tsunamis, volcanic eruptions, tornadoes, sinkholes, blizzards, floods, droughts, wildfires, and other calamities and disasters.

Many localized areas are subject to human-made pollution of the air and water, acid rain and toxic substances, loss of vegetation (overgrazing, deforestation, desertification), loss of wildlife, species extinction, soil degradation, soil depletion, erosion, and introduction of invasive species.

According to the United Nations, a scientific consensus exists linking human activities to global warming due to industrial carbon dioxide emissions. This is predicted to produce changes such as the melting of glaciers and ice sheets, more extreme temperature ranges, significant changes in weather and a global rise in average sea levels.<sup>[178]</sup>

## Human geography

Cartography, the study and practice of map making, and vicariously geography, have historically been the disciplines devoted to depicting the Earth. Surveying, the determination of locations and distances, and to a lesser extent navigation, the determination of position and direction, have developed alongside cartography and geography, providing and suitably quantifying the requisite information.

Earth has reached approximately 7,000,000,000 human inhabitants as of October 31, 2011.<sup>[179]</sup> Projections indicate that the world's human population will reach 9.2 billion in 2050.<sup>[180]</sup> Most of the growth is expected to take place in developing nations. Human population density varies widely around the world, but a majority live in Asia. By 2020, 60% of the world's population is expected to be living in urban, rather than rural, areas.<sup>[181]</sup>

It is estimated that only one-eighth of the surface of the Earth is suitable for humans to live on—three-quarters is covered by oceans, and half of the land area is either desert (14%),<sup>[182]</sup> high mountains (27%),<sup>[183]</sup> or other less suitable terrain. The northernmost permanent settlement in the world is Alert, on Ellesmere Island in Nunavut, Canada.<sup>[184]</sup> (82°28'N) The southernmost is the Amundsen-Scott South Pole Station, in Antarctica, almost exactly at the South Pole. (90°S)

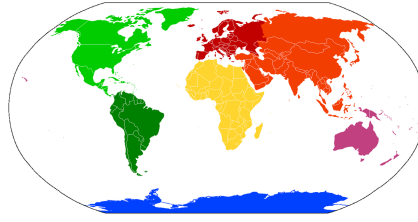
Independent sovereign nations claim the planet's entire land surface, except for some parts of Antarctica and the odd unclaimed area of Bir Tawil between Egypt and Sudan. As of 2011 there are 204 sovereign states, including the 193 United Nations member states. In addition, there are 59 dependent territories, and a number of autonomous areas, territories under dispute and other entities.<sup>[17]</sup> Historically, Earth has never had a sovereign government with authority over the entire globe, although a number of nation-states have striven for world domination and failed.<sup>[185]</sup>

The United Nations is a worldwide intergovernmental organization that was created with the goal of intervening in the disputes between nations, thereby avoiding armed conflict.<sup>[186]</sup> The U.N. serves primarily as a forum for international diplomacy and international law. When the consensus of the membership permits, it provides a mechanism for armed intervention.<sup>[187]</sup>

The first human to orbit the Earth was Yuri Gagarin on April 12, 1961.<sup>[188]</sup> In total, about 487 people have visited outer space and reached Earth orbit as of July 30, 2010, and, of these, twelve have walked on the Moon.<sup>[189][190][191]</sup> Normally the only humans in space are those on the International Space Station. The station's crew, currently six people, is usually replaced every six months.<sup>[192]</sup> The furthest humans have travelled from Earth is 400,171 km, achieved during the 1970 Apollo 13 mission.<sup>[193]</sup>

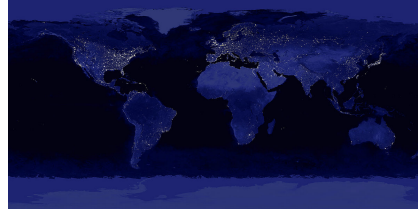
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The 7 continents of Earth:<sup>[194]</sup>

North America, South America, Antarctica, Africa, Europe, Asia, Australia



The Earth at night, a composite of DMSP/OLS ground illumination data on a simulated night-time image of the world. This image is not photographic and many features are brighter than they would appear to a direct observer.

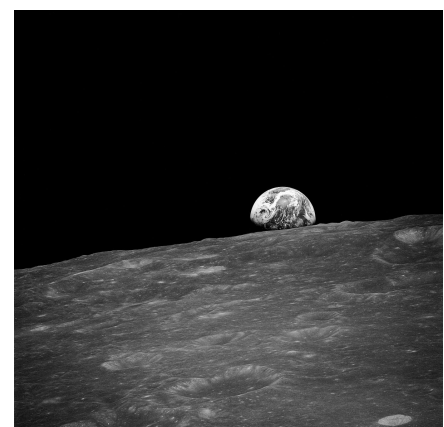


ISS video beginning just south-east of Alaska. The first city that the ISS passes over (seen approximately 10 seconds into the video) is San Francisco and the surrounding areas. A careful examination shows where the Golden Gate Bridge is located: a smaller strip of lights just before the city of San Francisco, nearest to the clouds on the right of the image. Very obvious lightning storms can be seen on the Pacific Ocean coastline, with clouds overhead. As the video continues, the ISS passes over Central America (green lights can be seen here), with the Yucatan Peninsula on the left. The pass ends as the ISS is over the capital city of Bolivia, La Paz.

## Cultural viewpoint

The name "Earth" derives from the Anglo-Saxon word *erda*, which means ground or soil, and is related to the German word *Erde*. It became *eorthe* later, and then *erthe* in Middle English.<sup>[195]</sup> The standard astronomical symbol of the Earth consists of a cross circumscribed by a circle.<sup>[196]</sup>

Unlike the rest of the planets in the Solar System, humankind did not begin to view the Earth as a moving object in orbit around the Sun until the 16th century.<sup>[197]</sup> Earth has often been personified as a deity, in particular a goddess. In many cultures the mother goddess is also portrayed as a fertility deity. Creation myths in many religions recall a story involving the creation of the Earth by a supernatural deity or deities. A variety of religious groups, often associated with fundamentalist branches of Protestantism<sup>[198]</sup> or Islam,<sup>[199]</sup> assert that their interpretations of these creation myths in sacred texts are literal truth and should be considered alongside or replace conventional scientific accounts of the formation of the Earth



The first photograph ever taken by astronauts of an "Earthrise", from Apollo 8

and the origin and development of life.<sup>[200]</sup> Such assertions are opposed by the scientific community<sup>[201][202]</sup> and by other religious groups.<sup>[203][204][205]</sup> A prominent example is the creation-evolution controversy.

In the past there were varying levels of belief in a flat Earth,<sup>[206]</sup> but this was displaced by the concept of a spherical Earth due to observation and circumnavigation.<sup>[207]</sup> The human perspective regarding the Earth has changed following the advent of spaceflight, and the biosphere is now widely viewed from a globally integrated perspective.<sup>[208][209]</sup> This is reflected in a growing environmental movement that is concerned about humankind's effects on the planet.<sup>[210]</sup>

## Notes

- [1] All astronomical quantities vary, both secularly and periodically. The quantities given are the values at the instant J2000.0 of the secular variation, ignoring all periodic variations.
- [2] aphelion =  $a \times (1 + e)$ ; perihelion =  $a \times (1 - e)$ , where  $a$  is the semi-major axis and  $e$  is the eccentricity.
- [3] Standish, E. Myles; Williams, James C.. "Orbital Ephemerides of the Sun, Moon, and Planets" (<http://iau-comm4.jpl.nasa.gov/XSChap8.pdf>) (PDF). International Astronomical Union Commission 4: (Ephemerides). . Retrieved 2010-04-03. See table 8.10.2. Calculation based upon 1 AU = 149,597,870,700(3) m.
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- [6] Allen, Clabon Walter; Cox, Arthur N. (2000). *Allen's Astrophysical Quantities* (<http://books.google.com/?id=w8PK2XFLH8C&pg=PA294>). Springer. p. 294. ISBN 0-387-98746-0. . Retrieved 2011-03-13.
- [7] The reference lists the longitude of the ascending node as  $-11.26064^\circ$ , which is equivalent to  $348.73936^\circ$  by the fact that any angle is equal to itself plus  $360^\circ$ .
- [8] The reference lists the longitude of perihelion, which is the sum of the longitude of the ascending node and the argument of perihelion. That is,  $114.20783^\circ + (-11.26064^\circ) = 102.94719^\circ$ .
- [9] US Space Command (March 1, 2001). "Reentry Assessment - US Space Command Fact Sheet" (<http://www.spaceref.com/news/viewpr.html?pid=4008>). SpaceRef Interactive. . Retrieved 2011-05-07.
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$$R_H = a \left( \frac{m}{3M} \right)^{\frac{1}{3}},$$

where  $m$  is the mass of the Earth,  $a$  is an Astronomical Unit, and  $M$  is the mass of the Sun. So the radius in A.U. is about:

$$\left( \frac{1}{3 \cdot 332,946} \right)^{\frac{1}{3}} = 0.01.$$

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- Earth - Profile (<http://solarsystem.nasa.gov/planets/profile.cfm?Object=Earth>) - Solar System Exploration (<http://solarsystem.nasa.gov/>) - NASA.
- Earth - Temperature and Precipitation Extremes (<http://www.ncdc.noaa.gov/oa/climate/globalextremes.html>) - NCDC.
- Earth - Climate Changes Cause Shape to Change (<http://www.nasa.gov/centers/goddard/earthandsun/earthshape.html>) - NASA.
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  - Video (01:02) (<http://www.youtube.com/watch?v=74mhQyuyELQ>) - Earth (Time-Lapse).
  - Video (00:27) (<http://www.youtube.com/watch?v=l6ahFFFQBZY>) - Earth and Auroras (Time-Lapse).

## Earth science

**Earth science** (also known as **geoscience**, **the geosciences** or **the Earth sciences**) is an all-embracing term for the sciences related to the planet Earth.<sup>[2]</sup> It is arguably a special case in planetary science, the Earth being the only known life-bearing planet. There are both reductionist and holistic approaches to Earth sciences. The formal discipline of Earth sciences may include the study of the atmosphere, hydrosphere, oceans and biosphere, as well as the solid earth. Typically Earth scientists will use tools from physics, chemistry, biology, chronology and mathematics to build a quantitative understanding of how the Earth system works, and how it evolved to its current state.



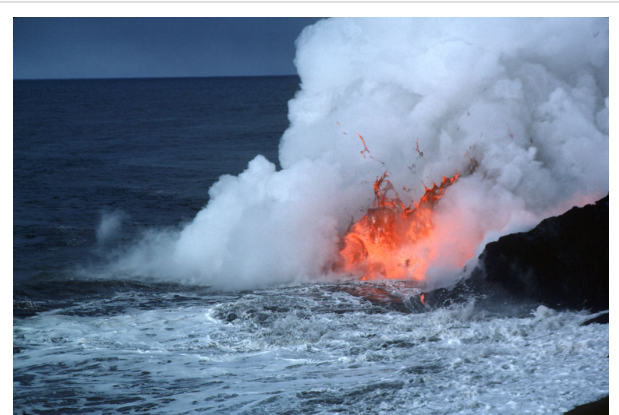
A volcano eruption is the release of stored energy from below the surface of Earth, originating from radioactive decay and gravitational sorting in the Earth's core and mantle, and residual energy gained during the Earth's formation.<sup>[1]</sup>



## Fields of study

The following fields of science are generally categorized within the geosciences:

- Geology describes the rocky parts of the Earth's crust (or lithosphere) and its historic development. Major subdisciplines are mineralogy and petrology, geochemistry, geomorphology, paleontology, stratigraphy, structural geology, engineering geology and sedimentology.<sup>[3][4]</sup>
- Physical geography covers aspects of geomorphology, soil study, hydrology, meteorology, climatology and biogeography.<sup>[5]</sup>
- Geophysics and geodesy investigate the shape of the Earth, its reaction to forces and its magnetic and gravity fields. Geophysicists explore the Earth's core and mantle as well as the tectonic and seismic activity of the lithosphere.<sup>[4][6][7]</sup>
- Soil science covers the outermost layer of the Earth's crust that is subject to soil formation processes (or pedosphere).<sup>[8]</sup> Major subdisciplines include edaphology and pedology.<sup>[9]</sup>
- Oceanography and hydrology (includes limnology) describe the marine and freshwater domains of the watery parts of the Earth (or hydrosphere). Major subdisciplines include hydrogeology and physical, chemical, and biological oceanography.
- Glaciology covers the icy parts of the Earth (or cryosphere).
- Atmospheric sciences cover the gaseous parts of the Earth (or atmosphere) between the surface and the exosphere (about 1000 km). Major subdisciplines are meteorology, climatology, atmospheric chemistry and atmospheric physics.



Lava flows from the Kilauea volcano into the ocean on the Island of Hawaii

## Earth's interior

Plate tectonics, mountain ranges, volcanoes, and earthquakes are geological phenomena that can be explained in terms of energy transformations in the Earth's crust.<sup>[10]</sup>

Beneath the Earth's crust lies the mantle which is heated by the radioactive decay of heavy elements. The mantle is not quite solid and consists of magma which is in a state of semi-perpetual convection. This convection process causes the lithospheric plates to move, albeit slowly. The resulting process is known as plate tectonics.<sup>[11][12][13][14]</sup>

Plate tectonics might be thought of as the process by which the earth is resurfaced. Through a process called seafloor spreading), new crust is created by the flow of magma from underneath the lithosphere to the surface, through fissures, where it cools and solidifies. Through a process called subduction, oceanic crust is pushed underground — beneath the rest of the lithosphere—where it comes into contact with magma and melts—rejoining the mantle from which it originally came.<sup>[12][14][15]</sup>

Areas of the crust where new crust is created are called *divergent boundaries*, those where it is brought back into the earth are *convergent boundaries* and those where plates slide past each other, but no new lithospheric material is created or destroyed, are referred to as transform (or conservative) boundaries.<sup>[12][14][16]</sup> Earthquakes result from the movement of the lithospheric plates, and they often occur near convergent boundaries where parts of the crust are forced into the earth as part of subduction.<sup>[17]</sup>

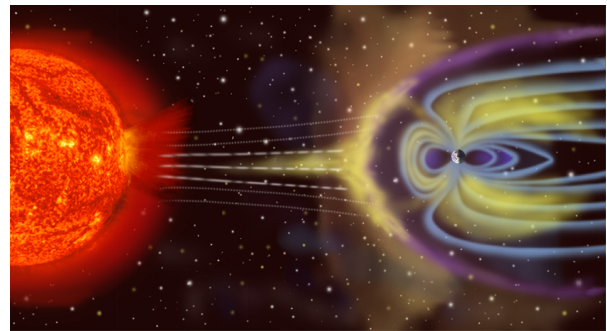
Volcanoes result primarily from the melting of subducted crust material. Crust material that is forced into the asthenosphere melts, and some portion of the melted material becomes light enough to rise to the surface—giving birth to volcanoes.<sup>[12][17]</sup>

## Earth's electromagnetic field

An electromagnet is a magnet that is created by a current that flows around a soft iron core.<sup>[18]</sup> Earth has a solid iron inner core surrounded by semi-liquid materials of the outer core that move in continuous currents around the inner core;<sup>[19]</sup> therefore, the Earth is an electromagnet. This is referred to as the dynamo theory of Earth's magnetism.<sup>[19][20]</sup>

## Atmosphere

The troposphere, stratosphere, mesosphere, thermosphere, and exosphere are the five layers which make up Earth's atmosphere. In all, the atmosphere is made up of about 78.0% nitrogen, 20.9% oxygen, and 0.92% argon. 75% of the gases in the atmosphere are located within the troposphere, the bottom-most layer. The remaining one percent of the atmosphere (all but the nitrogen, oxygen, and argon) contains small amounts of other gases including CO<sub>2</sub> and water vapors.<sup>[21]</sup> Water vapors and CO<sub>2</sub> allow the Earth's atmosphere to catch and hold the Sun's energy through a phenomenon called the greenhouse effect.<sup>[22]</sup> This allows Earth's surface to be warm enough to have liquid water and support life.



The magnetosphere shields the surface of Earth from the charged particles of the solar wind. It is compressed on the day (Sun) side due to the force of the arriving particles, and extended on the night side.

*Image not to scale.*

The magnetic field created by the internal motions of the core produces the magnetosphere which protects the Earth's atmosphere from the solar wind.<sup>[23]</sup> As the earth is 4.5 billion years old,<sup>[24]</sup> it would have lost its atmosphere by now if there were no protective magnetosphere.

In addition to storing heat, the atmosphere also protects living organisms by shielding the Earth's surface from cosmic rays. Note that the level of protection is high enough to prevent cosmic rays from destroying all life on Earth, yet low enough to aid the mutations that have an important role in pushing forward diversity in the biosphere.

## Methodology

Like all other scientists, Earth scientists apply the scientific method. They formulate hypotheses after observing events and gathering data about natural phenomena, and then they test hypotheses from such data.

A contemporary idea within earth science is uniformitarianism. Uniformitarianism says that "ancient geologic features are interpreted by understanding active processes that are readily observed". Simply stated, this means that geological processes occurring today happened in the past; The present is the key to the past. For example, a mountain need not be thought of as having been created in a moment, but instead it may be seen as the result of continuous subduction, causing magma to rise and form continental volcanic arcs.

## Earth's spheres

Earth science generally recognizes four spheres, the lithosphere, the hydrosphere, the atmosphere, and the biosphere;<sup>[25]</sup> these correspond to rocks, water, air, and life. Some practitioners include, as part of the spheres of the Earth, the cryosphere (corresponding to ice) as a distinct portion of the hydrosphere, as well as the pedosphere (corresponding to soil) as an active and intermixed sphere.

## Partial list of the major earth science topics

See: List of basic earth science topics

### Atmosphere

- Atmospheric chemistry
- Climatology
- Meteorology
  - Hydrometeorology
- Paleoclimatology

### Biosphere

- Biogeography
- Paleontology
  - Palynology
  - Micropaleontology
- Geomicrobiology
- Geoarchaeology

### Hydrosphere

- Hydrology
  - Geohydrology
- Limnology (freshwater science)
- Oceanography (marine science)
  - Chemical oceanography
  - Physical oceanography
  - Biological oceanography (marine biology)
  - Geological oceanography (marine geology)
    - Paleoceanography

### Lithosphere or geosphere

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- Geology
  - Economic geology
  - Engineering geology
  - Environmental geology
  - Historical geology
    - Quaternary geology
  - Planetary geology
  - Sedimentology
  - Stratigraphy
  - Structural geology
- Geography
  - Physical geography
- Geochemistry
- Geomorphology
- Geophysics
  - Geochronology
  - Geodynamics (see also Tectonics)
  - Geomagnetism
  - Gravimetry (also part of Geodesy)
  - Seismology
- Glaciology
- Hydrogeology
- Mineralogy
  - Crystallography
  - Gemology
- Petrology
- Speleology
- Volcanology

### **Pedosphere**

- Soil science
  - Edaphology
  - Pedology

### **Systems**

- Environmental science
- Geography
  - Human geography
  - Physical geography
- Gaia hypothesis

### **Others**

- Cartography
  - Geoinformatics (GIS)
  - Geostatistics
  - Geodesy and Surveying
  - NASA Earth Science Enterprise
-

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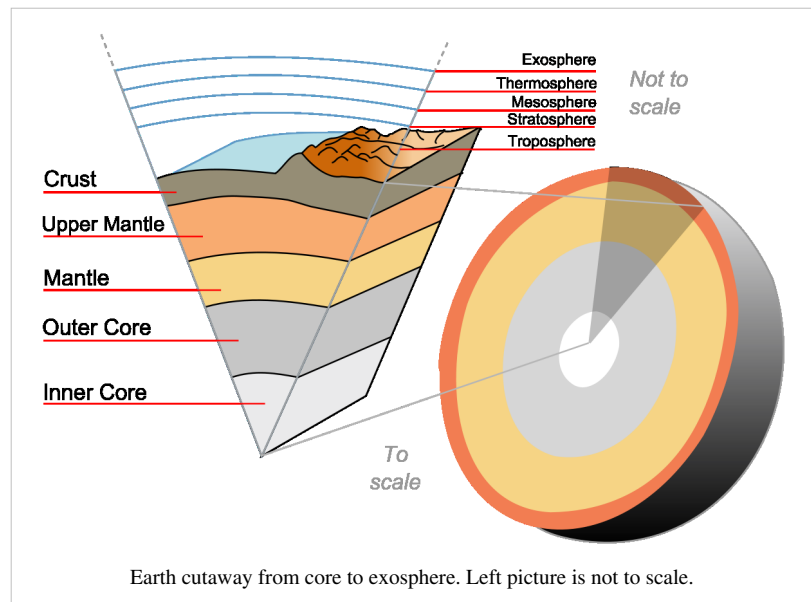
## External links

- Earth Science Picture of the Day (<http://epod.usra.edu/>), a service of Universities Space Research Association, sponsored by NASA Goddard Space Flight Center.
- Geoethics in Planetary and Space Exploration ([http://tierra.rediris.es/Geoethics\\_Planetary\\_Protection/](http://tierra.rediris.es/Geoethics_Planetary_Protection/)).
- (<http://www.nestanet.org/cms/content/welcome>), National Earth Science Teachers Association

# Structure of the Earth

The interior **structure of the Earth** is layered in spherical shells, like an onion. These layers can be defined by either their chemical or their rheological properties. The Earth has an outer silicate solid crust, a highly viscous mantle, a liquid outer core that is much less viscous than the mantle, and a solid inner core. Scientific understanding of Earth's internal structure is based on observations of topography and bathymetry, observations of rock in outcrop, samples brought to the surface from greater depths by volcanic activity, analysis of the seismic waves that pass

through the Earth, measurements of the gravity field of the Earth, and experiments with crystalline solids at pressures and temperatures characteristic of the Earth's deep interior.



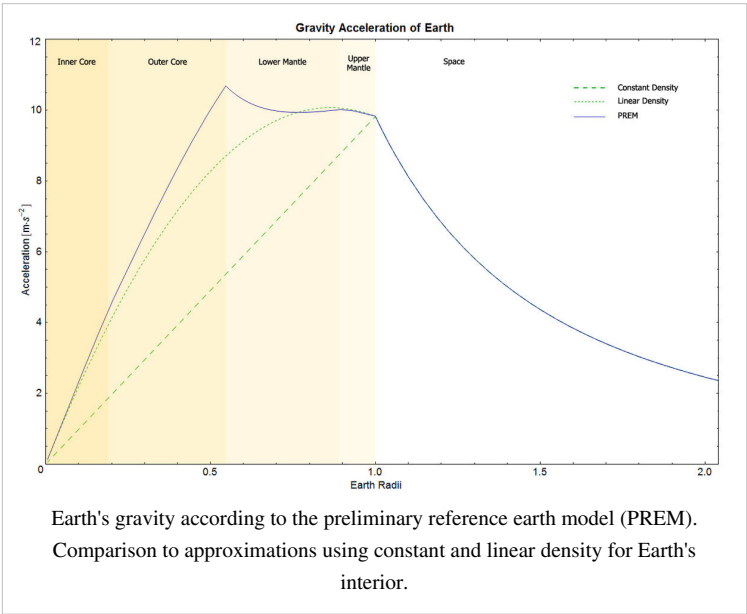
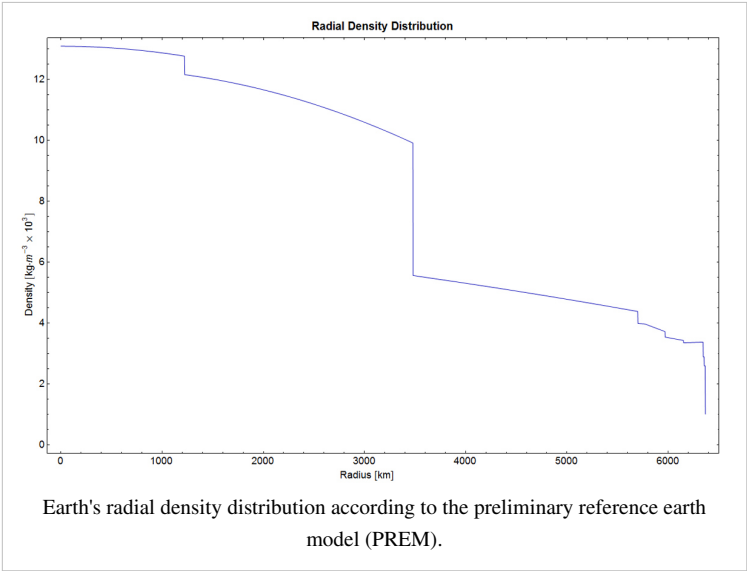


Assumptions

The force exerted by Earth's gravity can be used to calculate its mass, and by estimating the volume of the planet, its average density can be calculated. Astronomers can also calculate Earth's mass from its orbit and effects on nearby planetary bodies. Observations of rocks, bodies of water and atmosphere allow estimation of the mass, volume and density of rocks to a certain depth, so the remaining mass must be in the deeper layers.

Structure

The structure of Earth can be defined in two ways: by mechanical properties such as rheology, or chemically. Mechanically, it can be divided into lithosphere, asthenosphere, mesosphere, outer core, and the inner core. The interior of the earth is divided into 5 important layers. Chemically, Earth can be divided into the crust, upper mantle, lower mantle, outer core, and inner core. The geologic component layers of Earth<sup>[1]</sup> are at the following depths below the surface:



Depth		Layer
Kilometres	Miles	
0–60	0–37	Lithosphere (locally varies between 5 and 200 km)
0–35	0–22	... Crust (locally varies between 5 and 70 km)
35–60	22–37	... Uppermost part of mantle
35–2,890	22–1,790	Mantle
100–200	62–125	... Asthenosphere
35–660	22–410	... Upper mesosphere (upper mantle)
660–2,890	410–1,790	... Lower mesosphere (lower mantle)
2,890–5,150	1,790–3,160	Outer core
5,150–6,360	3,160–3,954	Inner core

The layering of Earth has been inferred indirectly using the time of travel of refracted and reflected seismic waves created by earthquakes. The core does not allow shear waves to pass through it, while the speed of travel (seismic velocity) is different in other layers. The changes in seismic velocity between different layers causes refraction owing to Snell's law. Reflections are caused by a large increase in seismic velocity and are similar to light reflecting from a mirror.

## Core

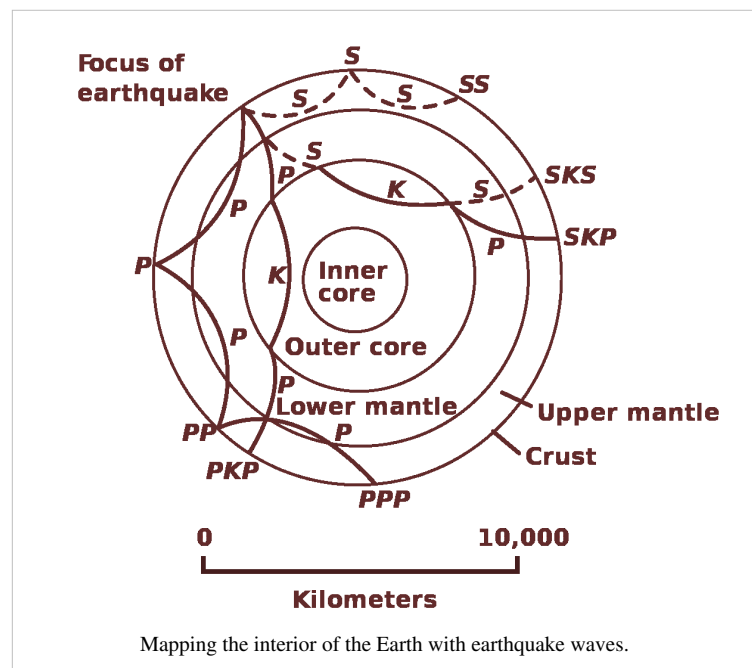
The average density of Earth is  $5,515 \text{ kg/m}^3$ .

Since the average density of surface material is only around  $3,000 \text{ kg/m}^3$ , we must

conclude that denser materials exist within Earth's core. Further evidence for the high density core comes from the study of seismology.

Seismic measurements show that the core is divided into two parts, a solid inner core with a radius of  $\sim 1,220 \text{ km}$ <sup>[2]</sup> and a liquid outer core extending beyond it to a radius of  $\sim 3,400 \text{ km}$ . The solid inner core was discovered in 1936 by Inge Lehmann and is generally believed to be composed primarily of iron and some nickel. In early stages of Earth's formation about 4.5 billion ( $4.5 \times 10^9$ ) years ago, melting would have caused denser substances to sink toward the center in a process called planetary differentiation (see also the iron catastrophe), while less-dense materials would have migrated to the crust. The core is thus believed to largely be composed of iron (80%), along with nickel and one or more light elements, whereas other dense elements, such as lead and uranium, either are too rare to be significant or tend to bind to lighter elements and thus remain in the crust (see felsic materials). Some have argued that the inner core may be in the form of a single iron crystal.<sup>[3][4]</sup>

On August 30, 2011, Professor Kei Hirose, professor of high-pressure mineral physics and petrology at the Tokyo Institute of Technology, became the first person to recreate conditions found at the earth's core under laboratory conditions, subjecting a sample of iron nickel alloy to the same type of pressure by gripping it in a vise between 2



diamond tips, and then heating the sample to approximately 4000 Kelvins with a laser. The sample was observed with x-rays, and strongly supported the theory that the earth's inner core was made of giant crystals running north to south.<sup>[5][6]</sup>

The liquid outer core surrounds the inner core and is believed to be composed of iron mixed with nickel and trace amounts of lighter elements.

Recent speculation suggests that the innermost part of the core is enriched in gold, platinum and other siderophile elements.<sup>[7]</sup>

The matter that comprises Earth is connected in fundamental ways to matter of certain chondrite meteorites, and to matter of outer portion of the Sun.<sup>[8][9]</sup> There is good reason to believe that Earth is, in the main, like a chondrite meteorite. Beginning as early as 1940, scientists, including Francis Birch, built geophysics upon the premise that Earth is like ordinary chondrites, the most common type of meteorite observed impacting Earth, while totally ignoring another, albeit less abundant type, called enstatite chondrites. The principal difference between the two meteorite types is that enstatite chondrites formed under circumstances of extremely limited available oxygen, leading to certain normally oxyphile elements existing either partially or wholly in the alloy portion that corresponds to the core of Earth.

Dynamo theory suggests that convection in the outer core, combined with the Coriolis effect, gives rise to Earth's magnetic field. The solid inner core is too hot to hold a permanent magnetic field (see Curie temperature) but probably acts to stabilize the magnetic field generated by the liquid outer core. The average magnetic field strength in the Earth's outer core is estimated to be 25 Gauss, 50 times stronger than the magnetic field at the surface.<sup>[10][11]</sup>

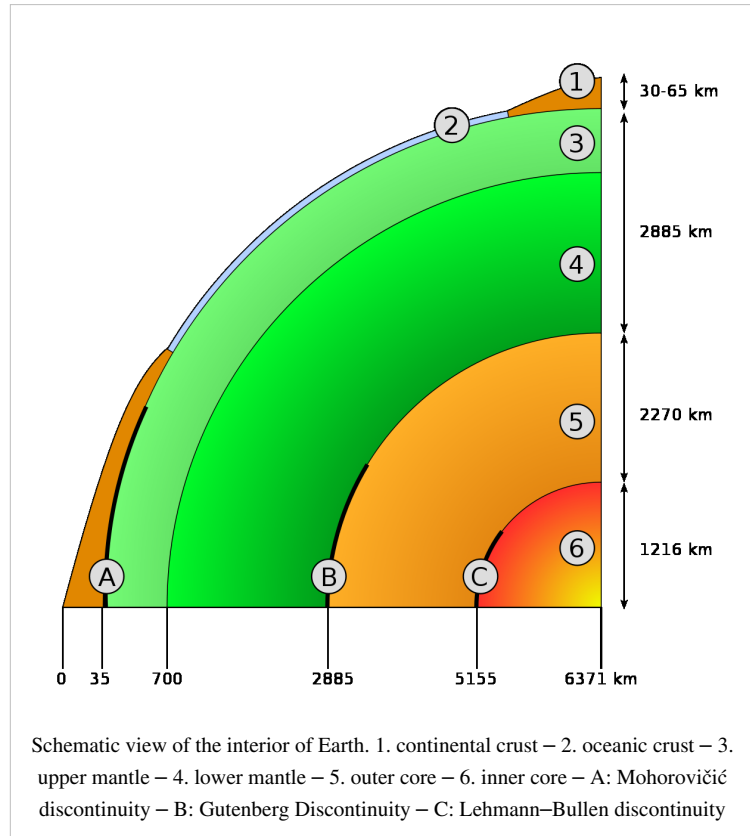
Recent evidence has suggested that the inner core of Earth may rotate slightly faster than the rest of the planet.<sup>[12]</sup> In August 2005 a team of geophysicists announced in the journal *Science* that, according to their estimates, Earth's inner core rotates approximately 0.3 to 0.5 degrees per year relative to the rotation of the surface.<sup>[13][14]</sup>

The current scientific explanation for the Earth's temperature gradient is a combination of heat left over from the planet's initial formation, decay of radioactive elements, and freezing of the inner core.

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## Mantle

Earth's mantle extends to a depth of 2,890 km, making it the thickest layer of the Earth. The pressure, at the bottom of the mantle, is  $\sim 140$  GPa (1.4 Matm). The mantle is composed of silicate rocks that are rich in iron and magnesium relative to the overlying crust. Although solid, the high temperatures within the mantle cause the silicate material to be sufficiently ductile that it can flow on very long timescales. Convection of the mantle is expressed at the surface through the motions of tectonic plates. The melting point and viscosity of a substance depends on the pressure it is under. As there is intense and increasing pressure as one travels deeper into the mantle, the lower part of the mantle flows less easily than does the upper mantle (chemical changes within the mantle may also be important). The viscosity of the mantle ranges between  $10^{21}$  and  $10^{24}$  Pa·s, depending on depth.<sup>[15]</sup> In comparison, the viscosity of water is approximately  $10^{-3}$  Pa·s and that of pitch is  $10^7$  Pa·s.



## Crust

The crust ranges from 5–70 km in depth and is the outermost layer. The thin parts are the oceanic crust, which underlie the ocean basins (5–10 km) and are composed of dense (mafic) iron magnesium silicate rocks, like basalt. The thicker crust is continental crust, which is less dense and composed of (felsic) sodium potassium aluminium silicate rocks, like granite. The rocks of the crust fall into two major categories – sial and sima (Suess, 1831–1914). It is estimated that sima starts about 11 km below the Conrad discontinuity (a second order discontinuity). The uppermost mantle together with the crust constitutes the lithosphere. The crust-mantle boundary occurs as two physically different events. First, there is a discontinuity in the seismic velocity, which is known as the Mohorovičić discontinuity or Moho. The cause of the Moho is thought to be a change in rock composition from rocks containing plagioclase feldspar (above) to rocks that contain no feldspars (below). Second, in oceanic crust, there is a chemical discontinuity between ultramafic cumulates and tectonized harzburgites, which has been observed from deep parts of the oceanic crust that have been obducted onto the continental crust and preserved as ophiolite sequences.

Many rocks now making up Earth's crust formed less than 100 million ( $1 \times 10^8$ ) years ago; however, the oldest known mineral grains are 4.4 billion ( $4.4 \times 10^9$ ) years old, indicating that Earth has had a solid crust for at least that long.<sup>[16]</sup>

## Historical development of alternative conceptions

In 1692 Edmund Halley (in a paper printed in *Philosophical Transactions of Royal Society of London*) put forth the idea of Earth consisting of a hollow shell about 500 miles thick, with two inner concentric shells around an innermost core, corresponding to the diameters of the planets Venus, Mars, and Mercury respectively.<sup>[17]</sup> Halley's construct was a method of accounting for the (flawed) values of the relative density of Earth and the Moon that had been given by Sir Isaac Newton, in *Principia* (1687). "Sir Isaac Newton has demonstrated the Moon to be more solid than our Earth, as 9 to 5," Halley remarked; "why may we not then suppose four ninths of our globe to be cavity?"<sup>[17]</sup>



Edmond Halley's hypothesis.

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## Further reading

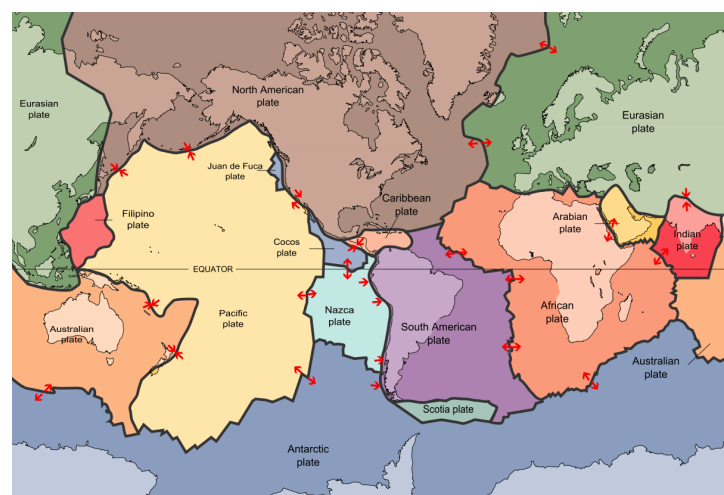
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## Plate tectonics

**Plate tectonics** (from the Late Latin *tectonicus*, from the Greek: τεκτονικός "pertaining to building")<sup>[2]</sup> is a scientific theory that describes the large-scale motions of Earth's lithosphere. The theory builds on the concepts of continental drift, developed during the first decades of the 20th century. It was accepted by the geoscientific community after the concepts of seafloor spreading were developed in the late 1950s and early 1960s.

The lithosphere is broken up into tectonic plates. On Earth, there are seven or eight major plates (depending on how they are defined) and many minor plates. Where plates meet, their relative motion determines the type of boundary: convergent, divergent, or transform. Earthquakes, volcanic activity, mountain-building, and oceanic trench formation occur along these plate boundaries. The lateral relative movement of the plates typically varies from zero to 100 mm annually.<sup>[3]</sup>

Tectonic plates are composed of oceanic lithosphere and thicker continental lithosphere, each topped by its own kind of crust. Along convergent boundaries, subduction carries plates into the mantle; the material lost is roughly balanced by the formation of new (oceanic) crust along divergent margins by seafloor spreading. In this way, the total surface of the globe remains the same. This prediction of plate tectonics is also referred to as the conveyor belt principle.



The tectonic plates of the world were mapped in the second half of the 20th century.

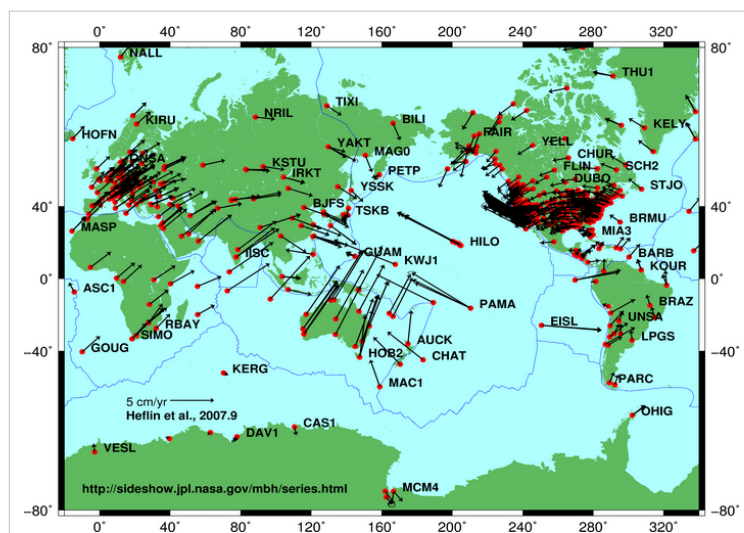


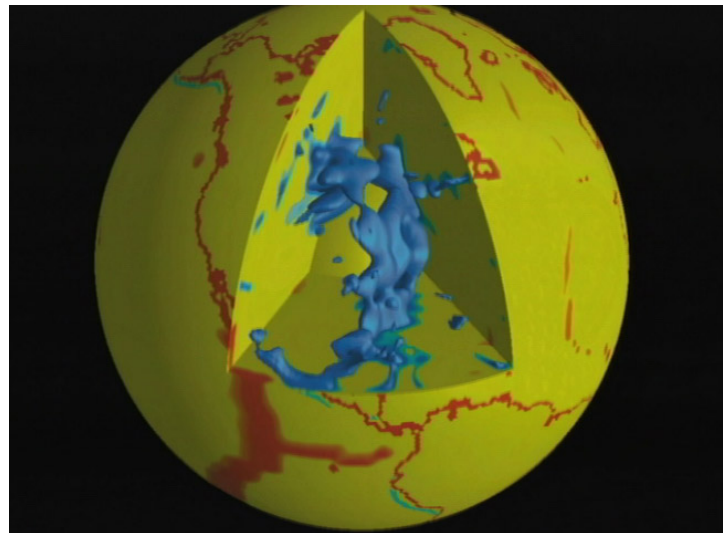
Plate motion based on Global Positioning System (GPS) satellite data from NASA JPL<sup>[1]</sup>. The vectors show direction and magnitude of motion.

Earlier theories (that still have



some supporters) proposed gradual shrinking (contraction) or gradual expansion of the globe.<sup>[4]</sup>

Tectonic plates are able to move because the Earth's lithosphere has a higher strength and lower density than the underlying asthenosphere. Lateral density variations in the mantle result in convection. Their movement is thought to be driven by a combination of the motion of seafloor away from the spreading ridge (due to variations in topography and density of the crust that result in differences in gravitational forces) and drag, downward suction, at the subduction zones. A different explanation lies in different forces generated by the rotation of the globe and tidal forces of the Sun and the Moon. The relative importance of each of these factors is unclear, and is still subject to debate (see also below).



Remnants of the Farallon Plate, deep in Earth's mantle. It is thought that much of the plate initially went under North America (particularly the western United States and southwest Canada) at a very shallow angle, creating much of the mountainous terrain in the area (particularly the southern Rocky Mountains).

## Key principles

The outer layers of the Earth are divided into lithosphere and asthenosphere. This is based on differences in mechanical properties and in the method for the transfer of heat. Mechanically, the lithosphere is cooler and more rigid, while the asthenosphere is hotter and flows more easily. In terms of heat transfer, the lithosphere loses heat by conduction whereas the asthenosphere also transfers heat by convection and has a nearly adiabatic temperature gradient. This division should not be confused with the *chemical* subdivision of these same layers into the mantle (comprising both the asthenosphere and the mantle portion of the lithosphere) and the crust: a given piece of mantle may be part of the lithosphere or the asthenosphere at different times, depending on its temperature and pressure.

The key principle of plate tectonics is that the lithosphere exists as separate and distinct *tectonic plates*, which ride on the fluid-like (visco-elastic solid) asthenosphere. Plate motions range up to a typical 10–40 mm/a (Mid-Atlantic Ridge; about as fast as fingernails grow), to about 160 mm/a (Nazca Plate; about as fast as hair grows).<sup>[5]</sup> The driving mechanism behind this movement is described separately below.

Tectonic lithosphere plates consist of lithospheric mantle overlain by either or both of two types of crustal material: oceanic crust (in older texts called *sima* from silicon and magnesium) and continental crust (*sial* from silicon and aluminium). Average oceanic lithosphere is typically 100 km (**unknown operator: u'strong'** mi) thick;<sup>[6]</sup> its thickness is a function of its age: as time passes, it conductively cools and becomes thicker. Because it is formed at mid-ocean ridges and spreads outwards, its thickness is therefore a function of its distance from the mid-ocean ridge where it was formed. For a typical distance oceanic lithosphere must travel before being subducted, the thickness varies from about 6 km (**unknown operator: u'strong'** mi) thick at mid-ocean ridges to greater than 100 km (**unknown operator: u'strong'** mi) at subduction zones; for shorter or longer distances, the subduction zone (and therefore also the mean) thickness becomes smaller or larger, respectively.<sup>[7]</sup> Continental lithosphere is typically ~200 km thick, though this also varies considerably between basins, mountain ranges, and stable cratonic interiors of continents. The two types of crust also differ in thickness, with continental crust being considerably thicker than oceanic (35 km vs. 6 km).<sup>[8]</sup>

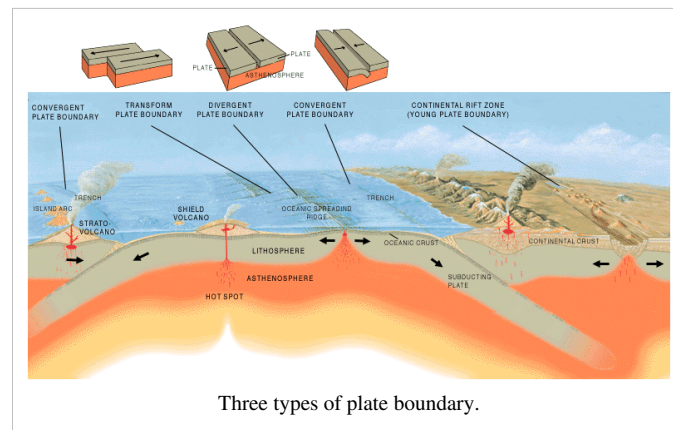
The location where two plates meet is called a *plate boundary*, and plate boundaries are commonly associated with geological events such as earthquakes and the creation of topographic features such as mountains, volcanoes, mid-ocean ridges, and oceanic trenches. The majority of the world's active volcanoes occur along plate boundaries, with the Pacific Plate's Ring of Fire being most active and most widely known. These boundaries are discussed in further detail below. Some volcanoes occur in the interiors of plates, and these have been variously attributed to internal plate deformation<sup>[9]</sup> and to mantle plumes.

As explained above, tectonic plates can include continental crust or oceanic crust, and many plates contain both. For example, the African Plate includes the continent and parts of the floor of the Atlantic and Indian Oceans. The distinction between oceanic crust and continental crust is based on their modes of formation. Oceanic crust is formed at sea-floor spreading centers, and continental crust is formed through arc volcanism and accretion of terranes through tectonic processes; though some of these terranes may contain ophiolite sequences, which are pieces of oceanic crust, these are considered part of the continent when they exit the standard cycle of formation and spreading centers and subduction beneath continents. Oceanic crust is also denser than continental crust owing to their different compositions. Oceanic crust is denser because it has less silicon and more heavier elements ("mafic") than continental crust ("felsic").<sup>[10]</sup> As a result of this density stratification, oceanic crust generally lies below sea level (for example most of the Pacific Plate), while the continental crust buoyantly projects above sea level (see the page isostasy for explanation of this principle).

## Types of plate boundaries

Basically, three types of plate boundaries exist,<sup>[11]</sup> with a fourth, mixed type, characterized by the way the plates move relative to each other. They are associated with different types of surface phenomena. The different types of plate boundaries are:<sup>[12][13]</sup>

1. *Transform boundaries (Conservative)* occur where plates slide or, perhaps more accurately, grind past each other along transform faults. The relative motion of the two plates is either sinistral (left side toward the observer) or dextral (right side toward the observer). The San Andreas Fault in California is an example of a transform boundary exhibiting dextral motion.
2. *Divergent boundaries (Constructive)* occur where two plates slide apart from each other. Mid-ocean ridges (e.g., Mid-Atlantic Ridge) and active zones of rifting (such as Africa's East African Rift) are both examples of divergent boundaries.
3. *Convergent boundaries (Destructive)* (or *active margins*) occur where two plates slide towards each other commonly forming either a subduction zone (if one plate moves underneath the other) or a continental collision (if the two plates contain continental crust). Deep marine trenches are typically associated with subduction zones, and the basins that develop along the active boundary are often called "foreland basins". The subducting slab contains many hydrous minerals, which release their water on heating; this water then causes the mantle to melt, producing volcanism. Examples of this are the Andes mountain range in South America and the Japanese island arc.
4. *Plate boundary zones* occur where the effects of the interactions are unclear and the boundaries, usually occurring along a broad belt, are not well defined, and may show various types of movements in different episodes.



## Driving forces of plate motion

Plate tectonics is basically a kinematic phenomenon: Earth scientists agree upon the observation and deduction that the plates have moved one with respect to the other, and debate and find agreements on how and when. But still a major question remains on what the motor behind this movement is; the geodynamic mechanism, and here science diverges in different theories.

Generally, it is accepted that tectonic plates are able to move because of the relative density of oceanic lithosphere and the relative weakness of the asthenosphere. Dissipation of heat from the mantle is acknowledged to be the original source of energy driving plate tectonics, through convection or large scale upwelling and doming. As a consequence, in the current view, although it is still a matter of some debate, because of the excess density of the oceanic lithosphere sinking in subduction zones a powerful source of plate motion is generated. When the new crust forms at mid-ocean ridges, this oceanic lithosphere is initially less dense than the underlying asthenosphere, but it becomes denser with age, as it conductively cools and thickens. The greater density of old lithosphere relative to the underlying asthenosphere allows it to sink into the deep mantle at subduction zones, providing most of the driving force for plate motions. The weakness of the asthenosphere allows the tectonic plates to move easily towards a subduction zone.<sup>[14]</sup> Although subduction is believed to be the strongest force driving plate motions, it cannot be the only force since there are plates such as the North American Plate which are moving, yet are nowhere being subducted. The same is true for the enormous Eurasian Plate. The sources of plate motion are a matter of intensive research and discussion among earth scientists. One of the main points is that the kinematic pattern of the movements itself should be separated clearly from the possible geodynamic mechanism that is invoked as the driving force of the observed movements, as some patterns may be explained by more than one mechanism.<sup>[15]</sup> Basically, the driving forces that are advocated at the moment, can be divided in three categories: mantle dynamics related, gravity related (mostly secondary forces), and Earth rotation related.

## Driving forces related to mantle dynamics

For a considerable period of around 25 years (last quarter of the twentieth century) the leading theory envisaged large scale convection currents in the upper mantle which are transmitted through the asthenosphere as the main driving force of the tectonic plates. This theory was launched by Arthur Holmes and some forerunners in the 1930s and was immediately recognized as the solution for the acceptance of the theory discussed since its occurrence in the papers of Alfred Wegener in the early years of the century. It was, though, long debated because the leading ("fixist") theory was still envisaging a static Earth without moving continents, up until the major break-throughs in the early sixties.

Two- and three-dimensional imaging of the Earth's interior (seismic tomography) shows that there is a laterally varying density distribution throughout the mantle. Such density variations can be material (from rock chemistry), mineral (from variations in mineral structures), or thermal (through thermal expansion and contraction from heat

energy). The manifestation of this varying lateral density is mantle convection from buoyancy forces.<sup>[16]</sup>

How mantle convection relates directly and indirectly to the motion of the plates is a matter of ongoing study and discussion in geodynamics. Somehow, this energy must be transferred to the lithosphere for tectonic plates to move. There are essentially two types of forces that are thought to influence plate motion: friction and gravity.

- **Basal drag** (friction): The plate motion is in this way driven by friction between the convection currents in the asthenosphere and the more rigid overlying floating lithosphere.
- **Slab suction** (gravity): Local convection currents exert a downward frictional pull on plates in subduction zones at ocean trenches. Slab suction may occur in a geodynamic setting wherein basal tractions continue to act on the plate as it dives into the mantle (although perhaps to a greater extent acting on both the under and upper side of the slab).

Lately, the convection theory is much debated as modern techniques based on 3D seismic tomography of imaging the internal structure of the Earth's mantle still fail to recognize these predicted large scale convection cells. Therefore, alternative views have been proposed:

In the theory of plume tectonics developed during the 1990s, a modified concept of mantle convection currents is used, related to super plumes rising from the deeper mantle which would be the drivers or the substitutes of the major convection cells. These ideas, which find their roots in the early 1930s with the so-called "fixistic" ideas of the European and Russian Earth Science Schools, find resonance in the modern theories which envisage hot spots/mantle plumes in the mantle which remain fixed and are overridden by oceanic and continental lithosphere plates during time, and leave their traces in the geological record (though these phenomena are not invoked as real driving mechanisms, but rather as a modulator). The modern theories that continue building on the older mantle doming concepts and see the movements of the plates a secondary phenomena, are beyond the scope of this page and are discussed elsewhere for example on the plume tectonics page.

Another suggestion is that the mantle flows neither in cells nor large plumes, but rather as a series of channels just below the Earth's crust which then provide basal friction to the lithosphere. This theory is called "surge tectonics" and became quite popular in geophysics and geodynamics during the 1980s and 1990s.<sup>[17]</sup>

## Driving forces related to gravity

Gravity related forces are usually invoked as secondary phenomena within the framework of a more general driving mechanism such as the various forms of mantle dynamics described above.

Gravitational sliding away from a spreading ridge: According to many authors, plate motion is driven by the higher elevation of plates at ocean ridges.<sup>[18]</sup> As oceanic lithosphere is formed at spreading ridges from hot mantle material, it gradually cools and thickens with age (and thus distance from the ridge). Cool oceanic lithosphere is significantly denser than the hot mantle material from which it is derived and so with increasing thickness it gradually subsides into the mantle to compensate the greater load. The result is a slight lateral incline with distance from the ridge axis.

This force is regarded as a secondary force often referred to as "ridge push". This is a misnomer as nothing is "pushing" horizontally and tensional features are dominant along ridges. It is more accurate to refer to this mechanism as gravitational sliding as variable topography across the totality of the plate can vary considerably and the topography of spreading ridges is only the most prominent feature. Other mechanisms generating this gravitational secondary force include flexural bulging of the lithosphere before it dives underneath an adjacent plate, which produces a clear topographical feature that can offset or at least affect the influence of topographical ocean ridges, and mantle plumes and hot spots, which are postulated to impinge on the underside of tectonic plates.

Slab-pull: Current scientific opinion is that the asthenosphere is insufficiently competent or rigid to directly cause motion by friction along the base of the lithosphere. Slab pull is therefore most widely thought to be the greatest force acting on the plates. In this current understanding, plate motion is mostly driven by the weight of cold, dense plates sinking into the mantle at trenches.<sup>[19]</sup> Recent models indicate that trench suction plays an important role as

well. However, as the North American Plate is nowhere being subducted, yet it is in motion presents a problem. The same holds for the African, Eurasian, and Antarctic plates.

Gravitational sliding away from mantle doming: According to older theories one of the driving mechanisms of the plates is the existence of large scale asthenosphere/mantle domes, which cause the gravitational sliding of lithosphere plates away from them. This gravitational sliding represents a secondary phenomenon of this, basically vertically oriented mechanism. This can act on various scales, from the small scale of one island arc up to the larger scale of an entire ocean basin.<sup>[20]</sup>

## Driving forces related to Earth rotation

Alfred Wegener, being a meteorologist, had proposed tidal forces and pole flight force as main driving mechanisms for continental drift. However, these forces were considered far too small to cause continental motion as the concept then was of continents plowing through oceanic crust.<sup>[21]</sup> Therefore, Wegener converted to convection currents as the main driving force in the last edition of his book in 1929.

In the plate tectonics context (accepted since the seafloor spreading proposals of Heezen, Hess, Dietz, Morley, Vine and Matthews (see below) during the early 1960s) though, oceanic crust is in motion *with* the continents which caused the proposals related to Earth rotation to be reconsidered. In more recent literature, these driving forces are:

1. Tidal drag due to the gravitational force the Moon (and the Sun) exerts on the crust of the Earth<sup>[22]</sup>
2. Shear strain of the Earth globe due to N-S compression related to the rotation and modulations of it;
3. Pole flight force: equatorial drift due to rotation and centrifugal effects: tendency of the plates to move from the poles to the equator ("*Polflucht*");
4. Coriolis effect acting on plates when they move around the globe;
5. Global deformation of the geoid due to small displacements of rotational pole with respect to the Earth crust;
6. Other smaller deformation effects of the crust due to wobbles and spin movements of the Earth rotation on a smaller time scale.

For these mechanisms to be overall valid, systematic relationships should exist all over the globe between the orientation and kinematics of deformation, and the geographical latitudinal and longitudinal grid of the Earth itself. Ironically, these systematic relations studies in the second half of the nineteenth century and the first half of the twentieth century do underline exactly the opposite: that the plates had not moved in time, that the deformation grid was fixed with respect to the Earth equator and axis, and that gravitational driving forces were generally acting vertically and caused only locally horizontal movements (the so-called pre-plate tectonic, "fixist theories"). Later studies (discussed below on this page) therefore invoked many of the relationships recognized during this pre-plate tectonics period, to support their theories (see the anticipations and reviews in the work of van Dijk and collaborators.<sup>[23]</sup>

Of the many forces discussed in this paragraph, tidal force is still highly debated and defended as a possible principle driving force, whereas the other forces are used or in global geodynamic models not using the plate tectonics concepts (therefore beyond the discussions treated in this section), or proposed as minor modulations within the overall plate tectonics model.

In 1973 George W. Moore<sup>[24]</sup> of the USGS and R. C. Bostrom<sup>[25]</sup> presented evidence for a general westward drift of the Earth's lithosphere with respect to the mantle, and, therefore, tidal forces or tidal lag or "friction" due to the Earth's rotation and the forces acting upon it by the Moon being a driving force for plate tectonics: as the Earth spins eastward beneath the moon, the moon's gravity ever so slightly pulls the Earth's surface layer back westward, just like proposed by Alfred Wegener (see above). In a more recent 2006 study,<sup>[26]</sup> scientists reviewed and advocated these earlier proposed ideas. It has also been suggested recently in Lovett (2006) that this observation may also explain why Venus and Mars have no plate tectonics, since Venus has no moon and Mars' moons are too small to have significant tidal effects on Mars. In a recent paper<sup>[27]</sup> it was suggested that, on the other hand, it can easily be observed that many plates are moving north and eastward, and that the dominantly westward motion of the Pacific

ocean basins derives simply from the eastward bias of the Pacific spreading center (which is not a predicted manifestation of such lunar forces). In the same paper the authors admit, however, that relative to the lower mantle, there is a slight westward component in the motions of all the plates. They demonstrated though that the westward drift, seen only for the past 30 Ma, is attributed to the increased dominance of the steadily growing and accelerating Pacific plate. The debate is still open.

### Relative significance of each driving force mechanism

The actual vector of a plate's motion must necessarily be a function of all the forces acting upon the plate. However, therein remains the problem regarding what degree each process contributes to the motion of each tectonic plate.

The diversity of geodynamic settings and properties of each plate must clearly result in differences in the degree to which such processes are actively driving the plates. One method of dealing with this problem is to consider the relative rate at which each plate is moving and to consider the available evidence of each driving force upon the plate as far as possible.

One of the most significant correlations found is that lithospheric plates attached to downgoing (subducting) plates move much faster than plates not attached to subducting plates. The Pacific plate, for instance, is essentially surrounded by zones of subduction (the so-called Ring of Fire) and moves much faster than the plates of the Atlantic basin, which are attached (perhaps one could say 'welded') to adjacent continents instead of subducting plates. It is thus thought that forces associated with the downgoing plate (slab pull and slab suction) are the driving forces which determine the motion of plates, except for those plates which are not being subducted.<sup>[19]</sup> The driving forces of plate motion continue to be active subjects of on-going research within geophysics and tectonophysics.

### Development of the theory

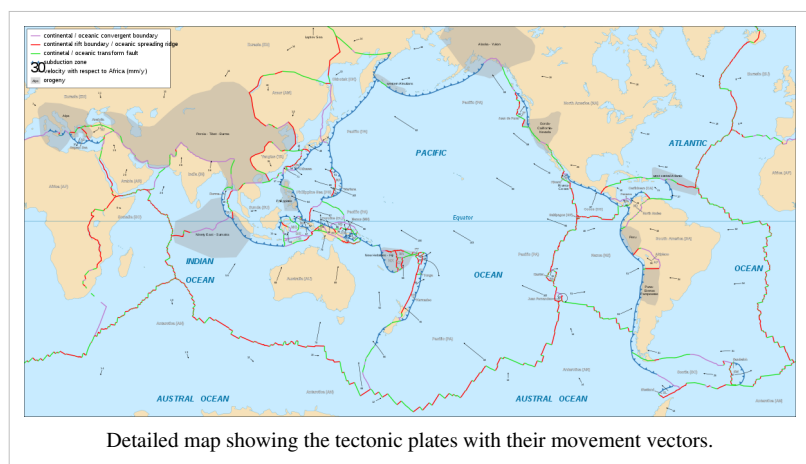
Further information: Timeline of the development of tectonophysics

Plate tectonics is the main current theory in Earth Sciences regarding the development of our planet Earth. It is, therefore, appropriate to dedicate some space to explain how the Earth Science community, step by step, has built this theory, from early speculations, through the gathering of proof and severe debates, up to the refinement and quantification, and still ongoing confrontations with alternative ideas.

### Summary

In line with other previous and contemporaneous proposals, in 1912 the meteorologist Alfred Wegener amply described what he called continental drift, expanded in his 1915 book *The Origin of Continents and Oceans*<sup>[28]</sup> and the scientific debate started that would end up fifty years later in the theory of plate tectonics.<sup>[29]</sup> Starting from the idea (also expressed by his forerunners) that the present continents once formed a single land mass (which was called Pangea later on) that drifted apart, thus releasing the continents from the Earth's mantle and likening them to "icebergs" of low

density granite floating on a sea of denser basalt.<sup>[30]</sup> Supporting evidence for the idea came from the dove-tailing outlines of South America's east coast and Africa's west coast, and from the matching of the rock formations along





these edges. Confirmation of their previous contiguous nature also came from the fossil plants *Glossopteris* and *Gangamopteris*, and the therapsid or mammal-like reptile *Lystrosaurus*, all widely distributed over South America, Africa, Antarctica, India and Australia. The evidence for such an erstwhile joining of these continents was patent to field geologists working in the southern hemisphere. The South African Alex du Toit put together a mass of such information in his 1937 publication *Our Wandering Continents*, and went further than Wegener in recognising the strong links between the Gondwana fragments.

But without detailed evidence and a force sufficient to drive the movement, the theory was not generally accepted: the Earth might have a solid crust and mantle and a liquid core, but there seemed to be no way that portions of the crust could move around. Distinguished scientists, such as Harold Jeffreys and Charles Schuchert, were outspoken critics of continental drift.

Despite much opposition, the view of continental drift gained support and a lively debate started between "drifters" or "mobilists" (proponents of the theory) and "fixists" (opponents). During the 1920s, 1930s and 1940s, the former reached important milestones proposing that convection currents might have driven the plate movements, and that spreading may have occurred below the sea within the oceanic crust. Concepts close to the elements now incorporated in plate tectonics were proposed by geophysicists and geologists (both fixists and mobilists) like Vening-Meinesz, Holmes, and Umbgrove.

One of the first pieces of geophysical evidence that was used to support the movement of lithospheric plates came from paleomagnetism. This is based on the fact that rocks of different ages show a variable magnetic field direction, evidenced by studies since the mid-nineteenth century. The magnetic north and south poles reverse through time, and, especially important in paleotectonic studies, the relative position of the magnetic north pole varies through time. Initially, during the first half of the twentieth century, the latter phenomenon was explained by introducing what was called "polar wander" (see apparent polar wander), i.e., it was assumed that the north pole location had been shifting through time. An alternative explanation, though, was that the continents had moved (shifted and rotated) relative to the north pole, and each continent, in fact, shows its own "polar wander path". During the late 1950s it was successfully shown on two occasions that these data could show the validity of continental drift: by Keith Runcorn in a paper in 1956,<sup>[31]</sup> and by Warren Carey in a symposium held in March 1956.<sup>[32]</sup>

The second piece of evidence in support of continental drift came during the late 1950s and early 60s from data on the bathymetry of the deep ocean floors and the nature of the oceanic crust such as magnetic properties and, more generally, with the development of marine geology<sup>[33]</sup> which gave evidence for the association of seafloor spreading along the mid-oceanic ridges and magnetic field reversals, published between 1959 and 1963 by Heezen, Dietz, Hess, Mason, Vine & Matthews, and Morley.<sup>[34]</sup>

Simultaneous advances in early seismic imaging techniques in and around Wadati-Benioff zones along the trenches bounding many continental margins, together with many other geophysical (e.g. gravimetric) and geological observations, showed how the oceanic crust could disappear into the mantle, providing the mechanism to balance the extension of the ocean basins with shortening along its margins.

All this evidence, both from the ocean floor and from the continental margins, made it clear around 1965 that continental drift was feasible and the theory of plate tectonics, which was defined in a series of papers between 1965 and 1967, was born, with all its extraordinary explanatory and predictive power. The theory revolutionized the Earth sciences, explaining a diverse range of geological phenomena and their implications in other studies such as paleogeography and paleobiology.

## Continental drift

In the late 19th and early 20th centuries, geologists assumed that the Earth's major features were fixed, and that most geologic features such as basin development and mountain ranges could be explained by vertical crustal movement, described in what is called the geosynclinal theory. Generally, this was placed in the context of a contracting planet Earth due to heat loss in the course of a relatively short geological time.

It was observed as early as 1596 that the opposite coasts of the Atlantic Ocean—or, more precisely, the edges of the continental shelves—have similar shapes and seem to have once fitted together.<sup>[35]</sup>

Since that time many theories were proposed to explain this apparent complementarity, but the assumption of a solid Earth made these various proposals difficult to accept.<sup>[36]</sup>

The discovery of radioactivity and its associated heating properties in 1895 prompted a re-examination of the apparent age of the Earth.<sup>[37]</sup> since this had previously been estimated by its cooling rate and assumption the Earth's surface radiated like a black body.<sup>[38]</sup> Those calculations

had implied that, even if it started at red heat, the Earth would have dropped to its present temperature in a few tens of millions of years. Armed with the knowledge of a new heat source, scientists realized that the Earth would be much older, and that its core was still sufficiently hot to be liquid.

By 1915, after having published a first article in 1912,<sup>[39]</sup> Alfred Wegener was making serious arguments for the idea of continental drift in the first edition of *The Origin of Continents and Oceans*.<sup>[28]</sup> In that book (re-issued in four successive editions up to the final one in 1936), he noted how the east coast of South America and the west coast of Africa looked as if they were once attached. Wegener was not the first to note this (Abraham Ortelius, Snider-Pellegrini, Eduard Suess, Roberto Mantovani and Frank Bursley Taylor preceded him just to mention a few), but he was the first to marshal significant fossil and paleo-topographical and climatological evidence to support this simple observation (and was supported in this by researchers such as Alex du Toit). Furthermore, when the rock strata of the margins of separate continents are very similar it suggests that these rocks were formed in the same way, implying that they were joined initially. For instance, parts of Scotland and Ireland contain rocks very similar to those found in Newfoundland and New Brunswick. Furthermore, the Caledonian Mountains of Europe and parts of the Appalachian Mountains of North America are very similar in structure and lithology.

However, his ideas were not taken seriously by many geologists, who pointed out that there was no apparent mechanism for continental drift. Specifically, they did not see how continental rock could plow through the much denser rock that makes up oceanic crust. Wegener could not explain the force that drove continental drift, and his vindication did not come until after his death in 1930.



Alfred Wegener in Greenland in the winter of 1912-13.

## Floating continents, paleomagnetism, and seismicity zones

As it was observed early that although granite existed on continents, seafloor seemed to be composed of denser basalt, the prevailing concept during the first half of the twentieth century was that there were two types of crust, named "sial" (continental type crust), and "sima" (oceanic type crust). Furthermore, it was supposed that a static shells of strata was present under the continents. It therefore looked apparent that a layer of basalt (sial) underlies the continental rocks.

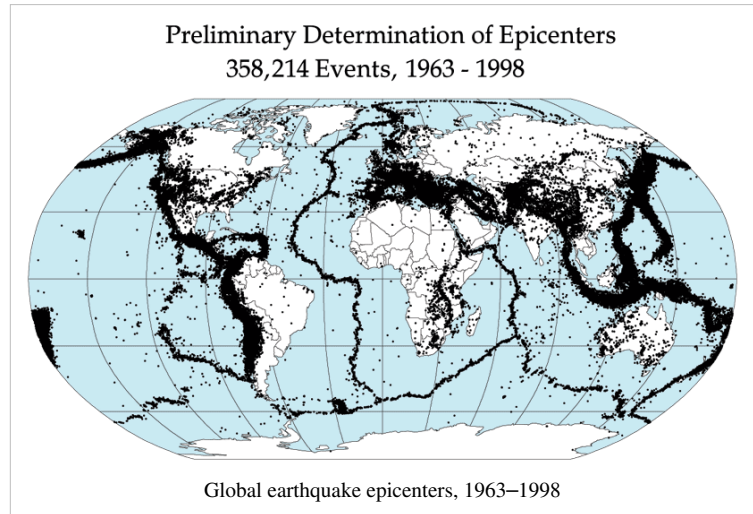
However, based upon abnormalities in plumb line deflection by the Andes in Peru, Pierre Bouguer had deduced that less-dense mountains must have a downward projection into the denser layer underneath. The concept that mountains had "roots" was confirmed by George B. Airy a hundred years later during study of Himalayan gravitation, and seismic studies detected corresponding density variations. Therefore, by the mid-1950s the question remained unresolved of whether mountain roots were clenched in surrounding basalt or were floating upon it like an iceberg.

During the 20th century, improvements in and greater use of seismic instruments such as seismographs enabled scientists to learn that earthquakes tend to be concentrated in specific areas, most notably along the oceanic trenches and spreading ridges. By the late 1920s, seismologists were beginning to identify several prominent earthquake zones parallel to the trenches that typically were inclined 40–60° from the horizontal and extended several hundred kilometers into the Earth. These zones later became known as Wadati-Benioff zones, or simply Benioff zones, in honor of the seismologists who first recognized them, Kiyoo Wadati of Japan and Hugo Benioff of the United States. The study of global seismicity greatly advanced in the 1960s with the establishment of the Worldwide Standardized Seismograph Network (WWSSN)<sup>[40]</sup> to monitor the compliance of the 1963 treaty banning above-ground testing of nuclear weapons. The much improved data from the WWSSN instruments allowed seismologists to map precisely the zones of earthquake concentration world wide.

Meanwhile, debates developed around the phenomena of polar wander. Since the early debates of continental drift, scientists had discussed and used evidence that polar drift had occurred because continents seemed to have moved through different climatic zones during the past. Furthermore, paleomagnetic data had shown that the magnetic pole had also shifted during time. Reasoning in an opposite way, the continents might have shifted and rotated, while the pole remained relatively fixed. The first time the evidence of magnetic polar wander was used to support the movements of continents was in a paper by Keith Runcorn in 1956,<sup>[31]</sup> and successive papers by him and his students Ted Irving (who was actually the first to be convinced of the fact that paleomagnetism supported continental drift) and Ken Creer.

This was immediately followed by a symposium in Tasmania in March 1956.<sup>[41]</sup> In this symposium, the evidence was used in the theory of an expansion of the global crust. In this hypothesis the shifting of the continents can be simply explained by a large increase in size of the Earth since its formation. However, this was unsatisfactory because its supporters could offer no convincing mechanism to produce a significant expansion of the Earth. Certainly there is no evidence that the moon has expanded in the past 3 billion years; other work would soon show that the evidence was equally in support of continental drift on a globe with a stable radius.

During the thirties up to the late fifties, works by Vening-Meinesz, Holmes, Umbgrove, and numerous others outlined concepts that were close or nearly identical to modern plate tectonics theory. In particular, the English



geologist Arthur Holmes proposed in 1920 that plate junctions might lie beneath the sea, and in 1928 that convection currents within the mantle might be the driving force.<sup>[42]</sup> Often, these contributions are forgotten because:

- At the time, continental drift was not accepted.
- Some of these ideas were discussed in the context of abandoned fixistic ideas of a deforming globe without continental drift or an expanding Earth.
- They were published during an episode of extreme political and economic instability that hampered scientific communication.
- Many were published by European scientists and at first not mentioned or given little credit in the papers on sea floor spreading published by the American researchers in the 1960s.

## Mid-oceanic ridge spreading and convection

In 1947, a team of scientists led by Maurice Ewing utilizing the Woods Hole Oceanographic Institution's research vessel *Atlantis* and an array of instruments, confirmed the existence of a rise in the central Atlantic Ocean, and found that the floor of the seabed beneath the layer of sediments consisted of basalt, not the granite which is the main constituent of continents. They also found that the oceanic crust was much thinner than continental crust. All these new findings raised important and intriguing questions.<sup>[43]</sup>

The new data that had been collected on the ocean basins also showed particular characteristics regarding the bathymetry. One of the major outcomes of these datasets was that all along the globe, a system of mid-oceanic ridges was detected. An important conclusion was that along this system, new ocean floor was being created, which led to the concept of the "Great Global Rift". This was described in the crucial paper of Bruce Heezen (1960)<sup>[44]</sup> which would trigger a real revolution in thinking. A profound consequence of seafloor spreading is that new crust was, and is now, being continually created along the oceanic ridges. Therefore, Heezen advocated the so-called "expanding Earth" hypothesis of S. Warren Carey (see above). So, still the question remained: how can new crust be continuously added along the oceanic ridges without increasing the size of the Earth? In reality, this question had been solved already by numerous scientists during the forties and the fifties, like Arthur Holmes, Vening-Meinesz, Coates and many others: The crust in excess disappeared along what were called the oceanic trenches where so-called "subduction" occurred. Therefore, when various scientists during the early sixties started to reason on the data at their disposal regarding the ocean floor, the pieces of the theory fell quickly into place.

The question particularly intrigued Harry Hammond Hess, a Princeton University geologist and a Naval Reserve Rear Admiral, and Robert S. Dietz, a scientist with the U.S. Coast and Geodetic Survey who first coined the term *seafloor spreading*. Dietz and Hess (the former published the same idea one year earlier in *Nature*,<sup>[45]</sup> but priority belongs to Hess who had already distributed an unpublished manuscript of his 1962 article by 1960)<sup>[46]</sup> were among the small handful who really understood the broad implications of sea floor spreading and how it would eventually agree with the, at that time, unconventional and unaccepted ideas of continental drift and the elegant and mobilistic models proposed by previous workers like Holmes.

In the same year, Robert R. Coates of the U.S. Geological Survey described the main features of island arc subduction in the Aleutian Islands. His paper, though little-noted (and even ridiculed) at the time, has since been called "seminal" and "prescient". In reality, it actually shows that the work by the European scientists on island arcs and mountain belts performed and published during the 1930s up until the 1950s was applied and appreciated also in the United States.

If the Earth's crust was expanding along the oceanic ridges, Hess and Dietz reasoned like Holmes and others before them, it must be shrinking elsewhere. Hess followed Heezen suggesting that new oceanic crust continuously spreads away from the ridges in a conveyor belt-like motion. And, using the mobilistic concepts developed before, he correctly concluded that many millions of years later, the oceanic crust eventually descends along the continental margins where oceanic trenches – very deep, narrow canyons – are formed, e.g. along the rim of the Pacific Ocean basin. The important step Hess made was that convection currents would be the driving force in this process, arriving

at the same conclusions as Holmes had decades before with the only difference that the thinning of the ocean crust was performed using the mechanism of Heezen of spreading along the ridges. Hess therefore concluded that the Atlantic Ocean was expanding while the Pacific Ocean was shrinking. As old oceanic crust is "consumed" in the trenches, (like Holmes and others, he believed this was done by thickening of the continental lithosphere, not, as nowadays believed, by underthrusting at a larger scale of the oceanic crust itself into the mantle) new magma rises and erupts along the spreading ridges to form new crust. In effect, the ocean basins are perpetually being "recycled," with the creation of new crust and the destruction of old oceanic lithosphere occurring simultaneously, in a way that later would be called the Wilson cycle (see below). Thus, the new mobilistic concepts neatly explained why the Earth does not get bigger with sea floor spreading, why there is so little sediment accumulation on the ocean floor, and why oceanic rocks are much younger than continental rocks.

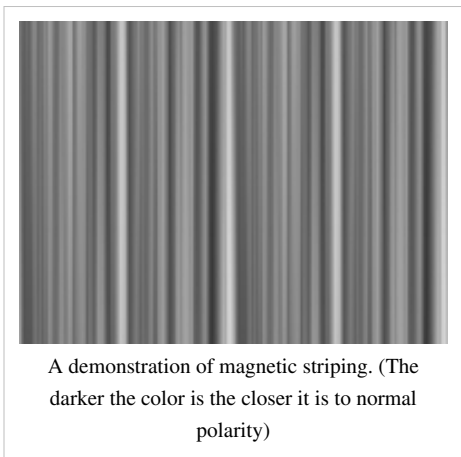
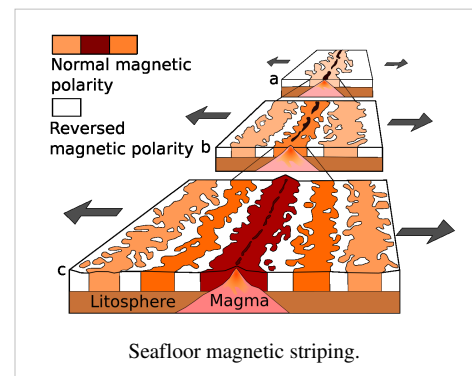
## Magnetic striping

Beginning in the 1950s, scientists like Victor Vacquier, using magnetic instruments (magnetometers) adapted from airborne devices developed during World War II to detect submarines, began recognizing odd magnetic variations across the ocean floor. This finding, though unexpected, was not entirely surprising because it was known that basalt—the iron-rich, volcanic rock making up the ocean floor—contains a strongly magnetic mineral (magnetite) and can locally distort compass readings. This distortion was recognized by Icelandic mariners as early as the late 18th century. More important, because the presence of magnetite gives the basalt measurable magnetic properties, these newly discovered magnetic variations provided another means to study the deep ocean floor. When newly formed rock cools, such magnetic materials recorded the Earth's magnetic field at the time.

As more and more of the seafloor was mapped during the 1950s, the magnetic variations turned out not to be random or isolated occurrences, but instead revealed recognizable patterns. When these magnetic patterns were mapped over a wide region, the ocean floor showed a zebra-like pattern: one stripe with normal polarity and the adjoining stripe with reversed polarity. The overall pattern, defined by these alternating bands of normally and reversely polarized rock, became known as magnetic striping, and was published by Ron G. Mason and co-workers in 1961, who did not find, though, an explanation for these data in terms of sea floor spreading, like Vine, Matthews and Morley a few years later.<sup>[47]</sup>

The discovery of magnetic striping called for an explanation. In the early 1960s scientists such as Heezen, Hess and Dietz had begun to theorise that mid-ocean ridges mark structurally weak zones where the ocean floor was being ripped in two lengthwise along the ridge crest (see the previous paragraph). New magma from deep within the Earth rises easily through these weak zones and eventually erupts along the crest of the ridges to create new oceanic crust. This process, at first denominated the "conveyer belt hypothesis" and later called seafloor spreading, operating over many millions of years continues to form new ocean floor all across the 50,000 km-long system of mid-ocean ridges.

Only four years after the maps with the "zebra pattern" of magnetic stripes were published, the link between sea floor spreading and these patterns was correctly placed, independently by Lawrence Morley, and by Fred Vine and Drummond Matthews, in 1963<sup>[48]</sup> now called the Vine-Matthews-Morley hypothesis. This hypothesis linked these



patterns to geomagnetic reversals and was supported by several lines of evidence:<sup>[49]</sup>

1. the stripes are symmetrical around the crests of the mid-ocean ridges; at or near the crest of the ridge, the rocks are very young, and they become progressively older away from the ridge crest;
2. the youngest rocks at the ridge crest always have present-day (normal) polarity;
3. stripes of rock parallel to the ridge crest alternate in magnetic polarity (normal-reversed-normal, etc.), suggesting that they were formed during different epochs documenting the (already known from independent studies) normal and reversal episodes of the Earth's magnetic field.

By explaining both the zebra-like magnetic striping and the construction of the mid-ocean ridge system, the seafloor spreading hypothesis (SFS) quickly gained converts and represented another major advance in the development of the plate-tectonics theory. Furthermore, the oceanic crust now came to be appreciated as a natural "tape recording" of the history of the geomagnetic field reversals (GMFR) of the Earth's magnetic field. Nowadays, extensive studies are dedicated to the calibration of the normal-reversal patterns in the oceanic crust on one hand and known timescales derived from the dating of basalt layers in sedimentary sequences (magnetostratigraphy) on the other, to arrive at estimates of past spreading rates and plate reconstructions.

### Definition and refining of the theory

After all these considerations, Plate Tectonics (or, as it was initially called "New Global Tectonics") became quickly accepted in the scientific world, and numerous papers followed that defined the concepts:

- In 1965, Tuzo Wilson who had been a promotor of the sea floor spreading hypothesis and continental drift from the very beginning<sup>[50]</sup> added the concept of transform faults to the model, completing the classes of fault types necessary to make the mobility of the plates on the globe work out.<sup>[51]</sup>
- A symposium on continental drift was held at the Royal Society of London in 1965 which must be regarded as the official start of the acceptance of plate tectonics by the scientific community, and which abstracts are issued as Blackett, Bullard & Runcorn (1965). In this symposium, Edward Bullard and co-workers showed with a computer calculation how the continents along both sides of the Atlantic would best fit to close the ocean, which became known as the famous "Bullard's Fit".
- In 1966 Wilson published the paper that referred to previous plate tectonic reconstructions, introducing the concept of what is now known as the "Wilson Cycle".<sup>[52]</sup>
- In 1967, at the American Geophysical Union's meeting, W. Jason Morgan proposed that the Earth's surface consists of 12 rigid plates that move relative to each other.<sup>[53]</sup>
- Two months later, Xavier Le Pichon published a complete model based on 6 major plates with their relative motions, which marked the final acceptance by the scientific community of plate tectonics.<sup>[54]</sup>
- In the same year, McKenzie and Parker independently presented a model similar to Morgan's using translations and rotations on a sphere to define the plate motions.<sup>[55]</sup>



## Implications for biogeography

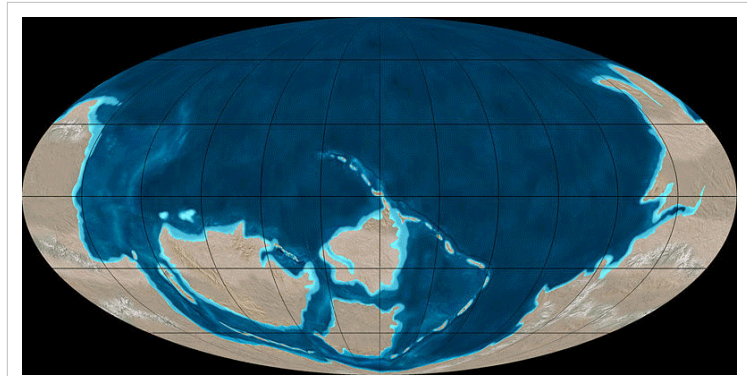
Continental drift theory helps biogeographers to explain the disjunct biogeographic distribution of present day life found on different continents but having similar ancestors.<sup>[56]</sup> In particular, it explains the Gondwanan distribution of ratites and the Antarctic flora.

## Plate reconstruction

Reconstruction is used to establish past (and future) plate configurations, helping determine the shape and make-up of ancient supercontinents and providing a basis for paleogeography.

### Defining plate boundaries

Current plate boundaries are defined by their seismicity.<sup>[57]</sup> Past plate boundaries within existing plates are identified from a variety of evidence, such as the presence of ophiolites that are indicative of vanished oceans.<sup>[58]</sup>



Reconstruction of plate configurations for the whole Phanerozoic

### Past plate motions

Tectonic motion first began around three billion years ago.<sup>[59]</sup>

Various types of quantitative and semi-quantitative information are available to constrain past plate motions. The geometric fit between continents, such as between west Africa and South America is still an important part of plate reconstruction. Magnetic stripe patterns provide a reliable guide to relative plate motions going back into the Jurassic period.<sup>[60]</sup> The tracks of hotspots give absolute reconstructions but these are only available back to the Cretaceous.<sup>[61]</sup> Older reconstructions rely mainly on paleomagnetic pole data, although these only constrain the latitude and rotation, but not the longitude. Combining poles of different ages in a particular plate to produce apparent polar wander paths provides a method for comparing the motions of different plates through time.<sup>[62]</sup> Additional evidence comes from the distribution of certain sedimentary rock types,<sup>[63]</sup> faunal provinces shown by particular fossil groups, and the position of orogenic belts.<sup>[61]</sup>

### Formation and break-up of continents

The movement of plates has caused the formation and break-up of continents over time, including occasional formation of a supercontinent that contains most or all of the continents. The supercontinent Columbia or Nuna formed during a period of 2000 to 1800.0<sup>[64]</sup> million years ago and broke up about 1500 to 1300<sup>[65]</sup> million years ago.<sup>[66]</sup> The supercontinent Rodinia is thought to have formed about 1 billion years ago and to have embodied most or all of Earth's continents, and broken up into eight continents around 600<sup>[67]</sup> million years ago. The eight continents later re-assembled into another supercontinent called Pangea; Pangea broke up into Laurasia (which became North America and Eurasia) and Gondwana (which became the remaining continents).

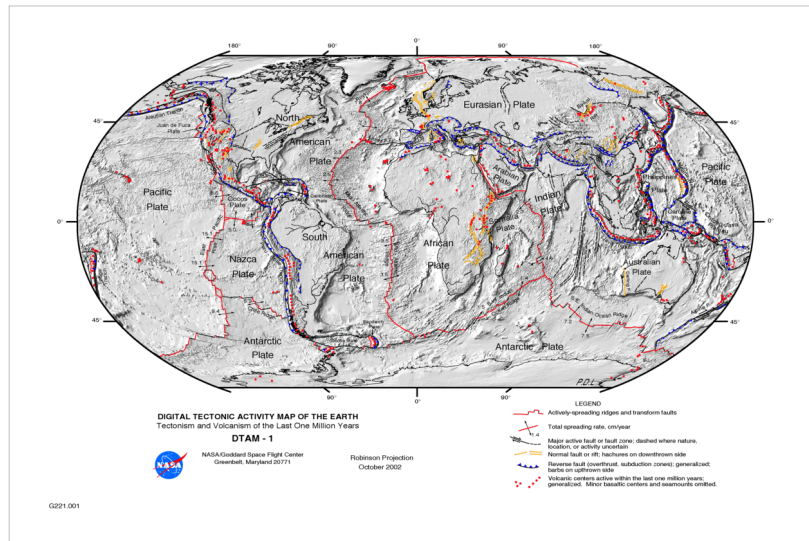
The Himalayas, the world's tallest mountain range, are assumed to be formed by the collision of two major plates. Before uplift, they were covered by the Tethys Ocean.

## Current plates

Depending on how they are defined, there are usually seven or eight "major" plates: African, Antarctic, Eurasian, North American, South American, Pacific, and Indo-Australian. The latter is sometimes subdivided into the Indian and Australian plates.

There are dozens of smaller plates, the seven largest of which are the Arabian, Caribbean, Juan de Fuca, Cocos, Nazca, Philippine Sea and Scotia.

The current motion of the tectonic plates is nowadays revealed from remote sensing satellite data sets, calibrated with ground station measurements.



## Other celestial bodies (planets, moons)

The appearance of plate tectonics on terrestrial planets is related to planetary mass, with more massive planets than Earth expected to exhibit plate tectonics. Earth may be a borderline case, owing its tectonic activity to abundant water<sup>[68]</sup> (Silica and water form a deep eutectic.)

### Venus

Venus shows no evidence of active plate tectonics. There is debatable evidence of active tectonics in the planet's distant past; however, events taking place since then (such as the plausible and generally accepted hypothesis that the Venusian lithosphere has thickened greatly over the course of several hundred million years) has made constraining the course of its geologic record difficult. However, the numerous well-preserved impact craters have been utilized as a dating method to approximately date the Venusian surface (since there are thus far no known samples of Venusian rock to be dated by more reliable methods). Dates derived are dominantly in the range 500 to 750<sup>[69]</sup> million years ago, although ages of up to 1200<sup>[70]</sup> million years ago have been calculated. This research has led to the fairly well accepted hypothesis that Venus has undergone an essentially complete volcanic resurfacing at least once in its distant past, with the last event taking place approximately within the range of estimated surface ages. While the mechanism of such an impressive thermal event remains a debated issue in Venusian geosciences, some scientists are advocates of processes involving plate motion to some extent.

One explanation for Venus' lack of plate tectonics is that on Venus temperatures are too high for significant water to be present.<sup>[71], [72]</sup> The Earth's crust is soaked with water, and water plays an important role in the development of shear zones. Plate tectonics requires weak surfaces in the crust along which crustal slices can move, and it may well be that such weakening never took place on Venus because of the absence of water. However, some researchers remain convinced that plate tectonics is or was once active on this planet.

## Mars

Mars is considerably smaller than Earth and Venus, and there is evidence for ice on its surface and in its crust.

In the 1990s, it was proposed that Martian Crustal Dichotomy was created by plate tectonic processes.<sup>[73]</sup> Scientists today disagree, and believe that it was created either by upwelling within the Martian mantle that thickened the crust of the Southern Highlands and formed Tharsis<sup>[74]</sup> or by a giant impact that excavated the Northern Lowlands.<sup>[75]</sup>

Valles Marineris is a tectonic boundary.<sup>[76]</sup>

Observations made of the magnetic field of Mars by the *Mars Global Surveyor* spacecraft in 1999 showed patterns of magnetic striping discovered on this planet. Some scientists interpreted these as requiring plate tectonic processes, such as seafloor spreading.<sup>[77]</sup> However, their data fail a "magnetic reversal test", which is used to see if they were formed by flipping polarities of a global magnetic field.<sup>[78]</sup>

## Galilean satellites of Jupiter

Some of the satellites of Jupiter have features that may be related to plate-tectonic style deformation, although the materials and specific mechanisms may be different from plate-tectonic activity on Earth.

## Titan, moon of Saturn

Titan, the largest moon of Saturn, was reported to show tectonic activity in images taken by the Huygens Probe, which landed on Titan on January 14, 2005.<sup>[79]</sup>

## Exoplanets

It is believed that many planets around other stars will have plate tectonics. On Earth-sized planets, plate tectonics is more likely if there are oceans of water,<sup>[80]</sup> but on larger super-earths plate tectonics is very likely even if the planet is dry.<sup>[68]</sup>

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## External links

- This Dynamic Earth: The Story of Plate Tectonics (<http://pubs.usgs.gov/gip/dynamic/dynamic.html>). USGS.
- Understanding Plate Tectonics (<http://pubs.usgs.gov/publications/text/understanding.html>). USGS.
- The PLATES Project (<http://www.ig.utexas.edu/research/projects/plates/>). Jackson School of Geosciences.
- An explanation of tectonic forces (<http://www.tectonic-forces.org>). Example of calculations to show that Earth Rotation could be a driving force.
- Bird, P. (2003); An updated digital model of plate boundaries ([http://peterbird.name/publications/2003\\_PB2002/2003\\_PB2002.htm](http://peterbird.name/publications/2003_PB2002/2003_PB2002.htm)).
- Map of tectonic plates (<http://snobear.colorado.edu/Markw/Mountains/03/week3.html>).
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- GPlates (<http://www.gplates.org/>), desktop software for the interactive visualization of plate-tectonics.
- MORVEL plate velocity estimates and information (<http://www.geology.wisc.edu/~chuck/MORVEL/>). C. DeMets, D. Argus, & R. Gordon.

## Videos

- Khan Academy Explanation of evidence (<http://www.youtube.com/watch?v=6EdsBabSZ4g>)
- 750 million years of global tectonic activity (<http://www.ucmp.berkeley.edu/geology/tectonics.html>). Movie.

# Geology

**Geology** (from the Greek γῆ, *gê*, "earth" and λόγος, *logos*, "study") is the science comprising the study of solid Earth, the rocks of which it is composed, and the processes by which it evolves. Geology gives insight into the history of the Earth, as it provides the primary evidence for plate tectonics, the evolutionary history of life, and past climates. In modern times, geology is commercially important for mineral and hydrocarbon exploration and for evaluating water resources; it is publicly important for the prediction and understanding of natural hazards, the remediation of environmental problems, and for providing insights into past climate change; plays a role in geotechnical engineering; and is a major academic discipline.



Students examining the Wasatch Fault near Salt Lake City, Utah.

## History

The study of the physical material of the Earth dates back at least to ancient Greece when Theophrastus (372–287 BCE) wrote the work *Peri Lithon* (*On Stones*). In the Roman period, Pliny the Elder wrote in detail of the many minerals and metals then in practical use, and correctly noted the origin of amber.

Some modern scholars, such as Fielding H. Garrison, are of the opinion that modern geology began in the medieval Islamic world.<sup>[2]</sup> Abu al-Rayhan al-Biruni (973–1048 CE) was one of the earliest Muslim geologists, whose works included the earliest writings on the geology of India, hypothesizing that the Indian subcontinent was once a sea.<sup>[3]</sup> Islamic Scholar Ibn Sina (Avicenna, 981–1037) proposed detailed explanations for the formation of mountains, the origin of earthquakes, and other topics central to modern Geology, which provided an essential foundation for the later development of the science.<sup>[4][5]</sup> In China, the polymath Shen Kuo (1031–1095) formulated a hypothesis for the process of land formation: based on his observation of fossil animal shells in a geological stratum in a mountain hundreds of miles from the ocean, he inferred that the land was formed by erosion of the mountains and by deposition of silt.<sup>[6]</sup>

Nicolas Steno (1638–1686) is credited with the law of superposition, the principle of original horizontality, and the principle of lateral continuity: three defining principles of stratigraphy. The word *geology* was first used by Ulisse Aldrovandi in 1603,<sup>[7]</sup> then by Jean-André Deluc in 1778 and introduced as a fixed term by Horace-Bénédict de Saussure in 1779. The word is derived from the Greek γῆ, *gê*, meaning "earth" and λόγος, *logos*, meaning "speech".<sup>[8]</sup> But according to another source, the word "Geology" comes from the Norwegian, Mikkel Pedersøn Escholt (1600–1699), who was a priest and scholar. Escholt was first used the definition in his book titled, *Geologica Norvegica* (1657).<sup>[9]</sup>

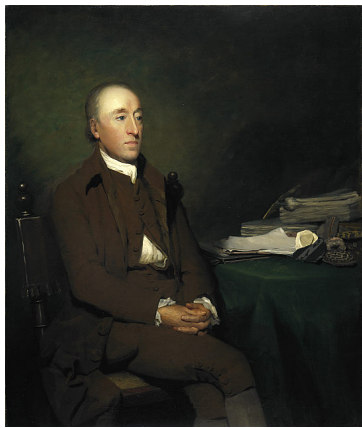
William Smith (1769–1839) drew some of the first geological maps and began the process of ordering rock strata (layers) by examining the fossils contained in them.<sup>[1]</sup>

James Hutton is often viewed as the first modern geologist.<sup>[10]</sup> In 1785 he presented a paper entitled *Theory of the Earth* to the Royal Society of Edinburgh. In his paper, he explained his theory that the Earth must be much older than had previously been supposed in order to allow enough time for mountains to be eroded and for sediments to form new rocks at the bottom of the sea, which in turn were raised up to become dry land. Hutton published a two-volume version of his ideas in 1795 (Vol. 1<sup>[11]</sup>, Vol. 2<sup>[12]</sup>).



William Smith's geologic map of England, Wales, and southern Scotland. Completed in 1815, it was the first national-scale geologic map, and by far the most accurate of its time.<sup>[1]</sup>





James Hutton, father of modern geology

Followers of Hutton were known as *Plutonists* because they believed that some rocks were formed by *vulcanism*, which is the deposition of lava from volcanoes, as opposed to the *Neptunists*, led by Abraham Werner, who believed that all rocks had settled out of a large ocean whose level gradually dropped over time.

Sir Charles Lyell first published his famous book, *Principles of Geology*,<sup>[13]</sup> in 1830. The book, which influenced the thought of Charles Darwin, successfully promoted the doctrine of uniformitarianism. This theory states that slow geological processes have occurred throughout the Earth's history and are still occurring today. In contrast, catastrophism is the theory that Earth's features formed in single, catastrophic events and remained unchanged thereafter. Though Hutton believed in uniformitarianism, the idea was not widely accepted at the time.

Much of 19th-century geology revolved around the question of the Earth's exact age. Estimates varied from a few 100,000 to billions of years.<sup>[14]</sup> By the early 20th century, radiometric dating allowed the Earth's age to be estimated at two billion years. The awareness of this vast amount of time opened the door to new theories about the processes that shaped the planet.

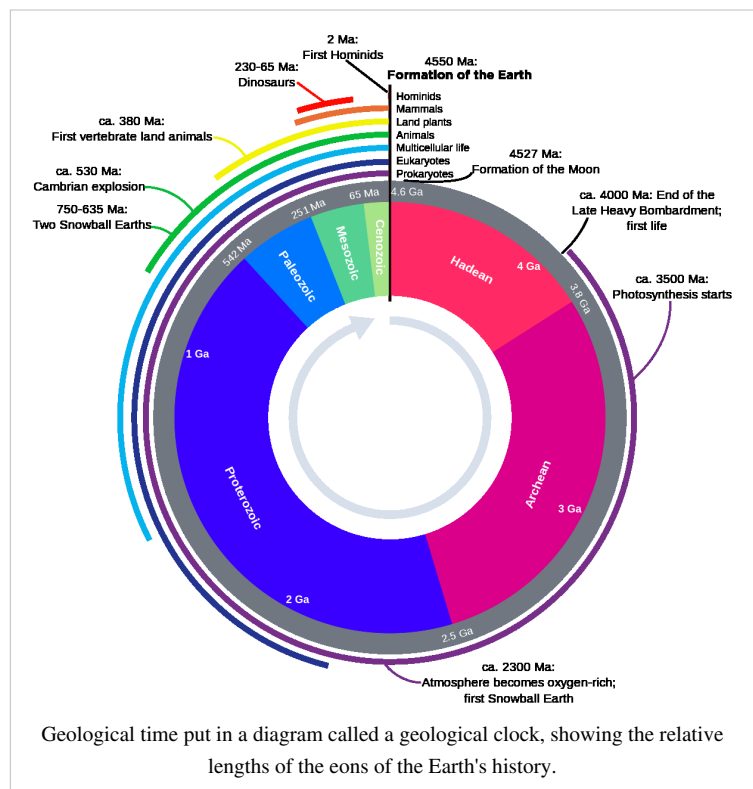
The most significant advances in 20th century geology have been the development of the theory of plate tectonics in the 1960s, and the refinement of estimates of the planet's age. Plate tectonic theory arose out of two separate geological observations: seafloor spreading and continental drift. The theory revolutionized the Earth sciences. Today the Earth is known to be approximately 4.5 billion years old.<sup>[15]</sup>

## Geologic time

The geologic time scale encompasses the history of the Earth.<sup>[16]</sup> It is bracketed at the old end by the dates of the earliest solar system material at 4.567 Ga,<sup>[17]</sup> (gigaannum: billion years ago) and the age of the Earth at 4.54 Ga<sup>[18][19]</sup> at the beginning of the informally recognized Hadean eon. At the young end of the scale, it is bracketed by the present day in the Holocene epoch.

### Important milestones

- 4.567 Ga: Solar system formation<sup>[17]</sup>
- 4.54 Ga: Accretion of Earth<sup>[18][19]</sup>
- c. 4 Ga: End of Late Heavy Bombardment, first life
- c. 3.5 Ga: Start of photosynthesis
- c. 2.3 Ga: Oxygenated atmosphere, first snowball Earth
- 730–635 Ma (megaannum: million years ago): two snowball Earths



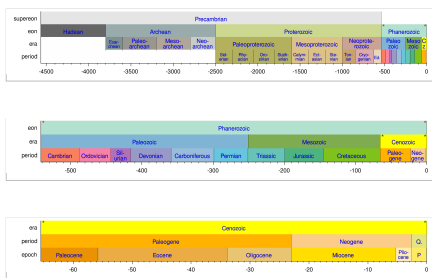
Geological time put in a diagram called a geological clock, showing the relative lengths of the eons of the Earth's history.



- 542± 0.3 Ma: Cambrian explosion – vast multiplication of hard-bodied life; first abundant fossils; start of the Paleozoic
- c. 380 Ma: First vertebrate land animals
- 250 Ma: Permian-Triassic extinction – 90% of all land animals die. End of Paleozoic and beginning of Mesozoic
- 65 Ma: Cretaceous–Paleogene extinction – Dinosaurs die; end of Mesozoic and beginning of Cenozoic
- c. 7 Ma – Present: Hominins
  - c. 7 Ma: First hominins appear
  - 3.9 Ma: First Australopithecus, direct ancestor to modern Homo sapiens, appear
  - 200 ka (kiloannum: thousand years ago): First modern Homo sapiens appear in East Africa

## Brief time scale

The second and third timelines are each subsections of their preceding timeline as indicated by asterisks. The Holocene (the latest epoch) is too small to be shown clearly on this timeline.



Millions of Years

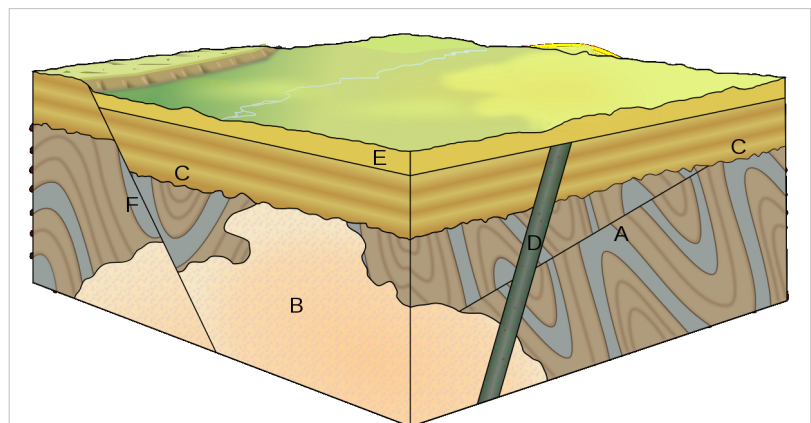
## Relative and absolute dating

Geological events can be given a precise date at a point in time, or they can be related to other events that came before and after them. Geologists use a variety of methods to give both relative and absolute dates to geological events. They then use these dates to find the rates at which processes occur.

### Relative dating

Methods for relative dating were developed when geology first emerged as a formal science. Geologists still use the following principles today as a means to provide information about geologic history and the timing of geologic events.

**The principle of Uniformitarianism** states that the geologic processes observed in operation that modify the Earth's crust at present have worked in much the same way over geologic time.<sup>[20]</sup> A fundamental principle of geology advanced by the 18th century Scottish physician and geologist James Hutton, is that "the present is the key to the past." In Hutton's words: "the past history of our globe must be explained by what can be seen to be happening now."<sup>[21]</sup>



Cross-cutting relations can be used to determine the relative ages of rock strata and other geological structures. Explanations: A - folded rock strata cut by a thrust fault; B - large intrusion (cutting through A); C - erosional angular unconformity (cutting off A & B) on which rock strata were deposited; D - volcanic dyke (cutting through A, B & C); E - even younger rock strata (overlying C & D); F - normal fault (cutting through A, B, C & E).

**The principle of intrusive relationships** concerns crosscutting intrusions. In geology, when an igneous intrusion cuts across a formation of sedimentary rock, it can be determined that the igneous intrusion is younger than the sedimentary rock. There are a number of different types of intrusions, including stocks, laccoliths, batholiths, sills and dikes.

**The principle of cross-cutting relationships** pertains to the formation of faults and the age of the sequences through which they cut. Faults are younger than the rocks they cut; accordingly, if a fault is found that penetrates some formations but not those on top of it, then the formations that were cut are older than the fault, and the ones that are not cut must be younger than the fault. Finding the key bed in these situations may help determine whether the fault is a normal fault or a thrust fault.<sup>[22]</sup>

**The principle of inclusions and components** states that, with sedimentary rocks, if inclusions (or *clasts*) are found in a formation, then the inclusions must be older than the formation that contains them. For example, in sedimentary rocks, it is common for gravel from an older formation to be ripped up and included in a newer layer. A similar situation with igneous rocks occurs when xenoliths are found. These foreign bodies are picked up as magma or lava flows, and are incorporated, later to cool in the matrix. As a result, xenoliths are older than the rock which contains them.

**The principle of original horizontality** states that the deposition of sediments occurs as essentially horizontal beds. Observation of modern marine and non-marine sediments in a wide variety of environments supports this generalization (although cross-bedding is inclined, the overall orientation of cross-bedded units is horizontal).<sup>[22]</sup>

**The principle of superposition** states that a sedimentary rock layer in a tectonically undisturbed sequence is younger than the one beneath it and older than the one above it. Logically a younger layer cannot slip beneath a layer previously deposited. This principle allows sedimentary layers to be viewed as a form of vertical time line, a partial or complete record of the time elapsed from deposition of the lowest layer to deposition of the highest bed.<sup>[22]</sup>



The Permian through Jurassic stratigraphy of the Colorado Plateau area of southeastern Utah is a great example of both Original Horizontality and the Law of Superposition. These strata make up much of the famous prominent rock formations in widely spaced protected areas such as Capitol Reef National Park and Canyonlands National Park. From top to bottom: Rounded tan domes of the Navajo Sandstone, layered red Kayenta Formation, cliff-forming, vertically jointed, red Wingate Sandstone, slope-forming, purplish Chinle Formation, layered, lighter-red Moenkopi Formation, and white, layered Cutler Formation sandstone. Picture from Glen Canyon National Recreation Area, Utah.

**The principle of faunal succession** is based on the appearance of fossils in sedimentary rocks. As organisms exist at the same time period throughout the world, their presence or (sometimes) absence may be used to provide a relative age of the formations in which they are found. Based on principles laid out by William Smith almost a hundred years before the publication of Charles Darwin's theory of evolution, the principles of succession were developed independently of evolutionary thought. The principle becomes quite complex, however, given the uncertainties of fossilization, the localization of fossil types due to lateral changes in habitat (facies change in sedimentary strata), and that not all fossils may be found globally at the same time.<sup>[23]</sup>

### Absolute dating

Geologists can also give precise absolute dates to geologic events. These dates are useful on their own, and can also be used in conjunction with relative dating methods or to calibrate relative dating methods.<sup>[24]</sup>

A large advance in geology in the advent of the 20th century was the ability to give precise absolute dates to geologic events through radioactive isotopes and other methods. The advent of radiometric dating changed the understanding of geologic time. Before, geologists could only use fossils to date sections of rock relative to one another. With isotopic dates, absolute dating became possible, and these absolute dates could be applied fossil sequences in which there was datable material, converting the old relative ages into new absolute ages.

For many geologic applications, isotope ratios are measured in minerals that give the amount of time that has passed since a rock passed through its particular closure temperature, the point at which different radiometric isotopes stop diffusing into and out of the crystal lattice.<sup>[25][26]</sup> These are used in geochronologic and thermochronologic studies. Common methods include uranium-lead dating, potassium-argon dating and argon-argon dating, and uranium-thorium dating. These methods are used for a variety of applications. Dating of lavas and ash layers can help to date stratigraphy and calibrate relative dating techniques. These methods can also be used to determine ages of pluton emplacement. Thermochemical techniques can be used to determine temperature profiles within the crust, the uplift of mountain ranges, and paleotopography.

Fractionation of the lanthanide series elements is used to compute ages since rocks were removed from the mantle.

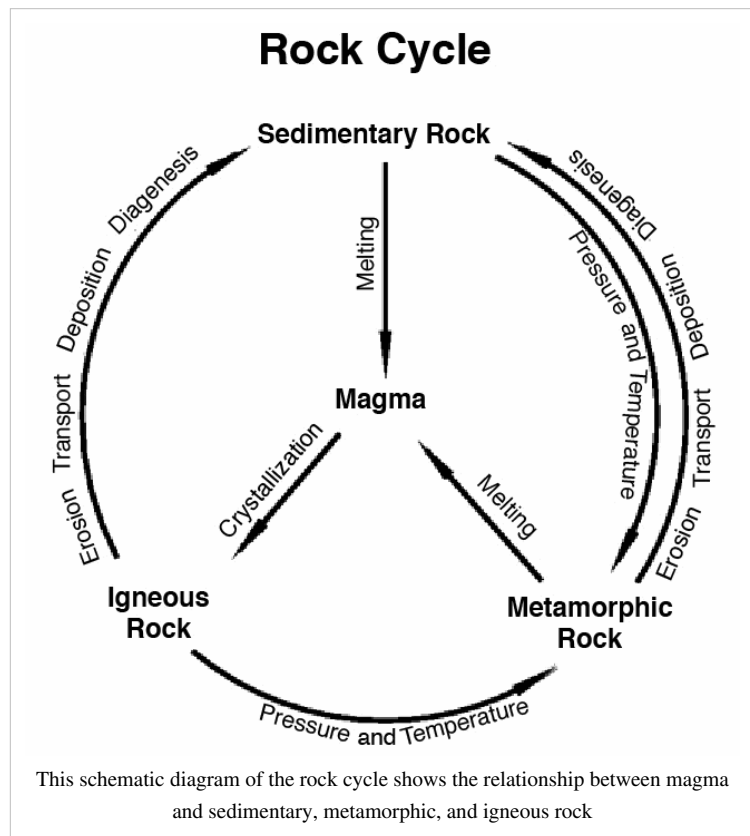
Other methods are used for more recent events. Optically stimulated luminescence and cosmogenic radionuclide dating are used to date surfaces and/or erosion rates. Dendrochronology can also be used for the dating of landscapes. Radiocarbon dating is used for young organic material.

## Geologic materials

The majority of geological data come from research on solid Earth materials. These typically fall into one of two categories: rock and unconsolidated material.

### Rock

There are three major types of rock: igneous, sedimentary, and metamorphic. The rock cycle is an important concept in geology which illustrates the relationships between these three types of rock, and magma. When a rock crystallizes from melt (magma and/or lava), it is an igneous rock. This rock can be weathered and eroded, and then redeposited and lithified into a sedimentary rock, or be turned into a metamorphic rock due to heat and pressure that change the mineral content of the rock and give it a characteristic fabric. The sedimentary rock can then be subsequently turned into a metamorphic rock due to heat and pressure, and the metamorphic rock can be weathered, eroded, deposited, and lithified, becoming a sedimentary rock. Sedimentary rock may also be re-eroded and redeposited, and metamorphic rock may also undergo additional metamorphism. All three



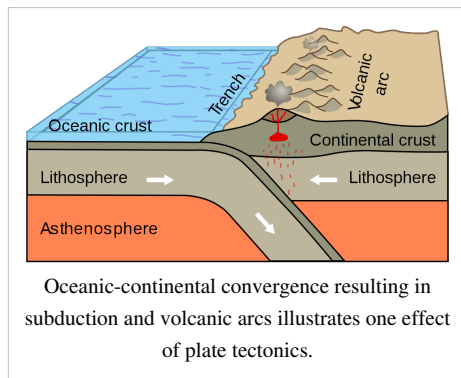
types of rocks may be re-melted; when this happens, a new magma is formed, from which an igneous rock may once again crystallize.

The majority of research in geology is associated with the study of rock, as rock provides the primary record of the majority of the geologic history of the Earth.

### Unconsolidated material

Geologists also study unlithified material, which typically comes from more recent deposits. Because of this, the study of such material is often known as Quaternary geology, after the recent Quaternary Period. This includes the study of sediment and soils, and is important to some (or many) studies in geomorphology, sedimentology, and paleoclimatology.

### Whole-Earth structure

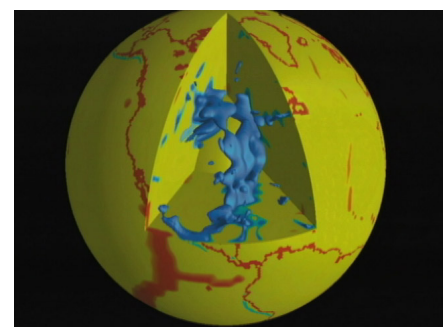


### Plate tectonics

In the 1960s, a series of discoveries, the most important of which was seafloor spreading,<sup>[27][28]</sup> showed that the Earth's lithosphere, which includes the crust and rigid uppermost portion of the upper mantle, is separated into a number of tectonic plates that move across the plastically deforming, solid, upper mantle, which is called the asthenosphere. There is an intimate coupling between the movement of the plates on the surface and the convection of the mantle: oceanic plate motions and mantle convection currents always move in the same direction, because the oceanic lithosphere is the rigid upper thermal boundary layer of the convecting mantle. This coupling between rigid plates moving on the surface of the Earth and the convecting mantle is called plate tectonics.

The development of plate tectonics provided a physical basis for many observations of the solid Earth. Long linear regions of geologic features could be explained as plate boundaries.<sup>[29]</sup> Mid-ocean ridges, high regions on the seafloor where hydrothermal vents and volcanoes exist, were explained as divergent boundaries, where two plates move apart. Arcs of volcanoes and earthquakes were explained as convergent boundaries, where one plate subducts under another. Transform boundaries, such as the San Andreas fault system, resulted in widespread powerful earthquakes.

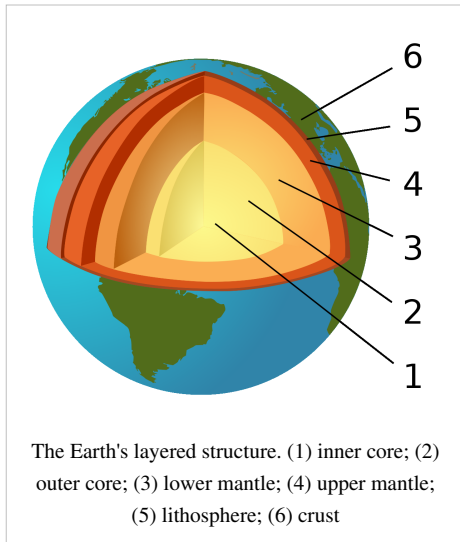
Plate tectonics also provided a mechanism for Alfred Wegener's theory of continental drift,<sup>[30]</sup> in which the continents move across the surface of the Earth over geologic time. They also provided a driving force for crustal



On this diagram, subducting slabs are in blue, and continental margins and a few plate boundaries are in red. The blue blob in the cutaway section is the seismically imaged Farallon Plate, which is subducting beneath North America. The remnants of this plate on the Surface of the Earth are the Juan de Fuca Plate and Explorer plate in the Northwestern USA / Southwestern Canada, and the Cocos Plate on the west coast of Mexico.

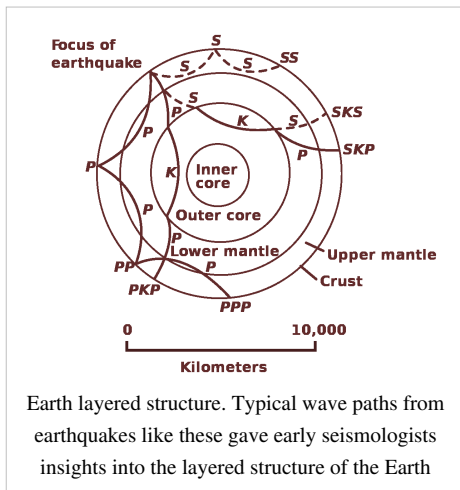
deformation, and a new setting for the observations of structural geology. The power of the theory of plate tectonics lies in its ability to combine all of these observations into a single theory of how the lithosphere moves over the convecting mantle.

## Earth structure



Advances in seismology, computer modeling, and mineralogy and crystallography at high temperatures and pressures give insights into the internal composition and structure of the Earth.

Seismologists can use the arrival times of seismic waves in reverse to image the interior of the Earth. Early advances in this field showed the existence of a liquid outer core (where shear waves were not able to propagate) and a dense solid inner core. These advances led to the development of a layered model of the Earth, with a crust and lithosphere on top, the mantle below (separated within itself by seismic discontinuities at 410 and 660 kilometers), and the outer core and inner core below that. More recently, seismologists have been able to create detailed images of wave speeds inside the earth in the same way a doctor images a body in a CT scan. These images have led to a much more detailed view of the interior of the Earth, and have replaced the simplified layered model with a much more dynamic model.



Mineralogists have been able to use the pressure and temperature data from the seismic and modelling studies alongside knowledge of the elemental composition of the Earth at depth to reproduce these conditions in experimental settings and measure changes in crystal structure. These studies explain the chemical changes associated with the major seismic discontinuities in the mantle, and show the crystallographic structures expected in the inner core of the Earth.

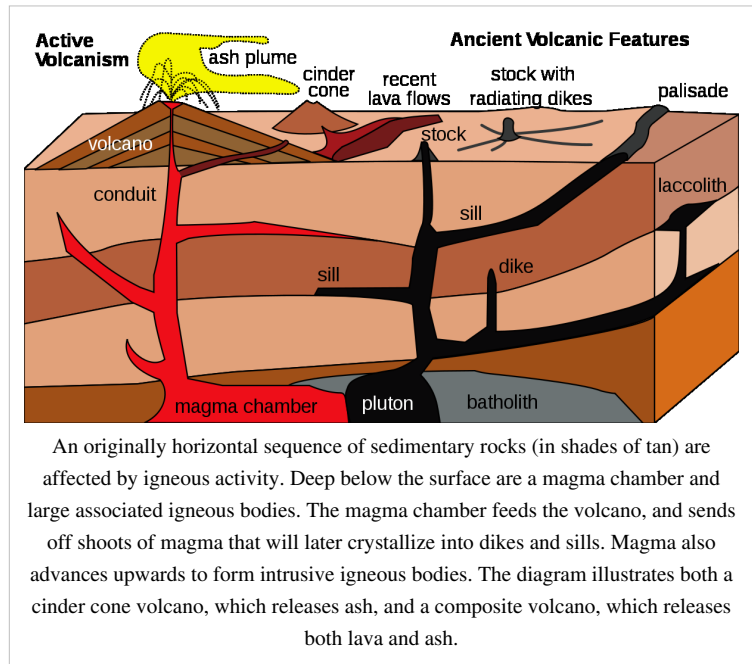
## Geological development of an area

The geology of an area changes through time as rock units are deposited and inserted and deformational processes change their shapes and locations.

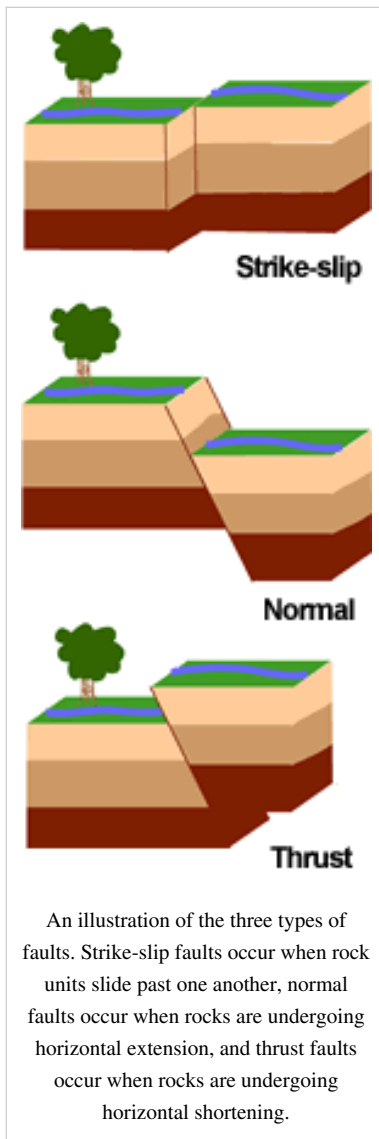
Rock units are first emplaced either by deposition onto the surface or intrusion into the overlying rock. Deposition can occur when sediments settle onto the surface of the Earth and later lithify into sedimentary rock, or when as volcanic material such as volcanic ash or lava flows blanket the surface. Igneous intrusions such as batholiths, laccoliths, dikes, and sills, push upwards into the overlying rock, and crystallize as they intrude.

After the initial sequence of rocks has been deposited, the rock units can be deformed and/or metamorphosed. Deformation typically occurs as a result of horizontal shortening, horizontal extension, or side-to-side (strike-slip) motion. These structural regimes broadly relate to convergent boundaries, divergent boundaries, and transform boundaries, respectively, between tectonic plates.

When rock units are placed under horizontal compression, they shorten and become thicker. Because rock units, other than muds, do not significantly change in volume, this is accomplished in two primary ways: through faulting and folding. In the shallow crust, where brittle deformation can occur, thrust faults form, which cause deeper rock to move on top of shallower rock. Because deeper rock is often older, as noted by the principle of superposition, this can result in older rocks moving on top of younger ones. Movement along faults can result in folding, either because the faults are not planar, or because the rock layers are dragged along, forming drag folds, as slip occurs along the



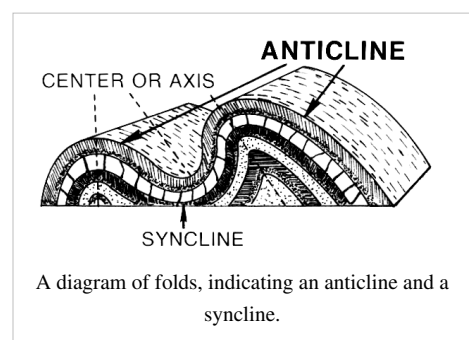




Earth, rocks behave plastically, and fold instead of faulting. These folds can either be those where the material in the center of the fold buckles upwards, creating "antiforms", or where it buckles downwards, creating "synforms". If the tops of the rock units within the folds remain pointing upwards, they are called anticlines and synclines, respectively. If some of the units in the fold are facing downward, the structure is called an overturned anticline or syncline, and if all of the rock units are overturned or the correct up-direction is unknown, they are simply called by the most general terms, antiforms and synforms.

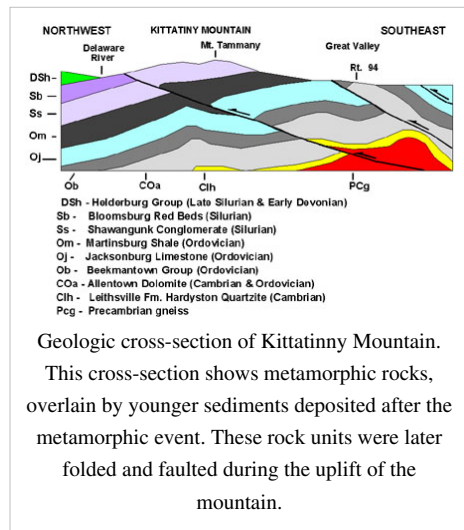
Even higher pressures and temperatures during horizontal shortening can cause both folding and metamorphism of the rocks. This metamorphism causes changes in the mineral composition of the rocks; creates a foliation, or planar surface, that is related to mineral growth under stress; and can remove signs of the original textures of the rocks, such as bedding in sedimentary rocks, flow features of lavas, and crystal patterns in crystalline rocks.

Extension causes the rock units as a whole to become longer and thinner. This is primarily accomplished through normal faulting and through the ductile stretching and thinning. Normal faults drop rock units that are higher below those that are lower. This typically results in younger units being placed below older units. Stretching of units can result in their thinning; in fact, there is a location within the Maria Fold and Thrust Belt in which the entire sedimentary sequence of the Grand Canyon can be seen over a length of less than a meter. Rocks at the depth to be ductilely stretched are often also metamorphosed. These stretched rocks can also pinch into lenses, known as boudins, after the French word for "sausage", because of their visual similarity.



Where rock units slide past one another, strike-slip faults develop in shallow regions, and become shear zones at deeper depths where the rocks deform ductilely.

The addition of new rock units, both depositionally and intrusively, often occurs during deformation. Faulting and other deformational processes result in the creation of topographic gradients, causing material on the rock unit that is increasing in elevation to be eroded by hillslopes and channels. These sediments are deposited on the rock unit that is going down. Continual motion along the fault maintains the topographic gradient in spite of the movement of sediment, and continues to create accommodation space for the material to deposit. Deformational events are often also associated with volcanism and igneous activity. Volcanic ashes and lavas accumulate on the surface, and igneous intrusions enter from below. Dikes, long, planar igneous intrusions, enter along cracks, and therefore often form in large numbers in areas that are being actively deformed. This can result in the emplacement of dike swarms, such as those that are observable across the Canadian shield, or rings of dikes around the lava tube of a volcano.



All of these processes do not necessarily occur in a single environment, and do not necessarily occur in a single order. The Hawaiian Islands, for example, consist almost entirely of layered basaltic lava flows. The sedimentary sequences of the mid-continental United States and the Grand Canyon in the southwestern United States contain almost-undeformed stacks of sedimentary rocks that have remained in place since Cambrian time. Other areas are much more geologically complex. In the southwestern United States, sedimentary, volcanic, and intrusive rocks have been metamorphosed, faulted, foliated, and folded. Even older rocks, such as the Acasta gneiss of the Slave craton in northwestern Canada, the oldest known rock in the world have been metamorphosed to the point where their origin is indiscernable without laboratory analysis. In addition, these processes can occur in stages. In many places, the Grand Canyon in the southwestern United States being a very visible example, the lower rock units were metamorphosed and deformed, and then deformation ended and the upper, undeformed units were deposited. Although any amount of rock emplacement and rock deformation can occur, and they can occur any number of times, these concepts provide a guide to understanding the geological history of an area.

## Methods of geology

Geologists use a number of field, laboratory, and numerical modeling methods to decipher Earth history and understand the processes that occur on and in the Earth. In typical geological investigations, geologists use primary information related to petrology (the study of rocks), stratigraphy (the study of sedimentary layers), and structural geology (the study of positions of rock units and their deformation). In many cases, geologists also study modern soils, rivers, landscapes, and glaciers; investigate past and current life and biogeochemical pathways, and use geophysical methods to investigate the subsurface.

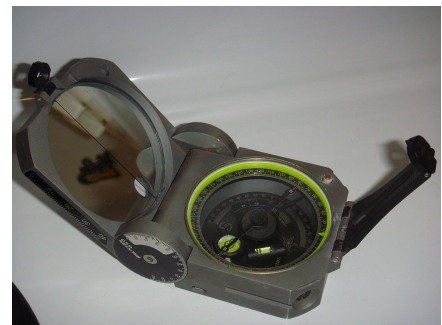


Washington State Land Forms

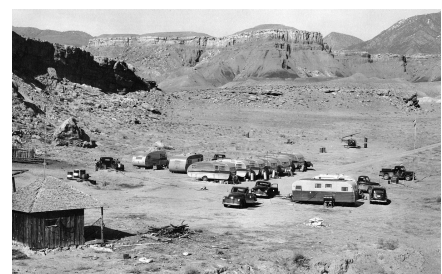
## Field methods

Geological field work varies depending on the task at hand. Typical fieldwork could consist of:

- Geological mapping<sup>[31]</sup>
  - Structural mapping: the locations of the major rock units and the faults and folds that led to their placement there.
  - Stratigraphic mapping: the locations of sedimentary facies (lithofacies and biofacies) or the mapping of isopachs of equal thickness of sedimentary rock
  - Surficial mapping: the locations of soils and surficial deposits
- Surveying of topographic features
  - Creation of topographic maps<sup>[32]</sup>
  - Work to understand change across landscapes, including:
    - Patterns of erosion and deposition
    - River channel change through migration and avulsion
    - Hillslope processes
- Subsurface mapping through geophysical methods<sup>[33]</sup>
  - These methods include:
    - Shallow seismic surveys
    - Ground-penetrating radar
    - Electrical resistivity tomography
  - They are used for:
    - Hydrocarbon exploration
    - Finding groundwater
    - Locating buried archaeological artifacts
- High-resolution stratigraphy
  - Measuring and describing stratigraphic sections on the surface
  - Well drilling and logging
- Biogeochemistry and geomicrobiology<sup>[34]</sup>
  - Collecting samples to:
    - Determine biochemical pathways
    - Identify new species of organisms. These organisms may help to show:
    - Identify new chemical compounds



A standard Brunton compass, used commonly by geologists in mapping and surveying



A typical USGS field mapping camp in the 1950s

- And to use these discoveries to
  - Understand early life on Earth and how it functioned and metabolized
  - Find important compounds for use in pharmaceuticals.
- Paleontology: excavation of fossil material
  - For research into past life and evolution
  - For museums and education



Today, handheld computers with GPS and geographic information systems software are often used in geological field work (digital geologic mapping).

- Collection of samples for geochronology and thermochronology<sup>[35]</sup>
- Glaciology: measurement of characteristics of glaciers and their motion<sup>[36]</sup>

## Laboratory methods

### Petrology

In addition to the field identification of rocks, petrologists identify rock samples in the laboratory. Two of the primary methods for identifying rocks in the laboratory are through optical microscopy and by using an electron microprobe. In an optical mineralogy analysis, thin sections of rock samples are analyzed through a petrographic microscope, where the minerals can be identified through their different properties in plane-polarized and cross-polarized light, including their birefringence, pleochroism, twinning, and interference properties with a conoscopic lens.<sup>[37]</sup> In the electron microprobe, individual locations are analyzed for their exact chemical compositions and variation in composition within individual crystals.<sup>[38]</sup> Stable<sup>[39]</sup> and radioactive isotope<sup>[40]</sup> studies provide insight into the geochemical evolution of rock units.

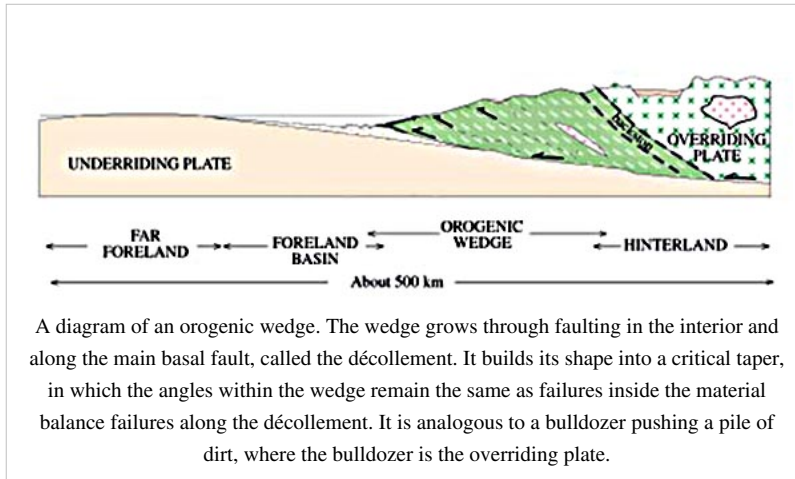
Petrologists use fluid inclusion data<sup>[41]</sup> and perform high temperature and pressure physical experiments<sup>[42]</sup> to understand the temperatures and pressures at which different mineral phases appear, and how they change through igneous<sup>[43]</sup> and metamorphic processes. This research can be extrapolated to the field to understand metamorphic processes and the conditions of crystallization of igneous rocks.<sup>[44]</sup> This work can also help to explain processes that occur within the Earth, such as subduction and magma chamber evolution.



A petrographic microscope, which is an optical microscope fitted with cross-polarizing lenses, a conoscopic lens, and compensators (plates of anisotropic materials; gypsum plates and quartz wedges are common), for crystallographic analysis.



## Structural geology



Structural geologists use microscopic analysis of oriented thin sections of geologic samples to observe the fabric within the rocks which gives information about strain within the crystal structure of the rocks. They also plot and combine measurements of geological structures in order to better understand the orientations of faults and folds in order to reconstruct the history of rock deformation in the area. In addition, they perform analog and numerical experiments of rock

deformation in large and small settings.

The analysis of structures is often accomplished by plotting the orientations various features onto stereonet. A stereonet is a stereographic projection of a sphere onto a plane, in which planes are projected as lines and lines are projected as points. These can be used to find the locations of fold axes, relationships between several faults, and relationships between other geologic structures.

Among the most well-known experiments in structural geology are those involving orogenic wedges, which are zones in which mountains are built along convergent tectonic plate boundaries.<sup>[45]</sup> In the analog versions of these experiments, horizontal layers of sand are pulled along a lower surface into a back stop, which results in realistic-looking patterns of faulting and the growth of a critically tapered (all angles remain the same) orogenic wedge.<sup>[46]</sup> Numerical models work in the same way as these analog models, though they are often more sophisticated and can include patterns of erosion and uplift in the mountain belt.<sup>[47]</sup> This helps to show the relationship between erosion and the shape of the mountain range. These studies can also give useful information about pathways for metamorphism through pressure, temperature, space, and time.<sup>[48]</sup>

## Stratigraphy

In the laboratory, stratigraphers analyze samples of stratigraphic sections that can be returned from the field, such as those from drill cores.<sup>[49]</sup> Stratigraphers also analyze data from geophysical surveys that show the locations of stratigraphic units in the subsurface.<sup>[50]</sup> Geophysical data and well logs can be combined to produce a better view of the subsurface, and stratigraphers often use computer programs to do this in three dimensions.<sup>[51]</sup> Stratigraphers can then use these data to reconstruct ancient processes occurring on the surface of the Earth,<sup>[52]</sup> interpret past environments, and locate areas for water, coal, and hydrocarbon extraction.



Exploration geologists examining a freshly recovered drill core.  
Chile, 1994

In the laboratory, biostratigraphers analyze rock samples from outcrop and drill cores for the fossils found in them.<sup>[49]</sup> These fossils help scientists to date the core and to understand the depositional environment in which the

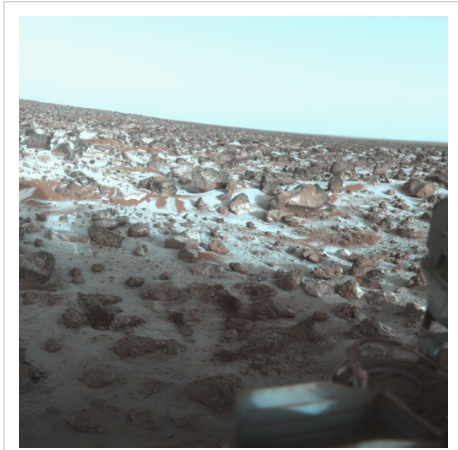
rock units formed. Geochronologists precisely date rocks within the stratigraphic section in order to provide better absolute bounds on the timing and rates of deposition.<sup>[53]</sup> Magnetic stratigraphers look for signs of magnetic reversals in igneous rock units within the drill cores.<sup>[49]</sup> Other scientists perform stable isotope studies on the rocks to gain information about past climate.<sup>[49]</sup>

## Planetary geology

With the advent of space exploration in the twentieth century, geologists have begun to look at other planetary bodies in the same way as the Earth. This led to the establishment of the field of planetary geology, sometimes known as astrogeology, in which geologic principles are applied to other bodies of the solar system.

Although the Greek-language-origin prefix *geo* refers to Earth, "geology" is often used in conjunction with the names of other planetary bodies when describing their composition and internal processes: examples are "the geology of Mars" and "Lunar geology". Specialised terms such as *selenology* (studies of the Moon), *areology* (of Mars), etc., are also in use.

Although planetary geologists are interested in all aspects of the planets, a significant focus is in the search for past or present life on other worlds. This has led to many missions whose purpose (or one of their purposes) is to examine planetary bodies for evidence of life. One of these is the Phoenix lander, which analyzed Martian polar soil for water and chemical and mineralogical constituents related to biological processes.



Surface of Mars as photographed by the Viking 2 lander December 9, 1977.

## Applied geology

### Economic geology

Economic geologists help locate and manage the Earth's natural resources, such as petroleum and coal, as well as mineral resources, which include metals such as iron, copper, and uranium.

### Mining geology

Mining geology consists of the extractions of mineral resources from the Earth. Some resources of economic interests include gemstones, metals, and many minerals such as asbestos, perlite, mica, phosphates, zeolites, clay, pumice, quartz, and silica, as well as elements such as sulfur, chlorine, and helium.



### Petroleum geology

Petroleum geologists study the locations of the subsurface of the Earth which can contain extractable hydrocarbons, especially petroleum and natural gas. Because many of these reservoirs are found in sedimentary basins,<sup>[54]</sup> they study the formation of these basins, as well as their sedimentary and tectonic evolution and the present-day positions of the rock units.

### Engineering geology

Engineering geology is the application of the geologic principles to engineering practice for the purpose of assuring that the geologic factors affecting the location, design, construction, operation and maintenance of engineering works are properly addressed.

In the field of civil engineering, geological principles and analyses are used in order to ascertain the mechanical principles of the material on which structures are built. This allows tunnels to be built without collapsing, bridges and skyscrapers to be built with sturdy foundations, and buildings to be built that will not settle in clay and mud.<sup>[55]</sup>

### Hydrology and environmental issues

Geology and geologic principles can be applied to various environmental problems, such as stream restoration, the restoration of brownfields, and the understanding of the interactions between natural habitat and the geologic environment. Groundwater hydrology, or hydrogeology, is used to locate groundwater,<sup>[56]</sup> which can often provide a ready supply of uncontaminated water and is especially important in arid regions,<sup>[57]</sup> and to monitor the spread of contaminants in groundwater wells.<sup>[56][58]</sup>

Geologists also obtain data through stratigraphy, boreholes, core samples, and ice cores. Ice cores<sup>[59]</sup> and sediment cores<sup>[60]</sup> are used to for paleoclimate reconstructions, which tell geologists about past and present temperature, precipitation, and sea level across the globe. These data are our primary source of information on global climate change outside of instrumental data.<sup>[61]</sup>

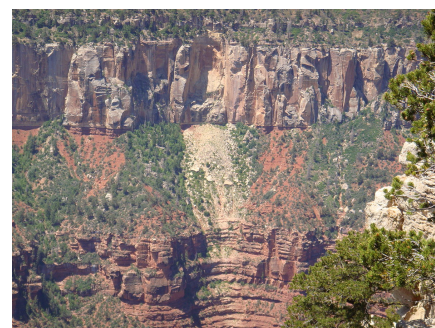
### Natural hazards

Geologists and geophysicists study natural hazards in order to enact safe building codes and warning systems that are used to prevent loss of property and life.<sup>[62]</sup> Examples of important natural hazards that are pertinent to geology (as opposed those that are mainly or only pertinent to meteorology) are:

- Avalanches
- Earthquakes
- Floods
- Landslides and debris flows
- River channel migration and avulsion
- Liquefaction
- Sinkholes
- Subsidence
- Tsunamis
- Volcanoes



Mud log in process, a common way to study the lithology when drilling oil wells.



Rockfall in the Grand Canyon

## Fields or related disciplines

- Earth science
  - Economic geology
    - Mining geology
    - Petroleum geology
  - Engineering geology
  - Environmental geology
  - Geoarchaeology
  - Geochemistry
    - Biogeochemistry
    - Isotope geochemistry
  - Geochronology
  - Geodetics
  - Geography
  - Geological modelling
  - Geometallurgy
  - Geomicrobiology
  - Geomorphology
  - Geomythology
  - Geophysics
  - Glaciology
  - Historical geology
  - Hydrogeology
  - Meteorology
  - Mineralogy
  - Oceanography
    - Marine geology
  - Paleoclimatology
  - Paleontology
    - Micropaleontology
    - Palynology
  - Petrology
  - Petrophysics
  - Plate tectonics
  - Sedimentology
  - Seismology
  - Soil science
    - Pedology (soil study)
  - Speleology
  - Stratigraphy
    - Biostratigraphy
    - Chronostratigraphy
    - Lithostratigraphy
  - Structural geology
  - Volcanology
-

## Regional geology

### By mountain range

- Geology of the Alps
- Geology of the Andes
- Geology of the Appalachians
- Geology of the Himalaya
- Geology of the Rocky Mountains

### By nations

- Geology of Australia
    - Geology of the Australian Capital Territory
    - Geology of Tasmania
    - Geology of Victoria
    - Geology of the Yilgarn Craton
  - Geology of China
  - Geology of Hong Kong
  - Geology of Europe
    - Geology of Iberia
    - Geology of Iceland
    - Geology of the Netherlands
    - Geology of Norway
    - Geology of Sweden
      - Geology of Gotland
  - Geology of the United Kingdom
    - Geology of England
      - Geology of Cheshire
      - Geology of Cornwall
        - Geology of Lizard, Cornwall
      - Geology of Dorset
      - Geology of Gloucestershire
      - Geology of Hampshire
      - Geology of East Sussex
      - Geology of Hertfordshire
      - Geology of Shropshire
      - Geology of Somerset
      - Geology of Yorkshire
    - Geology of Scotland
      - Geology of Orkney
    - Geology of Wales
    - Geology of Jersey
    - Geology of Guernsey
  - Geology of South America
    - Geology of Bolivia
    - Geology of Chile
-

- Geology of Colombia
- Geology of the Falkland Islands
- Geology of India
  - Geology of Sikkim
- Geology of Japan
- Geology of the Philippines
- Geology of New Zealand
- Geology of Vietnam
- Geology of the United States of America
  - *US geology by state:*
    - Geology of Alabama
    - Geology of Connecticut
    - Geology of Delaware
    - Geology of Florida
    - Geology of Georgia
    - Geology of Idaho
    - Geology of Illinois
    - Geology of Iowa
    - Geology of Kansas
    - Geology of Massachusetts
    - Geology of Minnesota
    - Geology of New Jersey
    - Geology of Oklahoma
    - Geology of Oregon
    - Geology of Pennsylvania
    - Geology of Tennessee
    - Geology of Texas
    - Geology of West Virginia
  - *US Geology by region or feature:*
    - Geology of the Appalachians
    - Geology of the Pacific Northwest
    - Geology of the Bryce Canyon area (*Utah*)
    - Geology of the Canyonlands area (*Utah*)
    - Geology of the Capitol Reef area (*Utah*)
    - Geology of the Death Valley area (*California*)
    - Geology of the Grand Canyon area (*Arizona*)
    - Geology of the Grand Teton area (*Wyoming*)
    - Geology of the Lassen area (*California*)
    - Geology of Mount Adams (*Washington*)
    - Geology of Mount Shasta (*California*)
    - Geology of the Yosemite area (*California*)
    - Geology of the Zion and Kolob canyons area (*Utah*)
    - Glacial geology of the Genesee River (*New York, Pennsylvania*)

## By planet

- Geology of Mars
- Geology of Mercury
- Geology of the Moon
- Geology of Venus

## Notes

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## External links

- (Petroleum) Geology Forum (<http://www.epgeology.com/>) a professional online community to discuss (petroleum) Geology!
  - Geology Rocks (<http://www.geologyrocks.co.uk/>) an Earth Sciences forum!
  - James Hutton's *Theory of the Earth* (<http://records.viu.ca/~johnstoi/essays/Hutton.htm>)
  - Charles Lyell's *Elements of Geology* ([http://books.google.com/books?id=-AcKAAAAIAAJ&printsec=frontcover&dq=Charles+Lyell&ei=YMOLSa\\_GE4uiyATW\\_Zy6BQ&client=firefox-a#PPR5,M1](http://books.google.com/books?id=-AcKAAAAIAAJ&printsec=frontcover&dq=Charles+Lyell&ei=YMOLSa_GE4uiyATW_Zy6BQ&client=firefox-a#PPR5,M1))
  - Charles Lyell's *Principles of Geology, or the Modern Changes of the Earth and its Inhabitants, Considered as Illustrative of Geology* ([http://books.google.com/books?id=O2YNAAAAYAAJ&printsec=frontcover&dq=Charles+Lyell&ei=YMOLSa\\_GE4uiyATW\\_Zy6BQ&client=firefox-a#PPR3,M1](http://books.google.com/books?id=O2YNAAAAYAAJ&printsec=frontcover&dq=Charles+Lyell&ei=YMOLSa_GE4uiyATW_Zy6BQ&client=firefox-a#PPR3,M1))
  - American Geophysical Union (<http://www.agu.org/>)
  - European Geosciences Union (<http://www.egu.eu/>)
  - Geological Society of America (<http://www.geosociety.org/>)
  - *Earth Science News, Maps, Dictionary, Articles, Jobs* (<http://geology.com/>)
  - *Geological Society of London* (<http://www.geolsoc.org.uk/>)
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# History of the Earth

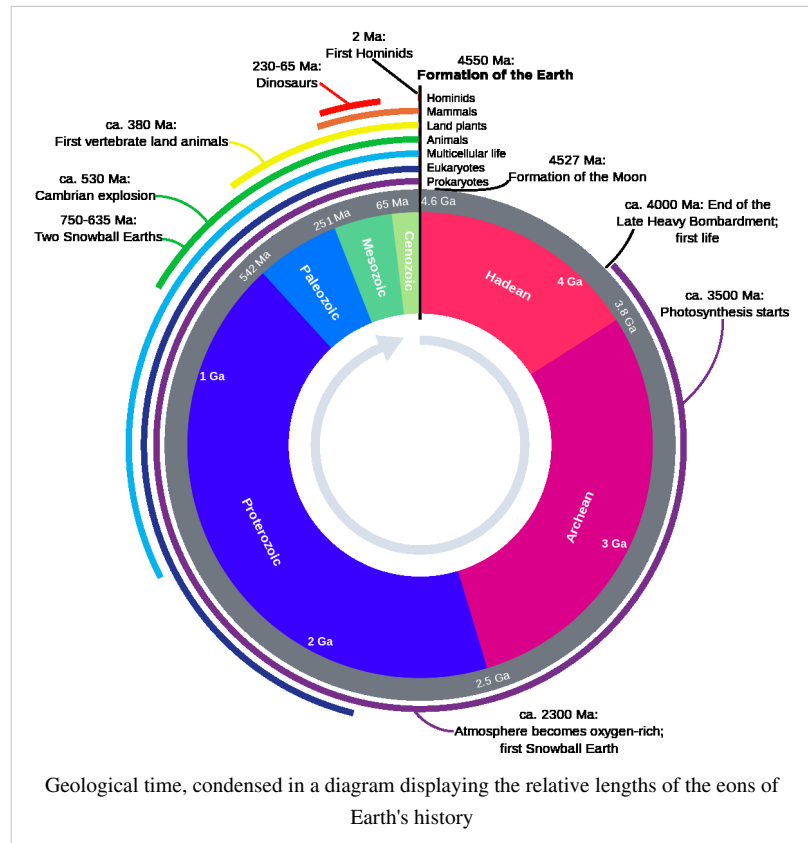
The **history of the Earth** describes the most important events and fundamental stages in the development of the planet Earth from its formation to the present day.<sup>[1][2]</sup> Nearly all branches of natural science have contributed to the understanding of the main events of the Earth's past. The age of Earth is approximately one-third of the age of the universe. An immense amount of biological and geological change has occurred in that time span.

Earth formed around 4.54 billion years ago by accretion from the solar nebula. Volcanic outgassing likely created the primordial atmosphere, but it contained almost no oxygen and would have been toxic to humans and most modern life. Much of the Earth was molten because of extreme volcanism and frequent collisions with other bodies.

One very large collision is thought to

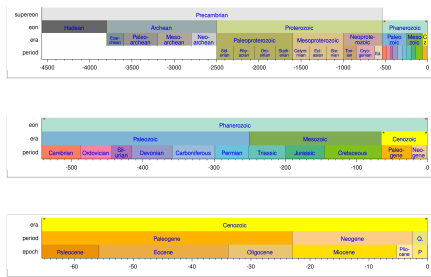
have been responsible for tilting the Earth at an angle and forming the Moon. Over time, such cosmic bombardments ceased, allowing the planet to cool and form a solid crust. Water that was brought here by comets and asteroids condensed into clouds and the oceans took shape. Earth was finally hospitable to life, and the earliest forms that arose enriched the atmosphere with oxygen. Life on Earth remained small and microscopic for at least one billion years. About 580 million years ago, complex multicellular life arose, and during the Cambrian period it experienced a rapid diversification into most major phyla. Around six million years ago, the primate lineage that would lead to chimpanzees (the closest relatives of humans) diverged from the lineage that would lead to modern humans.

Biological and geological change has been constantly occurring on our planet since the time of its formation. Organisms continuously evolve, taking on new forms or going extinct in response to an ever-changing planet. The process of plate tectonics has played a major role in the shaping of Earth's oceans and continents, as well as the life they harbor. The biosphere, in turn, has had a significant effect on the atmosphere and other abiotic conditions on the planet, such as the formation of the ozone layer, the proliferation of oxygen, and the creation of soil. Though humans are unable to perceive it due to their relatively brief life spans, this change is ongoing and will continue for the next few billion years.



## Geologic time scale

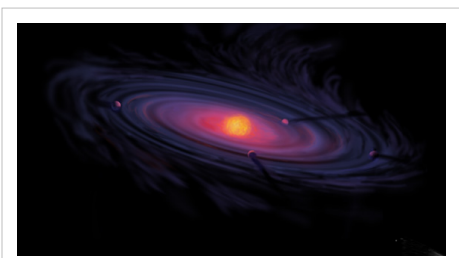
The history of the Earth is organized chronologically in a table known as the Geologic Time Scale, which is split into intervals based on stratigraphic analysis.<sup>[2][3]</sup> A full time scale can be found at the main article.



Millions of Years

## Solar System formation

The standard model for the formation of the Solar System (including the Earth) is the solar nebula hypothesis.<sup>[4]</sup> In this model, the Solar system formed from a large, rotating cloud of interstellar dust and gas called the solar nebula. It was composed of hydrogen and helium created shortly after the Big Bang 13.7 Ga (billion years ago) and heavier elements ejected by supernovae. About 4.5 Ga, the nebula began a contraction that may have been triggered by the shock wave of a nearby supernova.<sup>[5]</sup> A shock wave would have also made the nebula rotate. As the cloud began to accelerate, its angular momentum, gravity and inertia flattened it into a protoplanetary disk perpendicular to its axis of rotation. Small perturbations due to collisions and the angular momentum of other large debris created the means by which kilometer-sized protoplanets began to form, orbiting the nebular center.<sup>[6]</sup>



An artist's rendering of a protoplanetary disk

The center of the nebula, not having much angular momentum, collapsed rapidly, the compression heating it until nuclear fusion of hydrogen into helium began. After more contraction, a T Tauri star ignited and evolved into the Sun. Meanwhile, in the outer part of the nebula gravity caused matter to condense around density perturbations and dust particles, and the rest of the protoplanetary disk began separating into rings. In a process known as runaway accretion, successively larger fragments of dust and debris clumped together to form planets.<sup>[6]</sup> Earth formed in this manner about 4.54 billion years ago (with an uncertainty of 1%)<sup>[7][8][9][10]</sup> and was largely completed within 10–20 million years.<sup>[11]</sup> The solar wind of the newly formed T Tauri star cleared out most of the material in the disk that had not already condensed into larger bodies. The same process is expected to produce accretion disks around virtually all newly forming stars in the universe, some of which yield planets.<sup>[12]</sup>

The proto-Earth grew by accretion until its interior was hot enough to melt the heavy, siderophile metals. Having higher densities than the silicates, the metals sank. This *iron catastrophe* resulted in the separation of a primitive mantle and a (metallic) core only 10 million years after the Earth began to form, producing the layered structure of Earth and setting up the formation of Earth's magnetic field.<sup>[13]</sup> Earth's first atmosphere, captured from the solar nebula, was composed of light (atmophile) elements from the solar nebula, mostly hydrogen and helium. A combination of the solar wind and Earth's heat would have driven off this atmosphere, as a result of which the atmosphere is now depleted in these elements compared to cosmic abundances.<sup>[14]</sup>

## Hadean and Archean Eons

The first eon in Earth's history, the *Hadean*, begins with the Earth's formation and is followed by the *Archean* eon at 3.8 Ga.<sup>[2]:145</sup> The oldest rocks found on Earth date to about 4.0 Ga, and the oldest detrital zircon crystals in rocks to about 4.4 Ga,<sup>[15][16][17]</sup> soon after the formation of the Earth's crust and the Earth itself. The giant impact hypothesis for the Moon's formation states that shortly after formation of an initial crust, the proto-Earth was impacted by a smaller protoplanet, which ejected part of the mantle and crust into space and created the Moon.<sup>[18][19][20]</sup>

From crater counts on other celestial bodies it is inferred that a period of intense meteorite impacts, called the *Late Heavy Bombardment*, began about 4.1 Ga, and concluded around 3.8 Ga, at the end of the Hadean.<sup>[21]</sup> In addition, volcanism was severe due to the large heat flow and geothermal gradient.<sup>[22]</sup> Nevertheless, detrital zircon crystals dated to 4.4 Ga show evidence of having undergone contact with liquid water, suggesting that the planet already had oceans or seas at that time.<sup>[15]</sup>

By the beginning of the Archean, the Earth had cooled significantly. Most present life forms could not have survived in the Archean atmosphere, which lacked oxygen and an ozone layer. Nevertheless it is believed that primordial life began to evolve by the early Archean, with candidate fossils dated to around 3.5 Ga.<sup>[1]</sup> Some scientists even speculate that life could have begun during the early Hadean, as far back as 4.4 Ga, surviving the possible Late Heavy Bombardment period in hydrothermal vents below the Earth's surface.<sup>[23]</sup>

## Formation of the Moon

Earth's only natural satellite, the Moon, is larger relative to its planet than any other satellite in the solar system.<sup>[24]</sup> During the Apollo program, rocks from the Moon's surface were brought to Earth. Radiometric dating of these rocks has shown the Moon to be  $4.53 \pm .01$  billion years old,<sup>[25]</sup> at least 30 million years after the solar system was formed.<sup>[26]</sup> New evidence suggests the Moon formed even later,  $4.48 \pm 0.02$  Ga, or 70–110 million years after the start of the Solar System.<sup>[1]</sup>

Theories for the formation of the Moon must explain its late formation as well as the following facts. First, the Moon has a low density (3.3 times that of water, compared to 5.5 for the earth<sup>[27]</sup>) and a small metallic core. Second, there is virtually no water or other volatiles on the moon. Third, the Earth and Moon have the same oxygen isotopic signature (relative abundance of the oxygen isotopes). Of the theories that have been proposed to account for these phenomena, only one is widely accepted: The *giant impact hypothesis* proposes that the Moon originated after a body the size of Mars struck the proto-Earth a glancing blow.<sup>[1]:256[28][29]</sup>

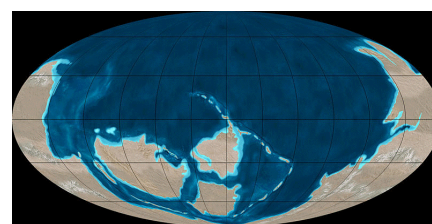
The collision between the impactor, sometimes named Theia,<sup>[26]</sup> and the Earth released about 100 million times more energy than the impact that caused the extinction of the dinosaurs. This was enough to vaporize some of the Earth's outer layers and melt both bodies.<sup>[28][1]:256</sup> A portion of the mantle material was ejected into orbit around the Earth. The giant impact hypothesis predicts that the Moon was depleted of metallic material,<sup>[30]</sup> explaining its abnormal composition.<sup>[31]</sup> The ejecta in orbit around the Earth could have condensed into a single body within a couple of weeks. Under the influence of its own gravity, the ejected material became a more spherical body: the Moon.<sup>[32]</sup>



Artist's impression of the enormous collision that likely formed the Moon

## First continents

Mantle convection, the process that drives plate tectonics today, is a result of heat flow from the Earth's interior to the Earth's surface.<sup>[33]:2</sup> It involves the creation of rigid tectonic plates at mid-oceanic ridges. These plates are destroyed by subduction into the mantle at subduction zones. During the early Archean (about 3.0 Ga) the mantle was much hotter than today, probably around 1600 °C,<sup>[34]:82</sup> so convection in the mantle was faster. While a process similar to present day plate tectonics did occur, this would have gone faster too. It is likely that during the Hadean and Archean, subduction zones were more common, and therefore tectonic plates were smaller.<sup>[1]:258</sup>



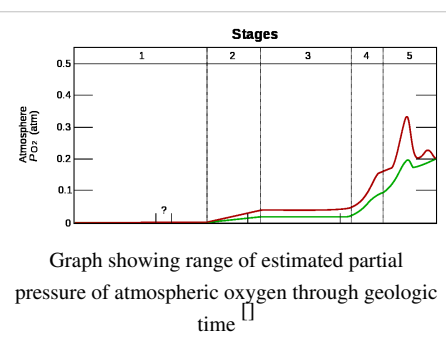
Animation showing the movement of Earth's continents throughout history, starting from the Cambrian period

The initial crust, formed when the Earth's surface first solidified, totally disappeared from a combination of this fast Hadean plate tectonics and the intense impacts of the Late Heavy Bombardment. However, it is thought that it was basaltic in composition, like today's oceanic crust, because little crustal differentiation had yet taken place.<sup>[1]:258</sup> The first larger pieces of continental crust, which is a product of differentiation of lighter elements during partial melting in the lower crust, appeared at the end of the Hadean, about 4.0 Ga. What is left of these first small continents are called cratons. These pieces of late Hadean and early Archean crust form the cores around which today's continents grew.<sup>[35]</sup>

The oldest rocks on Earth are found in the North American craton of Canada. They are tonalites from about 4.0 Ga. They show traces of metamorphism by high temperature, but also sedimentary grains that have been rounded by erosion during transport by water, showing rivers and seas existed then.<sup>[36]</sup> Cratons consist primarily of two alternating types of terranes. The first are so-called greenstone belts, consisting of low grade metamorphosed sedimentary rocks. These "greenstones" are similar to the sediments today found in oceanic trenches, above subduction zones. For this reason, greenstones are sometimes seen as evidence for subduction during the Archean. The second type is a complex of felsic magmatic rocks. These rocks are mostly tonalite, trondhjemite or granodiorite, types of rock similar in composition to granite (hence such terranes are called TTG-terranes). TTG-complexes are seen as the relicts of the first continental crust, formed by partial melting in basalt.<sup>[37]:Chapter 5</sup>

## Oceans and atmosphere

Earth is often described as having had three atmospheres. The first atmosphere, captured from the solar nebula, was composed of light (atmophile) elements from the solar nebula, mostly hydrogen and helium. A combination of the solar wind and Earth's heat would have driven off this atmosphere, as a result of which the atmosphere is now depleted in these elements compared to cosmic abundances.<sup>[14]</sup> After the impact, the molten Earth released volatile gases; and later more gases were released by volcanoes, completing a second atmosphere rich in greenhouse gases but poor in oxygen.<sup>[1]:256</sup> Finally, the third atmosphere, rich in oxygen, emerged when bacteria began to produce oxygen about 2.8 Ga.<sup>[38]:83–84,116–117</sup>



Graph showing range of estimated partial pressure of atmospheric oxygen through geologic time <sup>[1]</sup>

In early models for the formation of the atmosphere and ocean, the second atmosphere was formed by outgassing of volatiles from the Earth's interior. Now it is considered likely that many of the volatiles were delivered during accretion by a process known as *impact degassing* in which incoming bodies vaporize on impact. The ocean and atmosphere would therefore have started to form even as the Earth formed.<sup>[39]</sup> The new atmosphere probably contained water vapor, carbon dioxide, nitrogen, and smaller amounts of other gases.<sup>[1]</sup>

Planetesimals at a distance of 1 astronomical unit (AU), the distance of the Earth from the Sun, probably did not contribute any water to the Earth because the solar nebula was too hot for ice to form and the hydration of rocks by water vapor would have taken too long.<sup>[39][40]</sup> The water must have been supplied by meteorites from the outer asteroid belt and some large planetary embryos from beyond 2.5 AU.<sup>[39][41]</sup> Comets may also have contributed. Though most comets are today in orbits farther away from the Sun than Neptune, computer simulations show they were originally far more common in the inner parts of the solar system.<sup>[36]:130-132</sup>

As the planet cooled, clouds formed. Rain created the oceans. Recent evidence suggests the oceans may have begun forming as early as 4.4 Ga.<sup>[15]</sup> By the start of the Archean eon they already covered the Earth. This early formation has been difficult to explain because of a problem known as the faint young Sun paradox. Stars are known to get brighter as they age, and at the time of its formation the Sun would have been emitting only 70% of its current power. Many models predict that the Earth would have been covered in ice.<sup>[42][39]</sup> A likely solution is that there was enough carbon dioxide and methane to produce a greenhouse effect. The carbon dioxide would have been produced by volcanoes and the methane by early microbes. Another greenhouse gas, ammonium would have been ejected by volcanoes but quickly destroyed by ultraviolet radiation.<sup>[38]:83</sup>

## Origin of life

One of the reasons for interest in the early atmosphere and ocean is that they form the conditions under which life first arose. There are a lot of models, but little consensus, on how life emerged from non-living chemicals; chemical systems that have been created in the laboratory still fall well short of the minimum complexity for a living organism.<sup>[43][44]</sup>

The first step in the emergence of life may have been chemical reactions that produced many of the simpler organic compounds, including nucleobases and amino acids, that are the building blocks of life. An experiment in 1953 by Stanley Miller and Harold Urey showed that such molecules could form in an atmosphere of water, methane, ammonia and hydrogen with the aid of sparks to mimic the effect of lightning.<sup>[45]</sup> Although the atmospheric composition was likely different from the composition used by Miller and Urey, later experiments with more realistic compositions also managed to synthesize organic molecules.<sup>[46]</sup> Recent computer simulations have even shown that extraterrestrial organic molecules could have formed in the protoplanetary disk before the formation of the Earth.<sup>[47]</sup>

The next stage of complexity could have been reached from at least three possible starting points: self-replication, an organism's ability to produce offspring that are very similar to itself; metabolism, its ability to feed and repair itself; and external cell membranes, which allow food to enter and waste products to leave, but exclude unwanted substances.<sup>[48]</sup>

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### Replication first: RNA world

Even the simplest members of the three modern domains of life use DNA to record their "recipes" and a complex array of RNA and protein molecules to "read" these instructions and use them for growth, maintenance and self-replication.

The discovery that a kind of RNA molecule called a ribozyme can catalyze both its own replication and the construction of proteins led to the hypothesis that earlier life-forms were based entirely on RNA.<sup>[49]</sup> They could have formed an RNA world in which there were individuals but no species, as mutations and horizontal gene transfers would have meant that the offspring in each generation were quite likely to have different genomes from those that their parents started with.<sup>[50]</sup> RNA would later have been replaced by DNA, which is more stable and therefore can build longer genomes, expanding the range of capabilities a single organism can have.<sup>[51]</sup> Ribozymes remain as the main components of ribosomes, the "protein factories" of modern cells.<sup>[52]</sup>

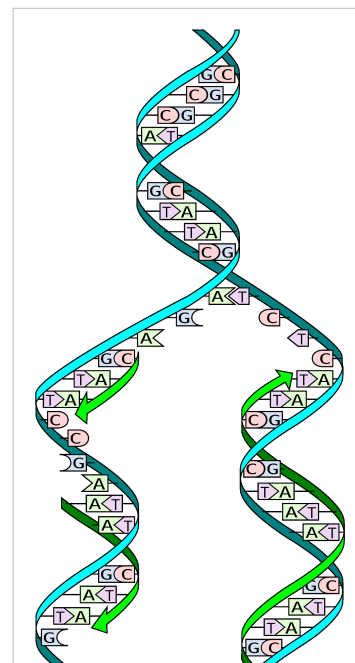
Although short, self-replicating RNA molecules have been artificially produced in laboratories,<sup>[53]</sup> doubts have been raised about whether natural non-biological synthesis of RNA is possible.<sup>[54][55][56]</sup> The earliest ribozymes may have been formed of simpler nucleic acids such as PNA, TNA or GNA, which would have been replaced later by RNA.<sup>[57][58]</sup> Other pre-RNA replicators have been posited, including crystals<sup>[59]:150</sup> and even quantum systems.<sup>[60]</sup>

In 2003 it was proposed that porous metal sulfide precipitates would assist RNA synthesis at about 100 °C (**unknown operator: u'strong'** °F) and ocean-bottom pressures near hydrothermal vents. In this hypothesis, lipid membranes would be the last major cell components to appear and until they did the proto-cells would be confined to the pores.<sup>[61]</sup>

### Metabolism first: Iron-sulfur world

Another long-standing hypothesis is that the first life was composed of protein molecules. Amino acids, the building blocks of proteins, are easily synthesized in plausible prebiotic conditions, as are small peptides (polymers of amino acids) that make good catalysts.<sup>[62]:295–297</sup> A series of experiments starting in 1997 showed that amino acids and peptides could form in the presence of carbon monoxide and hydrogen sulfide with iron sulfide and nickel sulfide as catalysts. Most of the steps in their assembly required temperatures of about 100 °C (**unknown operator: u'strong'** °F) and moderate pressures, although one stage required 250 °C (**unknown operator: u'strong'** °F) and a pressure equivalent to that found under 7 kilometres (**unknown operator: u'strong'** mi) of rock. Hence self-sustaining synthesis of proteins could have occurred near hydrothermal vents.<sup>[63]</sup>

A difficulty with the metabolism-first scenario is finding a way for organisms to evolve. Without the ability to replicate as individuals, aggregates of molecules would have "compositional genomes" (counts of molecular species in the aggregate) as the target of natural selection. However, a recent model shows that such a system is unable to evolve in response to natural selection.<sup>[64]</sup>



The replicator in virtually all known life is deoxyribonucleic acid. DNA is far more complex than the original replicator and its replication systems are highly elaborate.

### Membranes first: Lipid world

It has been suggested that double-walled "bubbles" of lipids like those that form the external membranes of cells may have been an essential first step.<sup>[65]</sup> Experiments that simulated the conditions of the early Earth have reported the formation of lipids, and these can spontaneously form liposomes, double-walled "bubbles", and then reproduce themselves. Although they are not intrinsically information-carriers as nucleic acids are, they would be subject to natural selection for longevity and reproduction. Nucleic acids such as RNA might then have formed more easily within the liposomes than they would have outside.<sup>[66]</sup>

### The clay theory

Some clays, notably montmorillonite, have properties that make them plausible accelerators for the emergence of an RNA world: they grow by self-replication of their crystalline pattern, are subject to an analog of natural selection (as the clay "species" that grows fastest in a particular environment rapidly becomes dominant), and can catalyze the formation of RNA molecules.<sup>[1]</sup> Although this idea has not become the scientific consensus, it still has active supporters.<sup>[67]:150–158[59]</sup>

Research in 2003 reported that montmorillonite could also accelerate the conversion of fatty acids into "bubbles", and that the bubbles could encapsulate RNA attached to the clay. Bubbles can then grow by absorbing additional lipids and dividing. The formation of the earliest cells may have been aided by similar processes.<sup>[68]</sup>

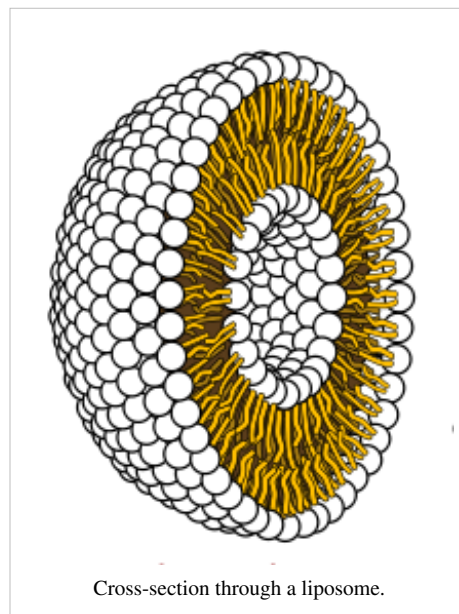
A similar hypothesis presents self-replicating iron-rich clays as the progenitors of nucleotides, lipids and amino acids.<sup>[69]</sup>

### Last common ancestor

It is believed that of this multiplicity of protocells, only one line survived. Current phylogenetic evidence suggests that the last universal common ancestor (LUCA) lived during the early Archean eon, perhaps 3.5 Ga or earlier.<sup>[70][71]</sup> This LUCA cell is the ancestor of all life on Earth today. It was probably a prokaryote, possessing a cell membrane and probably ribosomes, but lacking a nucleus or membrane-bound organelles such as mitochondria or chloroplasts. Like all modern cells, it used DNA as its genetic code, RNA for information transfer and protein synthesis, and enzymes to catalyze reactions. Some scientists believe that instead of a single organism being the last universal common ancestor, there were populations of organisms exchanging genes by lateral gene transfer.<sup>[70]</sup>

### Proterozoic Eon

The Proterozoic eon lasted from 2.5 Ga to 542 Ma (million years ago).<sup>[2]:130</sup> In this time span, cratons grew into continents with modern sizes. The change to an oxygen-rich atmosphere was a crucial development. Life developed from prokaryotes into eukaryotes and multicellular forms. The Proterozoic saw a couple of severe ice ages called snowball Earths. After the last Snowball Earth about 600 Ma, the evolution of life on Earth accelerated. About 580 Ma, the Ediacara biota formed the prelude for the Cambrian Explosion.



## Oxygen revolution

The earliest cells absorbed energy and food from the environment around them. They used fermentation, the breakdown of more complex compounds into less complex compounds with less energy, and used the energy so liberated to grow and reproduce. Fermentation can only occur in an *anaerobic* (oxygen-free) environment. The evolution of photosynthesis made it possible for cells to manufacture their own food.<sup>[72]:377</sup>

Most of the life that covers the surface of the Earth depends directly or indirectly on photosynthesis. The most common form, oxygenic photosynthesis, turns carbon dioxide, water and sunlight into food. It captures the energy of sunlight in energy-rich molecules such as ATP, which then provide the energy to make sugars. To supply the electrons in the circuit, hydrogen is stripped from water, leaving oxygen as a waste product.<sup>[73]</sup> Some organisms, including purple bacteria and green sulfur bacteria, use an anoxygenic form of photosynthesis that use alternatives to hydrogen stripped from water as electron donors; examples are hydrogen sulfide, sulfur and iron. Such organisms are mainly restricted to extreme environments such as hot springs and hydrothermal vents.<sup>[72]:379–382[74]</sup>

The simpler anoxygenic form arose about 3.8 Ga, not long after the appearance of life. The timing of oxygenic photosynthesis is more controversial; it had certainly appeared by about 2.4 Ga, but some researchers put it back as far as 3.2 Ga.<sup>[73]</sup> The latter "probably increased global productivity by at least two or three orders of magnitude."<sup>[75][1]</sup> Among the oldest remnants of oxygen-producing lifeforms are fossil stromatolites.<sup>[75][1][1]</sup>

At first, the released oxygen was bound up with limestone, iron, and other minerals. The oxidized iron appears as red layers in geological strata called banded iron formations that formed in abundance during the Siderian period (between 2500 Ma and 2300 Ma).<sup>[2]:133</sup> When most of the exposed readily reacting minerals were oxidized, oxygen finally began to accumulate in the atmosphere. Though each cell only produced a minute amount of oxygen, the combined metabolism of many cells over a vast time transformed Earth's atmosphere to its current state. This was Earth's third atmosphere.<sup>[76]:50–51[38]:83–84,116–117</sup>

Some of the oxygen was stimulated by incoming ultraviolet radiation to form ozone, which collected in a layer near the upper part of the atmosphere. The ozone layer absorbed, and still absorbs, a significant amount of the ultraviolet radiation that once had passed through the atmosphere. It allowed cells to colonize the surface of the ocean and eventually the land: without the ozone layer, ultraviolet radiation bombarding land and sea would have caused unsustainable levels of mutation in exposed cells.<sup>[77][36]:219–220</sup>

Photosynthesis had another major impact. Oxygen was toxic; much life on Earth probably died out as its levels rose in what is known as the *oxygen catastrophe*. Resistant forms survived and thrived, and some developed the ability to use oxygen to increase their metabolism and obtain more energy from the same food.<sup>[77]</sup>



Lithified stromatolites on the shores of Lake Thetis, Western Australia. Archean stromatolites are the first direct fossil traces of life on Earth.



A banded iron formation from the 3.15 Ga Moories Group, Barberton Greenstone Belt, South Africa. Red layers represent the times when oxygen was available, gray layers were formed in anoxic circumstances.

## Snowball Earth

The natural evolution of the Sun made it progressively more luminous during the Archean and Proterozoic eons; the Sun's luminosity increases 6% every billion years.<sup>[36]:165</sup> As a result, the Earth began to receive more heat from the Sun in the Proterozoic eon. However, the Earth did not get warmer. Instead, the geological record seems to suggest it cooled dramatically during the early Proterozoic. Glacial deposits found in South Africa date back to 2.2 Ga, at which time paleomagnetic evidence puts them near the equator. Thus, this glaciation, known as the Makganyene glaciation, may have been global. Some scientists suggest this and following Proterozoic ice ages were so severe that the planet was totally frozen over from the poles to the equator, a hypothesis called Snowball Earth.<sup>[78]</sup>

The ice age around 2.3 Ga could have been directly caused by the increased oxygen concentration in the atmosphere, which caused the decrease of methane (CH<sub>4</sub>) in the atmosphere. Methane is a strong greenhouse gas, but with oxygen it reacts to form CO<sub>2</sub>, a less effective greenhouse gas.<sup>[36]:172</sup> When free oxygen became available in the atmosphere, the concentration of methane could have decreased dramatically, enough to counter the effect of the increasing heat flow from the Sun.<sup>[79]</sup>

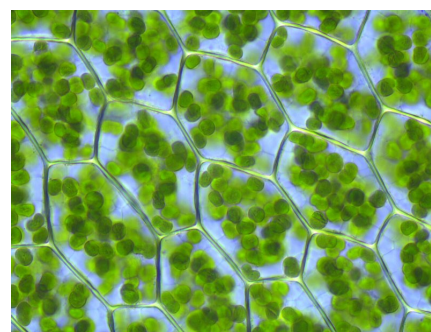
## Emergence of eukaryotes

Modern taxonomy classifies life into three domains. The time of the origin of these domains is uncertain. The Bacteria domain probably first split off from the other forms of life (sometimes called Neomura), but this supposition is controversial. Soon after this, by 2 Ga,<sup>[80]</sup> the Neomura split into the Archaea and the Eukarya. Eukaryotic cells (Eukarya) are larger and more complex than prokaryotic cells (Bacteria and Archaea), and the origin of that complexity is only now becoming known.

Around this time, the first proto-mitochondrion was formed. A bacterial cell related to today's *Rickettsia*,<sup>[81]</sup> which had evolved to metabolize oxygen, entered a larger prokaryotic cell, which lacked that capability. Perhaps the large cell attempted to digest the smaller one but failed (possibly due to the evolution of prey defenses). The smaller cell may have tried to parasitize the larger one. In any case, the smaller cell survived inside the larger cell. Using oxygen, it metabolized the larger cell's waste products and derived more energy. Part of this excess energy was returned to the host. The smaller cell replicated inside the larger one. Soon, a stable symbiosis developed between the large cell and the smaller cells inside it. Over time, the host cell acquired some of the genes of the smaller cells, and the two kinds became dependent on each other: the larger cell could not survive without the energy produced by the smaller ones, and these in turn could not survive without the raw materials provided by the larger cell. The whole cell is now considered a single organism, and the smaller cells are classified as organelles called mitochondria.<sup>[82]</sup>

A similar event occurred with photosynthetic cyanobacteria<sup>[83]</sup> entering large heterotrophic cells and becoming chloroplasts.<sup>[76]:60–61[84]:536–539</sup> Probably as a result of these changes, a line of cells capable of photosynthesis split off from the other eukaryotes more than 1 billion years ago. There were probably several such inclusion events. Besides the well-established endosymbiotic theory of the cellular origin of mitochondria and chloroplasts, there are theories that cells led to peroxisomes, spirochetes led to cilia and flagella, and that perhaps a DNA virus led to the cell nucleus,<sup>[85],[86]</sup> though none of them is widely accepted.<sup>[87]</sup>

Archaeans, bacteria, and eukaryotes continued to diversify and to become more complex and better adapted to their environments. Each domain repeatedly split into multiple lineages, although little is known about the history of the archaea and bacteria. Around 1.1 Ga, the supercontinent Rodinia was assembling.<sup>[88]</sup> The plant, animal, and fungi lines had split, though they still existed as solitary cells. Some of these lived in colonies, and gradually a division of labor began to take place; for instance, cells on the periphery might have started to assume different roles from those



Chloroplasts in the cells of a moss

in the interior. Although the division between a colony with specialized cells and a multicellular organism is not always clear, around 1 billion years ago<sup>[89]</sup> the first multicellular plants emerged, probably green algae.<sup>[90]</sup> Possibly by around 900 Ma<sup>[84]:488</sup> true multicellularity had also evolved in animals.

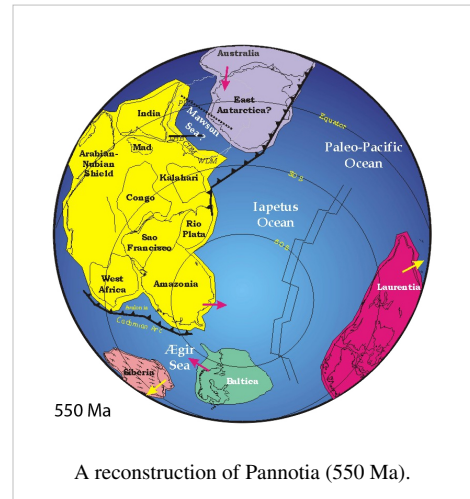
At first it probably resembled today's sponges, which have totipotent cells that allow a disrupted organism to reassemble itself.<sup>[84]:483-487</sup> As the division of labor was completed in all lines of multicellular organisms, cells became more specialized and more dependent on each other; isolated cells would die.

## Supercontinents in the Proterozoic

Reconstructions of tectonic plate movement in the past 250 million years (the Cenozoic and Mesozoic eras) can be made reliably using fitting of continental margins, ocean floor magnetic anomalies and paleomagnetic poles. No ocean crust dates back further than that, so earlier reconstructions are more difficult. Paleomagnetic poles are supplemented by geologic evidence such as orogenic belts, which mark the edges of ancient plates, and past distributions of flora and fauna. The further back in time, the scarcer and harder to interpret the data get and the more diverse the reconstructions.<sup>[91]:370</sup>

Throughout the history of the Earth, there have been times when continents collided and formed a supercontinent, which later broke up into new continents. About 1000 to 830 Ma, most continental mass was united in the supercontinent Rodinia.<sup>[91]:370[92]</sup> Rodinia may have been preceded by Early-Middle Proterozoic continents called Nuna and Columbia.<sup>[91]:374[93][94]</sup>

After the break-up of Rodinia about 800 Ma, the continents may have formed another short-lived supercontinent, Pannotia, around 550 Ma. The hypothetical supercontinent is sometimes referred to as Pannotia or Vendia.<sup>[95]:321-322</sup> The evidence for it is a phase of continental collision known as the Pan-African orogeny, which joined the continental masses of current-day Africa, South-America, Antarctica and Australia. The existence of Pannotia depends on the timing of the rifting between Gondwana (which included most of the landmass now in the Southern Hemisphere, as well as the Arabian Peninsula and the Indian subcontinent) and Laurentia (roughly equivalent to current-day North America).<sup>[91]:374</sup> It is at least certain that by the end of the Proterozoic eon, most of the continental mass lay united in a position around the south pole.<sup>[96]</sup>





## Late Proterozoic climate and life

The end of the Proterozoic saw at least two Snowball Earths, so severe that the surface of the oceans may have been completely frozen. This happened about 710 and 640 Ma, in the Cryogenian period.<sup>[97]</sup> These severe glaciations are less easy to explain than the early Proterozoic Snowball Earth. Most paleoclimatologists think the cold episodes were linked to the formation of the supercontinent Rodinia. Because Rodinia was centered on the equator, rates of chemical weathering increased and carbon dioxide (CO<sub>2</sub>) was taken from the atmosphere. Because CO<sub>2</sub> is an important greenhouse gas, climates cooled globally. In the same way, during the Snowball Earths most of the continental surface was covered with permafrost, which decreased chemical weathering again, leading to the end of the glaciations. An alternative hypothesis is that enough carbon dioxide escaped through volcanic outgassing that the resulting greenhouse effect raised global temperatures.<sup>[98]</sup> Increased volcanic activity resulted from the break-up of Rodinia at about the same time.



A 580 million year old fossil of *Spriggina floundensi*, an animal from the Ediacaran period. Such life forms could have been ancestors to the many new forms that originated in the Cambrian Explosion.

The Cryogenian period was followed by the Ediacaran period, which was characterized by a rapid development of new multicellular lifeforms.<sup>[99]</sup> Whether there is a connection between the end of the severe ice ages and the increase in diversity of life is not clear, but it does not seem coincidental. The new forms of life, called Ediacara biota, were larger and more diverse than ever. Though the taxonomy of most Ediacaran life forms is unclear, some were ancestors of groups of modern life.<sup>[100]</sup> Important developments were the origin of muscular and neural cells. None of the Ediacaran fossils had hard body parts like skeletons. These first appear after the boundary between the Proterozoic and Phanerozoic eons or Ediacaran and Cambrian periods.

## Phanerozoic Eon

The Phanerozoic is the major eon of life on Earth. It consists of three eras: The Paleozoic, Mesozoic, and Cenozoic,<sup>[3]</sup> and is the time when multi-cellular life greatly diversified into almost all of the organisms known today.<sup>[101]</sup>

### Paleozoic Era

The Paleozoic era (meaning: *era of old life forms*) was the first and longest era of the Phanerozoic eon, lasting from 542 to 251 Ma.<sup>[3]</sup> During the Paleozoic, many modern groups of life came into existence. Life colonized the land, first plants, then animals. Life usually evolved slowly. At times, however, there are sudden radiations of new species or mass extinctions. These bursts of evolution were often caused by unexpected changes in the environment resulting from natural disasters such as volcanic activity, meteorite impacts or climate changes.

The continents formed at the break-up of Pannotia and Rodinia at the end of the Proterozoic would slowly move together again during the Paleozoic. This would eventually result in phases of mountain building that created the supercontinent Pangaea in the late Paleozoic.



## Cambrian explosion

The rate of the evolution of life as recorded by fossils accelerated in the Cambrian period (542–488 Ma).<sup>[3]</sup> The sudden emergence of many new species, phyla, and forms in this period is called the Cambrian Explosion. The biological fomenting in the Cambrian Explosion was unprecedented before and since that time.<sup>[36]:229</sup> Whereas the Ediacaran life forms appear yet primitive and not easy to put in any modern group, at the end of the Cambrian most modern phyla were already present. The development of hard body parts such as shells, skeletons or exoskeletons in animals like molluscs, echinoderms, crinoids and arthropods (a well-known group of arthropods from the lower Paleozoic are the trilobites) made the preservation and fossilization of such life forms easier than those of their Proterozoic ancestors. For this reason, much more is known about life in and after the Cambrian than about that of older periods. Some of these Cambrian groups appear complex but are quite different from modern life; examples are *Anomalocaris* and *Haikouichthys*.



Trilobites first appeared during the Cambrian period and were among the most widespread and diverse groups of Paleozoic organisms.

During the Cambrian, the first vertebrate animals, among them the first fishes, had appeared.<sup>[84]:357</sup> A creature that could have been the ancestor of the fishes, or was probably closely related to it, was *Pikaia*. It had a primitive notochord, a structure that could have developed into a vertebral column later. The first fishes with jaws (Gnathostomata) appeared during the Ordovician. The colonisation of new niches resulted in massive body sizes. In this way, fishes with increasing sizes evolved during the early Paleozoic, such as the titanic placoderm *Dunkleosteus*, which could grow 7 meters long.

The diversity of life forms did not increase greatly because of a series of mass extinctions that define widespread biostratigraphic units called *biomeres*.<sup>[102]</sup> After each extinction pulse, the shelf regions were repopulated by similar life forms that may have been evolving slowly elsewhere.<sup>[103]</sup> By the late Cambrian, the trilobites had reached their greatest diversity and dominated nearly all fossil assemblages.<sup>[104]:34</sup> The boundary between the Cambrian and Ordovician (the following period, 488 to 444 <sup>[105]</sup> million years ago) is not associated with a recognized major extinction.<sup>[106]:3</sup>

### Paleozoic tectonics, paleogeography and climate

At the end of the Proterozoic, the supercontinent Pannotia had broken apart in the smaller continents Laurentia, Baltica, Siberia and Gondwana.<sup>[107]</sup> During periods when continents move apart, more oceanic crust is formed by volcanic activity. Because young volcanic crust is relatively hotter and less dense than old oceanic crust, the ocean floors will rise during such periods. This causes the sea level to rise. Therefore, in the first half of the Paleozoic, large areas of the continents were below sea level.

Early Paleozoic climates were warmer than today, but the end of the Ordovician saw a short ice age during which glaciers covered the south pole, where the huge continent Gondwana was situated. Traces of glaciation from this period are only found on former Gondwana. During the Late Ordovician ice age, a few mass extinctions took place, in which many brachiopods, trilobites, Bryozoa and corals disappeared. These marine species could probably not contend with the decreasing temperature of the sea water.<sup>[108]</sup> After the extinctions new species evolved, more diverse and better adapted. They would fill the niches left by the extinct species.

The continents Laurentia and Baltica collided between 450 and 400 Ma, during the Caledonian Orogeny, to form Laurussia (also known as Euramerica).<sup>[109]</sup> Traces of the mountain belt which resulted from this collision can be found in Scandinavia, Scotland and the northern Appalachians. In the Devonian period (416–359 Ma)<sup>[3]</sup> Gondwana and Siberia began to move towards Laurussia. The collision of Siberia with Laurussia caused the Uralian Orogeny, the collision of Gondwana with Laurussia is called the Variscan or Hercynian Orogeny in Europe or the Alleghenian Orogeny in North America. The latter phase took place during the Carboniferous period (359–299 Ma)<sup>[3]</sup> and resulted in the formation of the last supercontinent, Pangaea.<sup>[37]</sup>

### Colonization of land

Oxygen accumulation from photosynthesis resulted in the formation of an ozone layer that absorbed much of the Sun's ultraviolet radiation, meaning unicellular organisms that reached land were less likely to die, and prokaryotes began to multiply and become better adapted to survival out of the water. Prokaryote lineages<sup>[110]</sup> had probably colonized the land as early as 2.6 Ga<sup>[111]</sup> even before the origin of the eukaryotes. For a long time, the land remained barren of multicellular organisms. The supercontinent Pannotia formed around 600 Ma and then broke apart a short 50 million years later.<sup>[112]</sup> Fish, the earliest vertebrates, evolved in the oceans around 530 Ma.<sup>[84]:354</sup> A major

Cambrian–Ordovician, plants (probably resembling algae) and fungi started growing at the edges of the water, and then out of it.<sup>[113]:138–140</sup> The oldest fossils of land fungi and plants date to 480–460 Ma, though molecular evidence suggests the fungi may have colonized the land as early as 1000 Ma and the plants 700 Ma.<sup>[114]</sup> Initially remaining close to the water's edge, mutations and variations resulted in further colonization of this new environment. The timing of the first animals to leave the oceans is not precisely known: the oldest clear evidence is of arthropods on land around 450 Ma,<sup>[115]</sup> perhaps thriving and becoming better adapted due to the vast food source provided by the



Pangaea was a supercontinent that existed from about 300 to 180 Ma. The outlines of the modern continents and other landmasses are indicated on this map.



Artist's conception of Devonian flora

terrestrial plants. There is also unconfirmed evidence that arthropods may have appeared on land as early as 530 Ma.<sup>[116]</sup>

### Evolution of tetrapods

At the end of the Ordovician period, 443 Ma,<sup>[3]</sup> additional extinction events occurred, perhaps due to a concurrent ice age.<sup>[108]</sup> Around 380 to 375 Ma, the first tetrapods evolved from fish.<sup>[117]</sup> It is thought that perhaps fins evolved to become limbs which allowed the first tetrapods to lift their heads out of the water to breathe air. This would allow them to live in oxygen-poor water or pursue small prey in shallow water.<sup>[117]</sup> They may have later ventured on land for brief periods. Eventually, some of them became so well adapted to terrestrial life that they spent their adult lives on land, although they hatched in the water and returned to lay their eggs. This was the origin of the amphibians. About 365 Ma, another period of extinction occurred, perhaps as a result of global cooling.<sup>[118]</sup> Plants evolved seeds, which dramatically accelerated their spread on land, around this time (by approximately 360 Ma).<sup>[119][120]</sup>



*Tiktaalik*, a fish with limb-like fins and a predecessor of tetrapods. Reconstruction from fossils about 375 million years old.

About 20 million years later (340 Ma<sup>[84]:293–296</sup>), the amniotic egg evolved, which could be laid on land, giving a survival advantage to tetrapod embryos. This resulted in the divergence of amniotes from amphibians. Another 30 million years (310 Ma<sup>[84]:254–256</sup>) saw the divergence of the synapsids (including mammals) from the sauropsids (including birds and reptiles). Other groups of organisms continued to evolve, and lines diverged—in fish, insects, bacteria, and so on—but less is known of the details.

### Mesozoic Era

The Mesozoic ("middle life") era lasted from 251 Ma to 65.5 Ma.<sup>[3]</sup> It is subdivided into the Triassic, Jurassic, and Cretaceous periods. The era began with the Permian–Triassic extinction event, the most severe extinction event in the fossil record; 95% of the species on Earth died out.<sup>[121]</sup> It ended with the Cretaceous–Paleogene extinction event that wiped out the dinosaurs. The Permian–Triassic event was possibly caused by some combination of the Siberian Traps volcanic event, an asteroid impact, methane hydrate gasification, sea level fluctuations, and a major anoxic event. Either the proposed Wilkes Land crater<sup>[122]</sup> in Antarctica or Bedout structure off the northwest coast of Australia may indicate an impact connection with the Permian–Triassic extinction. But it remains uncertain whether either these or other proposed Permian–Triassic boundary craters are either real impact craters or even contemporaneous with the Permian–Triassic extinction event. Life persevered, and around 230 Ma, dinosaurs split off from their reptilian ancestors.<sup>[1]</sup> The Triassic–Jurassic extinction event at 200 Ma spared many of the dinosaurs,<sup>[3][123]</sup> and they soon became dominant among the vertebrates. Though some of the mammalian lines began to separate during this period, existing mammals were probably small animals resembling shrews.<sup>[84]:169</sup>



Dinosaurs were the dominant terrestrial vertebrates throughout most of the Mesozoic

By 180 Ma, Pangaea broke up into Laurasia and Gondwana. The boundary between avian and non-avian dinosaurs is not clear, but *Archaeopteryx*, traditionally considered one of the first birds, lived around 150 Ma.<sup>[124]</sup> The earliest evidence for the angiosperms evolving flowers is during the Cretaceous period, some 20 million years later (132 Ma).<sup>[125]</sup> Competition with birds drove many pterosaurs to extinction and the dinosaurs were probably already in decline when, 65 Ma, a 10-kilometre (**unknown operator: u'strong'** mi) asteroid struck Earth just off the

Yucatán Peninsula where the Chicxulub crater is today. This ejected vast quantities of particulate matter and vapor into the air that occluded sunlight, inhibiting photosynthesis. Most large animals, including the non-avian dinosaurs, became extinct,<sup>[126]</sup> marking the end of the Cretaceous period and Mesozoic era.

## Cenozoic Era

The Cenozoic era began at 65.6 Ma,<sup>[3]</sup> and is subdivided into the Paleogene and Neogene periods. Mammals and birds were able to survive the Cretaceous–Paleogene extinction event which killed off the dinosaurs and many other forms of life, and this is the era in which they diversified into their modern forms.

### Diversification of mammals

Mammals have existed since the late Triassic, but prior to the Cretaceous–Paleogene extinction event they remained small and generalized. During the Cenozoic, mammals rapidly diversified to fill the niches that the dinosaurs and other extinct animals had left behind, becoming the dominant vertebrates and creating many of the modern orders. With many marine reptiles extinct, some mammals began living in the oceans and became cetaceans. Others became felids and canids, swift and agile land predators. The dryer global climate of the Cenozoic led to the expansion of grasslands and the evolution of grazing and hooved mammals such as equids and bovids. Other mammals adapted to arboreal living and became the primates, of which one lineage would lead to modern humans.

### Human evolution

A small African ape living around 6 Ma was the last animal whose descendants would include both modern humans and their closest relatives, the chimpanzees.<sup>[84]:100–101</sup> Only two branches of its family tree have surviving descendants. Very soon after the split, for reasons that are still unclear, apes in one branch developed the ability to walk upright.<sup>[84]:95–99</sup> Brain size increased rapidly, and by 2 Ma, the first animals classified in the genus *Homo* had appeared.<sup>[113]:300</sup> Of course, the line between different species or even genera is somewhat arbitrary as organisms continuously change over generations. Around the same time, the other branch split into the ancestors of the common chimpanzee and the ancestors of the bonobo as evolution continued simultaneously in all life forms.<sup>[84]:100–101</sup>



Reconstruction of *Australopithecus africanus*, an early hominid.

The ability to control fire probably began in *Homo erectus* (or *Homo ergaster*), probably at least 790,000 years ago<sup>[127]</sup> but perhaps as early as 1.5 Ma.<sup>[84]:67</sup> The use and discovery of controlled fire may even predate *Homo erectus*. Fire was possibly used by the early Lower Paleolithic (Oldowan) hominid *Homo habilis* or strong australopithecines such as *Paranthropus*.<sup>[128]</sup>

It is more difficult to establish the origin of language; it is unclear whether *Homo erectus* could speak or if that capability had not begun until *Homo sapiens*.<sup>[84]:67</sup> As brain size increased, babies were born earlier, before their heads grew too large to pass through the pelvis. As a result, they exhibited more plasticity, and thus possessed an increased capacity to learn and required a longer period of dependence. Social skills became more complex, language became more sophisticated, and tools became more elaborate. This contributed to further cooperation and intellectual development.<sup>[129]:7</sup> Modern humans (*Homo sapiens*) are believed to have originated around 200,000 years ago or earlier in Africa; the oldest fossils date back to around 160,000 years ago.<sup>[130]</sup>

The first humans to show signs of spirituality are the Neanderthals (usually classified as a separate species with no surviving descendants); they buried their dead, often with no sign of food or tools.<sup>[131]:17</sup> However, evidence of more sophisticated beliefs, such as the early Cro-Magnon cave paintings (probably with magical or religious significance)<sup>[131]:17–19</sup> did not appear until 32,000 years ago.<sup>[132]</sup> Cro-Magnons also left behind stone figurines such



as Venus of Willendorf, probably also signifying religious belief.<sup>[131]:17–19</sup> By 11,000 years ago, *Homo sapiens* had reached the southern tip of South America, the last of the uninhabited continents (except for Antarctica, which remained undiscovered until 1820 AD).<sup>[133]</sup> Tool use and communication continued to improve, and interpersonal relationships became more intricate.

## Civilization

Further information: History of Africa, History of the Americas, History of Antarctica, and History of Eurasia

Throughout more than 90% of its history, *Homo sapiens* lived in small bands as nomadic hunter-gatherers.<sup>[129]:8</sup> As language became more complex, the ability to remember and communicate information resulted in a new replicator: the meme.<sup>[134]</sup> Ideas could be exchanged quickly and passed down the generations. Cultural evolution quickly outpaced biological evolution, and history proper began. Between 8500 and 7000 BC, humans in the Fertile Crescent in Middle East began the systematic husbandry of plants and animals: agriculture.<sup>[135]</sup> This spread to neighboring regions, and developed independently elsewhere, until most *Homo sapiens* lived sedentary lives in permanent settlements as farmers. Not all societies abandoned nomadism, especially those in isolated areas of the globe poor in domesticable plant species, such as Australia.<sup>[136]</sup> However, among those civilizations that did adopt agriculture, the relative stability and increased productivity provided by farming allowed the population to expand.

Agriculture had a major impact; humans began to affect the environment as never before. Surplus food allowed a priestly or governing class to arise, followed by increasing division of labor. This led to Earth's first civilization at Sumer in the Middle East, between 4000 and 3000 BC.<sup>[129]:15</sup> Additional civilizations quickly arose in ancient Egypt, at the Indus River valley and in China. The invention of writing enabled complex societies to arise: record-keeping and libraries served as a storehouse of knowledge and increased the cultural transmission of information. Humans no longer had to spend all their time working for survival—curiosity and education drove the pursuit of knowledge and wisdom.

Various disciplines, including science (in a primitive form), arose. New civilizations sprang up, traded with one another, and fought for territory and resources. Empires soon began to develop. By around 500 BC, there were advanced civilizations in the Middle East, Iran, India, China, and Greece, at times expanding, at times entering into decline.<sup>[129]:3</sup> The fundamentals of the Western world were largely shaped by the ancient Greco-Roman culture. The Roman Empire was Christianized by Emperor Constantine in the early fourth century and declined by the end of the fifth. Beginning with the seventh century, Christianization of Europe begin. In 1054 CE the Great Schism between the Roman Catholic Church and the Eastern Orthodox Church led to the prominent cultural differences between Western and Eastern Europe.

In the fourteenth century, the Renaissance began in Italy with advances in religion, art, and science.<sup>[129]:317–319</sup> At that time the Christian Church as a political entity lost much of its power. European civilization began to change beginning in 1500, leading to the scientific and industrial revolutions. That continent began to exert political and cultural dominance over human societies around the planet, a time known as the Colonial era (also see Age of Discovery).<sup>[129]:295–299</sup> In the eighteenth century a cultural movement known as the Age of Enlightenment further shaped the mentality of Europe and contributed to its secularization. From 1914 to 1918 and 1939 to 1945, nations around the world were embroiled in world wars. Established following World War I, the League of Nations was a first step in establishing international institutions to settle disputes peacefully. After failing to prevent World War II,



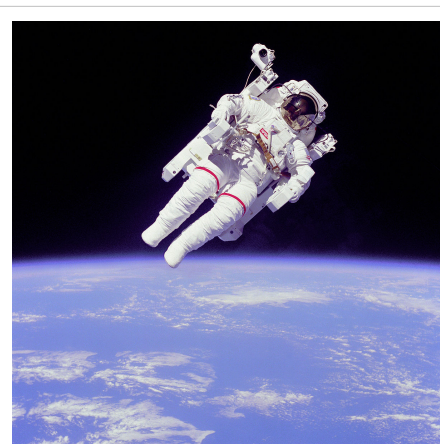
Vitruvian Man by Leonardo da Vinci epitomizes the advances in art and science seen during the Renaissance.

it was replaced by the United Nations. In 1992, several European nations joined in the European Union. As transportation and communication improved, the economies and political affairs of nations around the world have become increasingly intertwined. This globalization has often produced both conflict and cooperation.

### Recent events

Change has continued at a rapid pace from the mid-1940s to today. Technological developments include nuclear weapons, computers, genetic engineering, and nanotechnology. Economic globalization spurred by advances in communication and transportation technology has influenced everyday life in many parts of the world. Cultural and institutional forms such as democracy, capitalism, and environmentalism have increased influence. Major concerns and problems such as disease, war, poverty, violent radicalism, and recently, human-caused climate change have risen as the world population increases.

In 1957, the Soviet Union launched the first artificial satellite into orbit and, soon afterward, Yuri Gagarin became the first human in space. Neil Armstrong, an American, was the first to set foot on another astronomical object, the Moon. Unmanned probes have been sent to all the known planets in the solar system, with some (such as Voyager) having left the solar system. The Soviet Union and the United States were the earliest leaders in space exploration in the 20th century. Five space agencies, representing over fifteen countries,<sup>[137]</sup> have worked together to build the International Space Station. Aboard it, there has been a continuous human presence in space since 2000.<sup>[138]</sup> The World Wide Web was developed in the 1990s and since then has proved to be an indispensable source of information in the developed world.



Astronaut Bruce McCandless II outside of the space shuttle *Challenger* in 1984

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## External links

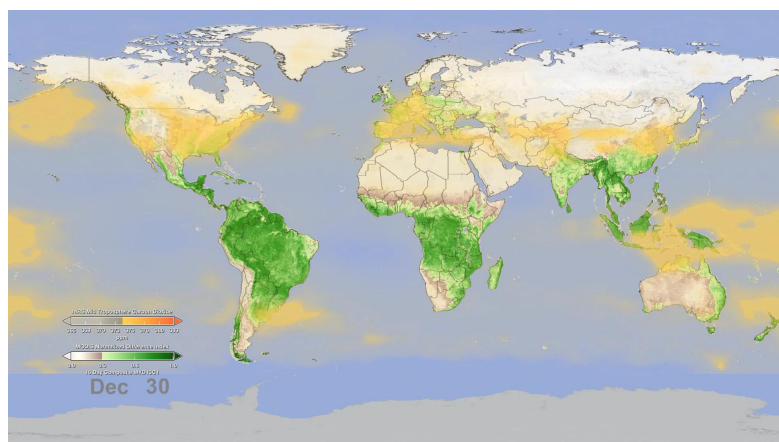
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# Atmosphere of Earth

The **atmosphere of Earth** is a layer of gases surrounding the planet Earth that is retained by Earth's gravity. The atmosphere protects life on Earth by absorbing ultraviolet solar radiation, warming the surface through heat retention (greenhouse effect), and reducing temperature extremes between day and night (the diurnal temperature variation).

Atmospheric stratification describes the structure of the atmosphere, dividing it into distinct layers, each with specific characteristics such as temperature or composition. The atmosphere has a mass of about  $5 \times 10^{18}$

kg, three quarters of which is within about 11 km (unknown operator: u'strong' mi; unknown operator: u'strong' ft) of the surface. The atmosphere becomes thinner and thinner with increasing altitude, with



This animation shows the buildup of tropospheric CO<sub>2</sub> in the Northern Hemisphere with a maximum around May. The maximum in the vegetation cycle follows, occurring in the late summer. Following the peak in vegetation, the drawdown of atmospheric CO<sub>2</sub> due to photosynthesis is apparent, particularly over the Boreal Forests.



no definite boundary between the atmosphere and outer space. An altitude of 120 km (**unknown operator: u'strong'** mi) is where atmospheric effects become noticeable during atmospheric reentry of spacecraft. The Kármán line, at 100 km (**unknown operator: u'strong'** mi), also is often regarded as the boundary between atmosphere and outer space.

**Air** is the name given to atmosphere used in breathing and photosynthesis. Dry air contains roughly (by volume) 78.09% nitrogen, 20.95% oxygen, 0.93% argon, 0.039% carbon dioxide, and small amounts of other gases. Air also contains a variable amount of water vapor, on average around 1%. While air content and atmospheric pressure vary at different layers, air suitable for the survival of terrestrial plants and terrestrial animals is currently only known to be found in Earth's troposphere and artificial atmospheres.



Blue light is scattered more than other wavelengths by the gases in the atmosphere, giving the Earth a blue halo when seen from space.



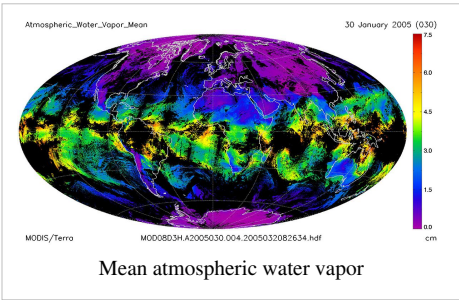
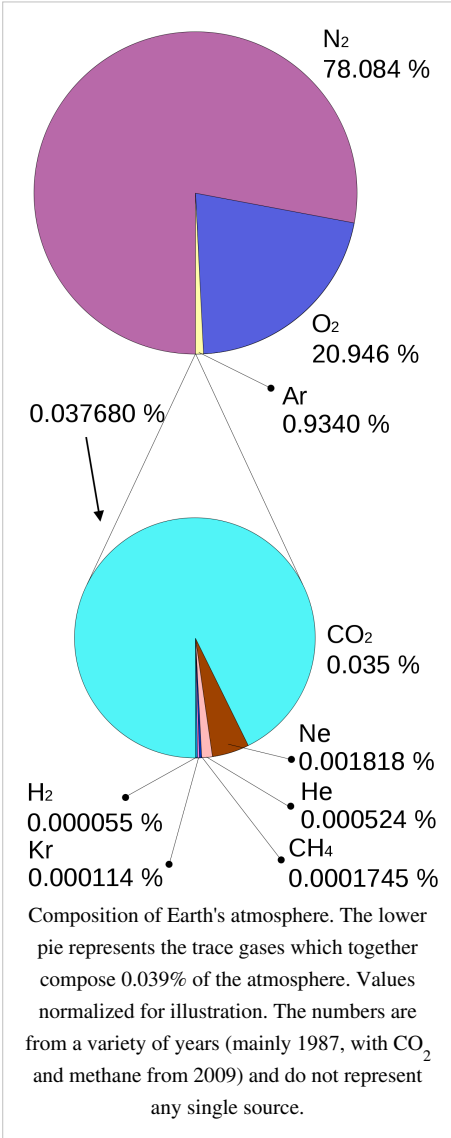
Limb view, of the Earth's atmosphere. Colours roughly denote the layers of the atmosphere.



This image shows the moon at centre, with the limb of Earth near the bottom transitioning into the orange-coloured troposphere, the lowest and most dense portion of the Earth's atmosphere. The troposphere ends abruptly at the tropopause, which appears in the image as the sharp boundary between the orange- and blue- coloured atmosphere. The silvery-blue noctilucent clouds extend far above the Earth's troposphere.

Composition

Air is mainly composed of nitrogen, oxygen, and argon, which together constitute the major gases of the atmosphere. The remaining gases are often referred to as trace gases,<sup>[1]</sup> among which are the greenhouse gases such as water vapor, carbon dioxide, methane, nitrous oxide, and ozone. Filtered air includes trace amounts of many other chemical compounds. Many natural substances may be present in tiny amounts in an unfiltered air sample, including dust, pollen and spores, sea spray, and volcanic ash. Various industrial pollutants also may be present, such as chlorine (elementary or in compounds), fluorine compounds, elemental mercury, and sulfur compounds such as sulfur dioxide [SO<sub>2</sub>].



### Composition of dry atmosphere, by volume<sup>[2]</sup>

<i>ppmv: parts per million by volume (note: volume fraction is equal to mole fraction for ideal gas only, see volume thermodynamics))</i>	
Gas	Volume
Nitrogen (N <sub>2</sub> )	780,840 ppmv (78.084%)
Oxygen (O <sub>2</sub> )	209,460 ppmv (20.946%)
Argon (Ar)	9,340 ppmv (0.9340%)
Carbon dioxide (CO <sub>2</sub> )	394.45 ppmv (0.039445%)
Neon (Ne)	18.18 ppmv (0.001818%)
Helium (He)	5.24 ppmv (0.000524%)
Methane (CH <sub>4</sub> )	1.79 ppmv (0.000179%)
Krypton (Kr)	1.14 ppmv (0.000114%)
Hydrogen (H <sub>2</sub> )	0.55 ppmv (0.000055%)
Nitrous oxide (N <sub>2</sub> O)	0.325 ppmv (0.0000325%)
Carbon monoxide (CO)	0.1 ppmv (0.00001%)
Xenon (Xe)	0.09 ppmv ( $9 \times 10^{-6}\%$ ) (0.000009%)
Ozone (O <sub>3</sub> )	0.0 to 0.07 ppmv (0 to $7 \times 10^{-6}\%$ )
Nitrogen dioxide (NO <sub>2</sub> )	0.02 ppmv ( $2 \times 10^{-6}\%$ ) (0.000002%)
Iodine (I <sub>2</sub> )	0.01 ppmv ( $1 \times 10^{-6}\%$ ) (0.000001%)
Ammonia (NH <sub>3</sub> )	trace
<b>Not included in above dry atmosphere:</b>	
Water vapor (H <sub>2</sub> O)	~0.40% over full atmosphere, typically 1%-4% at surface

## Structure of the atmosphere

### Principal layers

In general, air pressure and density decrease in the atmosphere as height increases. However, temperature has a more complicated profile with altitude. Because the general pattern of this profile is constant and recognizable through means such as balloon soundings, temperature provides a useful metric to distinguish between atmospheric layers. In this way, Earth's atmosphere can be divided into five main layers. From highest to lowest, these layers are:

#### Exosphere

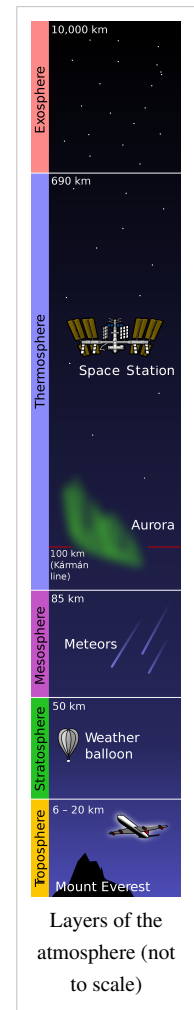
The outermost layer of Earth's atmosphere extends from the exobase upward. It is mainly composed of hydrogen and helium. The particles are so far apart that they can travel hundreds of kilometers without colliding with one another. Since the particles rarely collide, the atmosphere no longer behaves like a fluid. These free-moving particles follow ballistic trajectories and may migrate into and out of the magnetosphere or the solar wind.

#### Thermosphere

Temperature increases with height in the thermosphere from the mesopause up to the thermopause, then is constant with height. Unlike in the stratosphere, where the inversion is caused by absorption of radiation by ozone, in the thermosphere the inversion is a result of the extremely low density of molecules. The temperature of this layer can rise to 1500 °C (**unknown operator: u'strong' °F**), though the gas molecules are so far apart that temperature in the usual sense is not well defined. The air is so rarefied that an individual molecule (of oxygen, for example) travels an average of 1 kilometer between collisions with other molecules.<sup>[3]</sup> The International Space Station orbits in this layer, between 320 and 380 km (**unknown operator: u'strong' and unknown operator: u'strong' mi**). Because of the relative infrequency of molecular collisions, air above the mesopause is poorly mixed compared to air below. While the composition from the troposphere to the mesosphere is fairly constant, above a certain point, air is poorly mixed and becomes compositionally stratified. The point dividing these two regions is known as the turbopause. The region below is the homosphere, and the region above is the heterosphere. The top of the thermosphere is the bottom of the exosphere, called the exobase. Its height varies with solar activity and ranges from about 350–800 km (**unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'; unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'**).

#### Mesosphere

The mesosphere extends from the stratopause to 80–85 km (**unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'; unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'**). It is the layer where most meteors burn up upon entering the atmosphere. Temperature decreases with height in the mesosphere. The mesopause, the temperature minimum that marks the top of the mesosphere, is the coldest place on Earth and has an average temperature around −85 °C (**−unknown operator: u'strong' °F; unknown operator: u'strong' K**).<sup>[4]</sup> At the mesopause, temperatures may drop to −100 °C (**−unknown operator: u'strong' °F; unknown operator: u'strong' K**).<sup>[5]</sup> Due to the cold temperature of the mesosphere, water vapor is frozen, forming ice clouds (or Noctilucent clouds). A type of lightning referred to as either sprites or ELVES, form many miles above thunderclouds in the troposphere.



The stratosphere extends from the tropopause to about 51 km (unknown operator: u'strong' mi; unknown operator: u'strong' ft). Temperature increases with height due to increased absorption of ultraviolet radiation by the ozone layer, which restricts turbulence and mixing. While the temperature may be  $-60^{\circ}\text{C}$  (–unknown operator: u'strong' °F; unknown operator: u'strong' K) at the tropopause, the top of the stratosphere is much warmer, and may be near freezing. The stratopause, which is the boundary between the stratosphere and mesosphere, typically is at 50 to 55 km (unknown operator: u'strong' to unknown operator: u'strong' mi; unknown operator: u'strong' to unknown operator: u'strong' ft). The pressure here is 1/1000 sea level.

The troposphere begins at the surface and extends to between 9 km (**unknown operator: u'strong'** ft) at the poles and 17 km (**unknown operator: u'strong'** ft) at the equator,<sup>[6]</sup> with some variation due to weather. The troposphere is mostly heated by transfer of energy from the surface, so on average the lowest part of the troposphere is warmest and temperature decreases with altitude. This promotes vertical mixing (hence the origin of its name in the Greek word "τροπή", *trope*, meaning turn or overturn). The troposphere contains roughly 80% of the mass of the atmosphere.<sup>[7]</sup> The tropopause is the boundary between the troposphere and stratosphere.

Within the five principal layers determined by temperature are several layers determined by other properties:

- The ozone layer is contained within the stratosphere. In this layer ozone concentrations are about 2 to 8 parts per million, which is much higher than in the lower atmosphere but still very small compared to the main components of the atmosphere. It is mainly located in the lower portion of the stratosphere from about 15–35 km (**unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong' unknown operator: u'strong'; unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'**), though the thickness varies seasonally and geographically. About 90% of the ozone in our atmosphere is contained in the stratosphere.
- The ionosphere, the part of the atmosphere that is ionized by solar radiation, stretches from 50 to 1000 km (**unknown operator: u'strong'** to **unknown operator: u'strong'** mi; **unknown operator: u'strong'** to **unknown operator: u'strong'** ft) and typically overlaps both the exosphere and the thermosphere. It forms the inner edge of the magnetosphere. It has practical importance because it influences, for example, radio propagation on the Earth. It is responsible for auroras.
- The homosphere and heterosphere are defined by whether the atmospheric gases are well mixed. In the homosphere the chemical composition of the atmosphere does not depend on molecular weight because the gases are mixed by turbulence.<sup>[8]</sup> The homosphere includes the troposphere, stratosphere, and mesosphere. Above the *turbopause* at about 100 km (**unknown operator: u'strong'** mi; **unknown operator: u'strong'** ft) (essentially corresponding to the mesopause), the composition varies with altitude. This is because the distance that particles can move without colliding with one another is large compared with the size of motions that cause mixing. This allows the gases to stratify by molecular weight, with the heavier ones such as oxygen and nitrogen present only near the bottom of the heterosphere. The upper part of the heterosphere is composed almost completely of hydrogen, the lightest element.
- The planetary boundary layer is the part of the troposphere that is nearest the Earth's surface and is directly affected by it, mainly through turbulent diffusion. During the day the planetary boundary layer usually is well-mixed, while at night it becomes stably stratified with weak or intermittent mixing. The depth of the planetary boundary layer ranges from as little as about 100 m on clear, calm nights to 3000 m or more during the afternoon in dry regions.

The average temperature of the atmosphere at the surface of Earth is 14 °C (**unknown operator: u'strong'** °F; **unknown operator: u'strong'** K)<sup>[9]</sup> or 15 °C (**unknown operator: u'strong'** °F; **unknown operator: u'strong'** K),<sup>[10]</sup> depending on the reference.<sup>[11][12][13]</sup>

## Physical properties

### Pressure and thickness

The average atmospheric pressure at sea level is about 1 atmosphere (atm)=101.3 kPa (kilopascals)=14.7 psi (pounds per square inch)=760 torr=29.92 inches of mercury (symbol Hg). Total atmospheric mass is  $5.1480 \times 10^{18}$  kg ( $1.135 \times 10^{19}$  lb),<sup>[15]</sup> about 2.5% less than would be inferred from the average sea level pressure and the Earth's area of 51007.2 megahectares, this portion being displaced by the Earth's mountainous terrain. Atmospheric pressure is the total weight of the air above unit area at the point where the pressure is measured. Thus air pressure varies with location and weather.

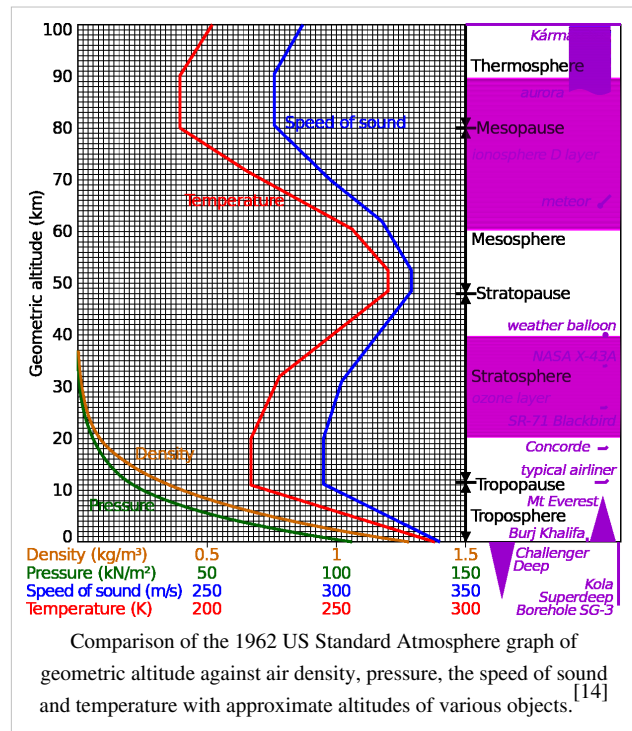
If the atmosphere had a uniform density, it would terminate abruptly at an altitude of 8.50 km (**unknown operator: u'strong'** ft). It actually decreases exponentially with altitude, dropping by half every 5.6 km (**unknown operator: u'strong'** ft) or by a factor of 1/e every 7.64 km (**unknown operator: u'sstrong'** ft), the average scale height of the atmosphere below 70 km (**unknown operator: u'sstrong'** mi; **unknown operator: u'sstrong'** ft). However, the atmosphere is more accurately modeled with a customized equation for each layer that takes gradients of temperature, molecular composition, solar radiation and gravity into account.

In summary, the mass of the earth's atmosphere is distributed approximately as follows:<sup>[16]</sup>

- 50% is below 5.6 km (**unknown operator: u'sstrong'** ft).
- 90% is below 16 km (**unknown operator: u'sstrong'** ft).
- 99.99997% is below 100 km (**unknown operator: u'sstrong'** mi; **unknown operator: u'sstrong'** ft), the Kármán line. By international convention, this marks the beginning of space where human travelers are considered astronauts.

By comparison, the summit of Mt. Everest is at 8848 m (**unknown operator: u'sstrong'** ft); commercial airliners typically cruise between 10 km (**unknown operator: u'sstrong'** ft) and 13 km (**unknown operator: u'sstrong'** ft) where the thinner air improves fuel economy; weather balloons reach 30.4 km (**unknown operator: u'sstrong'** ft) and above; and the highest X-15 flight in 1963 reached 108.0 km (**unknown operator: u'sstrong'** ft).

Even above the Kármán line, significant atmospheric effects such as auroras still occur. Meteors begin to glow in this region though the larger ones may not burn up until they penetrate more deeply. The various layers of the earth's ionosphere, important to HF radio propagation, begin below 100 km and extend beyond 500 km. By comparison, the International Space Station and Space Shuttle typically orbit at 350–400 km, within the F-layer of the ionosphere where they encounter enough atmospheric drag to require reboosts every few months. Depending on solar activity, satellites can still experience noticeable atmospheric drag at altitudes as high as 700–800 km.

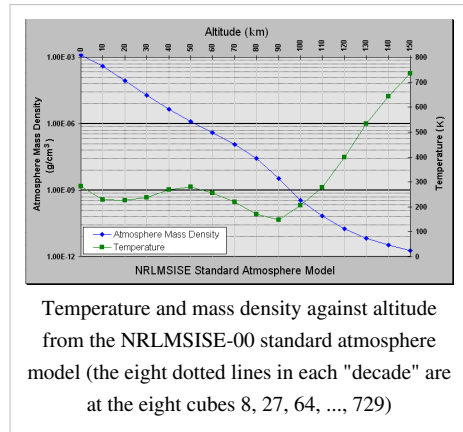




## Density and mass

The density of air at sea level is about  $1.2 \text{ kg/m}^3$  ( $1.2 \text{ g/L}$ ). Density is not measured directly but is calculated from measurements of temperature, pressure and humidity using the equation of state for air (a form of the ideal gas law). Atmospheric density decreases as the altitude increases. This variation can be approximately modeled using the barometric formula. More sophisticated models are used to predict orbital decay of satellites.

The average mass of the atmosphere is about 5 quadrillion ( $5 \times 10^{15}$ ) tonnes or  $1/1,200,000$  the mass of Earth. According to the American National Center for Atmospheric Research, "The total mean mass of the atmosphere is  $5.1480 \times 10^{18} \text{ kg}$  with an annual range due to water vapor of  $1.2$  or  $1.5 \times 10^{15} \text{ kg}$  depending on whether surface pressure or water vapor data are used; somewhat smaller than the previous estimate. The mean mass of water vapor is estimated as  $1.27 \times 10^{16} \text{ kg}$  and the dry air mass as  $5.1352 \pm 0.0003 \times 10^{18} \text{ kg}$ ."



## Optical properties

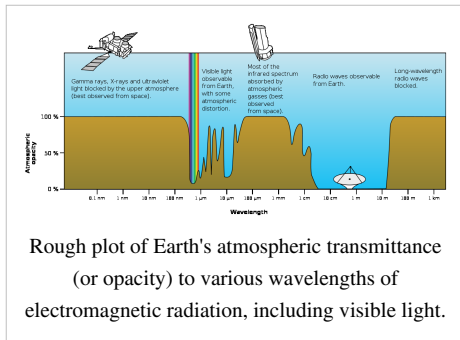
Solar radiation (or sunlight) is the energy the Earth receives from the Sun. The Earth also emits radiation back into space, but at longer wavelengths that we cannot see. Part of the incoming and emitted radiation is absorbed or reflected by the atmosphere.

## Scattering

When light passes through our atmosphere, photons interact with it through *scattering*. If the light does not interact with the atmosphere, it is called *direct radiation* and is what you see if you were to look directly at the Sun. *Indirect radiation* is light that has been scattered in the atmosphere. For example, on an overcast day when you cannot see your shadow there is no direct radiation reaching you, it has all been scattered. As another example, due to a phenomenon called Rayleigh scattering, shorter (blue) wavelengths scatter more easily than longer (red) wavelengths. This is why the sky looks blue; you are seeing scattered blue light. This is also why sunsets are red. Because the Sun is close to the horizon, the Sun's rays pass through more atmosphere than normal to reach your eye. Much of the blue light has been scattered out, leaving the red light in a sunset.

## Absorption

Different molecules absorb different wavelengths of radiation. For example,  $\text{O}_2$  and  $\text{O}_3$  absorb almost all wavelengths shorter than 300 nanometers. Water ( $\text{H}_2\text{O}$ ) absorbs many wavelengths above 700 nm. When a molecule absorbs a photon, it increases the energy of the molecule. We can think of this as heating the atmosphere, but the atmosphere also cools by emitting radiation, as discussed below.



The combined absorption spectra of the gases in the atmosphere leave "windows" of low opacity, allowing the transmission of only certain bands of light. The optical window runs from around 300 nm (ultraviolet-C) up into the range humans can see, the visible spectrum (commonly called light), at roughly 400–700 nm and continues to the infrared to around 1100 nm. There are also infrared and radio windows that transmit some infrared and radio waves at longer wavelengths. For example, the radio window runs from about one centimeter to about eleven-meter waves.

## Emission

*Emission* is the opposite of absorption, it is when an object emits radiation. Objects tend to emit amounts and wavelengths of radiation depending on their "black body" emission curves, therefore hotter objects tend to emit more radiation, with shorter wavelengths. Colder objects emit less radiation, with longer wavelengths. For example, the Sun is approximately 6000 K (**unknown operator: u'strong' °C; unknown operator: u'strong' °F**), its radiation peaks near 500 nm, and is visible to the human eye. The Earth is approximately 290 K (**unknown operator: u'strong' °C; unknown operator: u'strong' °F**), so its radiation peaks near 10,000 nm, and is much too long to be visible to humans.

Because of its temperature, the atmosphere emits infrared radiation. For example, on clear nights the Earth's surface cools down faster than on cloudy nights. This is because clouds ( $\text{H}_2\text{O}$ ) are strong absorbers and emitters of infrared radiation. This is also why it becomes colder at night at higher elevations. The atmosphere acts as a "blanket" to limit the amount of radiation the Earth loses into space.

The *greenhouse effect* is directly related to this absorption and emission (or "blanket") effect. Some chemicals in the atmosphere absorb and emit infrared radiation, but do not interact with sunlight in the visible spectrum. Common examples of these chemicals are  $\text{CO}_2$  and  $\text{H}_2\text{O}$ . If there are too much of these *greenhouse gases*, sunlight heats the Earth's surface, but the gases block the infrared radiation from exiting back to space. This imbalance causes the Earth to warm, and thus climate change.

## Refractive index

The refractive index of air is close to, but just greater than 1. Systematic variations in refractive index can lead to the bending of light rays over long optical paths. One example is that, under some circumstances, observers onboard ships can see other vessels just over the horizon because light is refracted in the same direction as the curvature of the Earth's surface.

The refractive index of air depends on temperature, giving rise to refraction effects when the temperature gradient is large. An example of such effects is the mirage.

## Circulation

*Atmospheric circulation* is the large-scale movement of air through the troposphere, and the means (with ocean circulation) by which heat is distributed around the Earth. The large-scale structure of the atmospheric circulation varies from year to year, but the basic structure remains fairly constant as it is determined by the Earth's rotation rate and the difference in solar radiation between the equator and poles.

## Evolution of Earth's atmosphere

### Earliest atmosphere

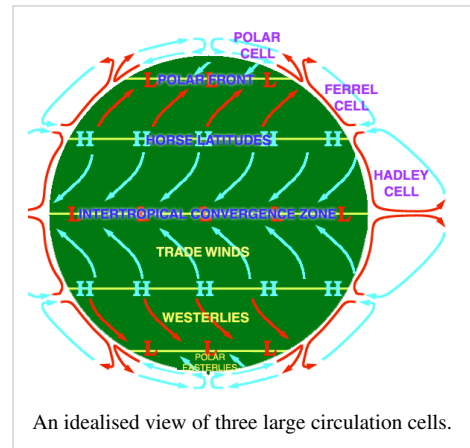
The first atmosphere would have consisted of gases in the solar nebula, primarily hydrogen. In addition there would probably have been simple hydrides such as are now found in gas-giant planets like Jupiter and Saturn, notably water vapor, methane and ammonia. As the solar nebula dissipated these gases would have escaped, partly driven off by the solar wind.<sup>[17]</sup>

### Second atmosphere

The next atmosphere, consisting largely of nitrogen plus carbon dioxide and inert gases, was produced by outgassing from volcanism, supplemented by gases produced during the late heavy bombardment of Earth by huge asteroids.<sup>[17]</sup> A major rainfall led to the buildup of a vast ocean. A major part of carbon dioxide exhalations were soon dissolved in water and built up carbonate sediments.

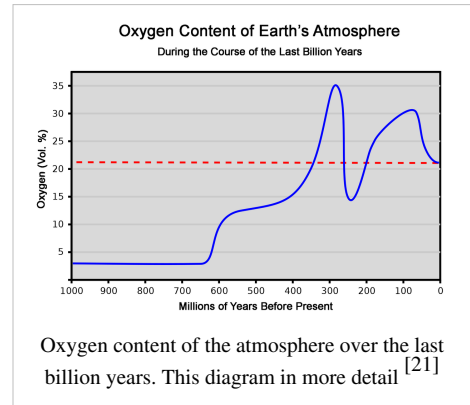
Water-related sediments have been found dating from as early as 3.8 billion years ago.<sup>[18]</sup> About 3.4 billion years ago, nitrogen was the major part of the then stable "second atmosphere". An influence of life has to be taken into account rather soon in the history of the atmosphere, since hints of early life forms are to be found as early as 3.5 billion years ago.<sup>[19]</sup> The fact that this is not perfectly in line with the 30% lower solar radiance (compared to today) of the early Sun has been described as the "faint young Sun paradox".

The geological record however shows a continually relatively warm surface during the complete early temperature record of the Earth with the exception of one cold glacial phase about 2.4 billion years ago. In the late Archaean eon an oxygen-containing atmosphere began to develop, apparently from photosynthesizing algae which have been found as stromatolite fossils from 2.7 billion years ago. The early basic carbon isotopy (isotope ratio proportions) is very much in line with what is found today,<sup>[20]</sup> suggesting that the fundamental features of the carbon cycle were established as early as 4 billion years ago.



### Third atmosphere

The accretion of continents about 3.5 billion years ago<sup>[22]</sup> added plate tectonics, constantly rearranging the continents and also shaping long-term climate evolution by allowing the transfer of carbon dioxide to large land-based carbonate storages. Free oxygen did not exist until about 1.7 billion years ago and this can be seen with the development of the red beds and the end of the banded iron formations. The Earth had a lot of iron in the beginning, and higher amounts of oxygen was not available in the atmosphere until all the iron had been oxidized. This signifies a shift from a reducing atmosphere to an oxidising atmosphere. O<sub>2</sub> showed major ups and downs until reaching a steady state of more than 15%.<sup>[23]</sup> The following time span was the Phanerozoic eon, during which oxygen-breathing metazoan life forms began to appear.



The amount of oxygen in the atmosphere has gone up and down during the last 600 million years. There was a peak 280 million years ago, when the amount of oxygen was about 30 %, much higher than today. Two main processes govern changes in the atmosphere: Plants convert carbon dioxide into the bodies of the plants, which emit oxygen into the atmosphere, and break down of pyrite rocks cause sulphur to be added to the oceans. Volcanos cause this sulphur to be oxidized, reducing the amount of oxygen in the atmosphere. But volcanos also emit carbon dioxide, so that plants can convert this to oxygen. The exact cause of the variation of oxygen in the atmosphere is not known. Periods with much oxygen in the atmosphere are believed to cause rapid development of animals. Even if the atmosphere today has only 21 percent oxygen, today is still regarded as a period with rapid development of animals because of a high amount of oxygen in the atmosphere.<sup>[24]</sup>

Currently, anthropogenic greenhouse gases are increasing in the atmosphere. According to the Intergovernmental Panel on Climate Change, this increase is the main cause of global warming.<sup>[25]</sup>

### Air pollution

*Air pollution* is the introduction of chemicals, particulate matter, or biological materials that cause harm or discomfort to organisms into the atmosphere.<sup>[26]</sup> Stratospheric ozone depletion is believed to be caused by air pollution (chiefly from chlorofluorocarbons).

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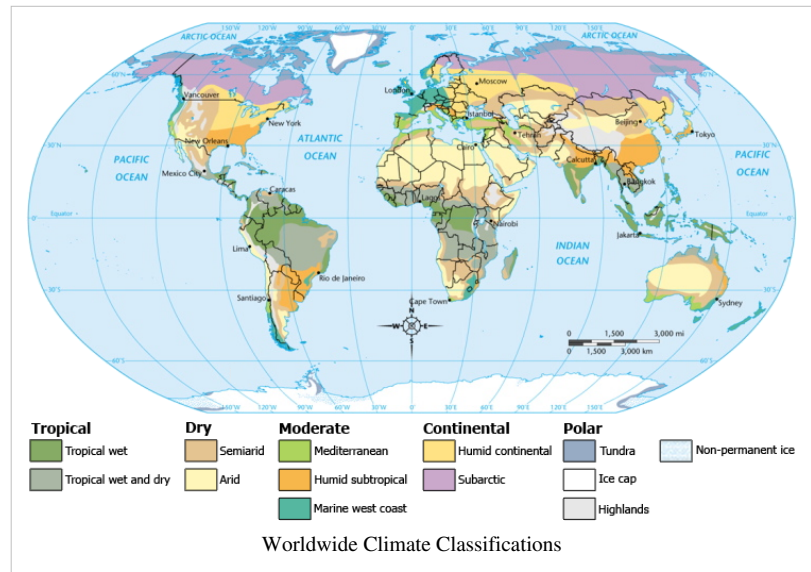
## External links

- NASA atmosphere models ([http://modelweb.gsfc.nasa.gov/spdf\\_models\\_home.html#atmo](http://modelweb.gsfc.nasa.gov/spdf_models_home.html#atmo))
- NASA's Earth Fact Sheet (<http://nssdc.gsfc.nasa.gov/planetary/factsheet/earthfact.html>)
- American Geophysical Union: Atmospheric Sciences (<http://atmospheres.agu.org/>)
- Outreach of the GEOMON project (<http://www.geomon.eu/outreach/>) See how Earth atmosphere is observed and monitored by a European project that combines many approaches.
- Stuff in the Air (<http://www.stuffintheair.com/>) Find out what the atmosphere contains.
- Layers of the Atmosphere (<http://www.srh.noaa.gov/srh/jetstream/atmos/layers.htm>)
- Answers to several questions of curious kids related to Air and Atmosphere (<http://www.scribd.com/doc/22854/Air-Atmosphere-and-Airplanes/>)
- The AMS Glossary of Meteorology (<http://amsglossary.allenpress.com/glossary>)
- Paul Crutzen Interview (<http://www.vega.org.uk/video/programme/111>) Free video of Paul Crutzen Nobel Laureate for his work on decomposition of ozone talking to Harry Kroto Nobel Laureate by the Vega Science Trust.

# Climate

**Climate** encompasses the statistics of temperature, humidity, atmospheric pressure, wind, precipitation, atmospheric particle count and other meteorological elemental measurements in a given region over long periods. Climate can be contrasted to weather, which is the present condition of these elements and their variations over shorter periods.

A region's climate is generated by the **climate system**, which has five components: atmosphere, hydrosphere, cryosphere, land surface, and biosphere.<sup>[1]</sup>



The climate of a location is affected by its latitude, terrain, and altitude, as well as nearby water bodies and their currents. Climates can be classified according to the average and the typical ranges of different variables, most commonly temperature and precipitation. The most commonly used classification scheme was originally developed by Wladimir Köppen. The Thornthwaite system,<sup>[2]</sup> in use since 1948, incorporates evapotranspiration along with temperature and precipitation information and is used in studying animal species diversity and potential effects of climate changes. The Bergeron and Spatial Synoptic Classification systems focus on the origin of air masses that define the climate of a region.

Paleoclimatology is the study of ancient climates. Since direct observations of climate are not available before the 19th century, paleoclimates are inferred from *proxy variables* that include non-biotic evidence such as sediments found in lake beds and ice cores, and biotic evidence such as tree rings and coral. Climate models are mathematical models of past, present and future climates. Climate change may occur over long and short timescales from a variety of factors; recent warming is discussed in global warming.

## Definition

Climate (from Ancient Greek *klima*, meaning *inclination*) is commonly defined as the weather averaged over a long period.<sup>[3]</sup> The standard averaging period is 30 years,<sup>[4]</sup> but other periods may be used depending on the purpose. Climate also includes statistics other than the average, such as the magnitudes of day-to-day or year-to-year variations. The Intergovernmental Panel on Climate Change (IPCC) glossary definition is:

*Climate in a narrow sense is usually defined as the "average weather," or more rigorously, as the statistical description in terms of the mean and variability of relevant quantities over a period ranging from months to thousands or millions of years. The classical period is 30 years, as defined by the World Meteorological Organization (WMO). These quantities are most often surface variables such as temperature, precipitation, and wind. Climate in a wider sense is the state, including a statistical description, of the climate system.* Intergovernmental Panel on Climate Change. Appendix I: Glossary. Retrieved on 2007-06-01.

The difference between climate and weather is usefully summarized by the popular phrase "Climate is what you expect, weather is what you get."<sup>[6]</sup> Over historical time spans there are a number of nearly constant variables that determine climate, including latitude, altitude, proportion of land to water, and proximity to oceans and mountains.



These change only over periods of millions of years due to processes such as plate tectonics. Other climate determinants are more dynamic: the thermohaline circulation of the ocean leads to a 5 °C (9 °F) warming of the northern Atlantic Ocean compared to other ocean basins.<sup>[7]</sup> Other ocean currents redistribute heat between land and water on a more regional scale. The density and type of vegetation coverage affects solar heat absorption,<sup>[8]</sup> water retention, and rainfall on a regional level. Alterations in the quantity of atmospheric greenhouse gases determines the amount of solar energy retained by the planet, leading to global warming or global cooling. The variables which determine climate are numerous and the interactions complex, but there is general agreement that the broad outlines are understood, at least insofar as the determinants of historical climate change are concerned.<sup>[9]</sup>

## Climate classification

There are several ways to classify climates into similar regimes. Originally, climates were defined in Ancient Greece to describe the weather depending upon a location's latitude. Modern climate classification methods can be broadly divided into *genetic* methods, which focus on the causes of climate, and *empiric* methods, which focus on the effects of climate. Examples of genetic classification include methods based on the relative frequency of different air mass types or locations within synoptic weather disturbances. Examples of empiric classifications include climate zones defined by plant hardiness,<sup>[10]</sup> evapotranspiration,<sup>[11]</sup> or more generally the Köppen climate classification which was originally designed to identify the climates associated with certain biomes. A common shortcoming of these classification schemes is that they produce distinct boundaries between the zones they define, rather than the gradual transition of climate properties more common in nature.

## Bergeron and Spatial Synoptic

The simplest classification is that involving air masses. The Bergeron classification is the most widely accepted form of air mass classification. Air mass classification involves three letters. The first letter describes its moisture properties, with c used for continental air masses (dry) and m for maritime air masses (moist). The second letter describes the thermal characteristic of its source region: T for tropical, P for polar, A for Arctic or Antarctic, M for monsoon, E for equatorial, and S for superior air (dry air formed by significant downward motion in the atmosphere). The third letter is used to designate the stability of the atmosphere. If the air mass is colder than the ground below it, it is labeled k. If the air mass is warmer than the ground below it, it is labeled w.<sup>[12]</sup> While air mass identification was originally used in weather forecasting during the 1950s, climatologists began to establish synoptic climatologies based on this idea in 1973.<sup>[13]</sup>

Based upon the Bergeron classification scheme is the Spatial Synoptic Classification system (SSC). There are six categories within the SSC scheme: Dry Polar (similar to continental polar), Dry Moderate (similar to maritime superior), Dry Tropical (similar to continental tropical), Moist Polar (similar to maritime polar), Moist Moderate (a hybrid between maritime polar and maritime tropical), and Moist Tropical (similar to maritime tropical, maritime monsoon, or maritime equatorial).<sup>[14]</sup>

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## Köppen

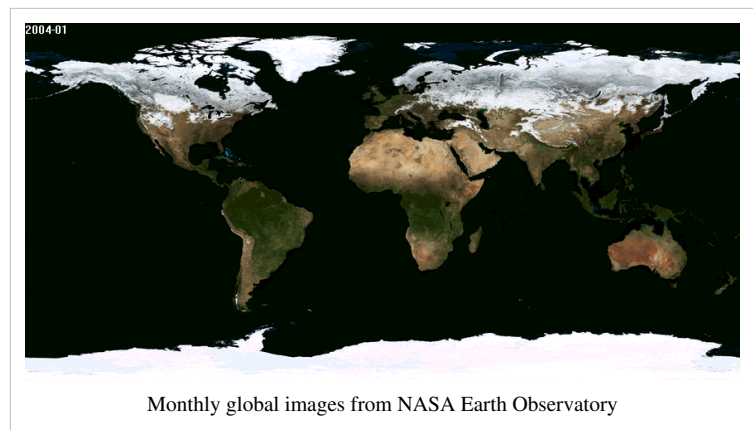
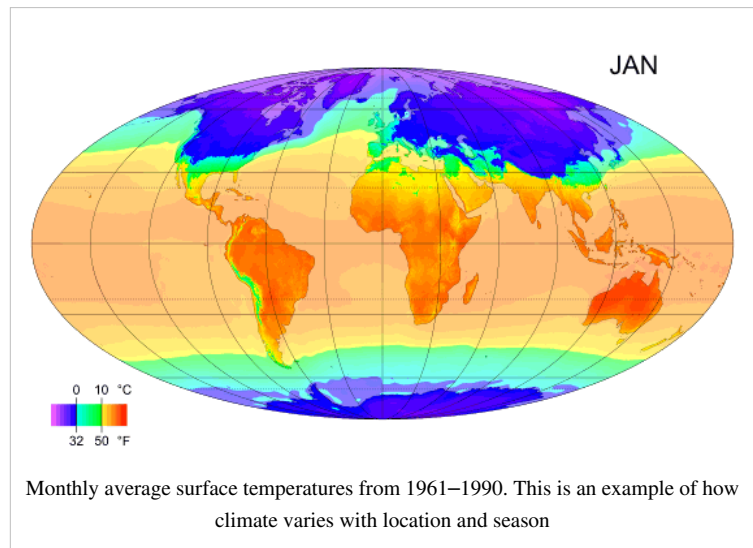
The Köppen classification depends on average monthly values of temperature and precipitation. The most commonly used form of the Köppen classification has five primary types labeled A through E. These primary types are A, tropical; B, dry; C, mild mid-latitude; D, cold mid-latitude; and E, polar. The five primary classifications can be further divided into secondary classifications such as rain forest, monsoon, tropical savanna, humid subtropical, humid continental, oceanic climate, Mediterranean climate, steppe, subarctic climate, tundra, polar ice cap, and desert.

**Rain forests** are characterized by high rainfall, with definitions setting minimum normal annual rainfall between 1750 millimetres (unknown operator: u'strong' in) and 2000 millimetres (unknown operator: u'strong' in). Mean monthly temperatures exceed 18 °C (unknown operator: u'strong' °F) during all months of the year.<sup>[15]</sup>

A **monsoon** is a seasonal prevailing wind which lasts for several months, ushering in a region's rainy season.<sup>[16]</sup> Regions within North America, South America, Sub-Saharan Africa, Australia and East Asia are monsoon regimes.<sup>[17]</sup>

A **tropical savanna** is a grassland biome located in semiarid to semi-humid climate regions of subtropical and tropical latitudes, with average temperatures remain at or above 18 °C (unknown operator: u'strong' °F) year round and rainfall between 750 millimetres (unknown operator: u'strong' in) and 1270 millimetres (unknown operator: u'strong' in) a year. They are widespread on Africa, and are found in India, the northern parts of South America, Malaysia, and Australia.<sup>[18]</sup>

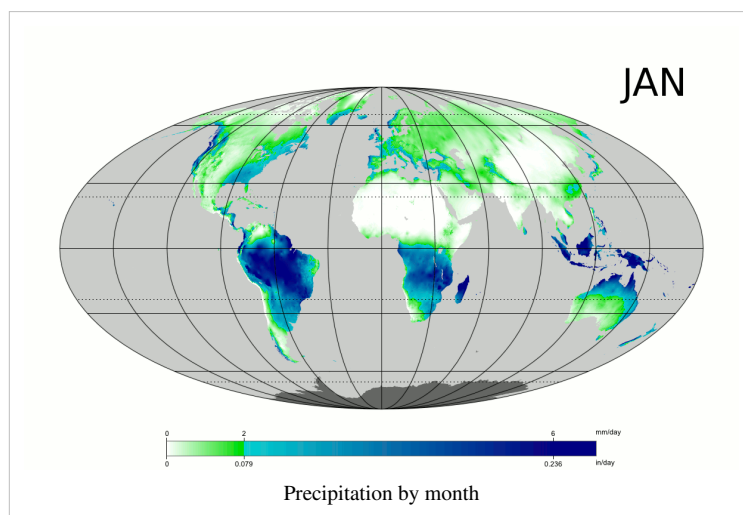
The **humid subtropical** climate zone where winter rainfall (and sometimes snowfall) is associated with large storms that the westerlies steer from west to east. Most summer rainfall occurs during thunderstorms and from occasional tropical cyclones.<sup>[19]</sup> Humid subtropical climates lie on the east side continents, roughly between latitudes 20° and 40° degrees away from the equator.<sup>[20]</sup>





## Thornthwaite

Devised by the American climatologist and geographer C. W. Thornthwaite, this climate classification method monitors the soil water budget using evapotranspiration.<sup>[11]</sup> It monitors the portion of total precipitation used to nourish vegetation over a certain area.<sup>[30]</sup> It uses indices such as a humidity index and an aridity index to determine an area's moisture regime based upon its average temperature, average rainfall, and average vegetation type.<sup>[31]</sup> The lower the value of the index in any given area, the drier the area is.



The moisture classification includes climatic classes with descriptors such as hyperhumid, humid, subhumid, subarid, semi-arid (values of  $-20$  to  $-40$ ), and arid (values below  $-40$ ).<sup>[32]</sup> Humid regions experience more precipitation than evaporation each year, while arid regions experience greater evaporation than precipitation on an annual basis. A total of 33 percent of the Earth's landmass is considered either arid or semi-arid, including southwest North America, southwest South America, most of northern and a small part of southern Africa, southwest and portions of eastern Asia, as well as much of Australia.<sup>[33]</sup> Studies suggest that precipitation effectiveness (PE) within the Thornthwaite moisture index is overestimated in the summer and underestimated in the winter.<sup>[34]</sup> This index can be effectively used to determine the number of herbivore and mammal species numbers within a given area.<sup>[35]</sup> The index is also used in studies of climate change.<sup>[34]</sup>

Thermal classifications within the Thornthwaite scheme include microthermal, mesothermal, and megathermal regimes. A microthermal climate is one of low annual mean temperatures, generally between  $0^{\circ}\text{C}$  (**unknown operator: u'strong'**  $^{\circ}\text{F}$ ) and  $14^{\circ}\text{C}$  (**unknown operator: u'strong'**  $^{\circ}\text{F}$ ) which experiences short summers and has a potential evaporation between 14 centimetres (**unknown operator: u'strong'** in) and 43 centimetres (**unknown operator: u'strong'** in).<sup>[36]</sup> A mesothermal climate lacks persistent heat or persistent cold, with potential evaporation between 57 centimetres (**unknown operator: u'strong'** in) and 114 centimetres (**unknown operator: u'strong'** in).<sup>[37]</sup> A megathermal climate is one with persistent high temperatures and abundant rainfall, with potential annual evaporation in excess of 114 centimetres (**unknown operator: u'strong'** in).<sup>[38]</sup>

## Record

### Modern

Details of the modern climate record are known through the taking of measurements from such weather instruments as thermometers, barometers, and anemometers during the past few centuries. The instruments used to study weather over the modern time scale, their known error, their immediate environment, and their exposure have changed over the years, which must be considered when studying the climate of centuries past.<sup>[39]</sup>

## Paleoclimatology

Paleoclimatology is the study of past climate over a great period of the Earth's history. It uses evidence from ice sheets, tree rings, sediments, coral, and rocks to determine the past state of the climate. It demonstrates periods of stability and periods of change and can indicate whether changes follow patterns such as regular cycles.<sup>[40]</sup>

## Climate change

Climate change is the variation in global or regional climates over time. It reflects changes in the variability or average state of the atmosphere over time scales ranging from decades to millions of years. These changes can be caused by processes internal to the Earth, external forces (e.g. variations in sunlight intensity) or, more recently, human activities.<sup>[41]</sup>

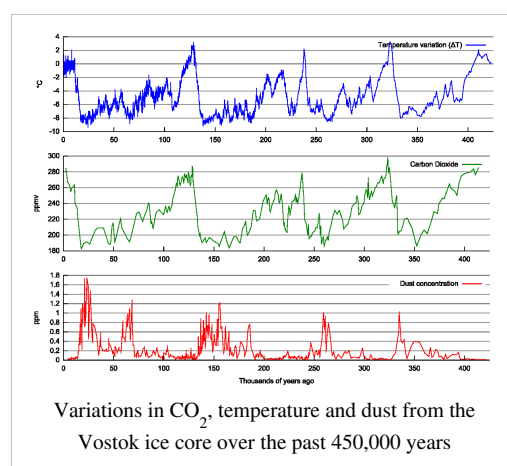
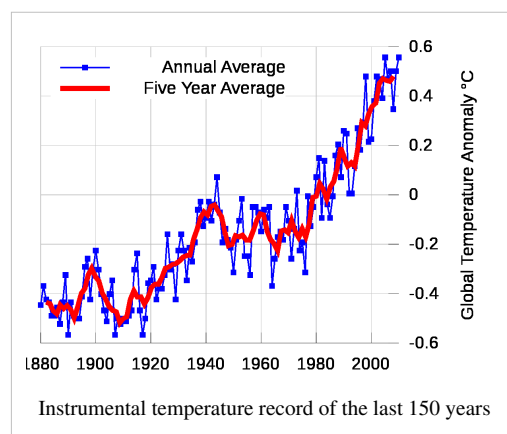
In recent usage, especially in the context of environmental policy, the term "climate change" often refers only to changes in modern climate, including the rise in average surface temperature known as global warming. In some cases, the term is also used with a presumption of human causation, as in the United Nations Framework Convention on Climate Change (UNFCCC). The UNFCCC uses "climate variability" for non-human caused variations.<sup>[42]</sup>

Earth has undergone periodic climate shifts in the past, including four major ice ages. These consisting of glacial periods where conditions are colder than normal, separated by interglacial periods. The accumulation of snow and ice during a glacial period increases the surface albedo, reflecting more of the Sun's energy into space and maintaining a lower atmospheric temperature. Increases in greenhouse gases, such as by volcanic activity, can increase the global temperature and produce an interglacial. Suggested causes of ice age periods include the positions of the continents, variations in the Earth's orbit,<sup>[43]</sup> changes in the solar output, and volcanism.<sup>[44]</sup>

## Climate models

Climate models use quantitative methods to simulate the interactions of the atmosphere,<sup>[45]</sup> oceans, land surface and ice. They are used for a variety of purposes from study of the dynamics of the weather and climate system to projections of future climate. All climate models balance, or very nearly balance, incoming energy as short wave (including visible) electromagnetic radiation to the earth with outgoing energy as long wave (infrared) electromagnetic radiation from the earth. Any imbalance results in a change in the average temperature of the earth.

The most talked-about applications of these models in recent years have been their use to infer the consequences of increasing greenhouse gases in the atmosphere, primarily carbon dioxide (see greenhouse gas). These models predict an upward trend in the global mean surface temperature, with the most rapid increase in temperature being projected



for the higher latitudes of the Northern Hemisphere.

Models can range from relatively simple to quite complex:

- Simple radiant heat transfer model that treats the earth as a single point and averages outgoing energy
- this can be expanded vertically (radiative-convective models), or horizontally
- finally, (coupled) atmosphere–ocean–sea ice global climate models discretise and solve the full equations for mass and energy transfer and radiant exchange.<sup>[46]</sup>

Climate forecasting is a way by some scientists are using to predict climate change. In 1997 the prediction division of the International Research Institute for Climate and Society at Columbia University began generating seasonal climate forecasts on a real-time basis. To produce these forecasts an extensive suite of forecasting tools was developed, including a multimodel ensemble approach that required thorough validation of each model's accuracy level in simulating interannual climate variability.<sup>[47]</sup>

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- IGBP Climate-change Index (<http://www.igbp.net/page.php?pid=504>)
- The Economics of Climate-based Data (<http://www.economics.noaa.gov/?goal=climate&file=home/>) NOAA Economics
- AgClimate (<http://www.agclimate.org>) IFAS
- Climate Models and modeling groups ([http://128.194.106.6/~baum/climate\\_modeling.html](http://128.194.106.6/~baum/climate_modeling.html))
- Climate Prediction Project (<http://climateapps2.oucs.ox.ac.uk/cpdnboinc/>)
- WorldClimate (<http://www.worldclimate.com>)
- ESPERE Climate Encyclopaedia (<http://www.atmosphere.mpg.de/enid/1442>)
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- A current view of the Bering Sea Ecosystem and Climate (<http://www.beringclimate.noaa.gov/>)
- Climate: Data and charts for world and US locations (<http://www.climate-charts.com/index.html>)
- MIL-HDBK-310, Global Climate Data ([http://www.everyspec.com/MIL-HDBK/MIL-HDBK-0300-0499/MIL\\_HDBK\\_310\\_1851/](http://www.everyspec.com/MIL-HDBK/MIL-HDBK-0300-0499/MIL_HDBK_310_1851/)) U.S. Department of Defense Data on natural environmental starting points for engineering analyses to derive environmental design criteria
- ClimateDiagrams.com (<http://www.climatediagrams.com>) Climate diagrams for over 3 000 weather stations and different climate periods from around the world. Users can create their own diagrams with their own data.
- IPCC Data Distribution Centre (<http://www.ipcc-data.org>) Climate data and guidance on its use.

## Further reading

- The Study of Climate on Alien Worlds; Characterizing atmospheres beyond our Solar System is now within our reach (<http://www.americanscientist.org/issues/feature/2012/4/the-study-of-climate-on-alien-worlds>) Kevin Heng July-August 2012 American Scientist

[[zh:氣候]

# Weather

**Weather** is the state of the atmosphere, to the degree that it is hot or cold, wet or dry, calm or stormy, clear or cloudy.<sup>[1]</sup> Most weather phenomena occur in the troposphere,<sup>[2][3]</sup> just below the stratosphere. Weather refers, generally, to day-to-day temperature and precipitation activity, whereas climate is the term for the average atmospheric conditions over longer periods of time.<sup>[4]</sup> When used without qualification, "weather" is understood to be the weather of Earth.

Weather is driven by density (temperature and moisture) differences between one place and another. These differences can occur due to the sun angle at any particular spot, which varies by latitude from the tropics. The strong temperature contrast between polar and tropical air gives rise to the jet stream. Weather systems in the mid-latitudes, such as extratropical cyclones, are caused by instabilities of the jet stream flow. Because the Earth's axis is tilted relative to its orbital plane, sunlight is incident at different angles at different times of the year. On Earth's surface, temperatures usually range  $\pm 40^{\circ}\text{C}$  ( $100^{\circ}\text{F}$  to  $-40^{\circ}\text{F}$ ) annually. Over thousands of years, changes in Earth's orbit affect the amount and distribution of solar energy received by the Earth and influence long-term climate and global climate change.

Surface temperature differences in turn cause pressure differences. Higher altitudes are cooler than lower altitudes due to differences in compressional heating. Weather forecasting is the application of science and technology to predict the state of the atmosphere for a future time and a given location. The atmosphere is a chaotic system, so small changes to one part of the system can grow to have large effects on the system as a whole. Human attempts to control the weather have occurred throughout human history, and there is evidence that human activity such as agriculture and industry has inadvertently modified weather patterns.

Studying how the weather works on other planets has been helpful in understanding how weather works on Earth. A famous landmark in the Solar System, Jupiter's *Great Red Spot*, is an anticyclonic storm known to have existed for at least 300 years. However, weather is not limited to planetary bodies. A star's corona is constantly being lost to space, creating what is essentially a very thin atmosphere throughout the Solar System. The movement of mass ejected from the Sun is known as the solar wind.



Thunderstorm near Garajau, Madeira

## Cause



Stratocumulus perlucidus clouds

On Earth, common weather phenomena include wind, cloud, rain, snow, fog and dust storms. Less common events include natural disasters such as tornadoes, hurricanes, typhoons and ice storms. Almost all familiar weather phenomena occur in the troposphere (the lower part of the atmosphere).<sup>[3]</sup> Weather does occur in the stratosphere and can affect weather lower down in the troposphere, but the exact mechanisms are poorly understood.<sup>[5]</sup>

Weather occurs primarily due to density (temperature and moisture) differences between one place to another.

These differences can occur due to the sun angle at any particular spot, which varies by latitude from the tropics. In other words, the farther from the tropics you lie, the lower the sun angle is, which causes those locations to be cooler due to the indirect sunlight.<sup>[6]</sup> The strong temperature contrast between polar and tropical air gives rise to the jet stream.<sup>[7]</sup> Weather systems in the mid-latitudes, such as extratropical cyclones, are caused by instabilities of the jet stream flow (see baroclinity).<sup>[8]</sup> Weather systems in the tropics, such as monsoons or organized thunderstorm systems, are caused by different processes.

Because the Earth's axis is tilted relative to its orbital plane, sunlight is incident at different angles at different times of the year. In June the Northern Hemisphere is tilted towards the sun, so at any given Northern Hemisphere latitude sunlight falls more directly on that spot than in December (see Effect of sun angle on climate).<sup>[9]</sup> This effect causes seasons. Over thousands to hundreds of thousands of years, changes in Earth's orbital parameters affect the amount and distribution of solar energy received by the Earth and influence long-term climate. (see Milankovitch cycles).<sup>[10]</sup>

The uneven solar heating (the formation of zones of temperature and moisture gradients, or frontogenesis) can also be due to the weather itself in the form of cloudiness and precipitation.<sup>[11]</sup> Higher altitudes are cooler than lower altitudes, which is explained by the lapse rate.<sup>[12][13]</sup> On local scales, temperature differences can occur because different surfaces (such as oceans, forests, ice sheets, or man-made objects) have differing physical characteristics such as reflectivity, roughness, or moisture content.

Surface temperature differences in turn cause pressure differences. A hot surface heats the air above it and the air expands, lowering the air pressure and its density.<sup>[14]</sup> The resulting horizontal pressure gradient accelerates the air from high to low pressure, creating wind, and Earth's rotation then causes curvature of the flow via the Coriolis effect.<sup>[15]</sup> The simple systems thus formed can then display emergent behaviour to produce more complex systems and thus other weather phenomena. Large scale examples include the Hadley cell while a smaller scale example would be coastal breezes.

The atmosphere is a chaotic system, so small changes to one part of the system can grow to have large effects on the system as a whole.<sup>[16]</sup> This makes it difficult to accurately predict weather more than a few days in advance, though weather forecasters are continually working to extend this limit through the scientific study of weather, meteorology. It is theoretically impossible to make useful day-to-day predictions more than about two weeks ahead, imposing an upper limit to potential for improved prediction skill.<sup>[17]</sup>

## Shaping the planet Earth

Weather is one of the fundamental processes that shape the Earth. The process of weathering breaks down the rocks and soils into smaller fragments and then into their constituent substances.<sup>[18]</sup> These are then free to take part in chemical reactions that can affect the surface further (such as acid rain) or are reformed into other rocks and soils. In this way, weather plays a major role in erosion of the surface.<sup>[19]</sup>

## Effect on humans

### Effects on populations

Weather has played a large and sometimes direct part in human history. Aside from climatic changes that have caused the gradual drift of populations (for example the desertification of the Middle East, and the formation of land bridges during glacial periods), extreme weather events have caused smaller scale population movements and intruded directly in historical events. One such event is the saving of Japan from invasion by the Mongol fleet of Kublai Khan by the Kamikaze winds in 1281.<sup>[20]</sup> French claims to Florida came to an end in 1565 when a hurricane destroyed the French fleet, allowing Spain to conquer Fort Caroline.<sup>[21]</sup> More recently, Hurricane Katrina redistributed over one million people from the central Gulf coast elsewhere across the United States, becoming the largest diaspora in the history of the United States.<sup>[22]</sup>



New Orleans, Louisiana, after being struck by Hurricane Katrina. Katrina was a Category 3 hurricane when it struck although it had been a category 5 hurricane in the Gulf of Mexico.

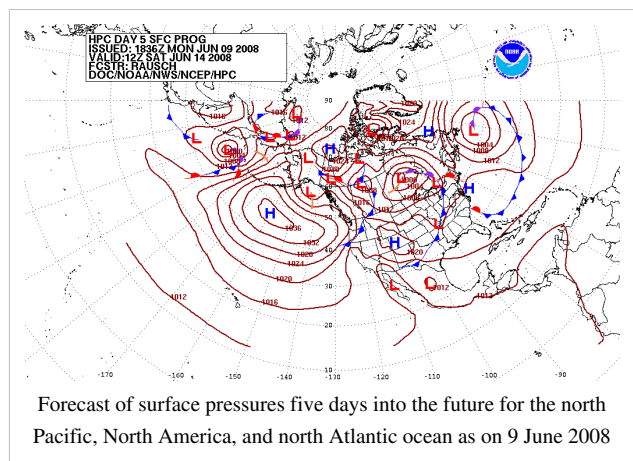
The Little Ice Age caused crop failures and famines in Europe. The 1690s saw the worst famine in France since the Middle Ages. Finland suffered a severe famine in 1696–1697, during which about one-third of the Finnish population died.<sup>[23]</sup>

### Effects on individuals

The human body is negatively affected by extremes in temperature, humidity, and wind.<sup>[24]</sup>

## Forecasting

Weather forecasting is the application of science and technology to predict the state of the atmosphere for a future time and a given location. Human beings have attempted to predict the weather informally for millennia, and formally since at least the nineteenth century.<sup>[25][26]</sup> Weather forecasts are made by collecting quantitative data about the current state of the atmosphere and using scientific understanding of atmospheric processes to project how the atmosphere will evolve.<sup>[27]</sup>



Forecast of surface pressures five days into the future for the north Pacific, North America, and north Atlantic ocean as on 9 June 2008

Once an all-human endeavor based mainly upon changes in barometric pressure, current weather conditions, and sky condition,<sup>[28][29]</sup> forecast models are now used to determine future conditions. Human input is still required to pick the best possible forecast model to base the forecast upon, which involves pattern recognition skills, teleconnections, knowledge of model performance, and knowledge of model biases. The chaotic nature of the atmosphere, the massive computational power required to solve the equations that describe the atmosphere, error involved in measuring the initial conditions, and an incomplete understanding of atmospheric processes mean that forecasts become less accurate as the difference in current time and the time for which the forecast is being made (the *range* of the forecast) increases. The use of ensembles and model consensus helps to narrow the error and pick the most likely outcome.<sup>[30][31][32]</sup>

There are a variety of end users to weather forecasts. Weather warnings are important forecasts because they are used to protect life and property.<sup>[33][34]</sup> Forecasts based on temperature and precipitation are important to agriculture,<sup>[35][36][37][38]</sup> and therefore to commodity traders within stock markets. Temperature forecasts are used by utility companies to estimate demand over coming days.<sup>[39][40][41]</sup> On an everyday basis, people use weather forecasts to determine what to wear on a given day. Since outdoor activities are severely curtailed by heavy rain, snow and the wind chill, forecasts can be used to plan activities around these events, and to plan ahead and survive them.

## Modification

The aspiration to control the weather is evident throughout human history: from ancient rituals intended to bring rain for crops to the U.S. Military Operation Popeye, an attempt to disrupt supply lines by lengthening the North Vietnamese monsoon. The most successful attempts at influencing weather involve cloud seeding; they include the fog- and low stratus dispersion techniques employed by major airports, techniques used to increase winter precipitation over mountains, and techniques to suppress hail.<sup>[42]</sup> A recent example of weather control was China's preparation for the 2008 Summer Olympic Games. China shot 1,104 rain dispersal rockets from 21 sites in the city of Beijing in an effort to keep rain away from the opening ceremony of the games on 8 August 2008. Guo Hu, head of the Beijing Municipal Meteorological Bureau (BMB), confirmed the success of the operation with 100 millimeters falling in Baoding City of Hebei Province, to the southwest and Beijing's Fangshan District recording a rainfall of 25 millimeters.<sup>[43]</sup>

Whereas there is inconclusive evidence for these techniques' efficacy, there is extensive evidence that human activity such as agriculture and industry results in inadvertent weather modification.<sup>[42]</sup>

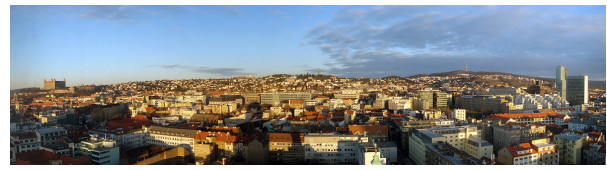
- Acid rain, caused by industrial emission of sulfur dioxide and nitrogen oxides into the atmosphere, adversely affects freshwater lakes, vegetation, and structures.
- Anthropogenic pollutants reduce air quality and visibility.
- Climate change caused by human activities that emit greenhouse gases into the air is expected to affect the frequency of extreme weather events such as drought, extreme temperatures, flooding, high winds, and severe storms.<sup>[44]</sup>

The effects of inadvertent weather modification may pose serious threats to many aspects of civilization, including ecosystems, natural resources, food and fiber production, economic development, and human health.<sup>[45]</sup>

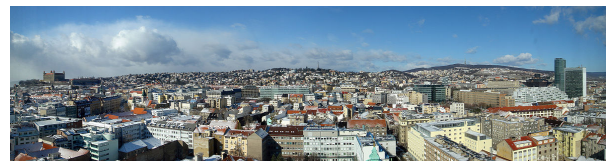


## Extremes on Earth

On Earth, temperatures usually range  $\pm 40^\circ\text{C}$  ( $100^\circ\text{F}$  to  $-40^\circ\text{F}$ ) annually. The range of climates and latitudes across the planet can offer extremes of temperature outside this range. The coldest air temperature ever recorded on Earth is  $-89.2^\circ\text{C}$  (**unknown operator: u'strong'**  $^\circ\text{F}$ ), at Vostok Station, Antarctica on 21 July 1983. The hottest air temperature ever recorded was  $57.7^\circ\text{C}$  (**unknown operator: u'strong'**  $^\circ\text{F}$ ) at 'Aziziya, Libya, on 13 September 1922,<sup>[46]</sup> but that reading is queried. The highest recorded average annual temperature was  $34.4^\circ\text{C}$  (**unknown operator: u'strong'**  $^\circ\text{F}$ ) at Dallol, Ethiopia.<sup>[47]</sup> The coldest recorded average annual temperature was  $-55.1^\circ\text{C}$  (**unknown operator: u'strong'**  $^\circ\text{F}$ ) at Vostok Station, Antarctica.<sup>[48]</sup> The coldest average annual temperature in a permanently inhabited location is at Eureka, Nunavut, in Canada, where the annual average temperature is  $-19.7^\circ\text{C}$  (**unknown operator: u'strong'**  $^\circ\text{F}$ ).<sup>[49]</sup>



Early morning sunshine over Bratislava, Slovakia.



The same area, just three hours later, after light snowfall.

## Extraterrestrial within the Solar System

Studying how the weather works on other planets has been seen as helpful in understanding how it works on Earth.<sup>[50]</sup> Weather on other planets follows many of the same physical principles as weather on Earth, but occurs on different scales and in atmospheres having different chemical composition. The Cassini–Huygens mission to Titan discovered clouds formed from methane or ethane which deposit rain composed of liquid methane and other organic compounds.<sup>[51]</sup> Earth's atmosphere includes six latitudinal circulation zones, three in each hemisphere.<sup>[52]</sup> In contrast, Jupiter's banded appearance shows many such zones,<sup>[53]</sup> Titan has a single jet stream near the 50th parallel north latitude,<sup>[54]</sup> and Venus has a single jet near the equator.<sup>[55]</sup>

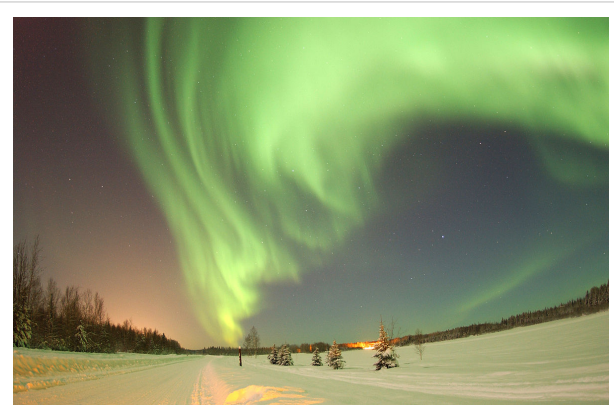


Jupiter's Great Red Spot in 1979.

One of the most famous landmarks in the Solar System, Jupiter's *Great Red Spot*, is an anticyclonic storm known to have existed for at least 300 years.<sup>[56]</sup> On other gas giants, the lack of a surface allows the wind to reach enormous speeds: gusts of up to 600 metres per second (about 2100 km/h or **unknown operator: u'strong'** mph) have been measured on the planet Neptune.<sup>[57]</sup> This has created a puzzle for planetary scientists. The weather is ultimately created by solar energy and the amount of energy received by Neptune is only about  $\frac{1}{900}$  of that received by Earth, yet the intensity of weather phenomena on Neptune is far greater than on Earth.<sup>[58]</sup> The strongest planetary winds discovered so far are on the extrasolar planet HD 189733 b, which is thought to have easterly winds moving at more than 9600 kilometres per hour (**unknown operator: u'strong'** mph).<sup>[59]</sup>

## Space weather

Weather is not limited to planetary bodies. Like all stars, the sun's corona is constantly being lost to space, creating what is essentially a very thin atmosphere throughout the Solar System. The movement of mass ejected from the Sun is known as the solar wind. Inconsistencies in this wind and larger events on the surface of the star, such as coronal mass ejections, form a system that has features analogous to conventional weather systems (such as pressure and wind) and is generally known as space weather. Coronal mass ejections have been tracked as far out in the solar system as Saturn.<sup>[60]</sup> The activity of this system can affect planetary atmospheres and occasionally surfaces. The interaction of the solar wind with the terrestrial atmosphere can produce spectacular aurorae,<sup>[61]</sup> and can play havoc with electrically sensitive systems such as electricity grids and radio signals.<sup>[62]</sup>



Aurora Borealis

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## External links

- Climate and Weather (<http://ucblibraries.colorado.edu/govpubs/us/climate.htm>) from *UCB Libraries GovPubs*
- The Economics of Extreme Weather Events on Society (<http://www.economics.noaa.gov/?goal=weather&file=events/>) NOAA Economics
- RainRadar: Worldwide radar directory (<http://rainradar.net>)
- National Weather Service (<http://www.weather.gov/>)



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# Life

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## Life

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Life (Biota / Vitae / Eobionti)

Plants in the Rwenzori Mountains, Uganda
Scientific classification 
Domains and kingdoms
Life on Earth: <ul style="list-style-type: none"><li>• Non-cellular life (viruses) <sup>[1]</sup></li><li>• Cellular life</li></ul> <div>Bacteria</div> <div>Archaea</div> <div>Eukarya</div> <div>Protista</div> <div>Fungi</div> <div>Plantae</div> <div>Animalia</div>

**Life** (cf. biota) is a characteristic that distinguishes objects that have signaling and self-sustaining processes from those that do not,<sup>[2][3]</sup> either because such functions have ceased (death), or else because they lack such functions and are classified as inanimate.<sup>[4][5]</sup> Biology is the science concerned with the study of life.

Any contiguous living system is called an organism. These animate entities undergo metabolism, maintain homeostasis, possess a capacity to grow, respond to stimuli, reproduce and, through natural selection, adapt to their environment in successive generations. More complex living organisms can communicate through various means.<sup>[2][6]</sup> A diverse array of living organisms can be found in the biosphere of Earth, and the properties common to these organisms—plants, animals, fungi, protists, archaea, and bacteria—are a carbon- and water-based cellular form with complex organization and heritable genetic information.

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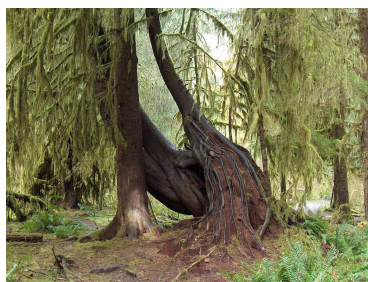


Scientific evidence suggests that life began on Earth some 3.7 billion years ago. The mechanism by which life emerged is still being investigated. Since then, life has evolved into a wide variety of forms, which biologists have classified into a hierarchy of taxa. Life can survive and thrive in a wide range of conditions. The meaning of life—its significance, purpose, and ultimate fate—is a central concept and question in philosophy and religion. Both philosophy and religion have offered interpretations as to how life relates to existence and consciousness, and on related issues such as life stance, purpose, conception of a god or gods, a soul or an afterlife. Different cultures throughout history have had widely varying approaches to these issues.

Though the existence of life is only confirmed on Earth, many scientists believe extraterrestrial life is not only plausible, but probable. Other planets and moons in the Solar System have been examined for evidence of having once supported simple life, and projects such as SETI have attempted to detect transmissions from possible alien civilizations. According to the panspermia hypothesis, life on Earth may have originated from meteorites that spread organic molecules or simple life that first evolved elsewhere.

## Early theories

### Materialism



Plant growth in the Hoh Rainforest



Herds of zebra and impala gathering on the Maasai Mara plain



An aerial photo of microbial mats around the Grand Prismatic Spring of Yellowstone National Park

Some of the earliest theories of life were materialist, holding that all that exists is matter, and that life is merely a complex form or arrangement of matter. Empedocles (430 BC) argued that every thing in the universe is made up of a combination of four eternal "elements" or "roots of all": earth, water, air, and fire. All change is explained by the arrangement and rearrangement of these four elements. The various forms of life are caused by an appropriate mixture of elements.<sup>[7]</sup>



Democritus (460 BC) thought that the essential characteristic of life is having a soul (*psyche*). Like other ancient writers, he was attempting to explain what makes something a *living* thing. His explanation was that fiery atoms make a soul in exactly the same way atoms and void account for any other thing. He elaborates on fire because of the apparent connection between life and heat, and because fire moves.<sup>[8]</sup>

Plato's world of eternal and unchanging Forms, imperfectly represented in matter by a divine Artisan, contrasts sharply with the various mechanistic Weltanschauungen, of which atomism was, by the fourth century at least, the most prominent... This debate persisted throughout the ancient world. Atomistic mechanism got a shot in the arm from Epicurus... while the Stoics adopted a divine teleology... The choice seems simple: either show how a structured, regular world could arise out of undirected processes, or inject intelligence into the system.<sup>[9]</sup>

—R. J. Hankinson, *Cause and Explanation in Ancient Greek Thought*

The mechanistic materialism that originated in ancient Greece was revived and revised by the French philosopher René Descartes, who held that animals and humans were assemblages of parts that together functioned as a machine. In the 19th century, the advances in cell theory in biological science encouraged this view and the evolutionary theory of Charles Darwin (1859) promised a mechanistic explanation for the origin of species on account of natural selection.<sup>[10]</sup>

## Hylomorphism

Hylomorphism is a theory (originating with Aristotle (322 BC)) that all things are a combination of matter and form. Biology was one of his main interests, and there is extensive biological material in his extant writings. In this view, all things in the material universe have both matter and form, and the form of a living thing is its soul (Greek *psyche*, Latin *anima*). There are three kinds of souls: the *vegetative soul* of plants, which causes them to grow and decay and nourish themselves, but does not cause motion and sensation; the *animal soul*, which causes animals to move and feel; and the *rational soul*, which is the source of consciousness and reasoning, which (Aristotle believed) is found only in man.<sup>[11]</sup> Each higher soul has all the attributes of the lower one. Aristotle believed that while matter can exist without form, form cannot exist without matter, and therefore the soul cannot exist without the body.<sup>[12]</sup>

This account is consistent with teleological explanations of life, which account for phenomena in terms of purpose or goal-directedness. Thus, the whiteness of the polar bear's coat is explained by its *purpose* of camouflage. The direction of causality (from the future to the past) is in contradiction with the scientific evidence for natural selection, which explains the consequence in terms of a prior cause. Biological features are explained not by looking at future optimal results, but by looking at the past evolutionary history of a species, which led to the natural selection of the features in question.<sup>[13]</sup>

## Vitalism

Vitalism is the belief that the life-principle is non-material. This originated with Stahl (17th century), and held sway until the middle of the 19th century. It appealed to philosophers such as Henri Bergson, Nietzsche, Wilhelm Dilthey,<sup>[14]</sup> anatomists like Bichat, and chemists like Liebig.<sup>[15]</sup> Vitalism included the idea that there was a fundamental difference between organic and inorganic material, and the belief that organic material can only be derived from living things. This was disproved in 1828, when Friedrich Wöhler prepared urea from inorganic materials.<sup>[16]</sup> This Wöhler synthesis is considered the starting point of modern organic chemistry. It is of historical significance because for the first time an organic compound was produced from inorganic reactants.<sup>[15]</sup>

During the 1850s, Helmholtz, anticipated by Mayer, demonstrated that no energy is lost in muscle movement, suggesting that there were no "vital forces" necessary to move a muscle.<sup>[17]</sup> These results led to the abandonment of scientific interest in vitalistic theories, although the belief lingered on in pseudoscientific theories such as homeopathy, which interprets diseases and sickness as caused by disturbances in a hypothetical vital force or life force.<sup>[18]</sup>

## Definitions

It is a challenge for scientists and philosophers to define life in unequivocal terms.<sup>[19][20][21]</sup> This is difficult partly because life is a process, not a pure substance.<sup>[22]</sup> Any definition must be sufficiently broad to encompass all life with which we are familiar, and must be sufficiently general to include life that may be fundamentally different from life on Earth.<sup>[23]</sup>

## Biology

Since there is no unequivocal definition of life, the current understanding is descriptive. Life is considered a characteristic of organisms that exhibit all or most of the following:<sup>[22][24]</sup>

1. **Homeostasis:** Regulation of the internal environment to maintain a constant state; for example, electrolyte concentration or sweating to reduce temperature.
2. **Organization:** Being structurally composed of one or more cells, which are the basic units of life.
3. **Metabolism:** Transformation of energy by converting chemicals and energy into cellular components (anabolism) and decomposing organic matter (catabolism). Living things require energy to maintain internal organization (homeostasis) and to produce the other phenomena associated with life.
4. **Growth:** Maintenance of a higher rate of anabolism than catabolism. A growing organism increases in size in all of its parts, rather than simply accumulating matter.
5. **Adaptation:** The ability to change over time in response to the environment. This ability is fundamental to the process of evolution and is determined by the organism's heredity, diet, and external factors.
6. **Response to stimuli:** A response can take many forms, from the contraction of a unicellular organism to external chemicals, to complex reactions involving all the senses of multicellular organisms. A response is often expressed by motion; for example, the leaves of a plant turning toward the sun (phototropism), and chemotaxis.
7. **Reproduction:** The ability to produce new individual organisms, either asexually from a single parent organism, or sexually from two parent organisms.

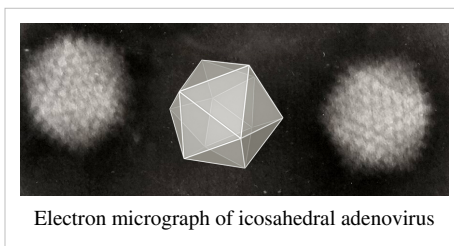
## Alternatives

To reflect the minimum phenomena required, other biological definitions of life have been proposed,<sup>[25]</sup> many of these are based upon chemical systems. Biophysicists have commented that living things function on negative entropy.<sup>[26][27]</sup> In other words, living processes can be viewed as a delay of the spontaneous diffusion or dispersion of the internal energy of biological molecules towards more potential microstates.<sup>[28]</sup> In more detail, according to physicists such as John Bernal, Erwin Schrödinger, Eugene Wigner, and John Avery, life is a member of the class of phenomena that are open or continuous systems able to decrease their internal entropy at the expense of substances or free energy taken in from the environment and subsequently rejected in a degraded form.<sup>[29][30][31]</sup> At a higher level, living beings are thermodynamic systems that have an organized molecular structure.<sup>[28]</sup> That is, life is matter that can reproduce itself and evolve as survival dictates.<sup>[32][33]</sup> Hence, life is a self-sustained chemical system capable of undergoing Darwinian evolution.<sup>[34]</sup>

Others take a systemic viewpoint that does not necessarily depend on molecular chemistry. One systemic definition of life is that living things are self-organizing and autopoietic (self-producing). Variations of this definition include Stuart Kauffman's definition as an autonomous agent or a multi-agent system capable of reproducing itself or themselves, and of completing at least one thermodynamic work cycle.<sup>[35]</sup> Life can be modeled as a network of inferior negative feedbacks of regulatory mechanisms subordinated to a superior positive feedback formed by the potential of expansion and reproduction.<sup>[36]</sup> Alternatively, life can be said to consist of things with the capacity for metabolism and motion,<sup>[22]</sup> or that life is self-reproduction "with variations"<sup>[37][38]</sup> or "with an error rate below the sustainability threshold."<sup>[38]</sup>

## Viruses

Viruses are most often considered replicators rather than forms of life. They have been described as "organisms at the edge of life,"<sup>[39]</sup> since they possess genes, evolve by natural selection,<sup>[40]</sup> and replicate by creating multiple copies of themselves through self-assembly. However, viruses do not metabolize and they require a host cell to make new products. Virus self-assembly within host cells has implications for the study of the origin of life, as it may support the hypothesis that life could have started as self-assembling organic molecules.<sup>[41][42]</sup>



Electron micrograph of icosahedral adenovirus

## Living systems theories

The idea that the Earth is alive is found in philosophy and religion, but the first scientific discussion was by the Scottish scientist James Hutton. In 1785, he stated that the Earth was a superorganism and that its proper study should be physiology. Hutton is considered the father of geology, but his idea of a living Earth was forgotten in the intense reductionism of the 19th century.<sup>[43]</sup> The Gaia hypothesis, proposed in the 1960s by scientist James Lovelock,<sup>[44][45]</sup> suggests that life on Earth functions as a single organism that defines and maintains environmental conditions necessary for its survival.<sup>[46]</sup>

A systems view of life treats environmental fluxes and biological fluxes together as a "reciprocity of influence",<sup>[47]</sup> and a reciprocal relation with environment is arguably as important for understanding life as it is for understanding ecosystems. As Harold J. Morowitz (1992) explains it, life is a property of an ecological system rather than a single organism or species.<sup>[48]</sup> He argues that an ecosystemic definition of life is preferable to a strictly biochemical or physical one. Robert Ulanowicz (2009) highlights mutualism as the key to understand the systemic, order-generating behavior of life and ecosystems.<sup>[49]</sup>

The first attempt at a general living systems theory for explaining the nature of life was in 1978, by American biologist James Grier Miller.<sup>[50]</sup> Such a general theory, arising out of the ecological and biological sciences, attempts to map general principles for how all living systems work. Instead of examining phenomena by attempting to break things down into component parts, a general living systems theory explores phenomena in terms of dynamic patterns of the relationships of organisms with their environment.<sup>[51]</sup> Robert Rosen (1991) built on this by defining a system component as "a unit of organization; a part with a function, i.e., a definite relation between part and whole." From this and other starting concepts, he developed a "relational theory of systems" that attempts to explain the special properties of life. Specifically, he identified the "nonfractionability of components in an organism" as the fundamental difference between living systems and "biological machines."<sup>[52]</sup>

## Origin

Evidence suggests that life on Earth has existed for about 3.7 billion years,<sup>[53]</sup> with the oldest traces of life found in fossils dating back 3.4 billion years.<sup>[54]</sup> All known life forms share fundamental molecular mechanisms, reflecting their common descent; based on these observations, hypotheses on the origin of life attempt to find a mechanism explaining the formation of a universal common ancestor, from simple organic molecules via pre-cellular life to protocells and metabolism. Models have been divided into "genes-first" and "metabolism-first" categories, but a recent trend is the emergence of hybrid models that combine both categories.<sup>[55]</sup>

There is no current scientific consensus as to how life originated. However, most accepted scientific models build on the following observations:

- The Miller-Urey experiment, and the work of Sidney Fox, show that conditions on the primitive Earth favored chemical reactions that synthesize amino acids and other organic compounds from inorganic precursors.
- Phospholipids spontaneously form lipid bilayers, the basic structure of a cell membrane.

Life synthesizes proteins, which are polymers of amino acids using instructions encoded by deoxyribonucleic acid (DNA). Protein synthesis entails intermediary ribonucleic acid (RNA) polymers. One possibility for how life began is that genes originated first, followed by proteins;<sup>[56]</sup> the alternative being that proteins came first and then genes.<sup>[57]</sup>

However, since genes and proteins are both required to produce the other, the problem of considering which came first is like that of the chicken or the egg. Most scientists have adopted the hypothesis that because of this, it is unlikely that genes and proteins arose independently.<sup>[58]</sup>

Therefore, a possibility, first suggested by Francis Crick,<sup>[59]</sup> is that the first life was based on RNA,<sup>[58]</sup> which has the DNA-like properties of information storage and the catalytic properties of some proteins. This is called the RNA world hypothesis, and it is supported by the observation that many of the most critical components of cells (those that evolve the slowest) are composed mostly or entirely of RNA. Also, many critical cofactors (ATP, Acetyl-CoA, NADH, etc.) are either nucleotides or substances clearly related to them. The catalytic properties of RNA had not yet been demonstrated when the hypothesis was first proposed,<sup>[60]</sup> but they were confirmed by Thomas Cech in 1986.<sup>[61]</sup>

One issue with the RNA world hypothesis is that synthesis of RNA from simple inorganic precursors is more difficult than for other organic molecules. One reason for this is that RNA precursors are very stable and react with each other very slowly under ambient conditions, and it has also been proposed that life consisted of other molecules before RNA.<sup>[62]</sup> However, the successful synthesis of certain RNA molecules under the conditions that existed prior to life on Earth has been achieved by adding alternative precursors in a specified order with the precursor phosphate present throughout the reaction.<sup>[63]</sup> This study makes the RNA world hypothesis more plausible.<sup>[64]</sup>

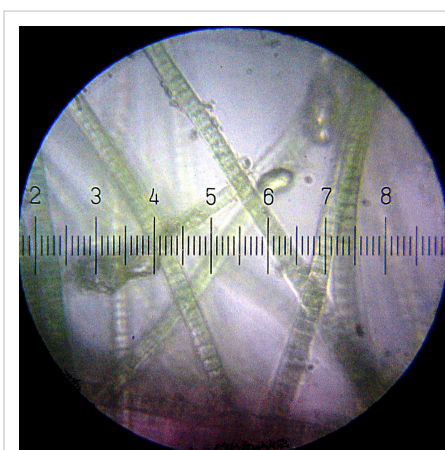
In 2009, experiments demonstrated Darwinian evolution of a two-component system of RNA enzymes (ribozymes) *in vitro*.<sup>[65]</sup> The work was performed in the laboratory of Gerald Joyce, who stated, "This is the first example, outside of biology, of evolutionary adaptation in a molecular genetic system."<sup>[66]</sup>

NASA findings in 2011, based on studies with meteorites found on Earth, suggest DNA and RNA components (adenine, guanine and related organic molecules) may be formed extraterrestrially in outer space.<sup>[67][68][69][70]</sup>

## Conditions

The diversity of life on Earth is a result of the dynamic interplay between genetic opportunity, metabolic capability, environmental challenges,<sup>[71]</sup> and symbiosis.<sup>[72][73][74]</sup> For most of its existence, Earth's habitable environment has been dominated by microorganisms and subjected to their metabolism and evolution. As a consequence of these microbial activities, the physical-chemical environment on Earth has been changing on a geologic time scale, thereby affecting the path of evolution of subsequent life.<sup>[71]</sup> For example, the release of molecular oxygen by cyanobacteria as a by-product of photosynthesis induced global changes in the Earth's environment. Since oxygen was toxic to most life on Earth at the time, this posed novel evolutionary challenges, and ultimately resulted in the formation of our planet's major animal and plant species. This interplay between organisms and their environment is an inherent feature of living systems.<sup>[71]</sup>

All life forms require certain core chemical elements needed for biochemical functioning. These include carbon, hydrogen, nitrogen, oxygen, phosphorus, and sulfur—the elemental macronutrients for all organisms<sup>[75]</sup>—often represented by the acronym CHNOPS. Together these make up nucleic acids, proteins and lipids, the bulk of living matter. Five of these six elements comprise the chemical components of DNA, the exception being sulfur. The latter is a component of the amino acids cysteine and methionine. The most



Cyanobacteria dramatically changed the composition of life forms on Earth by leading to the near-extinction of oxygen-intolerant organisms.

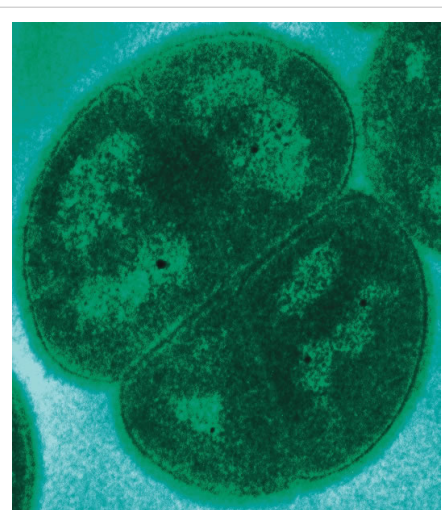
biologically abundant of these elements is carbon, which has the desirable attribute of forming multiple, stable covalent bonds. This allows carbon-based (organic) molecules to form an immense variety of chemical arrangements.<sup>[76]</sup> Alternative hypothetical types of biochemistry have been proposed that eliminate one or more of these elements, swap out an element for one not on the list, or change required chiralities or other chemical properties.<sup>[77][78]</sup>

## Range of tolerance

The inert components of an ecosystem are the physical and chemical factors necessary for life - energy (sunlight or chemical energy), water, temperature, atmosphere, gravity, nutrients, and ultraviolet solar radiation protection.<sup>[79]</sup> In most ecosystems, the conditions vary during the day and from one season to the next. To live in most ecosystems, then, organisms must be able to survive a range of conditions, called the "range of tolerance."<sup>[80]</sup> Outside that are the "zones of physiological stress," where the survival and reproduction are possible but not optimal. Beyond these zones are the "zones of intolerance," where survival and reproduction of that organism is unlikely or impossible. Organisms that have a wide range of tolerance are more widely distributed than organisms with a narrow range of tolerance.<sup>[80]</sup>

To survive, selected microorganisms can assume forms that enable them to withstand freezing, complete desiccation, starvation, high levels of radiation exposure, and other physical or chemical challenges. These microorganisms may survive exposure to such conditions for weeks, months, years, or even centuries.<sup>[71]</sup> Extremophiles are microbial life forms that thrive outside the ranges where life is commonly found. They excel at exploiting uncommon sources of energy. While all organisms are composed of nearly identical molecules, evolution has enabled such microbes to cope with this wide range of physical and chemical conditions. Characterization of the structure and metabolic diversity of microbial communities in such extreme environments is ongoing.<sup>[81]</sup>

Investigation of the tenacity and versatility of life on Earth, as well as an understanding of the molecular systems that some organisms utilize to survive such extremes, is important for the search for life beyond Earth.<sup>[71]</sup> In April 2012, scientists reported that lichen could survive and reproduce in a simulated Martian environment.<sup>[82][83]</sup>



*Deinococcus radiodurans* is an extremophile that can resist extremes of cold, dehydration, vacuum, acid, and radiation exposure.

## Form and function

Cells are the basic unit of structure in every living thing, and all cells arise from pre-existing cells by division. Cell theory was formulated by Henri Dutrochet, Theodor Schwann, Rudolf Virchow and others during the early nineteenth century, and subsequently became widely accepted.<sup>[84]</sup> The activity of an organism depends on the total activity of its cells, with energy flow occurring within and between them. Cells contain hereditary information that is carried forward as a genetic code during cell division.<sup>[85]</sup>

There are two primary types of cells. Prokaryotes lack a nucleus and other membrane-bound organelles, although they have circular DNA and ribosomes. Bacteria and Archaea are two domains of prokaryotes. The other primary type of cells are the eukaryotes, which have distinct nuclei bound by a nuclear membrane and membrane-bound organelles, including mitochondria, chloroplasts, lysosomes, rough and smooth endoplasmic reticulum, and vacuoles. In addition, they possess organized chromosomes that store genetic material. All species of large complex organisms are eukaryotes, including animals, plants and fungi, though most species of eukaryote are protist

microorganisms.<sup>[86]</sup> The conventional model is that eukaryotes evolved from prokaryotes, with the main organelles of the eukaryotes forming through endosymbiosis between bacteria and the progenitor eukaryotic cell.<sup>[87]</sup>

The molecular mechanisms of cell biology are based on proteins. Most of these are synthesized by the ribosomes through an enzyme-catalyzed process called protein biosynthesis. A sequence of amino acids is assembled and joined together based upon gene expression of the cell's nucleic acid.<sup>[88]</sup> In eukaryotic cells, these proteins may then be transported and processed through the Golgi apparatus in preparation for dispatch to their destination.

Cells reproduce through a process of cell division in which the parent cell divides into two or more daughter cells. For prokaryotes, cell division occurs through a process of fission in which the DNA is replicated, then the two copies are attached to parts of the cell membrane. In eukaryotes, a more complex process of mitosis is followed. However, the end result is the same; the resulting cell copies are identical to each other and to the original cell (except for mutations), and both are capable of further division following an interphase period.<sup>[89]</sup>

Multicellular organisms may have first evolved through the formation of colonies of like cells. These cells can form group organisms through cell adhesion. The individual members of a colony are capable of surviving on their own, whereas the members of a true multi-cellular organism have developed specialties, making them dependent on the remainder of the organism for survival. Such organisms are formed clonally or from a single germ cell that is capable of forming the various specialized cells that form the adult organism. This specialization allows multicellular organisms to exploit resources more efficiently than single cells.<sup>[90]</sup>

Cells have evolved methods to perceive and respond to their microenvironment, thereby enhancing their adaptability. Cell signaling coordinates cellular activities, and hence governs the basic functions of multicellular organisms. Signaling between cells can occur through direct cell contact using juxtacrine signalling, or indirectly through the exchange of agents as in the endocrine system. In more complex organisms, coordination of activities can occur through a dedicated nervous system.<sup>[91]</sup>

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## Classification

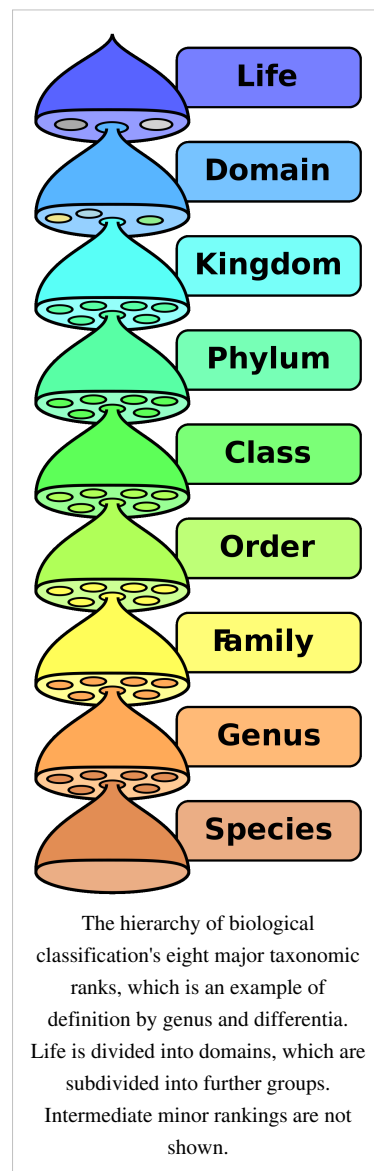
The first known attempt to classify organisms was conducted by the Greek philosopher Aristotle (384–322 BC), who classified all living organisms known at that time as either a plant or an animal, based mainly on their ability to move. He also distinguished animals with blood from animals without blood (or at least without red blood), which can be compared with the concepts of vertebrates and invertebrates respectively, and divided the blooded animals into five groups: viviparous quadrupeds (mammals), oviparous quadrupeds (reptiles and amphibians), birds, fishes and whales. The bloodless animals were also divided into five groups: cephalopods, crustaceans, insects (which included the spiders, scorpions, and centipedes, in addition to what we define as insects today), shelled animals (such as most molluscs and echinoderms) and "zoophytes." Though Aristotle's work in zoology was not without errors, it was the grandest biological synthesis of the time and remained the ultimate authority for many centuries after his death.<sup>[92]</sup>

The exploration of the American continent revealed large numbers of new plants and animals that needed descriptions and classification. In the latter part of the 16th century and the beginning of the 17th, careful study of animals commenced and was gradually extended until it formed a sufficient body of knowledge to serve as an anatomical basis for classification. In the late 1740s, Carolus Linnaeus introduced his system of binomial nomenclature for the classification of species.<sup>[93]</sup> Linnaeus attempted to improve the composition and reduce the length of the previously used many-worded names by abolishing unnecessary rhetoric, introducing new descriptive terms and precisely defining their meaning. By consistently using this system, Linnaeus separated nomenclature from taxonomy.

The fungi were originally treated as plants. For a short period Linnaeus had classified them in the taxon Vermes in Animalia, but later placed them back in Plantae. Copeland classified the Fungi in his Protoctista, thus partially avoiding the problem but acknowledging their special status.<sup>[94]</sup> The problem was eventually solved by Whittaker, when he gave them their own kingdom in his five-kingdom system. Evolutionary history shows that the fungi are more closely related to animals than to plants.<sup>[95]</sup>

As new discoveries enabled detailed study of cells and microorganisms, new groups of life were revealed, and the fields of cell biology and microbiology were created. These new organisms were originally described separately in protozoa as animals and protophyta/thallophyta as plants, but were united by Haeckel in the kingdom Protista; later, the prokaryotes were split off in the kingdom Monera, which would eventually be divided into two separate groups, the Bacteria and the Archaea. This led to the six-kingdom system and eventually to the current three-domain system, which is based on evolutionary relationships.<sup>[96]</sup> However, the classification of eukaryotes, especially of protists, is still controversial.<sup>[97]</sup>

As microbiology, molecular biology and virology developed, non-cellular reproducing agents were discovered, such as viruses and viroids. Whether these are considered alive has been a matter of debate; viruses lack characteristics of life such as cell membranes, metabolism and the ability to grow or respond to their environments. Viruses can still be classed into "species" based on their biology and genetics, but many aspects of such a classification remain controversial.<sup>[98]</sup>



In the 1960s a trend called cladistics emerged, arranging taxa based on clades in an evolutionary or phylogenetic tree.<sup>[99]</sup>

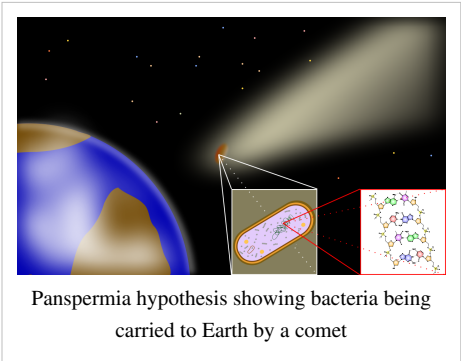
Linnaeus 1735 <sup>[100]</sup>	Haeckel 1866 <sup>[101]</sup>	Chatton 1925 <sup>[102][103]</sup>	Copeland 1938 <sup>[94][104]</sup>	Whittaker 1969 <sup>[105]</sup>	Woese et al. 1977 <sup>[106][107]</sup>	Woese et al. 1990 <sup>[96]</sup>	Cavalier-Smith 2004 <sup>[108]</sup>
2 kingdoms	3 kingdoms	2 empires	4 kingdoms	5 kingdoms	6 kingdoms	3 domains	6 kingdoms
(not treated)	Protista	Prokaryota	Monera	Monera	Eubacteria	Bacteria	Bacteria
					Archaeobacteria	Archaea	
		Eukaryota	Protoctista	Protista	Protista	Eukarya	Protozoa
							Chromista
Vegetabilia	Plantae		Plantae	Plantae	Plantae		Plantae
				Fungi	Fungi		Fungi
Animalia	Animalia		Animalia	Animalia	Animalia		Animalia

Extraterrestrial life

Earth is the only planet known to harbor life. Other locations within the Solar System that may host life include subsurface Mars, the atmosphere of Venus,<sup>[109]</sup> and subsurface oceans on some of the moons of the gas giant planets.<sup>[110]</sup> The Drake equation, which predicts the number of extraterrestrial civilizations in our galaxy with which we might come in contact, has been used to discuss the probability of life elsewhere, but many of the variables in this equation are difficult to estimate.<sup>[111]</sup>

The region around a main sequence star that could support Earth-like life on an Earth-like planet is known as the habitable zone. The inner and outer radii of this zone vary with the luminosity of the star, as does the time interval during which the zone survives. Stars more massive than the Sun have a larger habitable zone, but remain on the main sequence for a shorter time interval. Small red dwarf stars have the opposite problem, with a smaller habitable zone that is subject to higher levels of magnetic activity and the effects of tidal locking from close orbits. Hence, stars in the intermediate mass range such as the Sun may have a greater likelihood for Earth-like life to develop.<sup>[112]</sup> The location of the star within a galaxy may also have an impact on the likelihood of life forming. Stars in regions with a greater abundance of heavier elements that can form planets, in combination with a low rate of potentially habitat-damaging supernova events, are predicted to have a higher probability of hosting planets with complex life.<sup>[113]</sup>

Panspermia, also called exogenesis, is a hypothesis that life originated elsewhere in the universe and subsequently transferred to Earth in the form of spores via meteorites, comets, or cosmic dust. In October 2011, scientists found using spectroscopy that cosmic dust contains complex organic matter (specifically, aromatic-aliphatic organic solids) that could be created naturally, and rapidly, by stars.<sup>[114][115][116]</sup> The compounds are so complex that their chemical structures resemble the makeup of coal and petroleum; such chemical complexity was previously thought to arise only from living organisms.<sup>[114]</sup> These observations suggest that organic compounds introduced on Earth by interstellar dust particles could serve as basic ingredients for life due to their surface-catalytic activities.<sup>[70][117]</sup>



## Death

Death is the permanent termination of all vital functions or life processes in an organism or cell.<sup>[118][119]</sup> It can occur as a result of an accident, medical conditions, biological interaction, malnutrition, poisoning, senescence, or suicide. After death, the remains of an organism re-enter the biogeochemical cycle. Organisms may be consumed by a predator or a scavenger and leftover organic material may then be further decomposed by detritivores, organisms that recycle detritus, returning it to the environment for reuse in the food chain.

One of the challenges in defining death is in distinguishing it from life. Death would seem to refer to either the moment life ends, or when the state that follows life begins.<sup>[119]</sup> However, determining when death has occurred requires drawing precise conceptual boundaries between life and death. This is problematic, however, because there is little consensus over how to define life. The nature of death has for millennia been a central concern of the world's religious traditions and of philosophical inquiry. Many religions maintain faith in either a kind of afterlife or reincarnation for the soul, or resurrection of the body at a later date.

Extinction is the process by which a group of taxa or species dies out, reducing biodiversity.<sup>[120]</sup> The moment of extinction is generally considered the death of the last individual of that species. Because a species' potential range may be very large, determining this moment is difficult, and is usually done retrospectively after a period of apparent absence. Species become extinct when they are no longer able to survive in changing habitat or against superior competition. In Earth's history, over 99% of all the species that have ever lived have gone extinct;<sup>[121]</sup> however, mass extinctions may have accelerated evolution by providing opportunities for new groups of organisms to diversify.<sup>[122]</sup>

Fossils are the preserved remains or traces of animals, plants, and other organisms from the remote past. The totality of fossils, both discovered and undiscovered, and their placement in fossil-containing rock formations and sedimentary layers (strata) is known as the *fossil record*. A preserved specimen is called a fossil if it is older than the arbitrary date of 10,000 years ago.<sup>[123]</sup> Hence, fossils range in age from the youngest at the start of the Holocene Epoch to the oldest from the Archaean Eon, up to 3.4 billion years old.<sup>[124][125]</sup>

## Artificial life

Artificial life is a field of study that examines systems related to life, its processes, and its evolution through simulations using computer models, robotics, and biochemistry.<sup>[126]</sup> The study of artificial life imitates traditional biology by recreating some aspects of biological phenomena. Scientists study the logic of living systems by creating artificial environments—seeking to understand the complex information processing that defines such systems. While life is, by definition, alive, artificial life is generally referred to as data confined to a digital environment and existence.

Synthetic biology is a new area of biological research and technology that combines science and engineering. The common goal is the design and construction of new biological functions and systems not found in nature. Synthetic biology includes the broad redefinition and expansion of biotechnology, with the ultimate goals of being able to design and build engineered biological systems that process information, manipulate chemicals, fabricate materials and structures, produce energy, provide food, and maintain and enhance human health and our environment.<sup>[127]</sup>



Animal corpses, like this African buffalo, are recycled by the ecosystem, providing energy and nutrients for living creatures

## Notes

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## Further reading

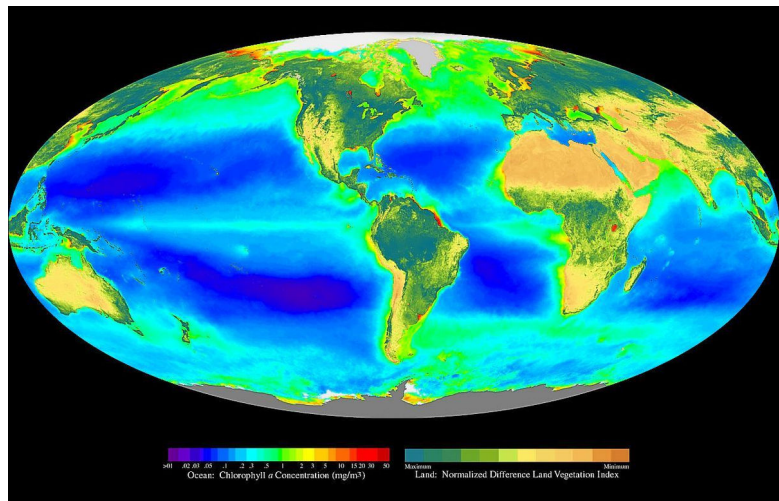
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## External links

- Wikispecies ([http://species.wikimedia.org/wiki/Main\\_Page](http://species.wikimedia.org/wiki/Main_Page)) – a free directory of life
- "The Adjacent Possible: A Talk with Stuart Kauffman" ([http://www.edge.org/3rd\\_culture/kauffman03/kauffman\\_index.html](http://www.edge.org/3rd_culture/kauffman03/kauffman_index.html))
- Stanford Encyclopedia of Philosophy entry (<http://plato.stanford.edu/entries/life/>)
- Life under extreme conditions (<http://www.larger-than-life.org/modules.php?name=Content&pa=showpage&pid=2>) An in depth look at how life can form under the most extreme conditions and circumstances.

# Biosphere

The **biosphere** is the global sum of all ecosystems. It can also be called the zone of life on Earth, a closed (apart from solar and cosmic radiation), and self-regulating system.<sup>[1]</sup> From the broadest biophysiological point of view, the biosphere is the global ecological system integrating all living beings and their relationships, including their interaction with the elements of the lithosphere, hydrosphere, and atmosphere. The biosphere is postulated to have evolved, beginning through a process of biogenesis or biopoesis, at least some 3.5 billion years ago.<sup>[2]</sup>



A false-color composite of global oceanic and terrestrial photoautotroph abundance, from September 1997 to August 2000. Provided by the SeaWiFS Project, NASA/Goddard Space Flight Center and ORBIMAGE.

In a broader sense; biospheres are any closed, self-regulating systems containing ecosystems; including artificial ones such as Biosphere 2 and BIOS-3; and, potentially, ones on other planets or moons.<sup>[3]</sup>

## Origin and use of the term

The term "biosphere" was coined by geologist Eduard Suess in 1875, which he defined as:<sup>[4]</sup>

"The place on Earth's surface where life dwells."

While this concept has a geological origin, it is an indication of the impact of both Charles Darwin and Matthew F. Maury on the earth sciences. The biosphere's ecological context comes from the 1920s (*see* Vladimir I. Vernadsky), preceding the 1935 introduction of the term "ecosystem" by Sir Arthur Tansley (*see* ecology history). Vernadsky defined ecology as the science of the biosphere. It is an interdisciplinary concept for integrating astronomy, geophysics, meteorology, biogeography, evolution, geology, geochemistry, hydrology and, generally speaking, all life and earth sciences.

## Narrow definition

Some life scientists and earth scientists use *biosphere* in a more limited sense. For example, geochemists define the biosphere as being the total sum of living organisms (the "biomass" or "biota" as referred to by biologists and ecologists). In this sense, the biosphere is but one of four separate components of the geochemical model, the other three being *lithosphere*, *hydrosphere*, and *atmosphere*. The narrow meaning used by geochemists is one of the consequences of specialization in modern science. Some might prefer the word *ecosphere*, coined in the 1960s, as all encompassing of both biological and physical components of the planet.



A familiar scene on Earth which simultaneously shows the lithosphere, hydrosphere and atmosphere.

The Second International Conference on Closed Life Systems defined *biospherics* as the science and technology of analogs and models of Earth's biosphere; i.e., artificial Earth-like biospheres. Others may include the creation of artificial non-Earth biospheres—for example, human-centered biospheres or a native Martian biosphere—in the field of biospherics.

## Gaia hypothesis

In the early 1970s, the British chemist James Lovelock and Lynn Margulis, a microbiologist from the United States, added to the hypothesis, specifically noting the ties between the biosphere and other Earth systems. For example, when carbon dioxide levels increase in the atmosphere, plants grow more quickly. As their growth continues, they remove more and more carbon dioxide from the atmosphere.

Many scientists are now involved in new fields of study that examine interactions between biotic and abiotic factors in the biosphere, such as geobiology and geomicrobiology.

Ecosystems occur when communities and their physical environment work together as a system. The difference between this and a biosphere is simple, the biosphere is everything in general terms.

## Extent of Earth's biosphere



Water covers 71% of the Earth's surface. Image is the Earth photographed from Apollo 17.

Every part of the planet, from the polar ice caps to the Equator, supports life of some kind. Recent advances in microbiology have demonstrated that microbes live deep beneath the Earth's terrestrial surface, and that the total mass of microbial life in so-called "uninhabitable zones" may, in biomass, exceed all animal and plant life on the surface. The actual thickness of the biosphere on earth is difficult to measure. Birds typically fly at altitudes of 650 to 1,800 metres, and fish that live deep underwater can be found down to -8,372 metres in the Puerto Rico Trench.<sup>[2]</sup>

There are more extreme examples for life on the planet: Rüppell's vulture has been found at altitudes of 11,300 metres; bar-headed geese migrate at altitudes of at least 8,300 metres; yaks live at elevations between 3,200 to 5,400 metres above sea level; mountain goats live up to 3,050 metres. Herbivorous animals at these elevations depend on lichens, grasses, and herbs.

Microscopic organisms live at such extremes that, taking them into consideration, makes the thickness of the biosphere much greater. Culturable microbes have been found in the Earth's upper atmosphere as high as 41 km (**unknown operator: u'strong'** mi) (Wainwright et al., 2003, in FEMS Microbiology Letters). It is unlikely, however, that microbes are active at such altitudes, where temperatures and air pressure are extremely low and ultraviolet radiation very high. More likely, these microbes were brought into the upper atmosphere by winds or possibly volcanic eruptions. Barophilic marine microbes have been found at more than 10 km (**unknown operator: u'strong'** mi) depth in the Marianas Trench (Takamia et al., 1997, in FEMS Microbiology Letters). Microbes are not limited to the air, water or the Earth's surface. Culturable thermophilic microbes have been extracted from cores drilled more than 5 km (**unknown operator: u'strong'** mi) into the Earth's crust in Sweden (Gold, 1992, and Szewzyk, 1994, both in PNAS), from rocks between 65-75 °C. Temperature increases with increasing depth into the Earth's crust. The speed at which the temperature increases depends on many factors, including type of crust (continental vs. oceanic), rock type, geographic location, etc. The upper known limit of temperature at which microbial life can exist is 122 °C (*Methanopyrus kandleri* Strain 116), and it is likely that the limit of life in the "deep biosphere" is defined by temperature rather than absolute depth.

Our biosphere is divided into a number of biomes, inhabited by broadly similar flora and fauna. On land, biomes are separated primarily by latitude. Terrestrial biomes lying within the Arctic and Antarctic Circles are relatively barren of plant and animal life, while most of the more populous biomes lie near the equator. Terrestrial organisms in temperate and Arctic biomes have relatively small amounts of total biomass, smaller energy budgets, and display prominent adaptations to cold, including world-spanning migrations, social adaptations, homeothermy, estivation and multiple layers of insulation.

## Specific biospheres

When the word is followed by a number, it is usually referring to a specific system or number. Thus:

- Biosphere 1, the planet Earth
- Biosphere 2, a laboratory in Arizona which contains 3.15 acres (13,000 m<sup>2</sup>) of closed ecosystem.
- BIOS-3, a closed ecosystem at the Institute of Biophysics in Krasnoyarsk, Siberia, in what was then the Soviet Union.
- Biosphere J (CEEFF, Closed Ecology Experiment Facilities), an experiment in Japan.<sup>[5][6]</sup>

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- <http://www.intechopen.com/books/the-biosphere>

## External links

- Article on the Biosphere at Encyclopedia of Earth (<http://www.eoearth.org/article/Biosphere>)
- GLOBIO.info (<http://www.globio.info/>), an ongoing programme to map the past, current and future impacts of human activities on the biosphere
- Paul Crutzen Interview (<http://www.vega.org.uk/video/programme/111>) Freeview video of Paul Crutzen Nobel Laureate for his work on decomposition of ozone talking to Harry Kroto Nobel Laureate by the Vega Science Trust.
- Atlas of the Biosphere (<http://www.sage.wisc.edu/atlas/>)



# Evolution

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**Evolution** is the change in the inherited characteristics of biological populations over successive generations. Evolutionary processes give rise to diversity at every level of biological organisation, including species, individual organisms and molecules such as DNA and proteins.<sup>[1]</sup>

Life on Earth originated and then evolved from a universal common ancestor approximately 3.7 billion years ago. Repeated speciation and the divergence of life can be inferred from shared sets of biochemical and morphological traits, or by shared DNA sequences. These homologous traits and sequences are more similar among species that share a more recent common ancestor, and can be used to reconstruct evolutionary histories, using both existing species and the fossil record. Existing patterns of biodiversity have been shaped both by speciation and by extinction.<sup>[2]</sup>

Charles Darwin was the first to formulate a scientific argument for the theory of evolution by means of natural selection. Evolution by natural selection is a process that is inferred from three facts about populations: 1) more offspring are produced than can possibly survive, 2) traits vary among individuals, leading to differential rates of survival and reproduction, and 3) trait differences are heritable.<sup>[3]</sup> Thus, when members of a population die they are replaced by the progeny of parents that were better adapted to survive and reproduce in the environment in which natural selection took place. This process creates and preserves traits that are seemingly fitted for the functional roles they perform.<sup>[4]</sup> Natural selection is the only known cause of adaptation, but not the only known cause of evolution. Other, nonadaptive causes of evolution include mutation and genetic drift.<sup>[5]</sup>

In the early 20th century, genetics was integrated with Darwin's theory of evolution by natural selection through the discipline of population genetics. The importance of natural selection as a cause of evolution was accepted into other branches of biology. Moreover, previously held notions about evolution, such as orthogenesis and "progress" became obsolete.<sup>[6]</sup> Scientists continue to study various aspects of evolution by forming and testing hypotheses, constructing scientific theories, using observational data, and performing experiments in both the field and the laboratory. Biologists agree that descent with modification is one of the most reliably established facts in science.<sup>[7]</sup> Discoveries in evolutionary biology have made a significant impact not just within the traditional branches of biology, but also in other academic disciplines (e.g., anthropology and psychology) and on society at large.<sup>[8][9]</sup>

## History of evolutionary thought

Further information: History of evolutionary thought

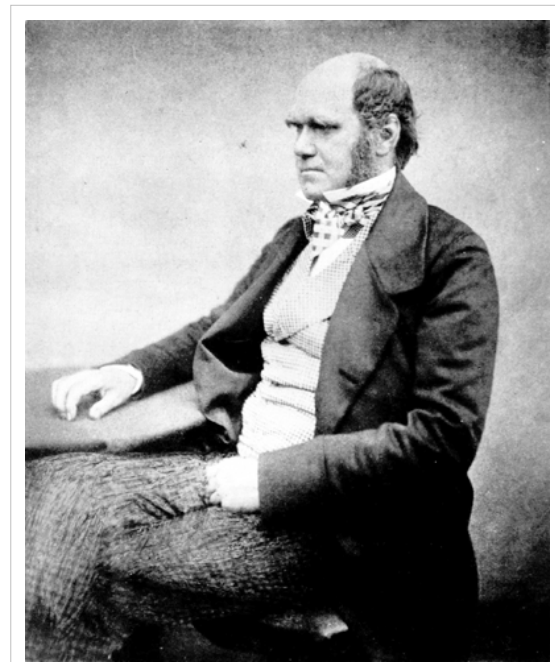
The proposal that one type of animal could descend from an animal of another type goes back to some of the first pre-Socratic Greek philosophers, such as Anaximander and Empedocles.<sup>[10][11]</sup> In contrast to these materialistic views, Aristotle understood all natural things, not only living things, as being imperfect actualisations of different fixed natural possibilities, known as "forms", "ideas", or (in Latin translations) "species".<sup>[12][13]</sup> This was part of his teleological understanding of nature in which all things have an intended role to play in a divine cosmic order. Variations of this idea became the standard understanding of the Middle Ages, and were integrated into Christian learning, but Aristotle did not demand that real types of animals corresponded one-for-one with exact metaphysical forms, and specifically gave examples of how new types of living things could come to be.<sup>[14]</sup>

In the 17th century the new method of modern science rejected Aristotle's approach, and sought explanations of natural phenomena in terms of laws of nature which were the same for all visible things, and did not need to assume any fixed natural categories, nor any divine cosmic order. But this new approach was slow to take root in the biological sciences, which became the last bastion of the concept of fixed natural types. John Ray used one of the previously more general terms for fixed natural types, "species", to apply to animal and plant types, but unlike Aristotle he strictly identified each type of living thing as a species, and proposed that each species can be defined by the features that perpetuate themselves each generation.<sup>[15]</sup> These species were designed by God, but showing

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differences caused by local conditions. The biological classification introduced by Carolus Linnaeus in 1735 also viewed species as fixed according to a divine plan.<sup>[16]</sup>

Other naturalists of this time speculated on evolutionary change of species over time according to natural laws. Maupertuis wrote in 1751 of natural modifications occurring during reproduction and accumulating over many generations to produce new species.<sup>[18]</sup> Buffon suggested that species could degenerate into different organisms, and Erasmus Darwin proposed that all warm-blooded animals could have descended from a single micro-organism (or "filament").<sup>[19]</sup> The first full-fledged evolutionary scheme was Lamarck's "transmutation" theory of 1809 which envisaged spontaneous generation continually producing simple forms of life developed greater complexity in parallel lineages with an inherent progressive tendency, and that on a local level these lineages adapted to the environment by inheriting changes caused by use or disuse in parents.<sup>[20][21]</sup> (The latter process was later called Lamarckism.)<sup>[20][22][23][24]</sup> These ideas were condemned by establishment naturalists as speculation lacking empirical support. In particular Georges Cuvier insisted that species were unrelated and fixed, their similarities reflecting divine design for functional needs. In the meantime, Ray's ideas of benevolent design had been developed by William Paley into a natural theology which proposed complex adaptations as evidence of divine design, and was admired by Charles Darwin.<sup>[25][26][27]</sup>



In 1842 Charles Darwin penned his first sketch of what became *On the Origin of Species*.<sup>[17]</sup>

The critical break from the concept of fixed species in biology began with the theory of evolution by natural selection, which was formulated by Charles Darwin. Partly influenced by *An Essay on the Principle of Population* by Thomas Robert Malthus, Darwin noted that population growth would lead to a "struggle for existence" where favorable variations could prevail as others perished. Each generation, many offspring fail to survive to an age of reproduction because of limited resources. This could explain the diversity of animals and plants from a common ancestry through the working of natural laws working the same for all types of thing.<sup>[28][29][30][31]</sup> Darwin was developing his theory of "natural selection" from 1838 onwards until Alfred Russel Wallace sent him a similar theory in 1858. Both men presented their separate papers to the Linnean Society of London.<sup>[32]</sup> At the end of 1859, Darwin's publication of *On the Origin of Species* explained natural selection in detail and in a way that led to an increasingly wide acceptance of Darwinian evolution. Thomas Henry Huxley applied Darwin's ideas to humans, using paleontology and comparative anatomy to provide strong evidence that humans and apes shared a common ancestry. Some were disturbed by this since it implied that humans did not have a special place in the universe.<sup>[33]</sup>

Precise mechanisms of reproductive heritability and the origin of new traits remained a mystery. Towards this end, Darwin developed his provisional theory of pangenesis.<sup>[34]</sup> In 1865 Gregor Mendel reported that traits were inherited in a predictable manner through the independent assortment and segregation of elements (later known as genes). Mendel's laws of inheritance eventually supplanted most of Darwin's pangenesis theory.<sup>[35]</sup> August Weismann made the important distinction between germ cells (sperm and eggs) and somatic cells of the body, demonstrating that heredity passes through the germ line only. Hugo de Vries connected Darwin's pangenesis theory to Weismann's germ/soma cell distinction and proposed that Darwin's pangenes were concentrated in the cell nucleus and when expressed they could move into the cytoplasm to change the cell's structure. De Vries was also one of the researchers who made Mendel's work well-known, believing that Mendelian traits corresponded to the transfer of heritable variations along the germline.<sup>[36]</sup> To explain how new variants originate, De Vries developed a mutation theory that

led to a temporary rift between those who accepted Darwinian evolution and biometricians who allied with de Vries.<sup>[21][37][38]</sup> At the turn of the 20th century, pioneers in the field of population genetics, such as J.B.S. Haldane, Sewall Wright, and Ronald Fisher, set the foundations of evolution onto a robust statistical philosophy. The false contradiction between Darwin's theory, genetic mutations, and Mendelian inheritance was thus reconciled.<sup>[39]</sup>

In the 1920s and 1930s a modern evolutionary synthesis connected natural selection, mutation theory, and Mendelian inheritance into a unified theory that applied generally to any branch of biology. The modern synthesis was able to explain patterns observed across species in populations, through fossil transitions in palaeontology, and even complex cellular mechanisms in developmental biology.<sup>[21][40]</sup> The publication of the structure of DNA by James Watson and Francis Crick in 1953 demonstrated a physical basis for inheritance.<sup>[41]</sup> Molecular biology improved our understanding of the relationship between genotype and phenotype. Advancements were also made in phylogenetic systematics, mapping the transition of traits into a comparative and testable framework through the publication and use of evolutionary trees.<sup>[42][43]</sup> In 1973, evolutionary biologist Theodosius Dobzhansky penned that "nothing in biology makes sense except in the light of evolution", because it has brought to light the relations of what first seemed disjointed facts in natural history into a coherent explanatory body of knowledge that describes and predicts many observable facts about life on this planet.<sup>[44]</sup>

Since then, the modern synthesis has been further extended to explain biological phenomena across the full and integrative scale of the biological hierarchy, from genes to species. This extension has been dubbed "eco-evo-devo".<sup>[45][45][46][47]</sup>

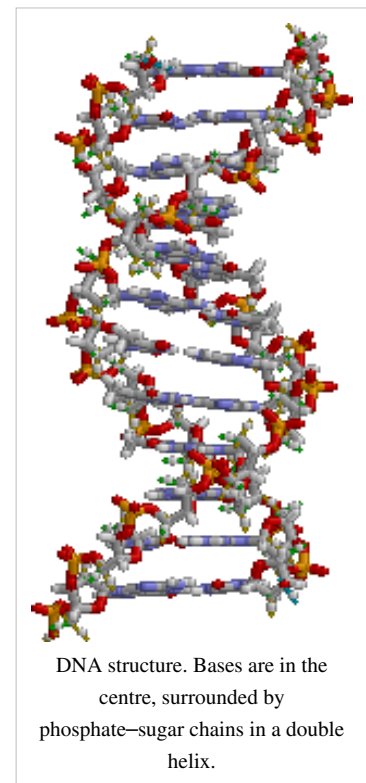
## Heredity

Further information: Introduction to genetics, Genetics, Heredity, and Norms of reaction

Evolution in organisms occurs through changes in heritable traits – particular characteristics of an organism. In humans, for example, eye colour is an inherited characteristic and an individual might inherit the "brown-eye trait" from one of their parents.<sup>[48]</sup> Inherited traits are controlled by genes and the complete set of genes within an organism's genome is called its genotype.<sup>[49]</sup>

The complete set of observable traits that make up the structure and behaviour of an organism is called its phenotype. These traits come from the interaction of its genotype with the environment.<sup>[50]</sup> As a result, many aspects of an organism's phenotype are not inherited. For example, suntanned skin comes from the interaction between a person's genotype and sunlight; thus, suntans are not passed on to people's children. However, some people tan more easily than others, due to differences in their genotype; a striking example are people with the inherited trait of albinism, who do not tan at all and are very sensitive to sunburn.<sup>[51]</sup>

Heritable traits are passed from one generation to the next via DNA, a molecule that encodes genetic information.<sup>[49]</sup> DNA is a long polymer composed of four types of bases. The sequence of bases along a particular DNA molecule specify the genetic information, in a manner similar to a sequence of letters spelling out a sentence. Before a cell divides, the DNA is copied, so that each of the resulting two cells will inherit the DNA sequence. Portions of a DNA molecule that specify a single functional unit are called genes; different genes have different sequences of bases. Within cells, the long strands of DNA form condensed structures called chromosomes. The specific location of a DNA sequence within a chromosome is known as a locus. If the DNA sequence at a locus varies between individuals, the different forms of this sequence are called alleles. DNA sequences can change through mutations, producing new alleles. If a



mutation occurs within a gene, the new allele may affect the trait that the gene controls, altering the phenotype of the organism.<sup>[52]</sup> However, while this simple correspondence between an allele and a trait works in some cases, most traits are more complex and are controlled by multiple interacting genes.<sup>[53][54]</sup>

Recent findings have confirmed important examples of heritable changes that cannot be explained by changes to the sequence of nucleotides in the DNA. These phenomena are classed as epigenetic inheritance systems.<sup>[55]</sup> DNA methylation marking chromatin, self-sustaining metabolic loops, gene silencing by RNA interference and the three dimensional conformation of proteins (such as prions) are areas where epigenetic inheritance systems have been discovered at the organismic level.<sup>[56][57]</sup> Developmental biologists suggest that complex interactions in genetic networks and communication among cells can lead to heritable variations that may underlay some of the mechanics in developmental plasticity and canalization.<sup>[58]</sup> Heritability may also occur at even larger scales. For example, ecological inheritance through the process of niche construction is defined by the regular and repeated activities of organisms in their environment. This generates a legacy of effects that modify and feed back into the selection regime of subsequent generations. Descendants inherit genes plus environmental characteristics generated by the ecological actions of ancestors.<sup>[59]</sup> Other examples of heritability in evolution that are not under the direct control of genes include the inheritance of cultural traits and symbiogenesis.<sup>[60][61]</sup>

## Variation



White peppered moth



Black morph in peppered moth evolution

Further information: Genetic diversity and Population genetics

An individual organism's phenotype results from both its genotype and the influence from the environment it has lived in. A substantial part of the variation in phenotypes in a population is caused by the differences between their genotypes.<sup>[54]</sup> The modern evolutionary synthesis defines evolution as the change over time in this genetic variation. The frequency of one particular allele will become more or less prevalent relative to other forms of that gene. Variation disappears when a new allele reaches the point of fixation — when it either disappears from the population or replaces the ancestral allele entirely.<sup>[62]</sup>

Natural selection will only cause evolution if there is enough genetic variation in a population. Before the discovery of Mendelian genetics, one common hypothesis was blending inheritance. But with blending inheritance, genetic variance would be rapidly lost, making evolution by natural selection implausible. The *Hardy-Weinberg principle* provides the solution to how variation is maintained in a population with Mendelian inheritance. The frequencies of alleles (variations in a gene) will remain constant in the absence of selection, mutation, migration and genetic drift.<sup>[63]</sup>

Variation comes from mutations in genetic material, reshuffling of genes through sexual reproduction and migration between populations (gene flow). Despite the constant introduction of new variation through mutation and gene flow, most of the genome of a species is identical in all individuals of that species.<sup>[64]</sup> However, even relatively

small differences in genotype can lead to dramatic differences in phenotype: for example, chimpanzees and humans differ in only about 5% of their genomes.<sup>[65]</sup>

## Mutation

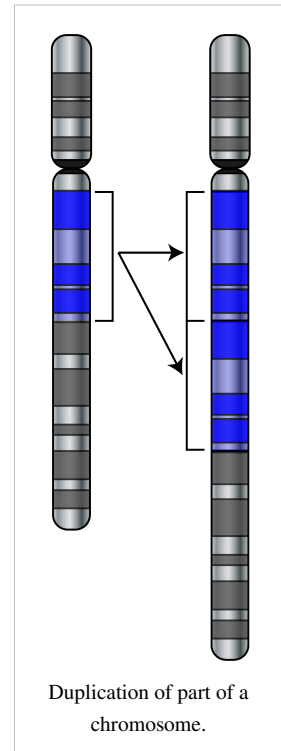
Further information: Mutation

Mutations are changes in the DNA sequence of a cell's genome. When mutations occur, they can either have no effect, alter the product of a gene, or prevent the gene from functioning. Based on studies in the fly *Drosophila melanogaster*, it has been suggested that if a mutation changes a protein produced by a gene, this will probably be harmful, with about 70% of these mutations having damaging effects, and the remainder being either neutral or weakly beneficial.<sup>[66]</sup>

Mutations can involve large sections of a chromosome becoming duplicated (usually by genetic recombination), which can introduce extra copies of a gene into a genome.<sup>[67]</sup> Extra copies of genes are a major source of the raw material needed for new genes to evolve.<sup>[68]</sup> This is important because most new genes evolve within gene families from pre-existing genes that share common ancestors.<sup>[69]</sup> For example, the human eye uses four genes to make structures that sense light: three for colour vision and one for night vision; all four are descended from a single ancestral gene.<sup>[70]</sup>

New genes can be generated from an ancestral gene when a duplicate copy mutates and acquires a new function. This process is easier once a gene has been duplicated because it increases the redundancy of the system; one gene in the pair can acquire a new function while the other copy continues to perform its original function.<sup>[71][72]</sup> Other types of mutations can even generate entirely new genes from previously noncoding DNA.<sup>[73][74]</sup>

The generation of new genes can also involve small parts of several genes being duplicated, with these fragments then recombining to form new combinations with new functions.<sup>[75][76]</sup> When new genes are assembled from shuffling pre-existing parts, domains act as modules with simple independent functions, which can be mixed together to produce new combinations with new and complex functions.<sup>[77]</sup> For example, polyketide synthases are large enzymes that make antibiotics; they contain up to one hundred independent domains that each catalyze one step in the overall process, like a step in an assembly line.<sup>[78]</sup>



## Sex and recombination

Further information: Sexual reproduction, Genetic recombination, and Evolution of sexual reproduction

In asexual organisms, genes are inherited together, or *linked*, as they cannot mix with genes of other organisms during reproduction. In contrast, the offspring of sexual organisms contain random mixtures of their parents' chromosomes that are produced through independent assortment. In a related process called homologous recombination, sexual organisms exchange DNA between two matching chromosomes.<sup>[79]</sup> Recombination and reassortment do not alter allele frequencies, but instead change which alleles are associated with each other, producing offspring with new combinations of alleles.<sup>[80]</sup> Sex usually increases genetic variation and may increase the rate of evolution.<sup>[81][82]</sup>

## Gene flow

Further information: Gene flow

Gene flow is the exchange of genes between populations and between species.<sup>[83]</sup> It can therefore be a source of variation that is new to a population or to a species. Gene flow can be caused by the movement of individuals between separate populations of organisms, as might be caused by the movement of mice between inland and coastal populations, or the movement of pollen between heavy metal tolerant and heavy metal sensitive populations of grasses.

Gene transfer between species includes the formation of hybrid organisms and horizontal gene transfer. Horizontal gene transfer is the transfer of genetic material from one organism to another organism that is not its offspring; this is most common among bacteria.<sup>[84]</sup> In medicine, this contributes to the spread of antibiotic resistance, as when one bacteria acquires resistance genes it can rapidly transfer them to other species.<sup>[85]</sup> Horizontal transfer of genes from bacteria to eukaryotes such as the yeast *Saccharomyces cerevisiae* and the adzuki bean beetle *Callosobruchus chinensis* has occurred.<sup>[86][87]</sup> An example of larger-scale transfers are the eukaryotic bdelloid rotifers, which have received a range of genes from bacteria, fungi and plants.<sup>[88]</sup> Viruses can also carry DNA between organisms, allowing transfer of genes even across biological domains.<sup>[89]</sup>

Large-scale gene transfer has also occurred between the ancestors of eukaryotic cells and bacteria, during the acquisition of chloroplasts and mitochondria. It is possible that eukaryotes themselves originated from horizontal gene transfers between bacteria and archaea.<sup>[90]</sup>

## Mechanisms

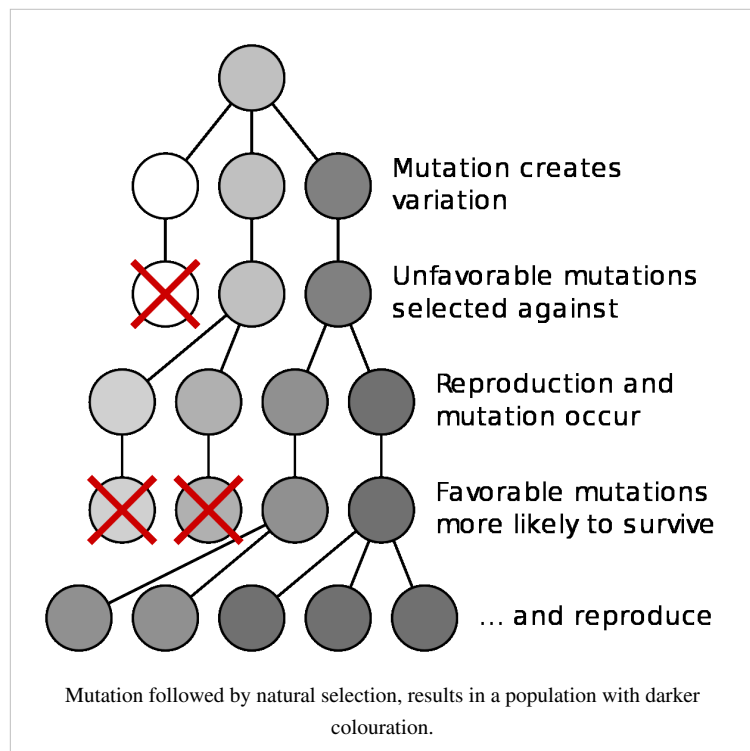
From a Neo-Darwinian perspective, evolution occurs when there are changes in the frequencies of alleles within a population of interbreeding organisms.<sup>[63]</sup> For example, the allele for black colour in a population of moths becoming more common. Mechanisms that can lead to changes in allele frequencies include natural selection, genetic drift, genetic hitchhiking, mutation and gene flow.

### Natural selection

Further information: Natural selection and Fitness (biology)

Evolution by means of natural selection is the process by which genetic mutations that enhance reproduction become and remain, more common in successive generations of a population. It has often been called a "self-evident" mechanism because it necessarily follows from three simple facts:

- Heritable variation exists within populations of organisms.
- Organisms produce more progeny than can survive.
- These offspring vary in their ability to survive and reproduce.

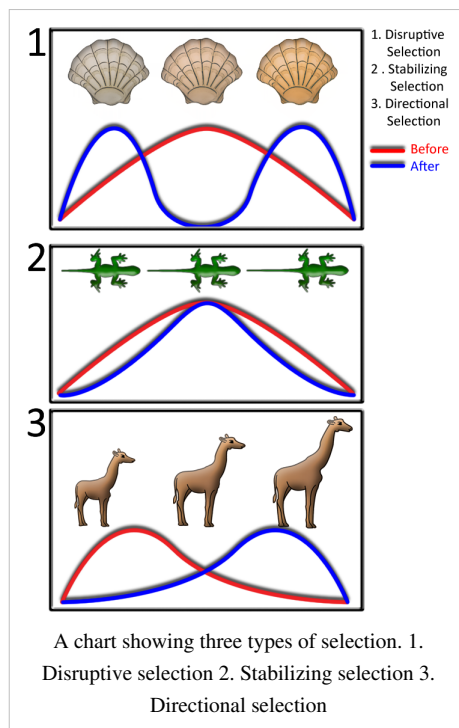




These conditions produce competition between organisms for survival and reproduction. Consequently, organisms with traits that give them an advantage over their competitors pass these advantageous traits on, while traits that do not confer an advantage are not passed on to the next generation.<sup>[91]</sup>

The central concept of natural selection is the evolutionary fitness of an organism.<sup>[92]</sup> Fitness is measured by an organism's ability to survive and reproduce, which determines the size of its genetic contribution to the next generation.<sup>[92]</sup> However, fitness is not the same as the total number of offspring; instead fitness is indicated by the proportion of subsequent generations that carry an organism's genes.<sup>[93]</sup> For example, if an organism could survive well and reproduce rapidly, but its offspring were all too small and weak to survive, this organism would make little genetic contribution to future generations and would thus have low fitness.<sup>[92]</sup>

If an allele increases fitness more than the other alleles of that gene, then with each generation this allele will become more common within the population. These traits are said to be "selected *for*". Examples of traits that can increase fitness are enhanced survival and increased fecundity. Conversely, the lower fitness caused by having a less beneficial or deleterious allele results in this allele becoming rarer — they are "selected *against*".<sup>[94]</sup> Importantly, the fitness of an allele is not a fixed characteristic; if the environment changes, previously neutral or harmful traits may become beneficial and previously beneficial traits become harmful.<sup>[52]</sup> However, even if the direction of selection does reverse in this way, traits that were lost in the past may not re-evolve in an identical form (see Dollo's law).<sup>[95][96]</sup>



Natural selection within a population for a trait that can vary across a range of values, such as height, can be categorised into three different types. The first is directional selection, which is a shift in the average value of a trait over time — for example, organisms slowly getting taller.<sup>[97]</sup> Secondly, disruptive selection is selection for extreme trait values and often results in two different values becoming most common, with selection against the average value. This would be when either short or tall organisms had an advantage, but not those of medium height. Finally, in stabilizing selection there is selection against extreme trait values on both ends, which causes a decrease in variance around the average value and less diversity.<sup>[91][98]</sup> This would, for example, cause organisms to slowly become all the same height.

A special case of natural selection is sexual selection, which is selection for any trait that increases mating success by increasing the attractiveness of an organism to potential mates.<sup>[99]</sup> Traits that evolved through sexual selection are particularly prominent in males of some animal species, despite traits such as cumbersome antlers, mating calls or bright colours that attract predators, decreasing the survival of

individual males.<sup>[100]</sup> This survival disadvantage is balanced by higher reproductive success in males that show these hard to fake, sexually selected traits.<sup>[101]</sup>

Natural selection most generally makes nature the measure against which individuals and individual traits, are more or less likely to survive. "Nature" in this sense refers to an ecosystem, that is, a system in which organisms interact with every other element, physical as well as biological, in their local environment. Eugene Odum, a founder of ecology, defined an ecosystem as: "Any unit that includes all of the organisms...in a given area interacting with the physical environment so that a flow of energy leads to clearly defined trophic structure, biotic diversity and material cycles (ie: exchange of materials between living and nonliving parts) within the system."<sup>[102]</sup> Each population within

an ecosystem occupies a distinct niche, or position, with distinct relationships to other parts of the system. These relationships involve the life history of the organism, its position in the food chain and its geographic range. This broad understanding of nature enables scientists to delineate specific forces which, together, comprise natural selection.

Natural selection can act at different levels of organisation, such as genes, cells, individual organisms, groups of organisms and species.<sup>[103][104][105]</sup> Selection can act at multiple levels simultaneously.<sup>[106]</sup> An example of selection occurring below the level of the individual organism are genes called transposons, which can replicate and spread throughout a genome.<sup>[107]</sup> Selection at a level above the individual, such as group selection, may allow the evolution of co-operation, as discussed below.<sup>[108]</sup>

## Biased mutation

In addition to being a major source of variation, mutation may also function as a mechanism of evolution when there are different probabilities at the molecular level for different mutations to occur, a process known as mutation bias.<sup>[109]</sup> If two genotypes, for example one with the nucleotide G and another with the nucleotide A in the same position, have the same fitness, but mutation from G to A happens more often than mutation from A to G, then genotypes with A will tend to evolve.<sup>[110]</sup> Different insertion vs. deletion mutation biases in different taxa can lead to the evolution of different genome sizes.<sup>[111][112]</sup> Developmental or mutational biases have also been observed in morphological evolution.<sup>[113][114]</sup> For example, according to the phenotype-first theory of evolution, mutations can eventually cause the genetic assimilation of traits that were previously induced by the environment.<sup>[115][116]</sup>

Mutation bias effects are superimposed on other processes. If selection would favor either one out of two mutations, but there is no extra advantage to having both, then the mutation that occurs the most frequently is the one that is most likely to become fixed in a population.<sup>[117][118]</sup> Mutations leading to the loss of function of a gene are much more common than mutations that produce a new, fully functional gene. Most loss of function mutations are selected against. But when selection is weak, mutation bias towards loss of function can affect evolution.<sup>[119]</sup> For example, pigments are no longer useful when animals live in the darkness of caves, and tend to be lost.<sup>[120]</sup> This kind of loss of function can occur because of mutation bias, and/or because the function had a cost, and once the benefit of the function disappeared, natural selection leads to the loss. Loss of sporulation ability in a bacterium during laboratory evolution appears to have been caused by mutation bias, rather than natural selection against the cost of maintaining sporulation ability.<sup>[121]</sup> When there is no selection for loss of function, the speed at which loss evolves depends more on the mutation rate than it does on the effective population size,<sup>[122]</sup> indicating that it is driven more by mutation bias than by genetic drift.

## Genetic drift

Further information: Genetic drift and Effective population size

Genetic drift is the change in allele frequency from one generation to the next that occurs because alleles are subject to sampling error.<sup>[123]</sup> As a result, when selective forces are absent or relatively weak, allele frequencies tend to "drift" upward or downward randomly (in a random walk). This drift halts when an allele eventually becomes fixed, either by disappearing from the population, or replacing the other alleles entirely. Genetic drift may therefore eliminate some alleles from a population due to chance alone. Even in the absence of selective forces, genetic drift can cause two separate populations that began with the same genetic structure to drift apart into two divergent populations with different sets of alleles.<sup>[124]</sup>

It is usually difficult to measure the relative importance of selection and neutral processes, including drift.<sup>[125]</sup> The comparative importance of adaptive and non-adaptive forces in driving evolutionary change is an area of current research.<sup>[126]</sup>

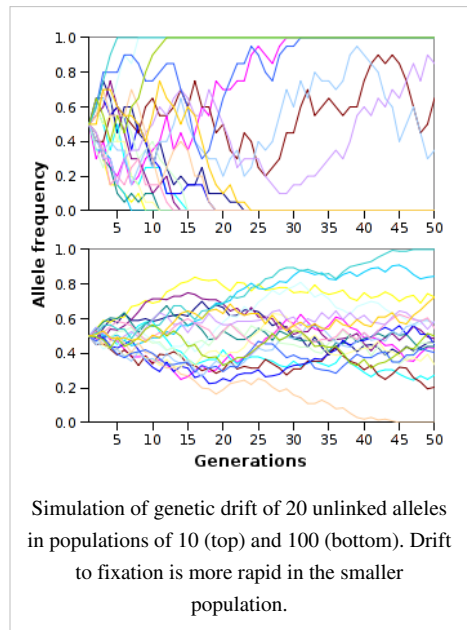
The neutral theory of molecular evolution proposed that most evolutionary changes are the result of the fixation of neutral mutations by genetic drift.<sup>[5]</sup> Hence, in this model, most genetic changes in a population are the result of constant mutation pressure and genetic drift.<sup>[127]</sup> This form of the neutral theory is now largely abandoned, since it does not seem to fit the genetic variation seen in nature.<sup>[128][129]</sup> However, a more recent and better-supported version of this model is the nearly neutral theory, where a mutation that would be neutral in a small population is not necessarily neutral in a large population.<sup>[91]</sup> Other alternative theories propose that genetic drift is dwarfed by other stochastic forces in evolution, such as genetic hitchhiking, also known as genetic draft.<sup>[123][130][131]</sup>

The time for a neutral allele to become fixed by genetic drift depends on population size, with fixation occurring more rapidly in smaller populations.<sup>[132]</sup> The number of individuals in a population is not critical, but instead a measure known as the effective population size.<sup>[133]</sup> The effective population is usually smaller than the total population since it takes into account factors such as the level of inbreeding and the stage of the lifecycle in which the population is the smallest.<sup>[133]</sup> The effective population size may not be the same for every gene in the same population.<sup>[134]</sup>

## Genetic hitchhiking

Further information: Genetic hitchhiking, Hill-Robertson effect, Selective sweep, and Genetic drift

Recombination allows alleles on the same strand of DNA to become separated. However, the rate of recombination is low (approximately two events per chromosome per generation). As a result, genes close together on a chromosome may not always be shuffled away from each other and genes that are close together tend to be inherited together, a phenomenon known as linkage.<sup>[135]</sup> This tendency is measured by finding how often two alleles occur together on a single chromosome compared to expectations, which is called their linkage disequilibrium. A set of alleles that is usually inherited in a group is called a haplotype. This can be important when one allele in a particular haplotype is strongly beneficial: natural selection can drive a selective sweep that will also cause the other alleles in the haplotype to become more common in the population; this effect is called genetic hitchhiking or genetic draft.<sup>[136]</sup> Genetic draft caused by the fact that some neutral genes are genetically linked to others that are under selection can be partially captured by an appropriate effective population size.<sup>[130]</sup>



## Gene flow

Further information: Gene flow, Hybrid (biology), and Horizontal gene transfer

Gene flow is the exchange of genes between populations and between species.<sup>[83]</sup> The presence or absence of gene flow fundamentally changes the course of evolution. Due to the complexity of organisms, any two completely isolated populations will eventually evolve genetic incompatibilities through neutral processes, as in the Bateson-Dobzhansky-Muller model, even if both populations remain essentially identical in terms of their adaptation to the environment.

If genetic differentiation between populations develops, gene flow between populations can introduce traits or alleles which are disadvantageous in the local population and this may lead to organism within these populations to evolve mechanisms that prevent mating with genetically distant populations, eventually resulting in the appearance of new species. Thus, exchange of genetic information between individuals is fundamentally important for the development of the biological species concept (BSC).

During the development of the modern synthesis, Sewall Wright's developed his shifting balance theory that gene flow between partially isolated populations was an important aspect of adaptive evolution.<sup>[137]</sup> However, recently there has been substantial criticism of the importance of the shifting balance theory.<sup>[138]</sup>

## Outcomes

Evolution influences every aspect of the form and behaviour of organisms. Most prominent are the specific behavioural and physical adaptations that are the outcome of natural selection. These adaptations increase fitness by aiding activities such as finding food, avoiding predators or attracting mates. Organisms can also respond to selection by co-operating with each other, usually by aiding their relatives or engaging in mutually beneficial symbiosis. In the longer term, evolution produces new species through splitting ancestral populations of organisms into new groups that cannot or will not interbreed.

These outcomes of evolution are sometimes divided into macroevolution, which is evolution that occurs at or above the level of species, such as extinction and speciation and microevolution, which is smaller evolutionary changes, such as adaptations, within a species or population.<sup>[139]</sup> In general, macroevolution is regarded as the outcome of long periods of microevolution.<sup>[140]</sup> Thus, the distinction between micro- and macroevolution is not a fundamental one – the difference is simply the time involved.<sup>[141]</sup> However, in macroevolution, the traits of the entire species may be important. For instance, a large amount of variation among individuals allows a species to rapidly adapt to new habitats, lessening the chance of it going extinct, while a wide geographic range increases the chance of speciation, by making it more likely that part of the population will become isolated. In this sense, microevolution and macroevolution might involve selection at different levels – with microevolution acting on genes and organisms, versus macroevolutionary processes such as species selection acting on entire species and affecting their rates of speciation and extinction.<sup>[142][143][144]</sup>

A common misconception is that evolution has goals or long-term plans; realistically however, evolution has no long-term goal and does not necessarily produce greater complexity.<sup>[145][146]</sup> Although complex species have evolved, they occur as a side effect of the overall number of organisms increasing and simple forms of life still remain more common in the biosphere.<sup>[147]</sup> For example, the overwhelming majority of species are microscopic prokaryotes, which form about half the world's biomass despite their small size,<sup>[148]</sup> and constitute the vast majority of Earth's biodiversity.<sup>[149]</sup> Simple organisms have therefore been the dominant form of life on Earth throughout its history and continue to be the main form of life up to the present day, with complex life only appearing more diverse because it is more noticeable.<sup>[150]</sup> Indeed, the evolution of microorganisms is particularly important to modern evolutionary research, since their rapid reproduction allows the study of experimental evolution and the observation of evolution and adaptation in real time.<sup>[151][152]</sup>

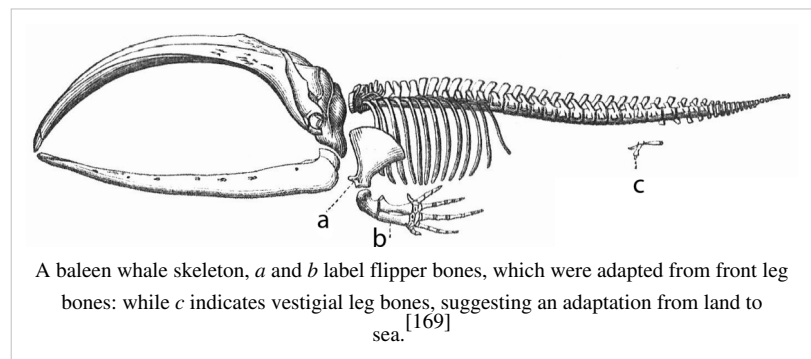
## Adaptation

Adaptation is the process that makes organisms better suited to their habitat.<sup>[153][154]</sup> Also, the term adaptation may refer to a trait that is important for an organism's survival. For example, the adaptation of horses' teeth to the grinding of grass. By using the term *adaptation* for the evolutionary process and *adaptive trait* for the product (the bodily part or function), the two senses of the word may be distinguished. Adaptations are produced by natural selection.<sup>[155]</sup> The following definitions are due to Theodosius Dobzhansky.

1. *Adaptation* is the evolutionary process whereby an organism becomes better able to live in its habitat or habitats.<sup>[156]</sup>
2. *Adaptedness* is the state of being adapted: the degree to which an organism is able to live and reproduce in a given set of habitats.<sup>[157]</sup>
3. An *adaptive trait* is an aspect of the developmental pattern of the organism which enables or enhances the probability of that organism surviving and reproducing.<sup>[158]</sup>

Adaptation may cause either the gain of a new feature, or the loss of an ancestral feature. An example that shows both types of change is bacterial adaptation to antibiotic selection, with genetic changes causing antibiotic resistance by both modifying the target of the drug, or increasing the activity of transporters that pump the drug out of the cell.<sup>[159]</sup> Other striking examples are the bacteria *Escherichia coli* evolving the ability to use citric acid as a nutrient in a long-term laboratory experiment,<sup>[160]</sup> *Flavobacterium* evolving a novel enzyme that allows these bacteria to grow on the by-products of nylon manufacturing,<sup>[161][162]</sup> and the soil bacterium *Sphingobium* evolving an entirely new metabolic pathway that degrades the synthetic pesticide pentachlorophenol.<sup>[163][164]</sup> An interesting but still controversial idea is that some adaptations might increase the ability of organisms to generate genetic diversity and adapt by natural selection (increasing organisms' evolvability).<sup>[165][166][167][168]</sup>

Adaptation occurs through the gradual modification of existing structures. Consequently, structures with similar internal organisation may have different functions in related organisms. This is the result of a single ancestral structure being adapted to function in different ways. The bones within bat wings, for example, are very similar to those in mice feet and



primate hands, due to the descent of all these structures from a common mammalian ancestor.<sup>[170]</sup> However, since all living organisms are related to some extent,<sup>[171]</sup> even organs that appear to have little or no structural similarity, such as arthropod, squid and vertebrate eyes, or the limbs and wings of arthropods and vertebrates, can depend on a common set of homologous genes that control their assembly and function; this is called deep homology.<sup>[172][173]</sup>

During evolution, some structures may lose their original function and become vestigial structures.<sup>[174]</sup> Such structures may have little or no function in a current species, yet have a clear function in ancestral species, or other closely related species. Examples include pseudogenes,<sup>[175]</sup> the non-functional remains of eyes in blind cave-dwelling fish,<sup>[176]</sup> wings in flightless birds,<sup>[177]</sup> and the presence of hip bones in whales and snakes.<sup>[169]</sup> Examples of vestigial structures in humans include wisdom teeth,<sup>[178]</sup> the coccyx,<sup>[174]</sup> the vermiform appendix,<sup>[174]</sup> and other behavioural vestiges such as goose bumps<sup>[179][180]</sup> and primitive reflexes.<sup>[181][182][183]</sup>

However, many traits that appear to be simple adaptations are in fact exaptations: structures originally adapted for one function, but which coincidentally became somewhat useful for some other function in the process.<sup>[184]</sup> One example is the African lizard *Holaspis guentheri*, which developed an extremely flat head for hiding in crevices, as can be seen by looking at its near relatives. However, in this species, the head has become so flattened that it assists in gliding from tree to tree—an exaptation.<sup>[184]</sup> Within cells, molecular machines such as the bacterial flagella<sup>[185]</sup>

and protein sorting machinery<sup>[186]</sup> evolved by the recruitment of several pre-existing proteins that previously had different functions.<sup>[139]</sup> Another example is the recruitment of enzymes from glycolysis and xenobiotic metabolism to serve as structural proteins called crystallins within the lenses of organisms' eyes.<sup>[187][188]</sup>

A critical principle of ecology is that of competitive exclusion: no two species can occupy the same niche in the same environment for a long time.<sup>[189]</sup> Consequently, natural selection will tend to force species to adapt to different ecological niches. This may mean that, for example, two species of cichlid fish adapt to live in different habitats, which will minimise the competition between them for food.<sup>[190]</sup>

An area of current investigation in evolutionary developmental biology is the developmental basis of adaptations and exaptations.<sup>[191]</sup> This research addresses the origin and evolution of embryonic development and how modifications of development and developmental processes produce novel features.<sup>[192]</sup> These studies have shown that evolution can alter development to produce new structures, such as embryonic bone structures that develop into the jaw in other animals instead forming part of the middle ear in mammals.<sup>[193]</sup> It is also possible for structures that have been lost in evolution to reappear due to changes in developmental genes, such as a mutation in chickens causing embryos to grow teeth similar to those of crocodiles.<sup>[194]</sup> It is now becoming clear that most alterations in the form of organisms are due to changes in a small set of conserved genes.<sup>[195]</sup>

## Co-evolution

Further information: Co-evolution

Interactions between organisms can produce both conflict and co-operation. When the interaction is between pairs of species, such as a pathogen and a host, or a predator and its prey, these species can develop matched sets of adaptations. Here, the evolution of one species causes adaptations in a second species. These changes in the second species then, in turn, cause new adaptations in the first species. This cycle of selection and response is called co-evolution.<sup>[196]</sup> An example is the production of tetrodotoxin in the rough-skinned newt and the evolution of tetrodotoxin resistance in its predator, the common garter snake. In this predator-prey pair, an evolutionary arms race has produced high levels of toxin in the newt and correspondingly high levels of toxin resistance in the snake.<sup>[197]</sup>



Common Garter Snake (*Thamnophis sirtalis*) which has evolved resistance to tetrodotoxin in its amphibian prey.

## Co-operation

Further information: Co-operation (evolution)

Not all co-evolved interactions between species involve conflict.<sup>[198]</sup> Many cases of mutually beneficial interactions have evolved. For instance, an extreme cooperation exists between plants and the mycorrhizal fungi that grow on their roots and aid the plant in absorbing nutrients from the soil.<sup>[199]</sup> This is a reciprocal relationship as the plants provide the fungi with sugars from photosynthesis. Here, the fungi actually grow inside plant cells, allowing them to exchange nutrients with their hosts, while sending signals that suppress the plant immune system.<sup>[200]</sup>

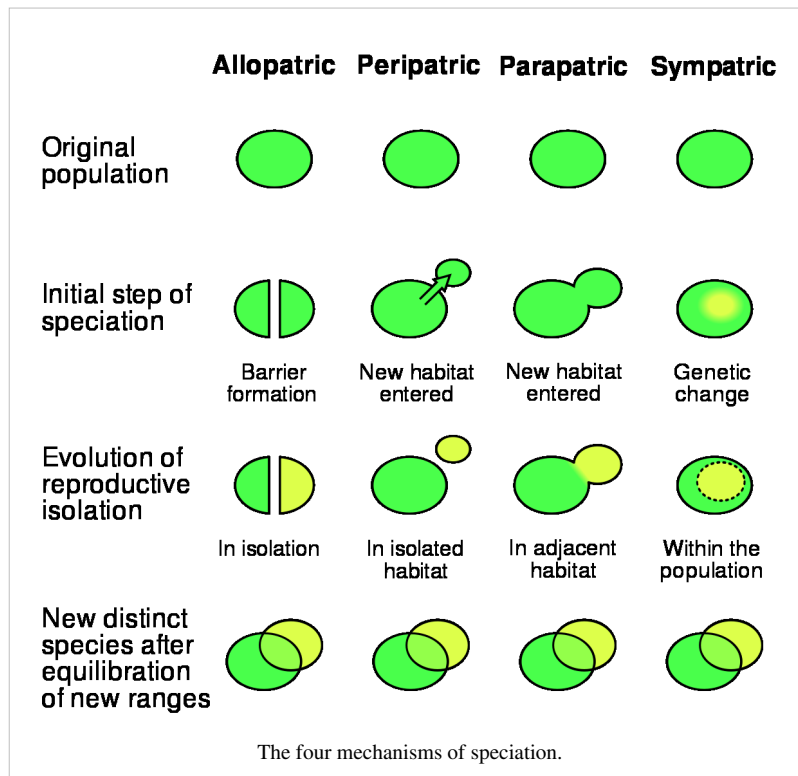
Coalitions between organisms of the same species have also evolved. An extreme case is the eusociality found in social insects, such as bees, termites and ants, where sterile insects feed and guard the small number of organisms in a colony that are able to reproduce. On an even smaller scale, the somatic cells that make up the body of an animal limit their reproduction so they can maintain a stable organism, which then supports a small number of the animal's germ cells to produce offspring. Here, somatic cells respond to specific signals that instruct them whether to grow, remain as they are, or die. If cells ignore these signals and multiply inappropriately, their uncontrolled growth causes cancer.<sup>[201]</sup>



Such cooperation within species may have evolved through the process of kin selection, which is where one organism acts to help raise a relative's offspring.<sup>[202]</sup> This activity is selected for because if the *helping* individual contains alleles which promote the helping activity, it is likely that its kin will *also* contain these alleles and thus those alleles will be passed on.<sup>[203]</sup> Other processes that may promote cooperation include group selection, where cooperation provides benefits to a group of organisms.<sup>[204]</sup>

## Speciation

Further information: Speciation



Speciation is the process where a species diverges into two or more descendant species.<sup>[205]</sup>

There are multiple ways to define the concept of "species". The choice of definition is dependent on the particularities of the species concerned.<sup>[206]</sup> For example, some species concepts apply more readily toward sexually reproducing organisms while others lend themselves better toward asexual organisms. Despite the diversity of various species concepts, these various concepts can be placed into one of three broad philosophical approaches: interbreeding, ecological and phylogenetic.<sup>[207]</sup> The biological species concept (BSC) is a classic example of the interbreeding approach.

Defined by Ernst Mayr in 1942, the

BSC states that "species are groups of actually or potentially interbreeding natural populations, which are reproductively isolated from other such groups"<sup>[208]:120</sup>. Despite its wide and long-term use, the BSC like others is not without controversy, for example because these concepts cannot be applied to prokaryotes,<sup>[209]</sup> and this is called the species problem.<sup>[206]</sup> Some researchers have attempted a unifying monistic definition of species, while others adopt a pluralistic approach and suggest that there may be a different ways to logically interpret the definition of a species.<sup>[206][207]</sup> "

Barriers to reproduction between two diverging sexual populations are required for the populations to become new species. Gene flow may slow this process by spreading the new genetic variants also to the other populations. Depending on how far two species have diverged since their most recent common ancestor, it may still be possible for them to produce offspring, as with horses and donkeys mating to produce mules.<sup>[210]</sup> Such hybrids are generally infertile. In this case, closely related species may regularly interbreed, but hybrids will be selected against and the species will remain distinct. However, viable hybrids are occasionally formed and these new species can either have properties intermediate between their parent species, or possess a totally new phenotype.<sup>[211]</sup> The importance of hybridisation in producing new species of animals is unclear, although cases have been seen in many types of animals,<sup>[212]</sup> with the gray tree frog being a particularly well-studied example.<sup>[213]</sup>

Speciation has been observed multiple times under both controlled laboratory conditions and in nature.<sup>[214]</sup> In sexually reproducing organisms, speciation results from reproductive isolation followed by genealogical divergence. There are four mechanisms for speciation. The most common in animals is allopatric speciation, which occurs in

populations initially isolated geographically, such as by habitat fragmentation or migration. Selection under these conditions can produce very rapid changes in the appearance and behaviour of organisms.<sup>[215][216]</sup> As selection and drift act independently on populations isolated from the rest of their species, separation may eventually produce organisms that cannot interbreed.<sup>[217]</sup>

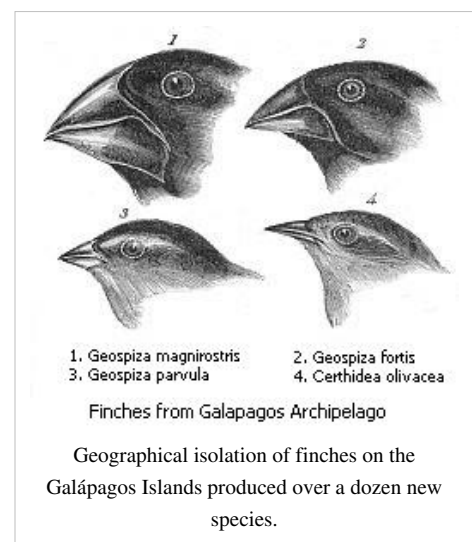
The second mechanism of speciation is peripatric speciation, which occurs when small populations of organisms become isolated in a new environment. This differs from allopatric speciation in that the isolated populations are numerically much smaller than the parental population. Here, the founder effect causes rapid speciation after an increase in inbreeding increases selection on homozygotes, leading to rapid genetic change.<sup>[218]</sup>

The third mechanism of speciation is parapatric speciation. This is similar to peripatric speciation in that a small population enters a new habitat, but differs in that there is no physical separation between these two populations. Instead, speciation results from the evolution of mechanisms that reduce gene flow between the two populations.<sup>[205]</sup> Generally this occurs when there has been a drastic change in the environment within the parental species' habitat. One example is the grass *Anthoxanthum odoratum*, which can undergo parapatric speciation in response to localised metal pollution from mines.<sup>[219]</sup> Here, plants evolve that have resistance to high levels of metals in the soil. Selection against interbreeding with the metal-sensitive parental population produced a gradual change in the flowering time of the metal-resistant plants, which eventually produced complete reproductive isolation. Selection against hybrids between the two populations may cause *reinforcement*, which is the evolution of traits that promote mating within a species, as well as character displacement, which is when two species become more distinct in appearance.<sup>[220]</sup>

Finally, in sympatric speciation species diverge without geographic isolation or changes in habitat. This form is rare since even a small amount of gene flow may remove genetic differences between parts of a population.<sup>[221]</sup> Generally, sympatric speciation in animals requires the evolution of both genetic differences and non-random mating, to allow reproductive isolation to evolve.<sup>[222]</sup>

One type of sympatric speciation involves cross-breeding of two related species to produce a new hybrid species. This is not common in animals as animal hybrids are usually sterile. This is because during meiosis the homologous chromosomes from each parent are from different species and cannot successfully pair. However, it is more common in plants because plants often double their number of chromosomes, to form polyploids.<sup>[223]</sup> This allows the chromosomes from each parental species to form matching pairs during meiosis, since each parent's chromosomes are represented by a pair already.<sup>[224]</sup> An example of such a speciation event is when the plant species *Arabidopsis thaliana* and *Arabidopsis arenosa* cross-bred to give the new species *Arabidopsis suecica*.<sup>[225]</sup> This happened about 20,000 years ago,<sup>[226]</sup> and the speciation process has been repeated in the laboratory, which allows the study of the genetic mechanisms involved in this process.<sup>[227]</sup> Indeed, chromosome doubling within a species may be a common cause of reproductive isolation, as half the doubled chromosomes will be unmatched when breeding with undoubled organisms.<sup>[228]</sup>

Speciation events are important in the theory of punctuated equilibrium, which accounts for the pattern in the fossil record of short "bursts" of evolution interspersed with relatively long periods of stasis, where species remain relatively unchanged.<sup>[229]</sup> In this theory, speciation and rapid evolution are linked, with natural selection and genetic drift acting most strongly on organisms undergoing speciation in novel habitats or small populations. As a result, the periods of stasis in the fossil record correspond to the parental population and the organisms undergoing speciation and rapid evolution are found in small populations or geographically restricted habitats and therefore rarely being preserved as fossils.<sup>[230]</sup>



## Extinction

Further information: Extinction



*Tyrannosaurus rex*. Non-avian dinosaurs died out in the Cretaceous–Paleogene extinction event at the end of the Cretaceous period.

Extinction is the disappearance of an entire species. Extinction is not an unusual event, as species regularly appear through speciation and disappear through extinction.<sup>[231]</sup> Nearly all animal and plant species that have lived on Earth are now extinct,<sup>[232]</sup> and extinction appears to be the ultimate fate of all species.<sup>[233]</sup> These extinctions have happened continuously throughout the history of life, although the rate of extinction spikes in occasional mass extinction events.<sup>[234]</sup> The Cretaceous–Paleogene extinction event, during which the non-avian dinosaurs went extinct, is the most well-known, but the earlier Permian–Triassic extinction event was even more severe, with approximately 96% of species driven to extinction.<sup>[234]</sup> The Holocene

extinction event is an ongoing mass extinction associated with humanity's expansion across the globe over the past few thousand years. Present-day extinction rates are 100–1000 times greater than the background rate and up to 30% of current species may be extinct by the mid 21st century.<sup>[235]</sup> Human activities are now the primary cause of the ongoing extinction event;<sup>[236]</sup> global warming may further accelerate it in the future.<sup>[237]</sup>

The role of extinction in evolution is not very well understood and may depend on which type of extinction is considered.<sup>[234]</sup> The causes of the continuous "low-level" extinction events, which form the majority of extinctions, may be the result of competition between species for limited resources (competitive exclusion).<sup>[45]</sup> If one species can out-compete another, this could produce species selection, with the fitter species surviving and the other species being driven to extinction.<sup>[104]</sup> The intermittent mass extinctions are also important, but instead of acting as a selective force, they drastically reduce diversity in a nonspecific manner and promote bursts of rapid evolution and speciation in survivors.<sup>[238]</sup>

## Evolutionary history of life

### Origin of life

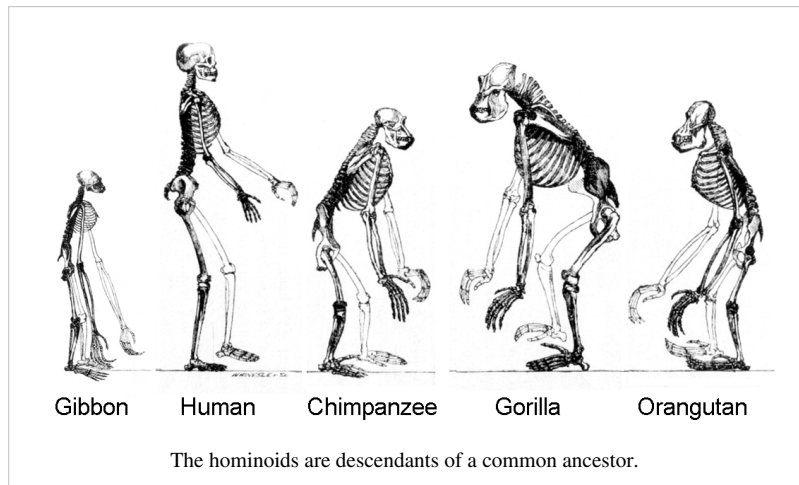
Further information: Abiogenesis and RNA world hypothesis

Highly energetic chemistry is thought to have produced a self-replicating molecule around 4 billion years ago and half a billion years later the last common ancestor of all life existed.<sup>[239]</sup> The current scientific consensus is that the complex biochemistry that makes up life came from simpler chemical reactions.<sup>[240]</sup> The beginning of life may have included self-replicating molecules such as RNA,<sup>[241]</sup> and the assembly of simple cells.<sup>[242]</sup>

## Common descent

Further information: Common descent and Evidence of common descent

All organisms on Earth are descended from a common ancestor or ancestral gene pool.<sup>[171][243]</sup> Current species are a stage in the process of evolution, with their diversity the product of a long series of speciation and extinction events.<sup>[244]</sup> The common descent of organisms was first deduced from four simple facts about organisms: First, they have geographic distributions that cannot be explained by local adaptation. Second, the diversity of life is not a set of completely unique organisms, but organisms that share morphological similarities. Third, vestigial traits with no clear purpose resemble functional ancestral traits and finally, that organisms can be classified using these similarities into a hierarchy of nested groups – similar to a family tree.<sup>[245]</sup> However, modern research has suggested that, due to horizontal gene transfer, this "tree of life" may be more complicated than a simple branching tree since some genes have spread independently between distantly related species.<sup>[246][247]</sup>

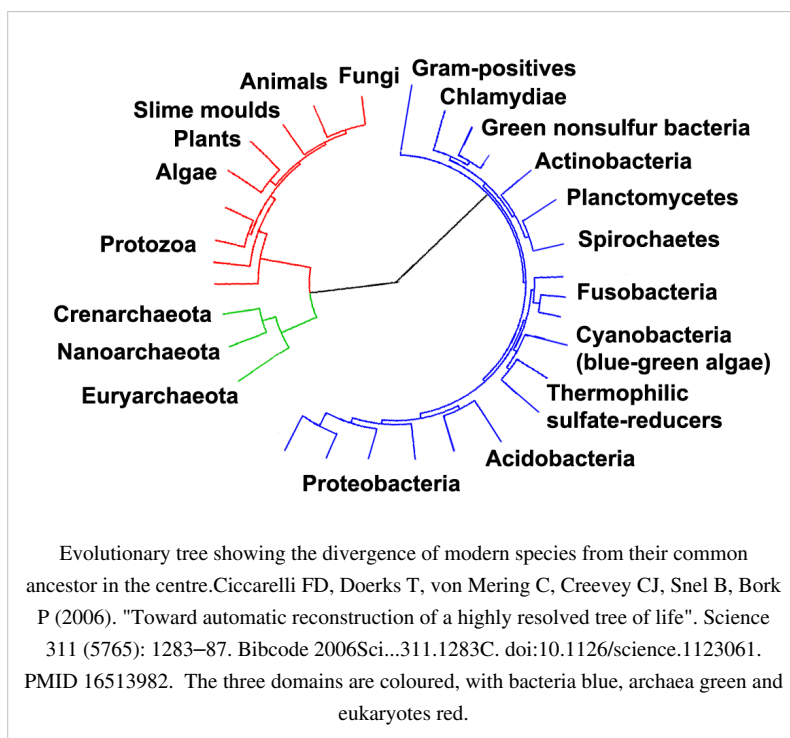


Past species have also left records of their evolutionary history. Fossils, along with the comparative anatomy of present-day organisms, constitute the morphological, or anatomical, record.<sup>[248]</sup> By comparing the anatomies of both modern and extinct species, paleontologists can infer the lineages of those species. However, this approach is most successful for organisms that had hard body parts, such as shells, bones or teeth. Further, as prokaryotes such as bacteria and archaea share a limited set of common morphologies, their fossils do not provide information on their ancestry.

More recently, evidence for common descent has come from the study of biochemical similarities between organisms. For example, all living cells use the same basic set of nucleotides and amino acids.<sup>[249]</sup> The development of molecular genetics has revealed the record of evolution left in organisms' genomes: dating when species diverged through the molecular clock produced by mutations.<sup>[250]</sup> For example, these DNA sequence comparisons have revealed that humans and chimpanzees share 96% of their genomes and analyzing the few areas where they differ helps shed light on when the common ancestor of these species existed.<sup>[251]</sup>

## Evolution of life

Prokaryotes inhabited the Earth from approximately 3–4 billion years ago.<sup>[252][253]</sup> No obvious changes in morphology or cellular organisation occurred in these organisms over the next few billion years.<sup>[254]</sup> The eukaryotic cells emerged between 1.6 – 2.7 billion years ago. The next major change in cell structure came when bacteria were engulfed by eukaryotic cells, in a cooperative association called endosymbiosis.<sup>[255][256]</sup> The engulfed bacteria and the host cell then underwent co-evolution, with the bacteria evolving into either mitochondria or hydrogenosomes.<sup>[257]</sup> Another engulfment of cyanobacterial-like organisms led to the formation of chloroplasts in algae and plants.<sup>[258]</sup>



The history of life was that of the unicellular eukaryotes, prokaryotes and archaea until about 610 million years ago when multicellular organisms began to appear in the oceans in the Ediacaran period.<sup>[252][259]</sup> The evolution of multicellularity occurred in multiple independent events, in organisms as diverse as sponges, brown algae, cyanobacteria, slime moulds and myxobacteria.<sup>[260]</sup>

Soon after the emergence of these first multicellular organisms, a remarkable amount of biological diversity appeared over approximately 10 million years, in an event called the Cambrian explosion. Here, the majority of types of modern animals appeared in the fossil record, as well as unique lineages that subsequently became extinct.<sup>[261]</sup> Various triggers for the Cambrian explosion have been proposed, including the accumulation of oxygen in the atmosphere from photosynthesis.<sup>[262]</sup>

About 500 million years ago, plants and fungi colonised the land and were soon followed by arthropods and other animals.<sup>[263]</sup> Insects were particularly successful and even today make up the majority of animal species.<sup>[264]</sup> Amphibians first appeared around 364 million years ago, followed by early amniotes and birds around 155 million years ago (both from "reptile"-like lineages), mammals around 129 million years ago, homininae around 10 million years ago and modern humans around 250,000 years ago.<sup>[265][266][267]</sup> However, despite the evolution of these large animals, smaller organisms similar to the types that evolved early in this process continue to be highly successful and dominate the Earth, with the majority of both biomass and species being prokaryotes.<sup>[149]</sup>

## Applications

Concepts and models used in evolutionary biology, such as natural selection, have many applications.<sup>[268]</sup>

Artificial selection is the intentional selection of traits in a population of organisms. This has been used for thousands of years in the domestication of plants and animals.<sup>[269]</sup> More recently, such selection has become a vital part of genetic engineering, with selectable markers such as antibiotic resistance genes being used to manipulate DNA. In repeated rounds of mutation and selection proteins with valuable properties have evolved, for example modified enzymes and new antibodies, in a process called directed evolution.<sup>[270]</sup>

Understanding the changes that have occurred during organism's evolution can reveal the genes needed to construct parts of the body, genes which may be involved in human genetic disorders.<sup>[271]</sup> For example, the mexican tetra is an albino cavefish that lost its eyesight during evolution. Breeding together different populations of this blind fish produced some offspring with functional eyes, since different mutations had occurred in the isolated populations that had evolved in different caves.<sup>[272]</sup> This helped identify genes required for vision and pigmentation.<sup>[273]</sup>

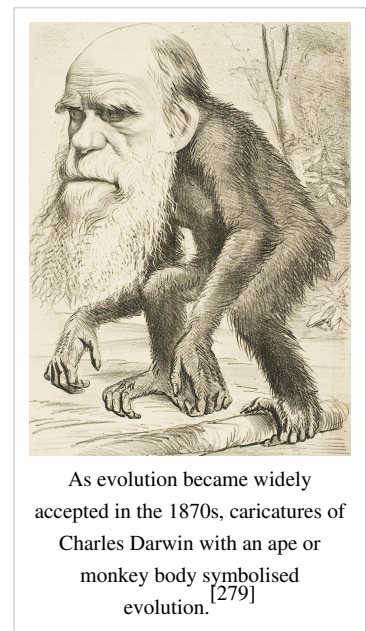
In computer science, simulations of evolution using evolutionary algorithms and artificial life started in the 1960s and was extended with simulation of artificial selection.<sup>[274]</sup> Artificial evolution became a widely recognised optimisation method as a result of the work of Ingo Rechenberg in the 1960s. He used evolution strategies to solve complex engineering problems.<sup>[275]</sup> Genetic algorithms in particular became popular through the writing of John Holland.<sup>[276]</sup> Practical applications also include automatic evolution of computer programs.<sup>[277]</sup> Evolutionary algorithms are now used to solve multi-dimensional problems more efficiently than software produced by human designers and also to optimise the design of systems.<sup>[278]</sup>

## Social and cultural responses

Further information: Social effect of evolutionary theory and Objections to evolution

In the 19th century, particularly after the publication of *On the Origin of Species* in 1859, the idea that life had evolved was an active source of academic debate centred on the philosophical, social and religious implications of evolution. Today, the modern evolutionary synthesis is accepted by a vast majority of scientists.<sup>[45]</sup> However, evolution remains a contentious concept for some theists.<sup>[280]</sup>

While various religions and denominations have reconciled their beliefs with evolution through concepts such as theistic evolution, there are creationists who believe that evolution is contradicted by the creation myths found in their religions and who raise various objections to evolution.<sup>[139][281][282]</sup> As had been demonstrated by responses to the publication of *Vestiges of the Natural History of Creation* in 1844, the most controversial aspect of evolutionary biology is the implication of human evolution that humans share common ancestry with apes and that the mental and moral faculties of humanity have the same types of natural causes as other inherited traits in animals.<sup>[283]</sup> In some countries, notably the United States, these tensions between science and religion have fuelled the current creation-evolution controversy, a religious conflict focusing on politics and public education.<sup>[284]</sup> While other scientific fields such as cosmology<sup>[285]</sup> and Earth science<sup>[286]</sup> also conflict with literal interpretations of many religious texts, evolutionary biology experiences significantly more opposition from religious literalists.



As evolution became widely accepted in the 1870s, caricatures of Charles Darwin with an ape or monkey body symbolised evolution.<sup>[279]</sup>

The teaching of evolution in American secondary school biology classes was uncommon in most of the first half of the 20th century. The Scopes Trial decision of 1925 caused the subject to become very rare in American secondary biology textbooks for a generation, but it was gradually re-introduced about a generation later and legally protected with the 1968 *Epperson v. Arkansas* decision. Since then, the competing religious belief of creationism was legally disallowed in secondary school curricula in various decisions in the 1970s and 1980s, but it returned in pseudoscientific form as intelligent design, to be excluded once again in the 2005 *Kitzmiller v. Dover Area School District* case.<sup>[287]</sup>



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### Introductory reading

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### General information

- Evolution (<http://www.bbc.co.uk/programmes/p00545gl>) on *In Our Time* at the BBC. ( listen now ([http://www.bbc.co.uk/iplayer/console/p00545gl/In\\_Our\\_Time\\_Evolution](http://www.bbc.co.uk/iplayer/console/p00545gl/In_Our_Time_Evolution)))
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# Microorganism

A **microorganism** (from the Greek: μικρός, *mikrós*, "small" and ὄργανισμός, *organismós*, "organism"; also spelled **micro-organism**, **micro organism** or **microörganism**) or **microbe** is a microscopic organism that comprises either a single cell (unicellular), cell clusters,<sup>[1]</sup> or multicellular relatively complex organisms. The study of microorganisms is called microbiology, a subject that began with Anton van Leeuwenhoek's discovery of microorganisms in 1675, using a microscope of his own design.

Microorganisms are very diverse; they include bacteria, fungi, algae, and protozoa; microscopic plants (green algae); and animals such as rotifers and planarians.

Some microbiologists also include viruses, but others consider these as nonliving.<sup>[2][3]</sup> Most microorganisms are unicellular (single-celled), but this is not universal, since some multicellular organisms are microscopic, while some unicellular protists and bacteria, like *Thiomargarita namibiensis*, are macroscopic and visible to the naked eye.<sup>[4]</sup>

Microorganisms live in all parts of the biosphere where there is liquid water, including soil, hot springs, on the ocean floor, high in the atmosphere and deep inside rocks within the Earth's crust. Microorganisms are critical to nutrient recycling in ecosystems as they act as decomposers. As some microorganisms can fix nitrogen, they are a vital part of the nitrogen cycle, and recent studies indicate that airborne microbes may play a role in precipitation and weather.<sup>[5]</sup>

Microbes are also exploited by people in biotechnology, both in traditional food and beverage preparation, and in modern technologies based on genetic engineering. However, pathogenic microbes are harmful, since they invade and grow within other organisms, causing diseases that kill humans, other animals and plants.<sup>[6]</sup>

## History

### Evolution

Single-celled microorganisms were the first forms of life to develop on Earth, approximately 3–4 billion years ago.<sup>[7][8][9]</sup> Further evolution was slow,<sup>[10]</sup> and for about 3 billion years in the Precambrian eon, all organisms were microscopic.<sup>[11]</sup> So, for most of the history of life on Earth the only forms of life were microorganisms.<sup>[12]</sup> Bacteria, algae and fungi have been identified in amber that is 220 million years old, which shows that the morphology of microorganisms has changed little since the Triassic period.<sup>[13]</sup>

Most microorganisms can reproduce rapidly, but slow when the environment is cold. And microbes such as bacteria can also freely exchange genes by conjugation, transformation and transduction between widely-divergent species.<sup>[14]</sup> This horizontal gene transfer, coupled with a high mutation rate and many other means of genetic variation, allows microorganisms to swiftly evolve (via natural selection) to survive in new environments and respond to environmental stresses. A nice "text book" example of this is the evolution of specialised Nylon-eating bacteria, it's also been studied in experimental evolution. This rapid evolution is important in medicine, as it has led to the recent development of 'super-bugs' — pathogenic bacteria that are resistant to modern antibiotics.<sup>[15]</sup>



A cluster of *Escherichia coli* Bacteria magnified 10,000 times.

## Pre-microbiology

The possibility that microorganisms exist was discussed for many centuries before their actual discovery in the 17th century. The existence of unseen microbiological life was postulated by Jainism, which is based on Mahavira's teachings as early as 6th century BCE.<sup>[16]</sup> Paul Dundas notes that Mahavira asserted existence of unseen microbiological creatures living in earth, water, air and fire.<sup>[17]</sup> Jain scriptures also describe nigodas, which are sub-microscopic creatures living in large clusters and having a very short life and are said to pervade each and every part of universe, even in tissues of plants and flesh of animals.<sup>[18]</sup> However, the earliest known idea to indicate the possibility of diseases spreading by yet unseen organisms was that of the Roman scholar Marcus Terentius Varro in a 1st century BC book titled *On Agriculture* in which he warns against locating a homestead near swamps:

... and because there are bred certain minute creatures that cannot be seen by the eyes, which float in the air and enter the body through the mouth and nose and there cause serious diseases.<sup>[19]</sup>

In *The Canon of Medicine* (1020), Abū Alī ibn Sīnā (Avicenna) hypothesized that tuberculosis and other diseases might be contagious<sup>[20][21]</sup>

In 1546, Girolamo Fracastoro proposed that epidemic diseases were caused by transferable seedlike entities that could transmit infection by direct or indirect contact, or even without contact over long distances.

All these early claims about the existence of microorganisms were speculative and were not based on any data or science. Microorganisms were neither proven, observed, nor correctly and accurately described until the 17th century. The reason for this was that all these early studies lacked the microscope.

## History of microorganisms' discovery



Antonie van Leeuwenhoek, the first microbiologist and the first to observe microorganisms using a microscope

Antonie Van Leeuwenhoek (1632–1723) was one of the first people to observe microorganisms, using a microscope of his own design, and made one of the most important contributions to biology.<sup>[22]</sup> Robert Hooke was the first to use a microscope to observe living things; his 1665 book *Micrographia* contained descriptions of plant cells.

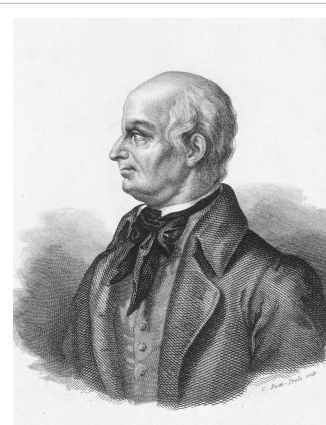
Before Leeuwenhoek's discovery of microorganisms in 1675, it had been a mystery why grapes could be turned into wine, milk into cheese, or why food would spoil. Leeuwenhoek did not make the connection between these processes and microorganisms, but using a microscope, he did establish that there were forms of life that were not visible to the naked eye.<sup>[23][24]</sup> Leeuwenhoek's discovery, along with subsequent observations by Spallanzani and Pasteur, ended the long-held belief that life spontaneously appeared from non-living substances during the process of spoilage.

Lazzaro Spallanzani (1729–1799) found that boiling broth would sterilise it and kill any microorganisms in it. He also found that new microorganisms could settle only in a broth if the broth was exposed to the air.

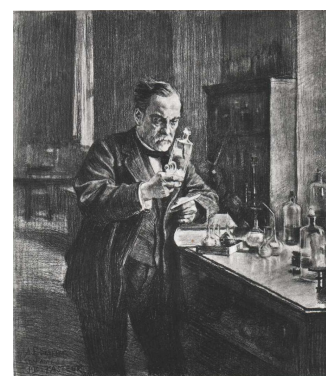
Louis Pasteur (1822–1895) expanded upon Spallanzani's findings by exposing boiled broths to the air, in vessels that contained a filter to prevent all particles from passing through to the growth medium, and also in vessels with no filter at all, with air being admitted via a curved tube that would not allow dust particles to come in contact with the broth. By boiling the broth beforehand, Pasteur ensured that no microorganisms survived within the broths at the beginning of his experiment. Nothing grew in the broths in the course of Pasteur's experiment.

This meant that the living organisms that grew in such broths came from outside, as spores on dust, rather than spontaneously generated within the broth. Thus, Pasteur dealt the death blow to the theory of spontaneous generation and supported germ theory.

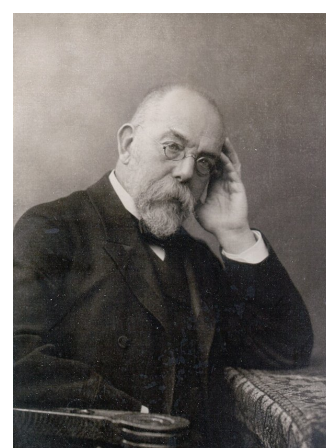
In 1876, Robert Koch (1843–1910) established that microbes can cause disease. He found that the blood of cattle who were infected with anthrax always had large numbers of *Bacillus anthracis*. Koch found that he could transmit anthrax from one animal to another by taking a small sample of blood from the infected animal and injecting it into a healthy one, and this caused the healthy animal to become sick. He also found that he could grow the bacteria in a nutrient broth, then inject it into a healthy animal, and cause illness. Based on these experiments, he devised criteria for establishing a causal link between a microbe and a disease and these are now known as Koch's postulates.<sup>[25]</sup> Although these postulates cannot be applied in all cases, they do retain historical importance to the development of scientific thought and are still being used today.<sup>[26]</sup>



Lazzaro Spallanzani showed that boiling a broth stopped it from decaying



Louis Pasteur showed that Spallanzani's findings held even if air could enter through a filter that kept particles out



Robert Koch showed that microorganisms caused disease

## Classification and structure

Microorganisms can be found almost anywhere in the taxonomic organization of life on the planet. Bacteria and archaea are almost always microscopic, while a number of eukaryotes are also microscopic, including most protists, some fungi, as well as some animals and plants. Viruses are generally regarded as not living and therefore are not microbes, although the field of microbiology also encompasses the study of viruses.

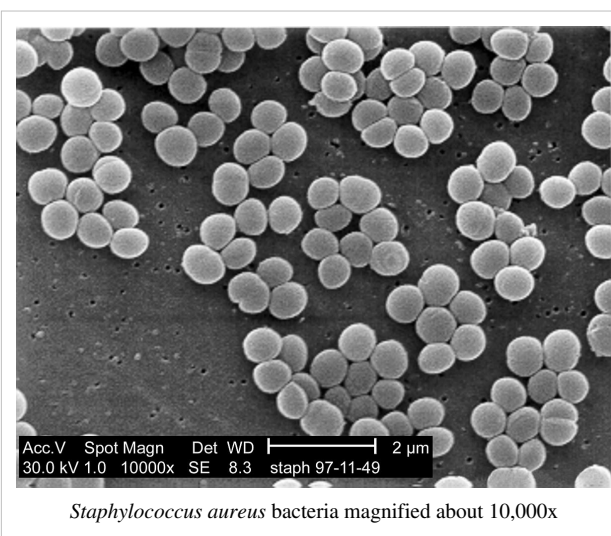
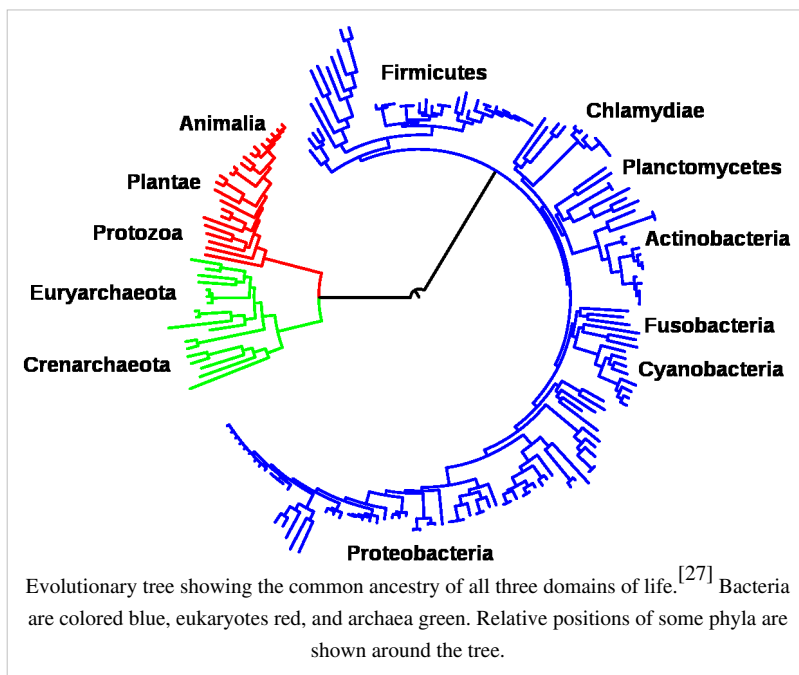
## Prokaryotes

Prokaryotes are organisms that lack a cell nucleus and the other membrane bound organelles. They are almost always unicellular, although some species such as myxobacteria can aggregate into complex structures as part of their life cycle.

Consisting of two domains, bacteria and archaea, the prokaryotes are the most diverse and abundant group of organisms on Earth and inhabit practically all environments where some liquid water is available and the temperature is below +140 °C. They are found in sea water, soil, air, animals' gastrointestinal tracts, hot springs and even deep beneath the Earth's crust in rocks.<sup>[28]</sup> Practically all surfaces that have not been specially sterilized are covered by prokaryotes. The number of prokaryotes on Earth is estimated to be around five million trillion trillion, or  $5 \times 10^{30}$ , accounting for at least half the biomass on Earth.<sup>[29]</sup>

## Bacteria

Almost all bacteria are invisible to the naked eye, with a few extremely rare exceptions, such as *Thiomargarita namibiensis*.<sup>[30]</sup> They lack membrane-bound organelles, and can function and reproduce as individual cells, but often aggregate in multicellular colonies.<sup>[31]</sup> Their genome is usually a single loop of DNA, although they can also harbor small pieces of DNA called plasmids. These plasmids can be transferred between cells through bacterial conjugation. Bacteria are surrounded by a cell wall, which provides strength and rigidity to their cells. They reproduce by binary fission or sometimes by budding, but do not undergo sexual reproduction. Some species form extraordinarily resilient spores, but for bacteria this is a mechanism for survival, not reproduction. Under optimal conditions bacteria can grow extremely rapidly and can double as quickly as every 10 minutes.<sup>[32]</sup>



## Archaea

Archaea are also single-celled organisms that lack nuclei. In the past, the differences between bacteria and archaea were not recognised and archaea were classified with bacteria as part of the kingdom Monera. However, in 1990 the microbiologist Carl Woese proposed the three-domain system that divided living things into bacteria, archaea and eukaryotes.<sup>[33]</sup> Archaea differ from bacteria in both their genetics and biochemistry. For example, while bacterial cell membranes are made from phosphoglycerides with ester bonds, archaean membranes are made of ether lipids.<sup>[34]</sup>

Archaea were originally described in extreme environments, such as hot springs, but have since been found in all types of habitats.<sup>[35]</sup> Only now are scientists beginning to realize how common archaea are in the environment, with crenarchaeota being the most common form of life in the ocean, dominating ecosystems below 150 m in depth.<sup>[36][37]</sup> These organisms are also common in soil and play a vital role in ammonia oxidation.<sup>[38]</sup>

## Eukaryotes

Most living things that are visible to the naked eye in their adult form are eukaryotes, including humans. However, a large number of eukaryotes are also microorganisms. Unlike bacteria and archaea, eukaryotes contain organelles such as the cell nucleus, the Golgi apparatus and mitochondria in their cells. The nucleus is an organelle that houses the DNA that makes up a cell's genome. DNA itself is arranged in complex chromosomes.<sup>[39]</sup> Mitochondria are organelles vital in metabolism as they are the site of the citric acid cycle and oxidative phosphorylation. They evolved from symbiotic bacteria and retain a remnant genome.<sup>[40]</sup> Like bacteria, plant cells have cell walls, and contain organelles such as chloroplasts in addition to the organelles in other eukaryotes. Chloroplasts produce energy from light by photosynthesis, and were also originally symbiotic bacteria.<sup>[40]</sup>

Unicellular eukaryotes are those eukaryotic organisms that consist of a single cell throughout their life cycle. This qualification is significant since most multicellular eukaryotes consist of a single cell called a zygote at the beginning of their life cycles. Microbial eukaryotes can be either haploid or diploid, and some organisms have multiple cell nuclei (see coenocyte). However, not all microorganisms are unicellular as some microscopic eukaryotes are made from multiple cells.

## Protists

Of eukaryotic groups, the protists are most commonly unicellular and microscopic. This is a highly diverse group of organisms that are not easy to classify.<sup>[41][42]</sup> Several algae species are multicellular protists, and slime molds have unique life cycles that involve switching between unicellular, colonial, and multicellular forms.<sup>[43]</sup> The number of species of protozoa is uncertain, since we may have identified only a small proportion of the diversity in this group of organisms.<sup>[44][45]</sup>

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## Animals

Most animals are multicellular,<sup>[46]</sup> but some are too small to be seen by the naked eye. Microscopic arthropods include dust mites and spider mites. Microscopic crustaceans include copepods and the cladocera, while many nematodes are too small to be seen with the naked eye. Another particularly common group of microscopic animals are the rotifers, which are filter feeders that are usually found in fresh water. Micro-animals reproduce both sexually and asexually and may reach new habitats as eggs that survive harsh environments that would kill the adult animal. However, some simple animals, such as rotifers and nematodes, can dry out completely and remain dormant for long periods of time.<sup>[47]</sup>

## Fungi

The fungi have several unicellular species, such as baker's yeast (*Saccharomyces cerevisiae*) and fission yeast (*Schizosaccharomyces pombe*). Some fungi, such as the pathogenic yeast *Candida albicans*, can undergo phenotypic switching and grow as single cells in some environments, and filamentous hyphae in others.<sup>[48]</sup> Fungi reproduce both asexually, by budding or binary fission, as well by producing spores, which are called conidia when produced asexually, or basidiospores when produced sexually.

## Plants

The green algae are a large group of photosynthetic eukaryotes that include many microscopic organisms. Although some green algae are classified as protists, others such as charophyta are classified with embryophyte plants, which are the most familiar group of land plants. Algae can grow as single cells, or in long chains of cells. The green algae include unicellular and colonial flagellates, usually but not always with two flagella per cell, as well as various colonial, coccoid, and filamentous forms. In the Charales, which are the algae most closely related to higher plants, cells differentiate into several distinct tissues within the organism. There are about 6000 species of green algae.<sup>[49]</sup>

## Habitats and ecology

Microorganisms are found in almost every habitat present in nature. Even in hostile environments such as the poles, deserts, geysers, rocks, and the deep sea. Some types of microorganisms have adapted to the extreme conditions and sustained colonies; these organisms are known as extremophiles. Extremophiles have been isolated from rocks as much as 7 kilometres below the Earth's surface,<sup>[50]</sup> and it has been suggested that the amount of living organisms below the Earth's surface may be comparable with the amount of life on or above the surface.<sup>[28]</sup> Extremophiles have been known to survive for a prolonged time in a vacuum, and can be highly resistant to radiation, which may even allow them to survive in space.<sup>[51]</sup> Many types of microorganisms have intimate symbiotic relationships with other larger organisms; some of which are mutually beneficial (mutualism), while others can be damaging to the host organism (parasitism). If microorganisms can cause disease in a host they are known as pathogens.



A microscopic mite *Lorryia formosa*.



## Extremophiles

Extremophiles are microorganisms that have adapted so that they can survive and even thrive in conditions that are normally fatal to most life-forms. For example, some species have been found in the following extreme environments:

- Temperature: as high as 130 °C (**unknown operator: u'strong' °F**),<sup>[52]</sup> as low as −17 °C (**unknown operator: u'strong' °F**)<sup>[53]</sup>
- Acidity/alkalinity: less than pH 0,<sup>[54]</sup> up to pH 11.5<sup>[55]</sup>
- Salinity: up to saturation<sup>[56]</sup>
- Pressure: up to 1,000–2,000 atm, down to 0 atm (e.g. vacuum of space)<sup>[57]</sup>
- Radiation: up to 5kGy<sup>[58]</sup>

Extremophiles are significant in different ways. They extend terrestrial life into much of the Earth's hydrosphere, crust and atmosphere, their specific evolutionary adaptation mechanisms to their extreme environment can be exploited in bio-technology, and their very existence under such extreme conditions increases the potential for extraterrestrial life.<sup>[59]</sup>

## Soil microbes

The nitrogen cycle in soils depends on the fixation of atmospheric nitrogen. One way this can occur is in the nodules in the roots of legumes that contain symbiotic bacteria of the genera *Rhizobium*, *Mesorhizobium*, *Sinorhizobium*, *Bradyrhizobium*, and *Azorhizobium*.<sup>[60]</sup>

## Symbiotic microbes

Symbiotic microbes such as fungi and algae form an association in lichen. Certain fungi form mycorrhizal symbioses with trees that increase the supply of nutrients to the tree.

## Importance

Microorganisms are vital to humans and the environment, as they participate in the Earth's element cycles such as the carbon cycle and nitrogen cycle, as well as fulfilling other vital roles in virtually all ecosystems, such as recycling other organisms' dead remains and waste products through decomposition. Microbes also have an important place in most higher-order multicellular organisms as symbionts. Many blame the failure of Biosphere 2 on an improper balance of microbes.<sup>[61]</sup>

## Use in food

Microorganisms are used in brewing, winemaking, baking, pickling and other food-making processes.

They are also used to control the fermentation process in the production of cultured dairy products such as yogurt and cheese. The cultures also provide flavour and aroma, and inhibit undesirable organisms.<sup>[62]</sup>

## Use in water treatment

Specially-cultured microbes are used in the biological treatment of sewage and industrial waste effluent, a process known as bioaugmentation.<sup>[63]</sup>

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## Use in energy

Microbes are used in fermentation to produce ethanol,<sup>[64]</sup> and in biogas reactors to produce methane.<sup>[65]</sup> Scientists are researching the use of algae to produce liquid fuels,<sup>[66]</sup> and bacteria to convert various forms of agricultural and urban waste into usable fuels.<sup>[67]</sup>

## Use in production of chemicals, enzymes etc.

Many microbes are used for commercial and industrial production of chemicals, enzymes and other bioactive molecules.

Examples of organic acid produced include

- **Acetic acid:** Produced by the bacterium *Acetobacter aceti* and other acetic acid bacteria (AAB)
- **Butyric acid** (butanoic acid): Produced by the bacterium *Clostridium butyricum*
- **Lactic acid:** *Lactobacillus* and others commonly called as lactic acid bacteria (LAB)
- **Citric acid:** Produced by the fungus *Aspergillus niger*

Microbes are used for preparation of bioactive molecules and enzymes.

- Streptokinase produced by the bacterium *Streptococcus* and modified by genetic engineering is used as a clot buster for removing clots from the blood vessels of patients who have undergone myocardial infarctions leading to heart attack.
- Cyclosporin A is a bioactive molecule used as an immunosuppressive agent in organ transplantation
- Stains produced by the yeast *Monascus purpureus* is commercialised as blood cholesterol lowering agents which acts by competitively inhibiting the enzyme responsible for synthesis of cholesterol.<sup>[68]</sup>

## Use in science

Microbes are also essential tools in biotechnology, biochemistry, genetics, and molecular biology. The yeasts (*Saccharomyces cerevisiae*) and fission yeast (*Schizosaccharomyces pombe*) are important model organisms in science, since they are simple eukaryotes that can be grown rapidly in large numbers and are easily manipulated.<sup>[69]</sup> They are particularly valuable in genetics, genomics and proteomics.<sup>[70][71]</sup> Microbes can be harnessed for uses such as creating steroids and treating skin diseases. Scientists are also considering using microbes for living fuel cells,<sup>[72]</sup> and as a solution for pollution.<sup>[73]</sup>

## Use in warfare

In the Middle Ages, diseased corpses were thrown into castles during sieges using catapults or other siege engines. Individuals near the corpses were exposed to the deadly pathogen and were likely to spread that pathogen to others.<sup>[74]</sup>

## Importance in human health

### Human digestion

Further information: Human flora#Human bacterial flora and human health

Microorganisms can form an endosymbiotic relationship with other, larger organisms. For example, the bacteria that live within the human digestive system contribute to gut immunity, synthesise vitamins such as folic acid and biotin, and ferment complex indigestible carbohydrates.<sup>[75]</sup>

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## Diseases and immunology

Microorganisms are the cause of many infectious diseases. The organisms involved include pathogenic bacteria, causing diseases such as plague, tuberculosis and anthrax; protozoa, causing diseases such as malaria, sleeping sickness and toxoplasmosis; and also fungi causing diseases such as ringworm, candidiasis or histoplasmosis. However, other diseases such as influenza, yellow fever or AIDS are caused by pathogenic viruses, which are not usually classified as living organisms and are not, therefore, microorganisms by the strict definition. As of 2007, no clear examples of archaean pathogens are known,<sup>[76]</sup> although a relationship has been proposed between the presence of some methanogens and human periodontal disease.<sup>[77]</sup>

## Importance in ecology

Further information: Decomposition

Microbes are critical to the processes of decomposition required to cycle nitrogen and other elements back to the natural world.

## Hygiene

Hygiene is the avoidance of infection or food spoiling by eliminating microorganisms from the surroundings. As microorganisms, in particular bacteria, are found virtually everywhere, the levels of harmful microorganisms can be reduced to acceptable levels. However, in some cases, it is required that an object or substance be completely sterile, i.e. devoid of all living entities and viruses. A good example of this is a hypodermic needle.

In food preparation microorganisms are reduced by preservation methods (such as the addition of vinegar), clean utensils used in preparation, short storage periods, or by cool temperatures. If complete sterility is needed, the two most common methods are irradiation and the use of an autoclave, which resembles a pressure cooker.

There are several methods for investigating the level of hygiene in a sample of food, drinking water, equipment, etc. Water samples can be filtrated through an extremely fine filter. This filter is then placed in a nutrient medium. Microorganisms on the filter then grow to form a visible colony. Harmful microorganisms can be detected in food by placing a sample in a nutrient broth designed to enrich the organisms in question. Various methods, such as selective media or PCR, can then be used for detection. The hygiene of hard surfaces, such as cooking pots, can be tested by touching them with a solid piece of nutrient medium and then allowing the microorganisms to grow on it.

There are no conditions where all microorganisms would grow, and therefore often several different methods are needed. For example, a food sample might be analyzed on three different nutrient mediums designed to indicate the presence of "total" bacteria (conditions where many, but not all, bacteria grow), molds (conditions where the growth of bacteria is prevented by, e.g., antibiotics) and coliform bacteria (these indicate a sewage contamination).

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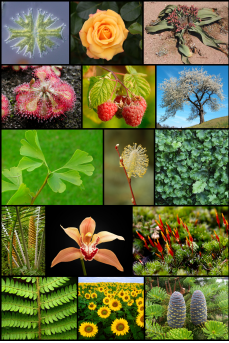
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## External links

- Our Microbial Planet (<http://dels.nas.edu/metagenomics>) A free poster from the National Academy of Sciences about the positive roles of microbes.
- "Uncharted Microbial World: Microbes and Their Activities in the Environment" ([http://www.asm.org/ASM/files/ccLibraryFiles/File/000000003691/Uncharted\\_Microbial\\_World.pdf](http://www.asm.org/ASM/files/ccLibraryFiles/File/000000003691/Uncharted_Microbial_World.pdf)) Report from the American Academy of Microbiology
- Understanding Our Microbial Planet: The New Science of Metagenomics ([http://dels.nas.edu/dels/rpt\\_briefs/metagenomics\\_final.pdf](http://dels.nas.edu/dels/rpt_briefs/metagenomics_final.pdf)) A 20-page educational booklet providing a basic overview of metagenomics and our microbial planet.
- Tree of Life Eukaryotes (<http://tolweb.org/Eukaryotes/3>)
- Microbe News from Genome News Network (<http://www.genomenewsnetwork.org/categories/index/microbes.php>)
- Microbes Patent List (<http://www.professorpatents.com/microbes.htm>) Microbes Related Patents
- Medical Microbiology (<http://gsbs.utmb.edu/microbook/toc.htm>) On-line textbook
- Through the microscope: A look at all things small ([http://www.microbiologytext.com/index.php?module=Book&func=toc&book\\_id=4](http://www.microbiologytext.com/index.php?module=Book&func=toc&book_id=4)) On-line microbiology textbook by Timothy Paustian and Gary Roberts, University of Wisconsin-Madison
- MicrobeID.com (<http://www.microbeid.com/>) Online Bacteria Identification Key and Probabilistic Identification Databases



# Plant

Plants	
Temporal range: Early Cambrian to recent, but see text,	
	
Scientific classification	
Domain:	Eukaryota
(unranked):	Archaeplastida
Kingdom:	<b>Plantae</b> Haeckel, 1866 <sup>[1]</sup>
Divisions	
<b>Green algae</b> <ul style="list-style-type: none"><li>• Chlorophyta</li><li>• Charophyta</li></ul>	
<b>Land plants (embryophytes)</b> <ul style="list-style-type: none"><li>• <b>Non-vascular land plants (bryophytes)</b><ul style="list-style-type: none"><li>• Marchantiophyta—liverworts</li><li>• Anthocerotophyta—hornworts</li><li>• Bryophyta—mosses</li><li>• †Horneophytopsida</li></ul></li><li>• <b>Vascular plants (tracheophytes)</b><ul style="list-style-type: none"><li>• †Rhyniophyta—rhyniophytes</li><li>• †Zosterophyllophyta—zosterophylls</li><li>• Lycopodiophyta—clubmosses</li><li>• †Trimerophytophyta—trimerophytes</li><li>• Pteridophyta—ferns and horsetails</li><li>• †Progymnospermophyta</li><li>• <b>Seed plants (spermatophytes)</b><ul style="list-style-type: none"><li>• †Pteridospermatophyta—seed ferns</li><li>• Pinophyta—conifers</li><li>• Cycadophyta—cycads</li><li>• Ginkgophyta—ginkgo</li><li>• Gnetophyta—gnetae</li><li>• Magnoliophyta—flowering plants</li></ul></li></ul></li></ul>	
†Nematophytes	

**Plants**, also called **green plants** (**Viridiplantae** in Latin), are living organisms of the kingdom **Plantae** including such multicellular groups as flowering plants, conifers, ferns and mosses, as well as, depending on definition, the green algae, but not red or brown seaweeds like kelp, nor fungi or bacteria.

Green plants have cell walls with cellulose and characteristically obtain most of their energy from sunlight via photosynthesis using chlorophyll contained in chloroplasts, which gives them their green color. Some plants are parasitic and may not produce normal amounts of chlorophyll or photosynthesize. Plants are also characterized by sexual reproduction, modular and indeterminate growth, and an alteration of generations, although asexual reproduction is common, and some plants bloom only once while others bear only one bloom.

Precise numbers are difficult to determine, but as of 2010, there are thought to be 300–315 thousand species of plants, of which the great majority, some 260–290 thousand, are seed plants (see the table below).<sup>[2]</sup> Green plants provide most of the world's free oxygen and are the basis of most of the earth's ecologies, especially on land. Plants described as grains, fruits and vegetables form mankind's basic foodstuffs, and have been domesticated for millennia. Plants enrich our lives as flowers and ornaments. Until recently and in great variety they have served as the source of most of our medicines and drugs. Their scientific study is known as botany.

## Definition

Plants are one of the two groups into which all living things have been traditionally divided; the other is animals. The division goes back at least as far as Aristotle (384 BC – 322 BC) who distinguished between plants which generally do not move, and animals which often are mobile to catch their food. Much later, when Linnaeus (1707–1778) created the basis of the modern system of scientific classification, these two groups became the kingdoms Vegetabilia (later Metaphyta or Plantae) and Animalia (also called Metazoa). Since then, it has become clear that the plant kingdom as originally defined included several unrelated groups, and the fungi and several groups of algae were removed to new kingdoms. However, these organisms are still often considered plants, particularly in popular contexts.

Outside of formal scientific contexts, the term "plant" implies an association with certain traits, such as being multicellular, possessing cellulose, and having the ability to carry out photosynthesis.<sup>[3][4]</sup>

## Current definitions of Plantae

When the name Plantae or plant is applied to a specific group of organisms or taxon, it usually refers to one of three concepts. From least to most inclusive, these three groupings are:

Name(s)	Scope	Description
Land plants, also known as Embryophyta or Metaphyta.	Plantae <i>sensu strictissimo</i>	This group includes the liverworts, hornworts, mosses, and vascular plants, as well as fossil plants similar to these surviving groups.
<b>Green plants</b> - also known as <b>Viridiplantae</b> , <b>Viridiphyta</b> or <b>Chlorobionta</b>	Plantae <i>sensu stricto</i>	This group includes the land plants plus various groups of green algae, including stoneworts. The names given to these groups vary considerably as of July 2011. Viridiplantae encompass a group of organisms that possess chlorophyll <i>a</i> and <i>b</i> , have plastids that are bound by only two membranes, are capable of storing starch, and have cellulose in their cell walls. It is this clade which is mainly the subject of this article.
Archaeplastida, Plastida or Primoplantae	Plantae <i>sensu lato</i>	This group comprises the green plants above plus Rhodophyta (red algae) and Glaucophyta (glaucophyte algae). This clade includes the organisms that eons ago acquired their chloroplasts directly by engulfing cyanobacteria.

Another way of looking at the relationships between the different groups which have been called "plants" is through a cladogram, which shows their evolutionary relationships. The evolutionary history of plants is not yet completely settled, but one accepted relationship between the three groups described above is shown below.<sup>[5]</sup> Those which have been called "plants" are in bold.

		Glaucophyta (glaucophyte algae)
		Rhodophyta (red algae)
		Chlorophyta (part of green algae)
		streptophyte algae (part of green algae)
Archaeplastida		
	Viridiplantae	Charales (stoneworts, often included in green algae)
	Streptophyta	land plants or embryophytes

The way in which the groups of green algae are combined and named varies considerably between authors.

Algae

Algae comprise several different groups of organisms which produce energy through photosynthesis and for that reason have been included in the plant kingdom in the past. Most conspicuous among the algae are the seaweeds, multicellular algae that may roughly resemble land plants, but are classified among the brown, red and green algae. Each of these algal groups also includes various microscopic and single-celled organisms. There is good evidence that some of these algal groups arose independently from separate non-photosynthetic ancestors, with the result that many groups of algae are no longer classified within the plant kingdom as it is defined here.<sup>[6][7]</sup>

The Viridiplantae, the green plants – green algae and land plants – form a clade, a group consisting of all the descendants of a common ancestor. With a few exceptions among the green algae, all green plants have many features in common, including cell walls containing cellulose, chloroplasts containing chlorophylls *a* and *b*, and food stores in the form of starch. They undergo closed mitosis without centrioles, and typically have mitochondria with flat cristae. The chloroplasts of green plants are surrounded by two membranes, suggesting they originated directly from endosymbiotic cyanobacteria.

Two additional groups, the Rhodophyta (red algae) and Glaucophyta (glaucophyte algae), also have chloroplasts which appear to be derived directly from endosymbiotic cyanobacteria, although they differ in the pigments which are used in photosynthesis and so are different in colour. All three groups together are generally believed to have a single common origin, and so are classified together in the taxon Archaeplastida, whose name implies that the chloroplasts or plastids of all the members of the taxon were derived from a single ancient endosymbiotic event. This is the broadest modern definition of the plants.



In contrast, most other algae (e.g. heterokonts, haptophytes, dinoflagellates, and euglenids) not only have different pigments but also have chloroplasts with three or four surrounding membranes. They are not close relatives of the Archaeplastida, presumably having acquired chloroplasts separately from ingested or symbiotic green and red algae. They are thus not included in even the broadest modern definition of the plant kingdom, although they were in the past.

The green plants or Viridiplantae were traditionally divided into the green algae (including the stoneworts) and the land plants. However, it is now known that the land plants evolved from within a group of green algae, so that the green algae by themselves are a paraphyletic group, i.e. a group which excludes some of the descendants of a common ancestor. Paraphyletic groups are generally avoided in modern classifications, so that in recent treatments the Viridiplantae have been divided into two clades, the Chlorophyta and the Streptophyta (or Charophyta).<sup>[8][9]</sup>

The Chlorophyta (a name that has also been used for *all* green algae) are the sister group to the group from which the land plants evolved. There are about 4,300 species<sup>[10]</sup> of mainly marine organisms, both unicellular and multicellular. The latter include the sea lettuce, *Ulva*.

The other group within the Viridiplantae are the mainly freshwater or terrestrial Streptophyta (or Charophyta), which consist of several groups of green algae plus the stoneworts and land plants. (The names have been used differently, e.g. Streptophyta to mean the group which excludes the land plants and Charophyta for the stoneworts alone or the stoneworts plus the land plants.) Streptophyte algae are either unicellular or form multicellular filaments, branched or unbranched.<sup>[9]</sup> The genus *Spirogyra* is a filamentous streptophyte alga familiar to many, as it is often used in teaching and is one of the organisms responsible for the algal "scum" which pond-owners so dislike. The freshwater stoneworts strongly resemble land plants and are believed to be their closest relatives. Growing underwater, they consist of a central stalk with whorls of branchlets, giving them a superficial resemblance to horsetails, species of the genus *Equisetum*, which are true land plants.

## Fungi

The classification of fungi has been controversial until quite recently in the history of biology. Linnaeus' original classification placed the fungi within the Plantae, since they were unquestionably not animals or minerals and these were the only other alternatives. With later developments in microbiology, in the 19th century Ernst Haeckel felt that another kingdom was required to classify newly discovered micro-organisms. The introduction of the new kingdom Protista in addition to Plantae and Animalia, led to uncertainty as to whether fungi truly were best placed in the Plantae or whether they ought to be reclassified as protists. Haeckel himself found it difficult to decide and it was not until 1969 that a solution was found whereby Robert Whittaker proposed the creation of the kingdom Fungi. Molecular evidence has since shown that the last common ancestor (concestor) of the Fungi was probably more similar to that of the Animalia than of any other kingdom, including the Plantae.

Whittaker's original reclassification was based on the fundamental difference in nutrition between the Fungi and the Plantae. Unlike plants, which generally gain carbon through photosynthesis, and so are called autotrophic phototrophs, fungi generally obtain carbon by breaking down and absorbing surrounding materials, and so are called heterotrophic saprotrophs. In addition, the substructure of multicellular fungi is different from that of plants, taking the form of many chitinous microscopic strands called hyphae, which may be further subdivided into cells or may form a syncytium containing many eukaryotic nuclei. Fruiting bodies, of which mushrooms are most familiar example, are the reproductive structures of fungi, and are unlike any structures produced by plants.

## Diversity

The table below shows some species count estimates of different green plant (Viridiplantae) divisions. It suggests there are about 300,000 species of living Viridiplantae, of which 85-90% are flowering plants. (Note: as these are from different sources and different dates, they are not necessarily comparable, and like all species counts, are subject to a degree of uncertainty in some cases.)

### Diversity of living green plant (Viridiplantae) divisions

Informal group	Division name	Common name	No. of living species	Approximate No. in informal group
Green algae	<b>Chlorophyta</b>	green algae (chlorophytes)	3,800 <sup>[11]</sup> – 4,300 <sup>[12]</sup>	8,500 (6,600 - 10,300)
	<b>Charophyta</b>	green algae (e.g. desmids & stoneworts)	2,800; <sup>[13]</sup> 4,000-6,000 <sup>[14]</sup>	
Bryophytes	<b>Marchantiophyta</b>	liverworts	6,000-8,000 <sup>[15]</sup>	19,000 (18,100 - 20,200)
	<b>Anthocerotophyta</b>	hornworts	100-200 <sup>[16]</sup>	
	<b>Bryophyta</b>	mosses	12,000 <sup>[17]</sup>	
Pteridophytes	<b>Lycopodiophyta</b>	club mosses	1,200 <sup>[7]</sup>	12,000 (12,200)
	<b>Pteridophyta</b>	ferns, whisk ferns & horsetails	11,000 <sup>[7]</sup>	
Seed plants	<b>Cycadophyta</b>	cycads	160 <sup>[18]</sup>	260,000 (259,511)
	<b>Ginkgophyta</b>	ginkgo	1 <sup>[19]</sup>	
	<b>Pinophyta</b>	conifers	630 <sup>[7]</sup>	
	<b>Gnetophyta</b>	gnetophytes	70 <sup>[7]</sup>	
	<b>Magnoliophyta</b>	flowering plants	258,650 <sup>[20]</sup>	

The naming of plants is governed by the International Code of Botanical Nomenclature and International Code of Nomenclature for Cultivated Plants (see cultivated plant taxonomy).

## Evolution

Further information: Evolutionary history of plants

The evolution of plants has resulted in increasing levels of complexity, from the earliest algal mats, through bryophytes, lycopods, ferns to the complex gymnosperms and angiosperms of today. The groups which appeared earlier continue to thrive, especially in the environments in which they evolved.

Evidence suggests that an algal scum formed on the land 1200<sup>[70]</sup> million years ago, but it was not until the Ordovician Period, around 450.0<sup>[21]</sup> million years ago, that land plants appeared.<sup>[22]</sup> However, new evidence from the study of carbon isotope ratios in Precambrian rocks has suggested that complex photosynthetic plants developed on the earth over 1000 m.y.a.<sup>[23]</sup> These began to diversify in the late Silurian Period, around 420<sup>[24]</sup> million years ago, and the fruits of their diversification are displayed in remarkable detail in an early Devonian fossil assemblage from the Rhynie chert. This chert preserved early plants in cellular detail, petrified in volcanic springs. By the middle of the Devonian Period most of the features recognised in plants today are present, including roots, leaves and secondary wood, and by late Devonian times seeds had evolved.<sup>[25]</sup> Late Devonian plants had thereby reached a degree of sophistication that allowed them to form forests of tall trees. Evolutionary innovation continued after the Devonian period. Most plant groups were relatively unscathed by the Permo-Triassic extinction event, although the

structures of communities changed. This may have set the scene for the evolution of flowering plants in the Triassic (~200<sup>[26]</sup> million years ago), which exploded in the Cretaceous and Tertiary. The latest major group of plants to evolve were the grasses, which became important in the mid Tertiary, from around 40<sup>[27]</sup> million years ago. The grasses, as well as many other groups, evolved new mechanisms of metabolism to survive the low CO<sub>2</sub> and warm, dry conditions of the tropics over the last 10<sup>[28]</sup> million years.

A proposed phylogenetic tree of Plantae, after Kenrick and Crane,<sup>[29]</sup> is as follows, with modification to the Pteridophyta from Smith et al.<sup>[30]</sup> The Prasinophyceae may be a paraphyletic basal group to all green plants.

Prasinophyceae (micromonads)





**Chlorophyta**

Trebouxiophyceae  
(Pleurostrophyceae)  
Chlorophyceae

Ulvophyceae

**Embryophytes**

The plants that are likely most familiar to us are the multicellular land plants, called embryophytes. They include the vascular plants, plants with full systems of leaves, stems, and roots. They also include a few of their close relatives, often called *bryophytes*, of which mosses and liverworts are the most common.

All of these plants have eukaryotic cells with cell walls composed of cellulose, and most obtain their energy through photosynthesis, using light and carbon dioxide to synthesize food. About three hundred plant species do not photosynthesize but are parasites on other species of photosynthetic plants. Plants are distinguished from green algae, which represent a mode of photosynthetic life similar to the kind modern plants are believed to have evolved from, by having specialized reproductive organs protected by non-reproductive tissues.



*Dicksonia antarctica*, a species of tree fern

Bryophytes first appeared during the early Paleozoic. They can only survive where moisture is available for significant periods, although some species are desiccation tolerant. Most species of bryophyte remain small throughout their life-cycle. This involves an alternation between two generations: a haploid stage, called the gametophyte, and a diploid stage, called the sporophyte. The sporophyte is short-lived and remains dependent on its parent gametophyte.

Vascular plants first appeared during the Silurian period, and by the Devonian had diversified and spread into many different land environments. They have a number of adaptations that allowed them to overcome the limitations of the bryophytes. These include a cuticle resistant to desiccation, and vascular tissues which transport water throughout the organism. In most the sporophyte acts as a separate individual, while the gametophyte remains small.

The first primitive seed plants, Pteridosperms (seed ferns) and Cordaites, both groups now extinct, appeared in the late Devonian and diversified through the Carboniferous, with further evolution through the Permian and Triassic periods. In these the gametophyte stage is completely reduced, and the sporophyte begins life inside an enclosure called a seed, which develops while on the parent plant, and with fertilisation by means of pollen grains. Whereas other vascular plants, such as ferns, reproduce by means of spores and so need moisture to develop, some seed plants can survive and reproduce in extremely arid conditions.

Early seed plants are referred to as gymnosperms (naked seeds), as the seed embryo is not enclosed in a protective structure at pollination, with the pollen landing directly on the embryo. Four surviving groups remain widespread now, particularly the conifers, which are dominant trees in several biomes. The angiosperms, comprising the flowering plants, were the last major group of plants to appear, emerging from within the gymnosperms during the Jurassic and diversifying rapidly during the Cretaceous. These differ in that the seed embryo (angiosperm) is enclosed, so the pollen has to grow a tube to penetrate the protective seed coat; they are the predominant group of

flora in most biomes today.

## Fossils

Plant fossils include roots, wood, leaves, seeds, fruit, pollen, spores, phytoliths, and amber (the fossilized resin produced by some plants). Fossil land plants are recorded in terrestrial, lacustrine, fluvial and nearshore marine sediments. Pollen, spores and algae (dinoflagellates and acritarchs) are used for dating sedimentary rock sequences. The remains of fossil plants are not as common as fossil animals, although plant fossils are locally abundant in many regions worldwide.

The earliest fossils clearly assignable to Kingdom Plantae are fossil green algae from the Cambrian. These fossils resemble calcified multicellular members of the Dasycladales. Earlier Precambrian fossils are known which resemble single-cell green algae, but definitive identity with that group of algae is uncertain.

The oldest known fossils of embryophytes date from the Ordovician, though such fossils are fragmentary. By the Silurian, fossils of whole plants are preserved, including the lycophyte *Baragwanathia longifolia*. From the Devonian, detailed fossils of rhyniophytes have been found. Early fossils of these ancient plants show the individual cells within the plant tissue. The Devonian period also saw the evolution of what many believe to be the first modern tree, *Archaeopteris*. This fern-like tree combined a woody trunk with the fronds of a fern, but produced no seeds.

The Coal measures are a major source of Paleozoic plant fossils, with many groups of plants in existence at this time. The spoil heaps of coal mines are the best places to collect; coal itself is the remains of fossilised plants, though structural detail of the plant fossils is rarely visible in coal. In the Fossil Forest at Victoria Park in Glasgow, Scotland, the stumps of *Lepidodendron* trees are found in their original growth positions.

The fossilized remains of conifer and angiosperm roots, stems and branches may be locally abundant in lake and inshore sedimentary rocks from the Mesozoic and Cenozoic eras. Sequoia and its allies, magnolia, oak, and palms are often found.

Petrified wood is common in some parts of the world, and is most frequently found in arid or desert areas where it is more readily exposed by erosion. Petrified wood is often heavily silicified (the organic material replaced by silicon dioxide), and the impregnated tissue is often preserved in fine detail. Such specimens may be cut and polished using lapidary equipment. Fossil forests of petrified wood have been found in all continents.

Fossils of seed ferns such as *Glossopteris* are widely distributed throughout several continents of the Southern Hemisphere, a fact that gave support to Alfred Wegener's early ideas regarding Continental drift theory.

## Structure, growth, and development

Most of the solid material in a plant is taken from the atmosphere. Through a process known as photosynthesis, most plants use the energy in sunlight to convert carbon dioxide from the atmosphere, plus water, into simple sugars. Parasitic plants, on the other hand, use the resources of its host to grow. These sugars are then used as building blocks and form the main structural component of the plant. Chlorophyll, a green-colored, magnesium-containing pigment is essential to this process; it is generally present in plant leaves, and often in other plant parts as well.



A petrified log in Petrified Forest National Park.

Plants usually rely on soil primarily for support and water (in quantitative terms), but also obtain compounds of nitrogen, phosphorus, and other crucial elemental nutrients. Epiphytic and lithophytic plants often depend on rainwater or other sources for nutrients and carnivorous plants supplement their nutrient requirements with insect prey that they capture. For the majority of plants to grow successfully they also require oxygen in the atmosphere and around their roots for respiration. However, some plants grow as submerged aquatics, using oxygen dissolved in the surrounding water, and a few specialized vascular plants, such as mangroves, can grow with their roots in anoxic conditions.

### Factors affecting growth

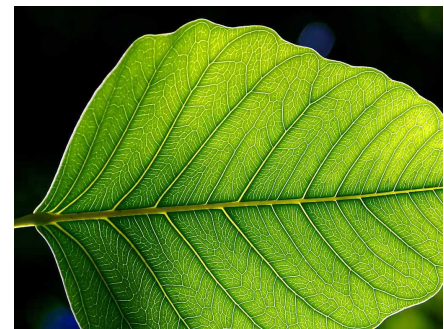
The genotype of a plant affects its growth. For example, selected varieties of wheat grow rapidly, maturing within 110 days, whereas others, in the same environmental conditions, grow more slowly and mature within 155 days.<sup>[31]</sup>

Growth is also determined by environmental factors, such as temperature, available water, available light, and available nutrients in the soil. Any change in the availability of these external conditions will be reflected in the plants growth.

Biotic factors are also capable of affecting plant growth. Plants compete with other plants for space, water, light and nutrients. Plants can be so crowded that no single individual produces normal growth, causing etiolation and chlorosis. Optimal plant growth can be hampered by grazing animals, suboptimal soil composition, lack of mycorrhizal fungi, and attacks by insects or plant diseases, including those caused by bacteria, fungi, viruses, and nematodes.<sup>[31]</sup>

Simple plants like algae may have short life spans as individuals, but their populations are commonly seasonal. Other plants may be organized according to their seasonal growth pattern: annual plants live and reproduce within one growing season, biennial plants live for two growing seasons and usually reproduce in second year, and perennial plants live for many growing seasons and continue to reproduce once they are mature. These designations often depend on climate and other environmental factors; plants that are annual in alpine or temperate regions can be biennial or perennial in warmer climates. Among the vascular plants, perennials include both evergreens that keep their leaves the entire year, and deciduous plants which lose their leaves for some part of it. In temperate and boreal climates, they generally lose their leaves during the winter; many tropical plants lose their leaves during the dry season.

The growth rate of plants is extremely variable. Some mosses grow less than 0.001 millimeters per hour (mm/h), while most trees grow 0.025-0.250 mm/h. Some climbing species, such as kudzu, which do not need to produce thick supportive tissue, may grow up to 12.5 mm/h.



The leaf is usually the primary site of photosynthesis in plants.



There is no photosynthesis in deciduous leaves in autumn.

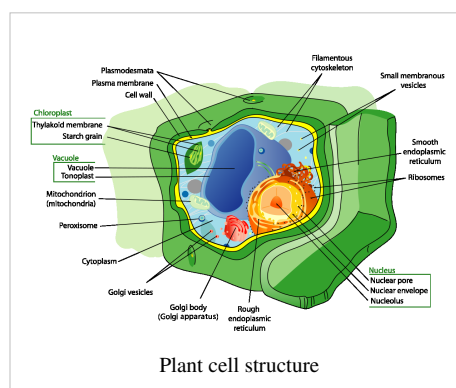
Plants protect themselves from frost and dehydration stress with antifreeze proteins, heat-shock proteins and sugars (sucrose is common). LEA (Late Embryogenesis Abundant) protein expression is induced by stresses and protects other proteins from aggregation as a result of desiccation and freezing.<sup>[32]</sup>



Dried dead plants

## Plant cell

Plant cells are typically distinguished by their large water-filled central vacuole, chloroplasts, and rigid cell walls that are made up of cellulose, hemicellulose, and pectin. Cell division is also characterized by the development of a phragmoplast for the construction of a cell plate in the late stages of cytokinesis. Just as in animals, plant cells differentiate and develop into multiple cell types. Totipotent meristematic cells can differentiate into vascular, storage, protective (e.g. epidermal layer), or reproductive tissues, with more primitive plants lacking some tissue types.<sup>[33]</sup>



Plant cell structure

## Physiology

### Photosynthesis

Plants are photosynthetic, which means that they manufacture their own food molecules using energy obtained from light. The primary mechanism plants have for capturing light energy is the pigment chlorophyll. All green plants contain two forms of chlorophyll, chlorophyll *a* and chlorophyll *b*. The latter of these pigments is not found in red or brown algae.

### Immune system

By means of cells that behave like nerves, plants receive and distribute within their systems information about incident light intensity and quality. Incident light which stimulates a chemical reaction in one leaf, will cause a chain reaction of signals to the entire plant via a type of cell termed a *bundle sheath cell*. Researchers from the Warsaw University of Life Sciences in Poland, found that plants have a specific memory for varying light conditions which prepares their immune systems against seasonal pathogens.<sup>[34]</sup> Plants use pattern-recognition receptors to recognize conserved microbial signatures. This recognition triggers an immune response. The first plant receptors of conserved microbial signatures were identified in rice (XA21, 1995)<sup>[35]</sup> and in *Arabidopsis* (FLS2, 2000).<sup>[36]</sup> Plants also carry immune receptors that recognize highly variable pathogen effectors. These include the NBS-LRR class of proteins.



## Internal distribution

Vascular plants differ from other plants in that they transport nutrients between different parts through specialized structures, called xylem and phloem. They also have roots for taking up water and minerals. The xylem moves water and minerals from the root to the rest of the plant, and the phloem provides the roots with sugars and other nutrient produced by the leaves.<sup>[33]</sup>

## Ecology

The photosynthesis conducted by land plants and algae is the ultimate source of energy and organic material in nearly all ecosystems. Photosynthesis radically changed the composition of the early Earth's atmosphere, which as a result is now 21% oxygen. Animals and most other organisms are aerobic, relying on oxygen; those that do not are confined to relatively rare anaerobic environments. Plants are the primary producers in most terrestrial ecosystems and form the basis of the food web in those ecosystems. Many animals rely on plants for shelter as well as oxygen and food.

Land plants are key components of the water cycle and several other biogeochemical cycles. Some plants have coevolved with nitrogen fixing bacteria, making plants an important part of the nitrogen cycle. Plant roots play an essential role in soil development and prevention of soil erosion.

## Distribution

Plants are distributed worldwide in varying numbers. While they inhabit a multitude of biomes and ecoregions, few can be found beyond the tundras at the northernmost regions of continental shelves. At the southern extremes, plants have adapted tenaciously to the prevailing conditions. (See Antarctic flora.)

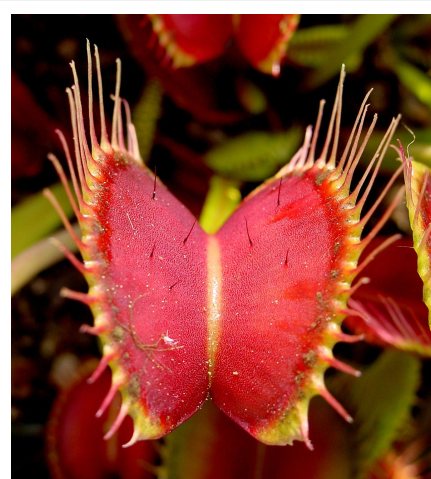
Plants are often the dominant physical and structural component of habitats where they occur. Many of the Earth's biomes are named for the type of vegetation because plants are the dominant organisms in those biomes, such as grasslands and forests.

## Ecological relationships

Numerous animals have coevolved with plants. Many animals pollinate flowers in exchange for food in the form of pollen or nectar. Many animals disperse seeds, often by eating fruit and passing the seeds in their feces. Myrmecophytes are plants that have coevolved with ants. The plant provides a home, and sometimes food, for the ants. In exchange, the ants defend the plant from herbivores and sometimes competing plants. Ant wastes provide organic fertilizer.

The majority of plant species have various kinds of fungi associated with their root systems in a kind of mutualistic symbiosis known as mycorrhiza. The fungi help the plants gain water and mineral nutrients from the soil, while the plant gives the fungi carbohydrates manufactured in photosynthesis. Some plants serve as homes for endophytic fungi that protect the plant from herbivores by producing toxins. The fungal endophyte, *Neotyphodium coenophialum*, in tall fescue (*Festuca arundinacea*) does tremendous economic damage to the cattle industry in the U.S.

Various forms of parasitism are also fairly common among plants, from the semi-parasitic mistletoe that merely takes some nutrients from its host, but still has photosynthetic leaves, to the fully parasitic broomrape and toothwort that acquire all their nutrients through connections to the roots of other plants, and so have no chlorophyll. Some



The Venus flytrap, a species of carnivorous plant.

plants, known as myco-heterotrophs, parasitize mycorrhizal fungi, and hence act as epiparasites on other plants.

Many plants are epiphytes, meaning they grow on other plants, usually trees, without parasitizing them. Epiphytes may indirectly harm their host plant by intercepting mineral nutrients and light that the host would otherwise receive. The weight of large numbers of epiphytes may break tree limbs. Hemiepiphytes like the strangler fig begin as epiphytes but eventually set their own roots and overpower and kill their host. Many orchids, bromeliads, ferns and mosses often grow as epiphytes. Bromeliad epiphytes accumulate water in leaf axils to form phytotelmata, complex aquatic food webs.<sup>[37]</sup>

Approximately 630 plants are carnivorous, such as the Venus Flytrap (*Dionaea muscipula*) and sundew (*Drosera* species). They trap small animals and digest them to obtain mineral nutrients, especially nitrogen and phosphorus.<sup>[38]</sup>

## Importance

The study of plant uses by people is termed economic botany or ethnobotany; some consider economic botany to focus on modern cultivated plants, while ethnobotany focuses on indigenous plants cultivated and used by native peoples. Human cultivation of plants is part of agriculture, which is the basis of human civilization. Plant agriculture is subdivided into agronomy, horticulture and forestry.

## Food

Much of human nutrition depends on land plants, either directly or indirectly.

Human nutrition depends to a large extent on cereals, especially maize (or corn), wheat and rice. Other staple crops include potato, cassava, and legumes. Human food also includes vegetables, spices, and certain fruits, nuts, herbs, and edible flowers.

Beverages produced from plants include coffee, tea, wine, beer and alcohol.

Sugar is obtained mainly from sugar cane and sugar beet.

Cooking oils and margarine come from maize, soybean, rapeseed, safflower, sunflower, olive and others.

Food additives include gum arabic, guar gum, locust bean gum, starch and pectin.

Livestock animals including cows, pigs, sheep, and goats are all herbivores; and feed primarily or entirely on cereal plants, particularly grasses.

## Nonfood products

Wood is used for buildings, furniture, paper, cardboard, musical instruments and sports equipment. Cloth is often made from cotton, flax or synthetic fibers derived from cellulose, such as rayon and acetate. Renewable fuels from plants include firewood, peat and many other biofuels. Coal and petroleum are fossil fuels derived from plants. Medicines derived from plants include aspirin, taxol, morphine,



Potato plant. Potatoes spread to the rest of the world after European contact with the Americas in the late 15th and early 16th centuries and have since become an important field crop.



Timber in storage for later processing at a sawmill.

quinine, reserpine, colchicine, digitalis and vincristine. There are hundreds of herbal supplements such as ginkgo, Echinacea, feverfew, and Saint John's wort. Pesticides derived from plants include nicotine, rotenone, strychnine and pyrethrins. Drugs obtained from plants include opium, cocaine and marijuana. Poisons from plants include ricin, hemlock and curare. Plants are the source of many natural products such as fibers, essential oils, natural dyes, pigments, waxes, tannins, latex, gums, resins, alkaloids, amber and cork. Products derived from plants include soaps, paints, shampoos, perfumes, cosmetics, turpentine, rubber, varnish, lubricants, linoleum, plastics, inks, chewing gum and hemp rope. Plants are also a primary source of basic chemicals for the industrial synthesis of a vast array of organic chemicals. These chemicals are used in a vast variety of studies and experiments.

### Aesthetic uses

Thousands of plant species are cultivated for aesthetic purposes as well as to provide shade, modify temperatures, reduce wind, abate noise, provide privacy, and prevent soil erosion. People use cut flowers, dried flowers and houseplants indoors or in greenhouses. In outdoor gardens, lawn grasses, shade trees, ornamental trees, shrubs, vines, herbaceous perennials and bedding plants are used. Images of plants are often used in art, architecture, humor, language, and photography and on textiles, money, stamps, flags and coats of arms. Living plant art forms include topiary, bonsai, ikebana and espalier. Ornamental plants have sometimes changed the course of history, as in tulipomania. Plants are the basis of a multi-billion dollar per year tourism industry which includes travel to arboretums, botanical gardens, historic gardens, national parks, tulip festivals, rainforests, forests with colorful autumn leaves and the National Cherry Blossom Festival. Venus Flytrap, sensitive plant and resurrection plant are examples of plants sold as novelties.

### Scientific and cultural uses

Tree rings are an important method of dating in archeology and serve as a record of past climates. Basic biological research has often been done with plants, such as the pea plants used to derive Gregor Mendel's laws of genetics. Space stations or space colonies may one day rely on plants for life support. Plants are used as national and state emblems, including state trees and state flowers. Ancient trees are revered and many are famous. Numerous world records are held by plants. Plants are often used as memorials, gifts and to mark special occasions such as births, deaths, weddings and holidays. Plants figure prominently in mythology, religion and literature. The field of ethnobotany studies plant use by indigenous cultures which helps to conserve endangered species as well as discover new medicinal plants. Gardening is the most popular leisure activity in the U.S. Working with plants or horticulture therapy is beneficial for rehabilitating people with disabilities. Certain plants contain psychotropic chemicals which are extracted and ingested, including tobacco, cannabis (marijuana), and opium.



A section of a Yew branch showing 27 annual growth rings, pale sapwood and dark heartwood, and pith (centre dark spot). The dark radial lines are longitudinal sections of small branches which became included by growth of the tree.



## Negative effects

Weeds are plants that grow where people do not want them. People have spread plants beyond their native ranges and some of these introduced plants become invasive, damaging existing ecosystems by displacing native species. Invasive plants cause billions of dollars in crop losses annually by displacing crop plants, they increase the cost of production and the use of chemical means to control them affects the environment.

Plants may cause harm to animals, including people. Plants that produce windblown pollen invoke allergic reactions in people who suffer from hay fever. A wide variety of plants are poisonous. Toxalbumins are plant poisons fatal to most mammals and act as a serious deterrent to consumption. Several plants cause skin irritations when touched, such as poison ivy. Certain plants contain psychotropic chemicals, which are extracted and ingested or smoked, including tobacco, cannabis (marijuana), cocaine and opium. Smoking causes damage to health or even death, while some drugs may also be harmful or fatal to people.<sup>[39][40]</sup> Both illegal and legal drugs derived from plants may have negative effects on the economy, affecting worker productivity and law enforcement costs.<sup>[41][42]</sup> Some plants cause allergic reactions when ingested, while other plants cause food intolerances that negatively affect health.

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- African Plants Initiative database ([http://www.aluka.org/action/doBrowse?sa=1&sa\\_sel=](http://www.aluka.org/action/doBrowse?sa=1&sa_sel=))
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- Chilean plants at *Chilebosque* (<http://www.chilebosque.cl/>)
- Dave's garden (<http://davesgarden.com/pf/>) plenty of information mostly about garden plants
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- Meet the Plants-National Tropical Botanical Garden ([http://www.ntbg.org/plants/choose\\_a\\_plant.php](http://www.ntbg.org/plants/choose_a_plant.php))
- Lady Bird Johnson Wildflower Center - Native Plant Information Network at University of Texas, Austin (<http://www.wildflower.org/>)
- The Plant List (<http://www.theplantlist.org/>)
- United States Department of Agriculture (<http://plants.usda.gov/>) not limited to continental US species

# Animal

Animals	
Temporal range: Ediacaran – Recent	
	
Scientific classification	
Domain:	Eukaryota
(unranked)	Opisthokonta
(unranked)	Holozoa
(unranked)	Filozoa
Kingdom:	<b>Animalia</b> Linnaeus, 1758
Phyla	

- **Subkingdom Parazoa**
  - Porifera
  - Placozoa
- **Subkingdom Eumetazoa**
  - **Radiata (unranked)**
    - Ctenophora
    - Cnidaria
  - **Bilateria (unranked)**
    - Orthonectida
    - Rhombozoa
    - Acoelomorpha
    - Chaetognatha
    - **Superphylum Deuterostomia**
      - Chordata
      - Hemichordata
      - Echinodermata
      - Xenoturbellida
      - Vetulicolia †
    - **Protostomia (unranked)**
      - **Superphylum Ecdysozoa**
        - Kinorhyncha
        - Loricifera
        - Priapulida
        - Nematoda
        - Nematomorpha
        - Lobopodia
        - Onychophora
        - Tardigrada
        - Arthropoda
      - **Superphylum Platyzoa**
        - Platyhelminthes
        - Gastrotricha
        - Rotifera
        - Acanthocephala
        - Gnathostomulida
        - Micrognathozoa
        - Cycliophora
      - **Superphylum Lophotrochozoa**
        - Sipuncula
        - Hyolitha †
        - Nemertea
        - Phoronida
        - Bryozoa
        - Entoprocta
        - Brachiopoda
        - Mollusca
        - Annelida
        - Echiura

**Animals** are a major group of multicellular, eukaryotic organisms of the kingdom **Animalia** or **Metazoa**. Their body plan eventually becomes fixed as they develop, although some undergo a process of metamorphosis later on in their life. Most animals are motile, meaning they can move spontaneously and independently. All animals are also heterotrophs, meaning they must ingest other organisms or their products for sustenance.

Most known animal phyla appeared in the fossil record as marine species during the Cambrian explosion, about 542 million years ago.

## Etymology

The word "animal" comes from the Latin word *animalis*, meaning "having breath".<sup>[1]</sup> In everyday colloquial usage, the word often refers to non-human members of kingdom Animalia. Sometimes, only closer relatives of humans such as mammals and other vertebrates are meant in colloquial use.<sup>[2]</sup> The biological definition of the word refers to all members of the kingdom Animalia, encompassing creatures as diverse as sponges, jellyfish, insects and humans.<sup>[3]</sup>

## Characteristics

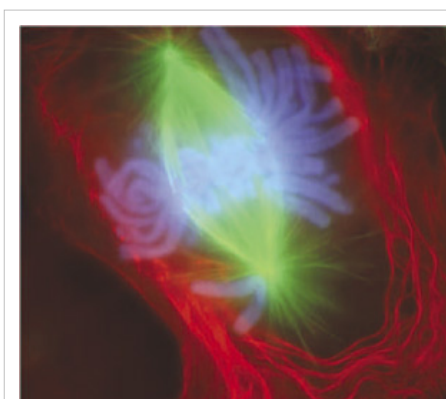
Animals have several characteristics that set them apart from other living things. Animals are eukaryotic and mostly multicellular,<sup>[4]</sup> which separates them from bacteria and most protists. They are heterotrophic,<sup>[5]</sup> generally digesting food in an internal chamber, which separates them from plants and algae.<sup>[6]</sup> They are also distinguished from plants, algae, and fungi by lacking rigid cell walls.<sup>[7]</sup> All animals are motile,<sup>[8]</sup> if only at certain life stages. In most animals, embryos pass through a blastula stage,<sup>[9]</sup> which is a characteristic exclusive to animals.

## Structure

With a few exceptions, most notably the sponges (Phylum Porifera) and Placozoa, animals have bodies differentiated into separate tissues. These include muscles, which are able to contract and control locomotion, and nerve tissues, which send and process signals. Typically, there is also an internal digestive chamber, with one or two openings.<sup>[10]</sup> Animals with this sort of organization are called metazoans, or eumetazoans when the former is used for animals in general.<sup>[11]</sup>

All animals have eukaryotic cells, surrounded by a characteristic extracellular matrix composed of collagen and elastic glycoproteins.<sup>[12]</sup> This may be calcified to form structures like shells, bones, and spicules.<sup>[13]</sup> During development, it forms a relatively flexible framework<sup>[14]</sup> upon which cells can move about and be reorganized, making complex structures possible. In contrast, other multicellular organisms, like plants and fungi, have cells held in place by cell walls, and so develop by progressive growth.<sup>[10]</sup> Also, unique to animal cells are the following intercellular junctions: tight junctions, gap junctions, and desmosomes.<sup>[15]</sup>

## Reproduction and development



A newt lung cell stained with fluorescent dyes undergoing the early anaphase stage of mitosis

Nearly all animals undergo some form of sexual reproduction.<sup>[16]</sup> They have a few specialized reproductive cells, which undergo meiosis to produce smaller, motile spermatozoa or larger, non-motile ova.<sup>[17]</sup> These fuse to form zygotes, which develop into new individuals.<sup>[18]</sup>

Many animals are also capable of asexual reproduction.<sup>[19]</sup> This may take place through parthenogenesis, where fertile eggs are produced without mating, budding, or fragmentation.<sup>[20]</sup>

A zygote initially develops into a hollow sphere, called a blastula,<sup>[21]</sup> which undergoes rearrangement and differentiation. In sponges, blastula larvae swim to a new location and develop into a new sponge.<sup>[22]</sup> In most other groups, the blastula undergoes more complicated rearrangement.<sup>[23]</sup> It first invaginates to form a gastrula

with a digestive chamber, and two separate germ layers — an external ectoderm and an internal endoderm.<sup>[24]</sup> In most cases, a mesoderm also develops between them.<sup>[25]</sup> These germ layers then differentiate to form tissues and organs.<sup>[26]</sup>

## Food and energy sourcing



All animals are heterotrophs, meaning that they feed directly or indirectly on other living things.<sup>[27]</sup> They are often further subdivided into groups such as carnivores, herbivores, omnivores, and parasites.<sup>[28]</sup>

Predation is a biological interaction where a predator (a heterotroph that is hunting) feeds on its prey (the organism that is attacked).<sup>[29]</sup> Predators may or may not kill their prey prior to feeding on them, but the act of predation always results in the death of the prey.<sup>[30]</sup> The other main category of consumption is detritivory, the consumption of dead organic matter.<sup>[31]</sup> It can at times be difficult to separate the two feeding behaviours, for example, where parasitic species prey on a host organism and then lay their eggs on it for their offspring to feed on its decaying corpse. Selective pressures imposed on one another has led to an evolutionary arms race between prey and predator, resulting in various antipredator adaptations.<sup>[32]</sup>

Most animals indirectly use the energy of sunlight by eating plants or plant-eating animals. Most plants use light to convert inorganic molecules in their environment into carbohydrates, fats, proteins and other biomolecules, characteristically containing reduced carbon in the form of carbon-hydrogen bonds. Starting with carbon dioxide (CO<sub>2</sub>) and water (H<sub>2</sub>O), photosynthesis converts the energy of sunlight into chemical energy in the form of simple sugars (e.g., glucose), with the release of molecular oxygen. These sugars are then used as the building blocks for plant growth, including the production of other biomolecules.<sup>[10]</sup> When an animal eats plants (or eats other animals which have eaten plants), the reduced carbon compounds in the food become a source of energy and building materials for the animal.<sup>[33]</sup> They are either used directly to help the animal grow, or broken down, releasing stored solar energy, and giving the animal the energy required for motion.<sup>[34]</sup> <sup>[35]</sup>

Animals living close to hydrothermal vents and cold seeps on the ocean floor are not dependent on the energy of sunlight.<sup>[36]</sup> Instead chemosynthetic archaea and bacteria form the base of the food chain.<sup>[37]</sup>

## Origin and fossil record

Further information: Urmetazoan

Animals are generally considered to have evolved from a flagellated eukaryote.<sup>[39]</sup> Their closest known living relatives are the choanoflagellates, collared flagellates that have a morphology similar to the choanocytes of certain sponges.<sup>[40]</sup> Molecular studies place animals in a supergroup called the opisthokonts, which also include the choanoflagellates, fungi and a few small parasitic protists.<sup>[41]</sup> The name comes from the posterior location of the flagellum in motile cells, such as most animal spermatozoa, whereas other eukaryotes tend to have anterior flagella.<sup>[42]</sup>

The first fossils that might represent animals appear in the Trezona Formation at Trezona Bore, West Central Flinders, South Australia.<sup>[43]</sup>

These fossils are interpreted as being early sponges. They were found in 665-million-year-old rock.<sup>[43]</sup>

The next oldest possible animal fossils are found towards the end of the Precambrian, around 610 million years ago, and are known as the Ediacaran or Vendian biota.<sup>[44]</sup> These are difficult to relate to later fossils, however. Some may represent precursors of modern phyla, but they may be separate groups, and it is possible they are not really animals at all.<sup>[45]</sup>

Aside from them, most known animal phyla make a more or less simultaneous appearance during the Cambrian period, about 542 million years ago.<sup>[46]</sup> It is still disputed whether this event, called the Cambrian explosion, represents a rapid divergence between different groups or a change in conditions that made fossilization possible.

Some paleontologists suggest that animals appeared much earlier than the Cambrian explosion, possibly as early as 1 billion years ago.<sup>[47]</sup> Trace fossils such as tracks and burrows found in the Tonian era indicate the presence of triploblastic worms, like metazoans, roughly as large (about 5 mm wide) and complex as earthworms.<sup>[48]</sup> During the



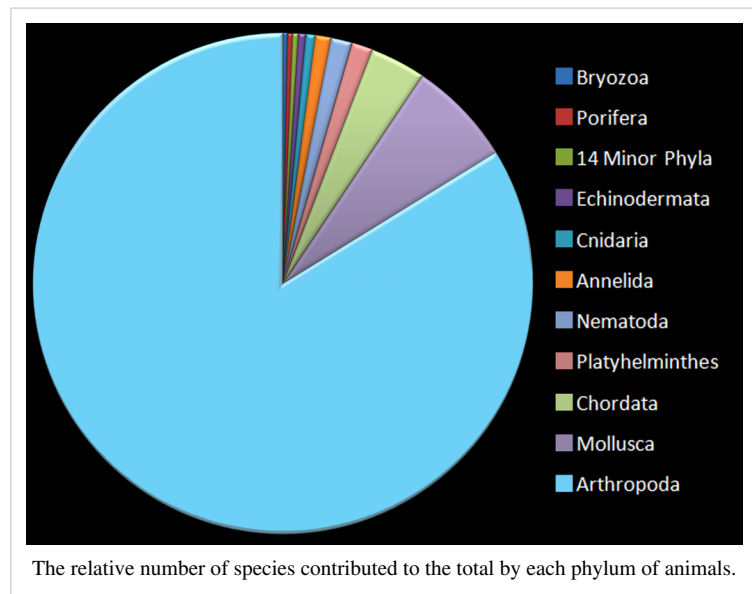
*Dunkleosteus* was a gigantic, 10-metre-long (unknown operator: u'strong' ft) prehistoric fish.<sup>[38]</sup>

beginning of the Tonian period around 1 billion years ago, there was a decrease in Stromatolite diversity, which may indicate the appearance of grazing animals, since stromatolite diversity increased when grazing animals went extinct at the End Permian and End Ordovician extinction events, and decreased shortly after the grazer populations recovered. However the discovery that tracks very similar to these early trace fossils are produced today by the giant single-celled protist *Gromia sphaerica* casts doubt on their interpretation as evidence of early animal evolution.<sup>[49][50]</sup>

## Groups of animals

### Porifera, Radiata and basal Bilateria

Phylogenetic analysis suggests that the Porifera and Ctenophora diverged before a clade that gave rise to the Bilateria, Cnidaria and Placozoa.<sup>[51]</sup>



Orange elephant ear sponge, *Agelas clathrodes*, in foreground. Two corals in the background: a sea fan, *Iciligorgia schrammi*, and a sea rod, *Plexaurella nutans*.

The sponges (Porifera) were long thought to have diverged from other animals early.<sup>[52]</sup> They lack the complex organization found in most other phyla.<sup>[53]</sup> Their cells are differentiated, but in most cases not organized into distinct tissues.<sup>[54]</sup> Sponges typically feed by drawing in water through pores.<sup>[55]</sup> Archaeocyatha, which have fused skeletons, may represent sponges or a separate phylum.<sup>[56]</sup> However, a phylogenomic study in 2008 of 150 genes in 29 animals across 21 phyla revealed that it is the Ctenophora or comb jellies which are the basal lineage of animals, at least among those 21 phyla. The authors speculate that sponges—or at least those lines of sponges they investigated—are not so primitive, but may instead be secondarily simplified.<sup>[57]</sup>

Among the other phyla, the Ctenophora and the Cnidaria, which includes sea anemones, corals, and jellyfish, are radially symmetric and have digestive chambers with a single opening, which serves as both the mouth and the anus.<sup>[58]</sup> Both have distinct tissues, but they are not organized into organs.<sup>[59]</sup> There are only two main germ layers, the ectoderm and endoderm, with only scattered cells between them. As such, these animals are sometimes called diploblastic.<sup>[60]</sup> The tiny placozoans are similar, but they do not have a permanent digestive

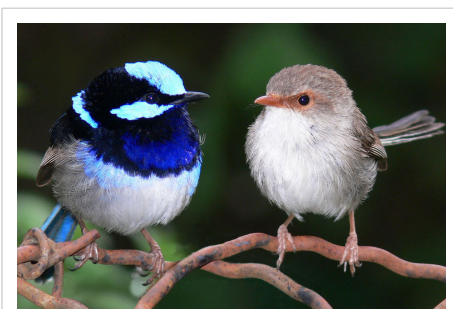
chamber.

The remaining animals form a monophyletic group called the Bilateria. For the most part, they are bilaterally symmetric, and often have a specialized head with feeding and sensory organs. The body is triploblastic, i.e. all three germ layers are well-developed, and tissues form distinct organs. The digestive chamber has two openings, a mouth and an anus, and there is also an internal body cavity called a coelom or pseudocoelom. There are exceptions to each of these characteristics, however — for instance adult echinoderms are radially symmetric, and certain parasitic worms have extremely simplified body structures.

Genetic studies have considerably changed our understanding of the relationships within the Bilateria. Most appear to belong to two major lineages: the deuterostomes and the protostomes, the latter of which includes the Ecdysozoa, Platyzoa, and Lophotrochozoa. In addition, there are a few small groups of bilaterians with relatively similar structure that appear to have diverged before these major groups. These include the Acoelomorpha, Rhombozoa, and Orthonectida. The Myxozoa, single-celled parasites that were originally considered Protozoa, are now believed to have developed from the Medusozoa as well.

## Deuterostomes

Deuterostomes differ from the other Bilateria, called protostomes, in several ways. In both cases there is a complete digestive tract. However, in protostomes, the first opening of the gut to appear in embryological development (the archenteron) develops into the mouth, with the anus forming secondarily. In deuterostomes the anus forms first, with the mouth developing secondarily.<sup>[61]</sup> In most protostomes, cells simply fill in the interior of the gastrula to form the mesoderm, called schizocoelous development, but in deuterostomes, it forms through invagination of the endoderm, called enterocoelic pouching.<sup>[62]</sup> Deuterostome embryos undergo radial cleavage during cell division, while protostomes undergo spiral cleavage.<sup>[63]</sup>



Superb Fairy-wren, *Malurus cyaneus*

All this suggests the deuterostomes and protostomes are separate, monophyletic lineages. The main phyla of deuterostomes are the Echinodermata and Chordata.<sup>[64]</sup> The former are radially symmetric and exclusively marine, such as starfish, sea urchins, and sea cucumbers.<sup>[65]</sup> The latter are dominated by the vertebrates, animals with backbones.<sup>[66]</sup> These include fish, amphibians, reptiles, birds, and mammals.<sup>[67]</sup>

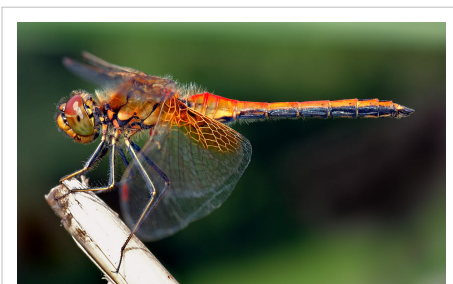
In addition to these, the deuterostomes also include the Hemichordata, or acorn worms.<sup>[68]</sup> Although they are not especially prominent today, the important fossil graptolites may belong to this group.<sup>[69]</sup>

The Chaetognatha or arrow worms may also be deuterostomes, but more recent studies suggest protostome affinities.

## Ecdysozoa

The Ecdysozoa are protostomes, named after the common trait of growth by moulting or ecdysis.<sup>[70]</sup> The largest animal phylum belongs here, the Arthropoda, including insects, spiders, crabs, and their kin. All these organisms have a body divided into repeating segments, typically with paired appendages. Two smaller phyla, the Onychophora and Tardigrada, are close relatives of the arthropods and share these traits.

The ecdysozoans also include the Nematoda or roundworms, perhaps the second largest animal phylum. Roundworms are typically



Yellow-winged darter, *Sympetrum flaveolum*

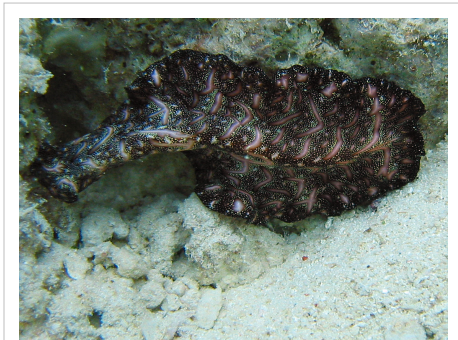
microscopic, and occur in nearly every environment where there is water.<sup>[71]</sup> A number are important parasites.<sup>[72]</sup> Smaller phyla related to them are the Nematomorpha or horsehair worms, and the Kinorhyncha, Priapulida, and Loricifera. These groups have a reduced coelom, called a pseudocoelom.

The remaining two groups of protostomes are sometimes grouped together as the Spiralia, since in both embryos develop with spiral cleavage.

## Platyzoa

The Platyzoa include the phylum Platyhelminthes, the flatworms.<sup>[73]</sup> These were originally considered some of the most primitive Bilateria, but it now appears they developed from more complex ancestors.<sup>[74]</sup> A number of parasites are included in this group, such as the flukes and tapeworms.<sup>[73]</sup> Flatworms are acoelomates, lacking a body cavity, as are their closest relatives, the microscopic Gastrotricha.<sup>[75]</sup>

The other platyzoan phyla are mostly microscopic and pseudocoelomate. The most prominent are the Rotifera or rotifers, which are common in aqueous environments. They also include the Acanthocephala or spiny-headed worms, the Gnathostomulida, Micrognathozoa, and possibly the Cycliophora.<sup>[76]</sup> These groups share the presence of complex jaws, from which they are called the Gnathiifera.



*Pseudobiceros bedfordi*, (Bedford's flatworm)

## Lophotrochozoa

The Lophotrochozoa include two of the most successful animal phyla, the Mollusca and Annelida.<sup>[77][78]</sup> The former, which is the second-largest animal phylum by number of described species, includes animals such as snails, clams, and squids, and the latter comprises the segmented worms, such as earthworms and leeches. These two groups have long been considered close relatives because of the common presence of trochophore larvae, but the annelids were considered closer to the arthropods because they are both segmented.<sup>[79]</sup> Now, this is generally considered convergent evolution, owing to many morphological and genetic differences between the two phyla.<sup>[80]</sup>



Roman snail, *Helix pomatia*

The Lophotrochozoa also include the Nemertea or ribbon worms, the Sipuncula, and several phyla that have a ring of ciliated tentacles around the mouth, called a lophophore.<sup>[81]</sup> These were traditionally grouped together as the lophophorates.<sup>[82]</sup> but it now appears that the lophophorate group may be paraphyletic,<sup>[83]</sup> with some closer to the nemerteans and some to the molluscs and annelids.<sup>[84][85]</sup> They include the Brachiopoda or lamp shells, which are prominent in the fossil record, the Entoprocta, the Phoronida, and possibly the Bryozoa or moss animals.<sup>[86]</sup>

## Model organisms

Because of the great diversity found in animals, it is more economical for scientists to study a small number of chosen species so that connections can be drawn from their work and conclusions extrapolated about how animals function in general. Because they are easy to keep and breed, the fruit fly *Drosophila melanogaster* and the nematode *Caenorhabditis elegans* have long been the most intensively studied metazoan model organisms, and were among the first life-forms to be genetically sequenced. This was facilitated by the severely reduced state of their genomes, but as many genes, introns, and linkages lost, these ecdysozoans can teach us little about the origins of animals in general. The extent of this type of evolution within the superphylum will be revealed by the crustacean,



annelid, and molluscan genome projects currently in progress. Analysis of the starlet sea anemone genome has emphasised the importance of sponges, placozoans, and choanoflagellates, also being sequenced, in explaining the arrival of 1500 ancestral genes unique to the Eumetazoa.<sup>[87]</sup>

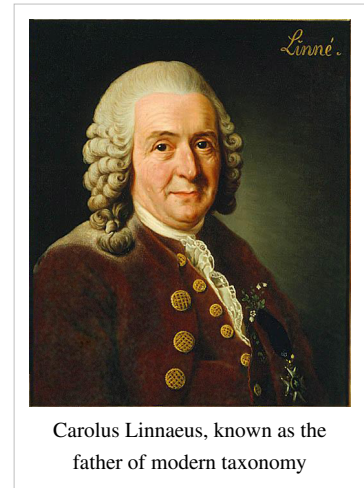
An analysis of the homoscleromorph sponge *Oscarella carmela* also suggests that the last common ancestor of sponges and the eumetazoan animals was more complex than previously assumed.<sup>[88]</sup>

Other model organisms belonging to the animal kingdom include the house mouse (*Mus musculus*) and zebrafish (*Danio rerio*).

## History of classification

Aristotle divided the living world between animals and plants, and this was followed by Carolus Linnaeus (Carl von Linné), in the first hierarchical classification.<sup>[89]</sup> Since then biologists have begun emphasizing evolutionary relationships, and so these groups have been restricted somewhat. For instance, microscopic protozoa were originally considered animals because they move, but are now treated separately.

In Linnaeus's original scheme, the animals were one of three kingdoms, divided into the classes of Vermes, Insecta, Pisces, Amphibia, Aves, and Mammalia. Since then the last four have all been subsumed into a single phylum, the Chordata, whereas the various other forms have been separated out. The above lists represent our current understanding of the group, though there is some variation from source to source.



Carolus Linnaeus, known as the father of modern taxonomy

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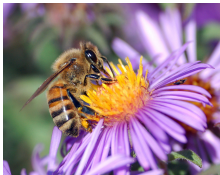
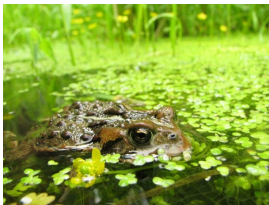

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## External links

- *Animal* (<http://www.eol.org/pages/1>) at the Encyclopedia of Life
- Tree of Life Project (<http://tolweb.org/>)
- Animal Diversity Web (<http://animaldiversity.ummz.umich.edu/site/index.html>) – University of Michigan's database of animals, showing taxonomic classification, images, and other information.
- ARKive (<http://www.arkive.org>) – multimedia database of worldwide endangered/protected species and common species of UK.
- Scientific American Magazine (December 2005 Issue) – Getting a Leg Up on Land (<http://www.sciam.com/article.cfm?chanID=sa006&articleID=000DC8B8-EA15-137C-AA1583414B7F0000>) About the evolution of four-limbed animals from fish.

# Ecology

## Ecology

	
	
	
<p>Ecology addresses the full scale of life, from tiny bacteria to processes that span the entire planet. Ecologists study many diverse and complex relations among species, such as predation and pollination. The diversity of life is organized into different habitats, from terrestrial (middle) to aquatic ecosystems.</p>	

**Ecology** (from Greek: οἶκος, "house"; -λογία, "study of"<sup>[A]</sup>) is the scientific study of the relations that living organisms have with respect to each other and their natural environment. Topics of interest to ecologists include the composition, distribution, amount (biomass), number, and changing states of organisms within and among ecosystems. Ecosystems are composed of dynamically interacting parts including organisms, the communities they make up, and the non-living components of their environment. Ecosystem processes, such as primary production, pedogenesis, nutrient cycling, and various niche construction activities, regulate the flux of energy and matter through an environment. These processes are sustained by the biodiversity within them. Biodiversity refers to the varieties of species in ecosystems, the genetic variations they contain, and the processes that are functionally enriched by the diversity of ecological interactions.

Ecology is an interdisciplinary branch of biology. The word "ecology" ("Ökologie") was coined in 1866 by the German scientist Ernst Haeckel (1834–1919). Ancient Greek philosophers such as Hippocrates and Aristotle laid the foundations of ecology in their studies on natural history. Modern ecology transformed into a more rigorous science in the late 19th century. Evolutionary concepts on adaptation and natural selection became cornerstones of modern ecological theory. Ecology is not synonymous with environment, environmentalism, natural history, or environmental science. It is closely related to physiology, evolutionary biology, genetics, and ethology. An understanding of how biodiversity affects ecological function is an important focus area in ecological studies. Ecologists seek to explain:

- Life processes and adaptations
- Distribution and abundance of organisms

- The movement of materials and energy through living communities
- The successional development of ecosystems, and
- The abundance and distribution of biodiversity in the context of the environment.

Ecology is a human science as well. There are many practical applications of ecology in conservation biology, wetland management, natural resource management (agriculture, forestry, fisheries), city planning (urban ecology), community health, economics, basic and applied science, and human social interaction (human ecology). Ecosystems maintain biophysical feedback mechanisms that modulate metabolic rates and evolutionary dynamics between living (biotic) and nonliving (abiotic) components of the planet. Ecosystems sustain life-supporting functions and produce natural capital through the regulation of continental climates, global biogeochemical cycles, water filtration, soils, food, fibres, medicines, erosion control, and many other natural features of scientific, historical, economic, or intrinsic value.

## History

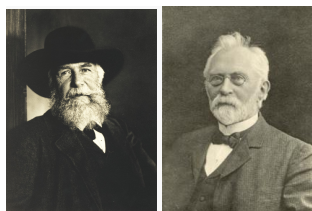
### Early beginnings

Ecology has a complex origin, due in large part to its interdisciplinary nature.<sup>[1]</sup> Ancient Greek philosophers such as Hippocrates and Aristotle were among the first to record observations on natural history. However, they viewed life in terms of essentialism, where species were conceptualized as static unchanging things while varieties were seen as aberrations of an idealized type. This contrasts against the modern understanding of ecological theory where varieties are viewed as the real phenomena of interest and having a role in the origins of adaptations by means of natural selection.<sup>[2][3][4]</sup> Early conceptions of ecology, such as a balance and regulation in nature can be traced to Herodotus (died *c.* 425 BC). Herodotus described one of the earliest accounts of mutualism in his observation of "natural dentistry". Basking Nile crocodiles, he noted, would open their mouths to give sandpipers safe access to pluck leeches out, giving nutrition to the sandpiper and oral hygiene for the crocodile.<sup>[1]</sup> Aristotle was an early influence on the philosophical development of ecology. He and his student Theophrastus made extensive observations on plant and animal migrations, biogeography, physiology, and on their behaviour, giving an early analogue to the modern concept of an ecological niche.<sup>[5][6]</sup>

Ecological concepts such as food chains, population regulation, and productivity were first developed in the 1700s, through the published works of microscopist Antoni van Leeuwenhoek (1632–1723) and botanist Richard Bradley (1688?–1732).<sup>[4]</sup> Biogeographer Alexander von Humbolt (1769–1859) was an early pioneer in ecological thinking and was among the first to recognize ecological gradients, where species are replaced or altered in form along environmental gradients, such as a cline forming along a rise in elevation. Humbolt drew inspiration from Isaac Newton as he developed a form of "terrestrial physics." In Newtonian fashion, he brought a scientific exactitude for measurement into natural history and even alluded to concepts that are the foundation of a modern ecological law on species-to-area relationships.<sup>[7][8][9]</sup> Natural historians, such as Humbolt, James Hutton and Jean-Baptiste Lamarck (among others) laid the foundations of the modern ecological sciences.<sup>[10]</sup> The term "ecology" (German: *Oekologie*) is of a more recent origin and was first coined by the German biologist Ernst Haeckel in his book *Generelle Morphologie der Organismen* (1866). Haeckel was a zoologist, artist, writer, and later in life a professor of comparative anatomy.<sup>[11][12]</sup>

By ecology, we mean the whole science of the relations of the organism to the environment including, in the broad sense, all the "conditions of existence."...Thus the theory of evolution explains the housekeeping relations of organisms mechanistically as the necessary consequences of effectual causes and so forms the monistic groundwork of ecology.

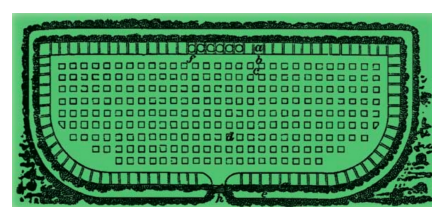
Ernst Haeckel (1866)<sup>[11]:140 [B]</sup>



Ernst Haeckel (left) and Eugenius Warming (right), two founders of ecology

Opinions differ on who was the founder of modern ecological theory. Some mark Haeckel's definition as the beginning,<sup>[13]</sup> others say it was Eugenius Warming with the writing of *Oecology of Plants: An Introduction to the Study of Plant Communities* (1895),<sup>[14]</sup> or Carl Linnaeus' principles on the economy of nature that matured in the early 18th century.<sup>[15][16]</sup> Linnaeus founded an early branch of ecology that he called the economy of nature.<sup>[15]</sup> His works influenced Charles Darwin, who adopted Linnaeus' phrase on the *economy or polity of nature* in *The Origin of Species*.<sup>[11]</sup> Linnaeus was the first to frame the balance of nature as a testable hypothesis. Haeckel, who admired Darwin's work, defined ecology in reference to the economy of nature, which has led some to question whether ecology and the economy of nature are synonymous.<sup>[16]</sup>

From Aristotle until Darwin, the natural world was predominantly considered static and unchanging. Prior to *The Origin of Species*, there was little appreciation or understanding of the dynamic and reciprocal relations between organisms, their adaptations, and the environment.<sup>[2]</sup> An exception is the 1789 publication *Natural History of Selborne* by Gilbert White (1720–1793), considered by some to be one of the earliest texts on ecology.<sup>[19]</sup> While Charles Darwin is mainly noted for his treatise on evolution,<sup>[20]</sup> he was one of the founders of soil ecology,<sup>[21]</sup> and he made note of the first ecological experiment in *The Origin of Species*.<sup>[17]</sup> Evolutionary theory changed the way that researchers approached the ecological sciences.<sup>[22]</sup>



The layout of the first ecological experiment, carried out in a grass garden at Woburn Abbey in 1816, was noted by Charles Darwin in *The Origin of Species*. The experiment studied the performance of different mixtures of species planted in different kinds of soils.<sup>[17][18]</sup>

Nowhere can one see more clearly illustrated what may be called the sensibility of such an organic complex,—expressed by the fact that whatever affects any species belonging to it, must speedily have its influence of some sort upon the whole assemblage. He will thus be made to see the impossibility of studying any form completely, out of relation to the other forms,—the necessity for taking a comprehensive survey of the whole as a condition to a satisfactory understanding of any part.

Stephen Forbes (1887)<sup>[23]</sup>

## Since 1900

Modern ecology is a young science that first attracted substantial scientific attention toward the end of the 19th century (around the same time that evolutionary studies were gaining scientific interest). In the early 20th century, ecology transitioned from a more descriptive form of natural history to a more analytical form of *scientific natural history*.<sup>[10][7]</sup> Frederic Clements published the first American ecology book in 1905,<sup>[24]</sup> presenting the idea of plant communities as a superorganism. This publication launched a debate between ecological holism and individualism that lasted until the 1970s. Clements' superorganism concept proposed that ecosystems progress through regular and determined stages of seral development that are analogous to the developmental stages of an organism. The Clementsian paradigm was challenged by Henry Gleason,<sup>[25]</sup> who stated that ecological communities develop from the unique and coincidental association of individual organisms. This perceptual shift placed the focus back onto the life histories of individual organisms and how this relates to the development of community associations.<sup>[26]</sup>

The Clementsian superorganism theory was an overextended application of an idealistic form of holism.<sup>[27][28]</sup> The term "holism" was coined in 1926 by Jan Christian Smuts, a South African general and polarizing historical figure who was inspired by Clements' superorganism concept.<sup>[29][C]</sup> Around the same time, Charles Elton pioneered the

concept of food chains in his classical book *Animal Ecology*.<sup>[30]</sup> Elton<sup>[30]</sup> defined ecological relations using concepts of food chains, food cycles, and food size, and described numerical relations among different functional groups and their relative abundance. Elton's 'food cycle' was replaced by 'food web' in a subsequent ecological text.<sup>[31]</sup> Alfred J. Lotka brought in many theoretical concepts applying thermodynamic principles to ecology. In 1942, Raymond Lindeman wrote a landmark paper on the trophic dynamics of ecology, which was published posthumously after initially being rejected for its theoretical emphasis. Trophic dynamics became the foundation for much of the work to follow on energy and material flow through ecosystems. Robert E. MacArthur advanced mathematical theory, predictions and tests in ecology in the 1950s, which inspired a resurgent school of theoretical mathematical ecologists.<sup>[10][32][33]</sup> Ecology also has developed through contributions from other nations, including Russia's Vladimir Vernadsky and his founding of the biosphere concept in the 1920s<sup>[34]</sup> and Japan's Kinji Imanishi and his concepts of harmony in nature and habitat segregation in the 1950s.<sup>[35]</sup> The scientific recognition of contributions to ecology from non-English-speaking cultures is hampered by language and translation barriers.<sup>[34]</sup>

This whole chain of poisoning, then, seems to rest on a base of minute plants which must have been the original concentrators. But what of the opposite end of the food chain—the human being who, in probable ignorance of all this sequence of events, has rigged his fishing tackle, caught a string of fish from the waters of Clear Lake, and taken them home to fry for his supper?

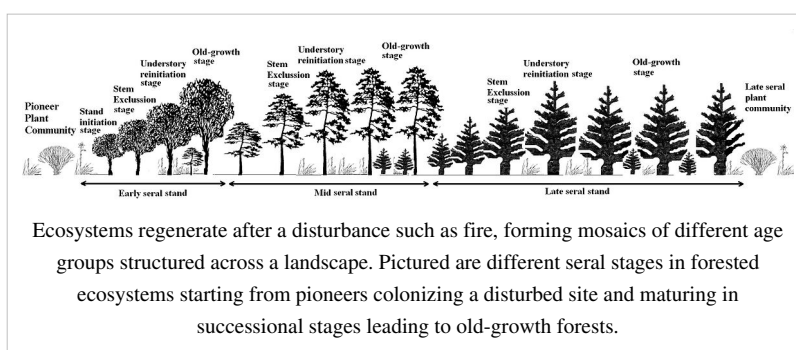
Rachel Carson (1962)<sup>[36]:48</sup>

Ecology surged in popular and scientific interest during the 1960–1970s environmental movement. There are strong historical and scientific ties between ecology, environmental management, and protection.<sup>[10]</sup> The historic emphasis and poetic naturalist writings for protection was on wild places, from notable ecologists in the history of conservation biology, such as Aldo Leopold and Arthur Tansley, were far removed from urban centres where the concentration of pollution and environmental degradation is located.<sup>[10][37]</sup> Palamar (2008)<sup>[37]</sup> notes an overshadowing by mainstream environmentalism of pioneering women in the early 1900s who fought for urban health ecology and brought about changes in environmental legislation. These women were precursors to the more popularized environmental movements after the 1950s. In 1962, marine biologist and ecologist Rachel Carson's book *Silent Spring* helped to mobilize the environmental movement by alerting the public to toxic pesticides, such as DDT, bioaccumulating in the environment. Carson used ecological science to link the release of environmental toxins to human and ecosystem health. More recently, ecologists have worked to bridge their understanding of the degradation of the planet's ecosystems with environmental politics, law, restoration, and natural resources management.<sup>[10][37][38][39]</sup>

## Integrative levels, scope, and scale of organization

The scope of ecology covers a wide array of interacting levels of organization spanning micro-level (e.g., cells) to planetary scale (e.g., ecosphere) phenomena. Ecosystems, for example, contain populations of individuals that aggregate into distinct ecological communities. It can take thousands of years for ecological processes to bring about the final

successional stages of a forest. An ecosystem's area can vary greatly, from tiny to vast. A single tree is of little consequence to the classification of a forest ecosystem, but critically relevant to organisms living in and on it.<sup>[40]</sup> Several generations of an aphid population can exist over the lifespan of a single leaf. Each of those aphids, in turn,



support diverse bacterial communities.<sup>[41]</sup> The nature of connections in ecological communities cannot be explained by knowing the details of each species in isolation, because the emergent pattern is neither revealed nor predicted until the ecosystem is studied as an integrated whole. Some ecological principles, however, do exhibit collective properties where the sum of the components explain the properties of the whole, such as birth rates of a population being equal to the sum of individual births over a designated time frame.<sup>[4]</sup>

## Hierarchical ecology

System behaviors must first be arrayed into levels of organization. Behaviors corresponding to higher levels occur at slow rates. Conversely, lower organizational levels exhibit rapid rates. For example, individual tree leaves respond rapidly to momentary changes in light intensity, CO<sub>2</sub> concentration, and the like. The growth of the tree responds more slowly and integrates these short-term changes.

O'Neill et al. (1986)<sup>[42]:76</sup>

The scale of ecological dynamics can operate like a closed island with respect to local site variables, such as aphids migrating on a tree, while at the same time remain open with regard to broader scale influences, such as atmosphere or climate. Hence, ecologists classify ecosystems hierarchically by analyzing data collected from finer scale units, such as vegetation associations, climate, and soil types, and integrate this information to identify emergent patterns of uniform organization and processes that operate on local to regional, landscape, and chronological scales.

To structure the study of ecology into a conceptually manageable framework, the biological world is organized into a nested hierarchy, ranging in scale from genes, to cells, to tissues, to organs, to organisms, to species, and up to the level of the biosphere.<sup>[43]</sup> This framework forms a panarchy<sup>[44]</sup> and exhibits non-linear behaviours; this means that "effect and cause are disproportionate, so that small changes in critical variables, such as the numbers of nitrogen fixers, can lead to disproportionate, perhaps irreversible, changes in the system properties."<sup>[45]:14</sup>

## Biodiversity

Biodiversity is the variety of life and its processes. It includes the variety of living organisms, the genetic differences among them, the communities and ecosystems in which they occur, and the ecological and evolutionary processes that keep them functioning, yet ever changing and adapting.

Noss & Carpenter (1994)<sup>[46]:5</sup>

Biodiversity (an abbreviation of "biological diversity") describes the diversity of life from genes to ecosystems and spans every level of biological organization. The term has several interpretations, and there are many ways to index, measure, characterize, and represent its complex organization.<sup>[47][48]</sup> Biodiversity includes species diversity, ecosystem diversity, genetic diversity and the complex processes operating at and among these respective levels.<sup>[48][49][50]</sup> Biodiversity plays an important role in ecological health as much as it does for human health.<sup>[51][52]</sup>

Preventing species extinctions is one way to preserve biodiversity, but factors such as genetic diversity and migration routes are equally important and are threatened on global scales. Conservation priorities and management techniques require different approaches and considerations to address the full ecological scope of biodiversity. Populations and species migration, for example, are sensitive indicators of ecosystem services that sustain and contribute natural capital toward the well-being of humanity.<sup>[53][54][55][56]</sup> An understanding of biodiversity has practical application for ecosystem-based conservation planners as they make ecologically responsible decisions in management recommendations to consultant firms, governments, and industry.<sup>[38]</sup>



## Habitat

The habitat of a species describes the environment over which a species is known to occur and the type of community that is formed as a result.<sup>[57]</sup> More specifically, "habitats can be defined as regions in environmental space that are composed of multiple dimensions, each representing a biotic or abiotic environmental variable; that is, any component or characteristic of the environment related directly (e.g. forage biomass and quality) or indirectly (e.g. elevation) to the use of a location by the animal."<sup>[58]:745</sup> For example, a habitat might be an aquatic or terrestrial environment that can be further categorized as a montane or alpine ecosystem. Habitat shifts provide important evidence of competition in nature where one population changes relative to the habitats that most other individuals of the species occupy. For example, one population of a species of tropical lizards (*Tropidurus hispidus*) has a flattened body relative to the main populations that live in open savanna. The population that lives in an isolated rock outcrop hides in crevasses where its flattened body offers a selective advantage. Habitat shifts also occur in the developmental life history of amphibians and in insects that transition from aquatic to terrestrial habitats. Biotope and habitat are sometimes used interchangeably, but the former applies to a community's environment, whereas the latter applies to a species' environment.<sup>[57][59][60]</sup>



Biodiversity of a coral reef. Corals adapt and modify their environment by forming calcium carbonate skeletons. This provides growing conditions for future generations and forms a habitat for many other species.<sup>[61]</sup>

## Niche

Definitions of the niche date back to 1917,<sup>[64]</sup> but G. Evelyn Hutchinson made conceptual advances in 1957<sup>[65][66]</sup> by introducing a widely adopted definition: "the set of biotic and abiotic conditions in which a species is able to persist and maintain stable population sizes."<sup>[64]:519</sup> The ecological niche is a central concept in the ecology of organisms and is sub-divided into the *fundamental* and the *realized* niche. The fundamental niche is the set of environmental conditions under which a species is able to persist. The realized niche is the set of environmental plus ecological conditions under which a species persists.<sup>[64][66][67]</sup> The Hutchinsonian niche is defined more technically as a "Euclidean hyperspace whose *dimensions* are defined as environmental variables and whose *size* is a function of the number of values that the environmental values may assume for which an organism has *positive fitness*."<sup>[68]:71</sup>

Biogeographical patterns and range distributions are explained or predicted through knowledge of a species' traits and niche requirements.<sup>[69]</sup> Species have functional traits that are uniquely adapted to the ecological niche. A trait is a measurable property, phenotype, or characteristic of an organism that may influence its survival. Genes play an important role in the interplay of development and environmental expression of traits.<sup>[28]</sup> Resident species evolve traits that are fitted to the selection pressures of their local environment. This tends to afford them a competitive advantage and discourages similarly adapted species from having an overlapping geographic range. The competitive exclusion principle states that two species cannot coexist indefinitely by living off the same limiting resource; one will always outcompete the other. When similarly adapted species overlap geographically, closer inspection reveals subtle ecological differences in their habitat or dietary requirements.<sup>[70]</sup> Some models and empirical studies, however, suggest that disturbances can stabilize the coevolution and shared niche occupancy of similar species inhabiting species-rich communities.<sup>[71]</sup> The habitat plus the niche is called the ecotop, which is defined as the full range of environmental and biological variables affecting an entire species.<sup>[57]</sup>



Termite mounds with varied heights of chimneys regulate gas exchange, temperature and other environmental parameters that are needed to sustain the internal physiology of the entire colony.<sup>[62][63]</sup>

## Niche construction

Organisms are subject to environmental pressures, but they also modify their habitats. The regulatory feedback between organisms and their environment can affect conditions from local (e.g., a beaver pond) to global scales, over time and even after death, such as decaying logs or silica skeleton deposits from marine organisms.<sup>[72]</sup> The process and concept of ecosystem engineering has also been called niche construction. Ecosystem engineers are defined as: "organisms that directly or indirectly modulate the availability of resources to other species, by causing physical state changes in biotic or abiotic materials. In so doing they modify, maintain and create habitats."<sup>[73]:373</sup>

The ecosystem engineering concept has stimulated a new appreciation for the influence that organisms have on the ecosystem and evolutionary process. The term "niche construction" is more often used in reference to the under-appreciated feedback mechanism of natural selection imparting forces on the abiotic niche.<sup>[62][74]</sup> An example of natural selection through ecosystem engineering occurs in the nests of social insects, including ants, bees, wasps, and termites. There is an emergent homeostasis or homeorhesis in the structure of the nest that regulates, maintains and defends the physiology of the entire colony. Termite mounds, for example, maintain a constant internal temperature through the design of air-conditioning chimneys. The structure of the nests themselves are subject to the

forces of natural selection. Moreover, a nest can survive over successive generations, so that progeny inherit both genetic material and a legacy niche that was constructed before their time.<sup>[4][62][63]</sup>

## Biome

Biomes are larger units of organization that categorize regions of the Earth's ecosystems, mainly according to the structure and composition of vegetation.<sup>[75]</sup> There are different methods to define the continental boundaries of biomes dominated by different functional types of vegetative communities that are limited in distribution by climate, precipitation, weather and other environmental variables. Biomes include tropical rainforest, temperate broadleaf and mixed forest, temperate deciduous forest, taiga, tundra, hot desert, and polar desert.<sup>[76]</sup> Other researchers have recently categorized other biomes, such as the human and oceanic microbiomes. To a microbe, the human body is a habitat and a landscape.<sup>[77]</sup> Microbiomes were discovered largely through advances in molecular genetics, which have revealed a hidden richness of microbial diversity on the planet. The oceanic microbiome plays a significant role in the ecological biogeochemistry of the planet's oceans.<sup>[78]</sup>

## Biosphere

The largest scale of ecological organization is the biosphere: the total sum of ecosystems on the planet. Ecological relationships regulate the flux of energy, nutrients, and climate all the way up to the planetary scale. For example, the dynamic history of the planetary atmosphere's CO<sub>2</sub> and O<sub>2</sub> composition has been affected by the biogenic flux of gases coming from respiration and photosynthesis, with levels fluctuating over time in relation to the ecology and evolution of plants and animals.<sup>[79]</sup> Ecological theory has also been used to explain self-emergent regulatory phenomena at the planetary scale: for example, the Gaia hypothesis is an example of holism applied in ecological theory.<sup>[80]</sup> The Gaia hypothesis states that there is an emergent feedback loop generated by the metabolism of living organisms that maintains the temperature of the Earth and atmospheric conditions within a narrow self-regulating range of tolerance.<sup>[81]</sup>

## Population ecology

Population ecology studies the dynamics of species populations and how these populations interact with the environment.<sup>[4]</sup> A population consists of individuals of the same species that live, interact and migrate through the same niche and habitat.<sup>[82]</sup>

A primary law of population ecology is the Malthusian growth model<sup>[83]</sup> which states, "a population will grow (or decline) exponentially as long as the environment experienced by all individuals in the population remains constant."<sup>[83]:18</sup> Simplified population models usually start with four variables: death, birth, immigration, and emigration.

An example of an introductory population model describes a closed population, such as on an island, where immigration and emigration does not take place. Hypotheses are evaluated with reference to a null hypothesis which states that random processes create the observed data. In these island models, the rate of population change is described by:

$$\frac{dN}{dT} = B - D = bN - dN = (b - d)N = rN,$$

where  $N$  is the total number of individuals in the population,  $B$  is the number of births,  $D$  is the number of deaths,  $b$  and  $d$  are the per capita rates of birth and death respectively, and  $r$  is the per capita rate of population change. The formula states that the rate of change in population size ( $dN/dT$ ) is equal to births minus deaths ( $B - D$ ).<sup>[83][84]</sup>

Using these modelling techniques, Malthus' population principle of growth was later transformed into a model known as the logistic equation:

$$\frac{dN}{dT} = aN \left(1 - \frac{N}{K}\right),$$

where  $N$  is the number of individuals measured as biomass density,  $a$  is the maximum per-capita rate of change, and  $K$  is the carrying capacity of the population. The formula states that the rate of change in population size ( $dN/dT$ ) is equal to growth ( $aN$ ) that is limited by carrying capacity ( $1 - N/K$ ).

Population ecology builds upon these introductory models to further understand demographic processes in real study populations. Commonly used types of data include life history, fecundity, and survivorship, and these are analysed using mathematical techniques such as matrix algebra. The information is used for managing wildlife stocks and setting harvest quotas.<sup>[84][85]</sup> In cases when the use of null hypotheses is not appropriate, ecologists may adopt different kinds of statistical methods, such as the Akaike information criterion,<sup>[86]</sup> or use models that can become mathematically complex as "several competing hypotheses are simultaneously confronted with the data."<sup>[87]</sup>

### Metapopulations and migration

The concept of metapopulations was defined in 1969<sup>[88]</sup> as "a population of populations which go extinct locally and recolonize."<sup>[89]:105</sup> Metapopulation ecology is another statistical approach that is often used in conservation research.<sup>[90]</sup> Metapopulation models simplify the landscape into patches of varying levels of quality,<sup>[91]</sup> and metapopulations are linked by the migratory behaviours of organisms. Animal migration is set apart from other kinds of movement because it involves the seasonal departure and return of individuals from a habitat.<sup>[92]</sup> Migration is also a population-level phenomenon, as with the migration routes followed by plants as they occupied northern post-glacial environments. Plant ecologists use pollen records that accumulate and stratify in wetlands to reconstruct the timing of plant migration and dispersal relative to historic and contemporary climates. These migration routes involved an expansion of the range as plant populations expanded from one area to another. There is a larger taxonomy of movement, such as commuting, foraging, territorial behaviour, stasis, and ranging. Dispersal is usually distinguished from migration because it involves the one way permanent movement of individuals from their birth population into another population.<sup>[93][94]</sup>

In metapopulation terminology, migrating individuals are classed as emigrants (when they leave a region) or immigrants (when they enter a region), and sites are classed either as sources or sinks. A site is a generic term that refers to places where ecologists sample populations, such as ponds or defined sampling areas in a forest. Source patches are productive sites that generate a seasonal supply of juveniles that migrate to other patch locations. Sink patches are unproductive sites that only receive migrants and will go extinct unless rescued by an adjacent source patch or environmental conditions become more favourable. Metapopulation models examine patch dynamics over time to answer questions about spatial and demographic ecology. The ecology of metapopulations is a dynamic process of extinction and colonization. Small patches of lower quality (i.e., sinks) are maintained or rescued by a seasonal influx of new immigrants. A dynamic metapopulation structure evolves from year to year, where some patches are sinks in dry years and are sources when conditions are more favourable. Ecologists use a mixture of computer models and field studies to explain metapopulation structure.<sup>[95][96]</sup>

### Community ecology

Community ecology examines how interactions among species and their environment affect the abundance, distribution and diversity of species within communities.

Johnson & Stinchcomb (2007)<sup>[97]:250</sup>

Community ecology is the study of the interactions among a collections of species that inhabit the same geographic area. Research in community ecology might measure primary production in a wetland in relation to decomposition and consumption rates. This requires an understanding of the community connections between plants (i.e., primary producers) and the decomposers (e.g., fungi and bacteria),<sup>[98]</sup>

or the analysis of predator-prey dynamics affecting amphibian biomass.<sup>[99]</sup> Food webs and trophic levels are two widely employed conceptual models used to explain the linkages among species.<sup>[4]</sup>



Interspecific interactions such as predation are a key aspect of community ecology.

## Ecosystem ecology

These ecosystems, as we may call them, are of the most various kinds and sizes. They form one category of the multitudinous physical systems of the universe, which range from the universe as a whole down to the atom.

Tansley (1935)<sup>[100]:299</sup>

Ecosystems are habitats within biomes that form an integrated whole and a dynamically responsive system having both physical and biological complexes. The underlying concept can be traced back to 1864 in the published work of George Perkins Marsh ("Man and Nature").<sup>[101][102]</sup> Within an ecosystem, organisms are linked to the physical and biological components of their environment to which they are adapted.<sup>[100]</sup> Ecosystems are complex adaptive systems where the interaction of life processes form self-organizing patterns across different scales of time and space.<sup>[103]</sup> Ecosystems are broadly categorized as terrestrial, freshwater, atmospheric, or marine.



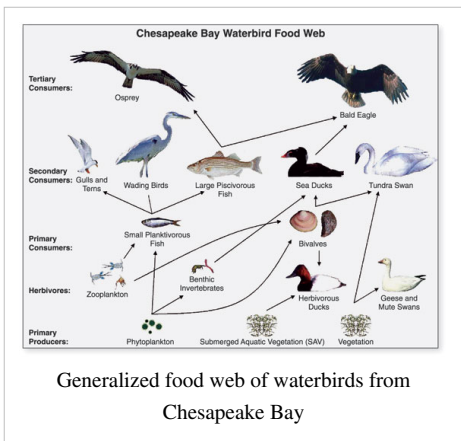
A riparian forest in the White Mountains, New Hampshire (USA), an example of ecosystem ecology

Differences stem from the nature of the unique physical environments that shapes the biodiversity within each. A more recent addition to ecosystem ecology are technoecosystems, which are affected by or primarily the result of human activity.<sup>[4]</sup>

## Food webs

A food web is the archetypal ecological network. Plants capture solar energy and use it to synthesize simple sugars during photosynthesis. As plants grow, they accumulate nutrients and are eaten by grazing herbivores, and the energy is transferred through a chain of organisms by consumption. The simplified linear feeding pathways that move from a basal trophic species to a top consumer is called the food chain. The larger interlocking pattern of food chains in an ecological community creates a complex food web. Food webs are a type of concept map or a heuristic device that is used to illustrate and study pathways of energy and material flows.<sup>[42][104][105]</sup>



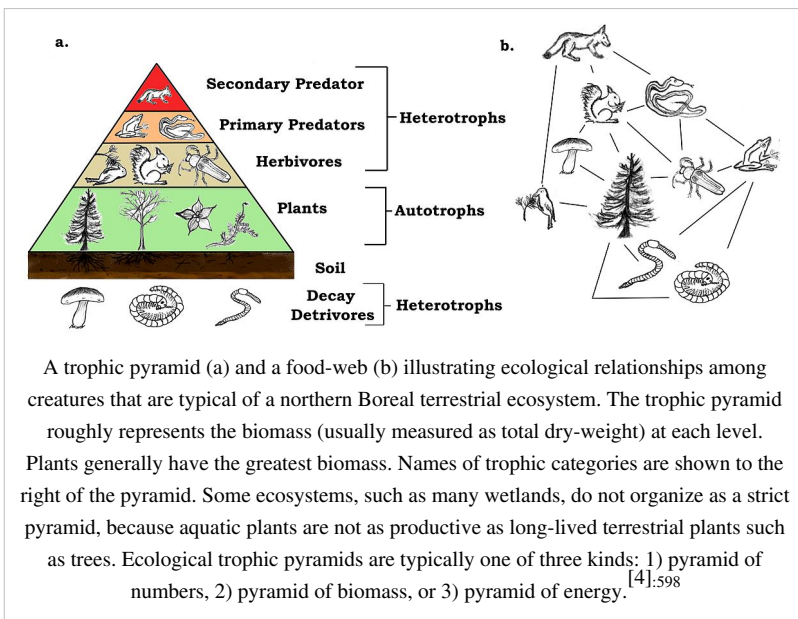


Food webs are often limited relative to the real world. Complete empirical measurements are generally restricted to a specific habitat, such as a cave or a pond, and principles gleaned from food web microcosm studies are extrapolated to larger systems.<sup>[106]</sup> Feeding relations require extensive investigations into the gut contents of organisms, which can be difficult to decipher, or (more recently) stable isotopes can be used to trace the flow of nutrient diets and energy through a food web.<sup>[107]</sup> Despite these limitations, food webs remain a valuable tool in understanding community ecosystems.<sup>[108]</sup>

Generalized food web of waterbirds from  
Chesapeake Bay

Food webs exhibit principles of ecological emergence through the nature of trophic relationships: some species have many weak feeding links (e.g., omnivores) while some are more specialized with fewer stronger feeding links (e.g., primary predators). Theoretical and empirical studies identify non-random emergent patterns of few strong and many weak linkages that explain how ecological communities remain stable over time.<sup>[109]</sup> Food webs are composed of subgroups where members in a community are linked by strong interactions, and the weak interactions occur between these subgroups. This increases food web stability.<sup>[110]</sup> Step by step lines or relations are drawn until a web of life is illustrated.<sup>[105][111][112][113]</sup>

### Trophic levels



A trophic level (from Greek *troph*, τροφή, trophē, meaning "food" or "feeding") is "a group of organisms acquiring a considerable majority of its energy from the adjacent level nearer the abiotic source."<sup>[114]:383</sup> Links in food webs primarily connect feeding relations or trophism among species. Biodiversity within ecosystems can be organized into trophic pyramids, in which the vertical dimension represents feeding relations that become further removed from the base of the food chain up toward top predators, and the horizontal dimension represents the abundance or

biomass at each level.<sup>[115]</sup> When the relative abundance or biomass of each species is sorted into its respective trophic level, they naturally sort into a 'pyramid of numbers'.<sup>[30]</sup>

Species are broadly categorized as autotrophs (or primary producers), heterotrophs (or consumers), and detritivores (or decomposers). Autotrophs are organisms that produce their own food (production is greater than respiration) by photosynthesis or chemosynthesis. Heterotrophs are organisms that must feed on others for nourishment and energy (respiration exceeds production).<sup>[4]</sup> Heterotrophs can be further sub-divided into different functional groups, including primary consumers (strict herbivores), secondary consumers (carnivorous predators that feed exclusively on herbivores) and tertiary consumers (predators that feed on a mix of herbivores and predators).<sup>[116]</sup> Omnivores do not fit neatly into a functional category because they eat both plant and animal tissues. It has been suggested that



omnivores have a greater functional influence as predators, because compared to herbivores they are relatively inefficient at grazing.<sup>[117]</sup>

Trophic levels are part of the holistic or complex systems view of ecosystems.<sup>[118][119]</sup> Each trophic level contains unrelated species that are grouped together because they share common ecological functions, giving a macroscopic view of the system.<sup>[120]</sup> While the notion of trophic levels provides insight into energy flow and top-down control within food webs, it is troubled by the prevalence of omnivory in real ecosystems. This has led some ecologists to "reiterate that the notion that species clearly aggregate into discrete, homogeneous trophic levels is fiction."<sup>[121]:815</sup> Nonetheless, recent studies have shown that real trophic levels do exist, but "above the herbivore trophic level, food webs are better characterized as a tangled web of omnivores."<sup>[122]:612</sup>

### Keystone species

A keystone species is a species that is connected to a disproportionately large number of other species in the food-web. Keystone species have lower levels of biomass in the trophic pyramid relative to the importance of their role. The many connections that a keystone species holds means that it maintains the organization and structure of entire communities. The loss of a keystone species results in a range of dramatic cascading effects that alters trophic dynamics, other food web connections, and can cause the extinction of other species.<sup>[123][124]</sup>

Sea otters (*Enhydra lutris*) are commonly cited as an example of a keystone species because they limit the density of sea urchins that feed on kelp. If sea otters are removed from the system, the urchins graze until the kelp beds disappear and this has a dramatic effect on community structure.<sup>[125]</sup> Hunting of sea otters, for example, is thought to have indirectly led to the extinction of the Steller's Sea Cow (*Hydrodamalis gigas*).<sup>[126]</sup> While the keystone species concept has been used extensively as a conservation tool, it has been criticized for being poorly defined from an operational stance. It is difficult to experimentally determine what species may hold a keystone role in each ecosystem. Furthermore, food web theory suggests that keystone species may not be common, so it is unclear how generally the keystone species model can be applied.<sup>[125][127]</sup>

### Ecological complexity

Complexity is easily understood as a large computational effort needed to piece together numerous interacting parts exceeding the iterative memory capacity of the human mind. Global patterns of biological diversity are complex. This biocomplexity stems from the interplay among ecological processes that operate and influence patterns at different scales that grade into each other, such as transitional areas or ecotones spanning landscapes. Complexity stems from the interplay among levels of biological organization as energy and matter is integrated into larger units that superimpose onto the smaller parts. "What were wholes on one level become parts on a higher one."<sup>[128]:209</sup> Small scale patterns do not necessarily explain large scale phenomena, otherwise captured in the expression (coined by Aristotle) 'the sum is greater than the parts'.<sup>[129][130][E]</sup>

"Complexity in ecology is of at least six distinct types: spatial, temporal, structural, process, behavioral, and geometric."<sup>[131]:3</sup> From these principles, ecologists have identified emergent and self-organizing phenomena that operate at different environmental scales of influence, ranging from molecular to planetary, and these require different explanations at each integrative level.<sup>[81][132]</sup> Ecological complexity relates to the dynamic resilience of ecosystems that transition to multiple shifting steady-states directed by random fluctuations of history.<sup>[44][133]</sup> Long-term ecological studies provide important track records to better understand the complexity and resilience of ecosystems over longer temporal and broader spatial scales. The International Long Term Ecological Network<sup>[134]</sup> manages and exchanges scientific information among research sites. The longest experiment in existence is the Park Grass Experiment, which was initiated in 1856.<sup>[135]</sup> Another example is the Hubbard Brook study, which has been in operation since 1960.<sup>[136]</sup>

## Holism

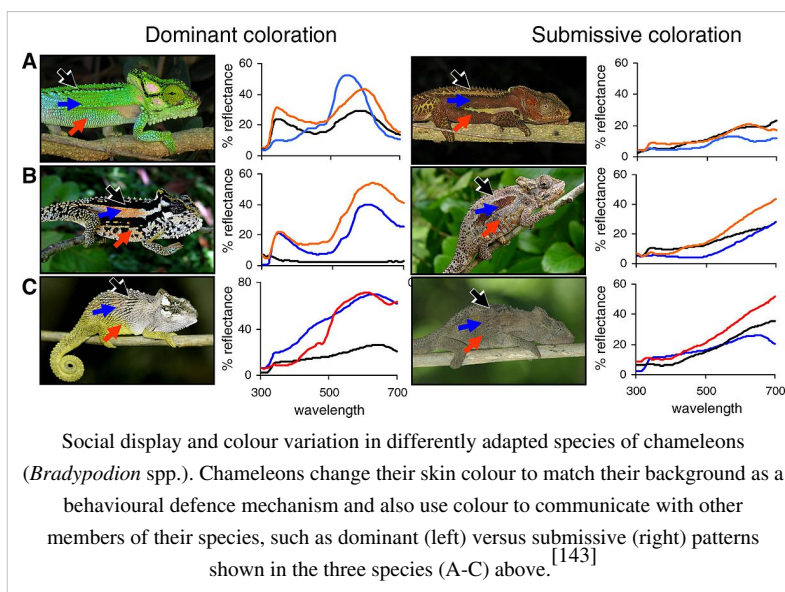
Holism remains a critical part of the theoretical foundation in contemporary ecological studies. Holism addresses the biological organization of life that self-organizes into layers of emergent whole systems that function according to nonreducible properties. This means that higher order patterns of a whole functional system, such as an ecosystem, cannot be predicted or understood by a simple summation of the parts.<sup>[137]</sup> "New properties emerge because the components interact, not because the basic nature of the components is changed."<sup>[4]:8</sup>

Ecological studies are necessarily holistic as opposed to reductionistic.<sup>[28][132][138]</sup> Holism has three scientific meanings or uses that identify with ecology: 1) the mechanistic complexity of ecosystems, 2) the practical description of patterns in quantitative reductionist terms where correlations may be identified but nothing is understood about the causal relations without reference to the whole system, which leads to 3) a metaphysical hierarchy whereby the causal relations of larger systems are understood without reference to the smaller parts. Scientific holism differs from mysticism that has appropriated the same term. An example of metaphysical holism is identified in the trend of increased exterior thickness in shells of different species. The reason for a thickness increase can be understood through reference to principles of natural selection via predation without need to reference or understand the biomolecular properties of the exterior shells.<sup>[27]</sup>

## Relation to evolution

Ecology and evolution are considered sister disciplines of the life sciences. Natural selection, life history, development, adaptation, populations, and inheritance are examples of concepts that thread equally into ecological and evolutionary theory. Morphological, behavioural and genetic traits, for example, can be mapped onto evolutionary trees to study the historical development of a species in relation to their functions and roles in different ecological circumstances. In this framework, the analytical tools of ecologists and evolutionists overlap as they organize, classify and investigate life through common systematic principals, such as phylogenetics or the Linnaean system of taxonomy.<sup>[139]</sup> The two disciplines often appear together, such as in the title of the journal *Trends in Ecology and Evolution*.<sup>[140]</sup> There is no sharp boundary separating ecology from evolution and they differ more in their areas of applied focus. Both disciplines discover and explain emergent and unique properties and processes operating across different spatial or temporal scales of organization.<sup>[28][81]</sup> While the boundary between ecology and evolution is not always clear, it is understood that ecologists study the abiotic and biotic factors that influence the evolutionary process.<sup>[141][142]</sup>

## Behavioural ecology



All organisms are motile to some extent. Even plants express complex behaviour, including memory and communication.<sup>[144]</sup>

Behavioural ecology is the study of an organism's behaviour in its environment and its ecological and evolutionary implications. Ethology is the study of observable movement or behaviour in animals. This could include investigations of motile sperm of plants, mobile phytoplankton, zooplankton swimming toward the female egg, the cultivation of fungi by weevils, the mating dance of a salamander, or social gatherings of

amoeba.<sup>[145][146][147][148][149]</sup>

Adaptation is the central unifying concept in behavioural ecology.<sup>[150]</sup> Behaviours can be recorded as traits and inherited in much the same way that eye and hair colour can. Behaviours can evolve by means of natural selection as adaptive traits conferring functional utilities that increases reproductive fitness.<sup>[151][152]</sup>

Predator-prey interactions are an introductory concept into food-web studies as well as behavioural ecology.<sup>[153]</sup> Prey species can exhibit different kinds of behavioural adaptations to predators, such as avoid, flee or defend. Many prey species are faced with multiple predators that differ in the degree of danger posed. To be adapted to their environment and face predatory threats, organisms must balance their energy budgets as they invest in different aspects of their life history, such as growth, feeding, mating, socializing, or modifying their habitat. Hypotheses posited in behavioural ecology are generally based on adaptive principles of conservation, optimization or efficiency.<sup>[67][141][154]</sup> For example, "[t]he threat-sensitive predator avoidance hypothesis predicts that prey should assess the degree of threat posed by different predators and match their behaviour according to current levels of risk"<sup>[155]</sup> or "[t]he optimal flight initiation distance occurs where expected postencounter fitness is maximized, which depends on the prey's initial fitness, benefits obtainable by not fleeing, energetic escape costs, and expected fitness loss due to predation risk."<sup>[156]</sup>

Elaborate sexual displays and posturing are encountered in the behavioural ecology of animals. The birds of paradise, for example, sing and display elaborate ornaments during courtship. These displays serve a dual purpose of signalling healthy or well-adapted individuals and desirable genes. The displays are driven by sexual selection as an advertisement of quality of traits among male suitors.<sup>[158]</sup>

## Social ecology

Social ecological behaviours are notable in the social insects, slime moulds, social spiders, human society, and naked mole rats where eusocialism has evolved. Social behaviours include reciprocally beneficial behaviours among kin and nest mates<sup>[147][152][159]</sup> and evolve from kin and group selection. Kin selection explains altruism through genetic relationships, whereby an altruistic behaviour leading to death is rewarded by the survival of genetic copies distributed among surviving relatives. The social insects, including ants, bees and wasps are most famously studied for this type of relationship because the male drones are clones that share the same genetic make-up as every other male in the colony.<sup>[152]</sup> In contrast, group selectionists find examples of altruism among non-genetic relatives and explain this through selection acting on the group, whereby it becomes selectively advantageous for groups if their members express altruistic behaviours to one another. Groups with predominantly altruistic members beat groups with predominantly selfish members.<sup>[152][160]</sup>



**Symbiosis:** Leafhoppers (*Eurymela fenestrata*) are protected by ants (*Iridomyrmex purpureus*) in a symbiotic relationship. The ants protect the leafhoppers from predators and in return the leafhoppers feeding on plants exude honeydew from their anus that provides energy and nutrients to tending ants.<sup>[157]</sup>

## Coevolution

Ecological interactions can be classified broadly into a host and an associate relationship. A host is any entity that harbours another that is called the associate.<sup>[161]</sup> Relationships within a species that are mutually or reciprocally beneficial are called mutualisms. Examples of mutualism include fungus-growing ants employing agricultural symbiosis, bacteria living in the guts of insects and other organisms, the fig wasp and yucca moth pollination complex, lichens with fungi and photosynthetic algae, and corals with photosynthetic algae.<sup>[162][163]</sup> If there is a physical connection between host and associate, the relationship is called symbiosis. Approximately 60% of all plants, for example, have a symbiotic relationship with arbuscular mycorrhizal fungi living in their roots forming an exchange network of carbohydrates for mineral nutrients.<sup>[164]</sup>

Indirect mutualisms occur where the organisms live apart. For example, trees living in the equatorial regions of the planet supply oxygen into the atmosphere that sustains species living in distant polar regions of the planet. This relationship is called commensalism because many others receive the benefits of clean air at no cost or harm to trees supplying the oxygen.<sup>[4][165]</sup> If the associate benefits while the host suffers, the relationship is called parasitism. Although parasites impose a cost to their host (e.g., via damage to their reproductive organs or propagules, denying the services of a beneficial partner), their net effect on host fitness is not necessarily negative and, thus, becomes difficult to forecast.<sup>[166][167]</sup> Coevolution is also driven by competition among species or among members of the same species under the banner of reciprocal antagonism, such as grasses competing for growth space. The Red Queen Hypothesis, for example, posits that parasites track down and specialize on the locally common genetic defence systems of its host that drives the evolution of sexual reproduction to diversify the genetic constituency of populations responding to the antagonistic pressure.<sup>[168][169]</sup>

## Biogeography

Biogeography (an amalgamation of *biology* and *geography*) is the comparative study of the geographic distribution of organisms and the corresponding evolution of their traits in space and time.<sup>[170]</sup> The *Journal of Biogeography* was established in 1974.<sup>[171]</sup> Biogeography and ecology share many of their disciplinary roots. For example, the theory of island biogeography, published by the mathematician Robert MacArthur and ecologist Edward O. Wilson in 1967<sup>[172]</sup> is considered one of the fundamentals of ecological theory.<sup>[173]</sup>

Biogeography has a long history in the natural sciences concerning the spatial distribution of plants and animals. Ecology and evolution provide the explanatory context for biogeographical studies.<sup>[170]</sup> Biogeographical patterns result from ecological processes that influence range distributions, such as migration and dispersal.<sup>[173]</sup> and from historical processes

that split populations or species into different areas. The biogeographic processes that result in the natural splitting of species explains much of the modern distribution of the Earth's biota. The splitting of lineages in a species is called vicariance biogeography and it is a sub-discipline of biogeography.<sup>[174]</sup> There are also practical applications in the field of biogeography concerning ecological systems and processes. For example, the range and distribution of biodiversity and invasive species responding to climate change is a serious concern and active area of research in the context of global warming.<sup>[54][175]</sup>

### *r/K*-Selection theory

A population ecology concept is *r/K* selection theory,<sup>[D]</sup> one of the first predictive models in ecology used to explain life-history evolution. The premise behind the *r/K* selection model is that natural selection pressures change according to population density. For example, when an island is first colonized, density of individuals is low. The initial increase in population size is not limited by competition, leaving an abundance of available resources for rapid population growth. These early phases of population growth experience *density-independent* forces of natural selection, which is called *r*-selection. As the population becomes more crowded, it approaches the island's carrying capacity, thus forcing individuals to compete more heavily for fewer available resources. Under crowded conditions, the population experiences density-dependent forces of natural selection, called *K*-selection.<sup>[176]</sup>

In the *r/K*-selection model, the first variable *r* is the intrinsic rate of natural increase in population size and the second variable *K* is the carrying capacity of a population.<sup>[67]</sup> Different species evolve different life-history strategies spanning a continuum between these two selective forces. An *r*-selected species is one that has high birth rates, low levels of parental investment, and high rates of mortality before individuals reach maturity. Evolution favours high rates of fecundity in *r*-selected species. Many kinds of insects and invasive species exhibit *r*-selected characteristics. In contrast, a *K*-selected species has low rates of fecundity, high levels of parental investment in the young, and low



**Parasitism:** A harvestman arachnid being parasitized by mites. The harvestman is being consumed, while the mites benefit from traveling on and feeding off of their host.



rates of mortality as individuals mature. Humans and elephants are examples of species exhibiting *K*-selected characteristics, including longevity and efficiency in the conversion of more resources into fewer offspring.<sup>[172][177]</sup>

## Molecular ecology

The important relationship between ecology and genetic inheritance predates modern techniques for molecular analysis. Molecular ecological research became more feasible with the development of rapid and accessible genetic technologies, such as the polymerase chain reaction (PCR). The rise of molecular technologies and influx of research questions into this new ecological field resulted in the publication *Molecular Ecology* in 1992.<sup>[178]</sup> Molecular ecology uses various analytical techniques to study genes in an evolutionary and ecological context. In 1994, John Avise also played a leading role in this area of science with the publication of his book, *Molecular Markers, Natural History and Evolution*.<sup>[179]</sup> Newer technologies opened a wave of genetic analysis into organisms once difficult to study from an ecological or evolutionary standpoint, such as bacteria, fungi and nematodes. Molecular ecology engendered a new research paradigm for investigating ecological questions considered otherwise intractable. Molecular investigations revealed previously obscured details in the tiny intricacies of nature and improved resolution into probing questions about behavioural and biogeographical ecology.<sup>[179]</sup> For example, molecular ecology revealed promiscuous sexual behaviour and multiple male partners in tree swallows previously thought to be socially monogamous.<sup>[180]</sup> In a biogeographical context, the marriage between genetics, ecology and evolution resulted in a new sub-discipline called phylogeography.<sup>[181]</sup>

## Human ecology

Perhaps the most important implication involves our view of human society. *Homo sapiens* is not an external disturbance, it is a keystone species within the system. In the long term, it may not be the magnitude of extracted goods and services that will determine sustainability. It may well be our disruption of ecological recovery and stability mechanisms that determines system collapse.

O'Neil (2001)<sup>[102]:3282</sup>

Ecology is as much a biological science as it is a human science.<sup>[4]</sup> Human ecology is an interdisciplinary investigation into the ecology of our species. "Human ecology may be defined: (1) from a bio-ecological standpoint as the study of man as the ecological dominant in plant and animal communities and systems; (2) from a bio-ecological standpoint as simply another animal affecting and being affected by his physical environment; and (3) as a human being, somehow different from animal life in general, interacting with physical and modified environments in a distinctive and creative way. A truly interdisciplinary human ecology will most likely address itself to all three."<sup>[182]:3</sup> The term was formally introduced in 1921, but many sociologists, geographers, psychologists, and other disciplines were interested in human relations to natural systems centuries prior, especially in the late 19th century.<sup>[182][183]</sup>

The ecological complexities human beings are facing through the technological transformation of the planetary biome has brought on the Anthropocene. The unique set of circumstances has generated the need for a new unifying science called coupled human and natural systems that builds upon, but moves beyond the field of human ecology.<sup>[137]</sup> Ecosystems tie into human societies through the critical and all encompassing life-supporting functions they sustain. In recognition of these functions and the incapability of traditional economic valuation methods to see the value in ecosystems, there has been a surge of interest in social-natural capital, which provides the means to put a value on the stock and use of information and materials stemming from ecosystem goods and services. Ecosystems produce, regulate, maintain, and supply services of critical necessity and beneficial to human health (cognitive and physiological), economies, and they even provide an information or reference function as a living library giving opportunities for science and cognitive development in children engaged in the complexity of the natural world. Ecosystems relate importantly to human ecology as they are the ultimate base foundation of global economics as every commodity and the capacity for exchange ultimately stems from the ecosystems on Earth.<sup>[184][185][186][137]</sup>



## Restoration and management

Ecosystem management is not just about science nor is it simply an extension of traditional resource management; it offers a fundamental reframing of how humans may work with nature.

Grumbine (1994)<sup>[187]:27</sup>

Ecology is an employed science of restoration, repairing disturbed sites through human intervention, in natural resource management, and in environmental impact assessments. Edward O. Wilson predicted in 1992 that this 21st century "will be the era of restoration in ecology".<sup>[188]</sup> Ecological science has boomed in the industrial investment of restoring ecosystems and their processes in abandoned sites after disturbance. Natural resource managers, in forestry, for example, employ ecologists to develop, adapt, and implement ecosystem based methods into the planning, operation, and restoration phases of land-use. Ecological science is used in the methods of sustainable harvesting, disease and fire outbreak management, in fisheries stock management, for integrating land-use with protected areas and communities, and conservation in complex geo-political landscapes.<sup>[187][38][189][187][190]</sup>

## Relation to the environment

The environment of ecosystems includes both physical parameters and biotic attributes. It is dynamically interlinked, imposed upon, modified, contextualized, and constraining while at the same time containing resources for organisms at any time throughout their life cycle.<sup>[4][191]</sup> Like "ecology," the term "environment" has different conceptual meanings and overlaps with the concept of "nature." Environment "...includes the physical world, the social world of human relations and the built world of human creation."<sup>[192]:62</sup> The physical environment is external to the level of biological organization under investigation, including abiotic factors such as temperature, radiation, light, chemistry, climate and geology. The biotic environment includes genes, cells, organisms, members of the same species (conspecifics) and other species that share a habitat.<sup>[193]</sup>

The distinction between external and internal environments, however, is an abstraction parsing life and environment into units or facts that are inseparable in reality. There is an interpenetration of cause and effect between the environment and life. The laws of thermodynamics, for example, apply to ecology by means of its physical state. With an understanding of metabolic and thermodynamic principles, a complete accounting of energy and material flow can be traced through an ecosystem. In this way, the environmental and ecological relations are studied through reference to conceptually manageable and isolated material parts. After the effective environmental components are understood through reference to their causes, however, they conceptually link back together as an integrated whole, or *holocoenotic* system as it was once called. This is known as the dialectical approach to ecology. The dialectical approach examines the parts, but integrates the organism and the environment into a dynamic whole (or *umwelt*). Change in one ecological or environmental factor can concurrently affect the dynamic state of an entire ecosystem.<sup>[28][194]</sup>

## Disturbance and resilience

Ecosystems are regularly confronted with natural environmental variations and disturbances over time and geographic space. A disturbance is any process that removes biomass from a community, such as a fire, flood, drought, or predation.<sup>[195]</sup> Disturbances occur over vastly different ranges in terms of magnitudes as well as distances and time periods,<sup>[196]</sup> and are both the cause and product of natural fluctuations in death rates, species assemblages, and biomass densities within an ecological community. These disturbances create places of renewal where new directions emerge out of the patchwork of natural experimentation and opportunity.<sup>[195][197][198]</sup> Ecological resilience is a cornerstone theory in ecosystem management. Biodiversity fuels the resilience of ecosystems acting as a kind of regenerative insurance.<sup>[198]</sup>

## Metabolism and the early atmosphere

Metabolism – the rate at which energy and material resources are taken up from the environment, transformed within an organism, and allocated to maintenance, growth and reproduction – is a fundamental physiological trait.

Ernest et al.<sup>[199]:991</sup>

The Earth was formed approximately 4.5 billion years ago.<sup>[200]</sup> As it cooled and a crust and oceans formed, its atmosphere transformed from being dominated by hydrogen to one composed mostly of methane and ammonia. Over the next billion years, the metabolic activity of life transformed the atmosphere into a mixture of carbon dioxide, nitrogen, and water vapor. These gases changed the way that light from the sun hit the Earth's surface and greenhouse effects trapped heat. There were untapped sources of free energy within the mixture of reducing and oxidizing gasses that set the stage for primitive ecosystems to evolve and, in turn, the atmosphere also evolved.<sup>[201]</sup>



The leaf is the primary site of photosynthesis in most plants.

Throughout history, the Earth's atmosphere and biogeochemical cycles have been in a dynamic equilibrium with planetary ecosystems. The history is characterized by periods of significant transformation followed by millions of years of stability.<sup>[202]</sup> The evolution of the earliest organisms, likely anaerobic methanogen microbes, started the process by converting atmospheric hydrogen into methane ( $4\text{H}_2 + \text{CO}_2 \rightarrow \text{CH}_4 + 2\text{H}_2\text{O}$ ). Anoxygenic photosynthesis reduced hydrogen concentrations and increased atmospheric methane, by converting hydrogen sulfide into water or other sulfur compounds (for example,  $2\text{H}_2\text{S} + \text{CO}_2 + h\nu \rightarrow \text{CH}_2\text{O} + \text{H}_2\text{O} + 2\text{S}$ ). Early forms of fermentation also increased levels of atmospheric methane. The transition to an oxygen-dominant atmosphere (the *Great Oxidation*) did not begin until

approximately 2.4–2.3 billion years ago, but photosynthetic processes started 0.3 to 1 billion years prior.<sup>[202][203]</sup>

## Radiation: heat, temperature and light

The biology of life operates within a certain range of temperatures. Heat is a form of energy that regulates temperature. Heat affects growth rates, activity, behaviour and primary production. Temperature is largely dependent on the incidence of solar radiation. The latitudinal and longitudinal spatial variation of temperature greatly affects climates and consequently the distribution of biodiversity and levels of primary production in different ecosystems or biomes across the planet. Heat and temperature relate importantly to metabolic activity. Poikilotherms, for example, have a body temperature that is largely regulated and dependent on the temperature of the external environment. In contrast, homeotherms regulate their internal body temperature by expending metabolic energy.<sup>[141][142][194]</sup>

There is a relationship between light, primary production, and ecological energy budgets. Sunlight is the primary input of energy into the planet's ecosystems. Light is composed of electromagnetic energy of different wavelengths. Radiant energy from the sun generates heat, provides photons of light measured as active energy in the chemical reactions of life, and also acts as a catalyst for genetic mutation.<sup>[141][142][194]</sup> Plants, algae, and some bacteria absorb light and assimilate the energy through photosynthesis. Organisms capable of assimilating energy by photosynthesis or through inorganic fixation of  $\text{H}_2\text{S}$  are autotrophs. Autotrophs—responsible for primary production—assimilate light energy which becomes metabolically stored as potential energy in the form of biochemical enthalpic bonds.<sup>[141][142][194]</sup>

## Physical environments

### Water

Wetland conditions such as shallow water, high plant productivity, and anaerobic substrates provide a suitable environment for important physical, biological, and chemical processes. Because of these processes, wetlands play a vital role in global nutrient and element cycles.

Cronk & Fennessy (2001)<sup>[204]:29</sup>

Diffusion of carbon dioxide and oxygen is approximately 10,000 times slower in water than in air. When soils are flooded, they quickly lose oxygen, becoming hypoxic (an environment with O<sub>2</sub> concentration below 2 mg/liter) and eventually completely anoxic where anaerobic bacteria thrive among the roots. Water also influences the intensity and spectral composition of light as it reflects off the water surface and submerged particles.<sup>[204]</sup> Aquatic plants exhibit a wide variety of morphological and physiological adaptations that allow them to survive, compete and diversify in these environments. For example, their roots and stems contain large air spaces (aerenchyma) that regulate the efficient transportation of gases (for example, CO<sub>2</sub> and O<sub>2</sub>) used in respiration and photosynthesis. Salt water plants (halophytes) have additional specialized adaptations, such as the development of special organs for shedding salt and osmoregulating their internal salt (NaCl) concentrations, to live in estuarine, brackish, or oceanic environments. Anaerobic soil microorganisms in aquatic environments use nitrate, manganese ions, ferric ions, sulfate, carbon dioxide and some organic compounds; other microorganisms are facultative anaerobes and use oxygen during respiration when the soil becomes drier. The activity of soil microorganisms and the chemistry of the water reduces the oxidation-reduction potentials of the water. Carbon dioxide, for example, is reduced to methane (CH<sub>4</sub>) by methanogenic bacteria.<sup>[204]</sup> The physiology of fish is also specially adapted to compensate for environmental salt levels through osmoregulation. Their gills form electrochemical gradients that mediate salt excretion in salt water and uptake in fresh water.<sup>[205]</sup>

### Gravity

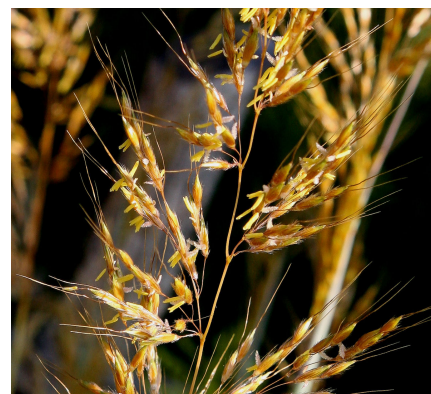
The shape and energy of the land is significantly affected by gravitational forces. On a large scale, the distribution of gravitational forces on the earth is uneven and influences the shape and movement of tectonic plates as well as influencing geomorphic processes such as orogeny and erosion. These forces govern many of the geophysical properties and distributions of ecological biomes across the Earth. On the organismal scale, gravitational forces provide directional cues for plant and fungal growth (gravitropism), orientation cues for animal migrations, and influence the biomechanics and size of animals.<sup>[141]</sup> Ecological traits, such as allocation of biomass in trees during growth are subject to mechanical failure as gravitational forces influence the position and structure of branches and leaves.<sup>[206]</sup> The cardiovascular systems of animals are functionally adapted to overcome pressure and gravitational forces that change according to the features of organisms (e.g., height, size, shape), their behaviour (e.g., diving, running, flying), and the habitat occupied (e.g., water, hot deserts, cold tundra).<sup>[207]</sup>

### Pressure

Climatic and osmotic pressure places physiological constraints on organisms, especially those that fly and respire at high altitudes, or dive to deep ocean depths. These constraints influence vertical limits of ecosystems in the biosphere, as organisms are physiologically sensitive and adapted to atmospheric and osmotic water pressure differences.<sup>[141]</sup> For example, oxygen levels decrease with decreasing pressure and are a limiting factor for life at higher altitudes.<sup>[208]</sup> Water transportation by plants is another important ecophysiological parameter affected by osmotic pressure gradients.<sup>[209][210][211]</sup> Water pressure in the depths of oceans requires that organisms adapt to these conditions. For example, diving animals such as whales, dolphins and seals are specially adapted to deal with changes in sound due to water pressure differences.<sup>[212]</sup> Differences between hagfish species provide another example of adaptation to deep-sea pressure through specialized protein adaptations.<sup>[213]</sup>

## Wind and turbulence

Turbulent forces in air and water affect environment and ecosystem distribution, form and dynamics. On a planetary scale, ecosystems are affected by circulation patterns in the global trade winds. Wind power and the turbulent forces it creates can influence heat, nutrient, and biochemical profiles of ecosystems.<sup>[141]</sup> For example, wind running over the surface of a lake creates turbulence, mixing the water column and influencing the environmental profile to create thermally layered zones, affecting how fish, algae, and other parts of the aquatic ecology are structured.<sup>[216][217]</sup> Wind speed and turbulence also influence evapotranspiration rates and energy budgets in plants and animals.<sup>[204][218]</sup> Wind speed, temperature and moisture content can vary as winds travel across different land features and elevations. For example, the westerlies come into contact with the coastal and interior mountains of western North America to produce a rain shadow on the leeward side of the mountain. The air expands and moisture condenses as the winds increase in elevation; this is called orographic lift and can cause precipitation. This environmental process produces spatial divisions in biodiversity, as species adapted to wetter conditions are range-restricted to the coastal mountain valleys and unable to migrate across the xeric ecosystems (e.g., of the Columbia Basin in western North America) to intermix with sister lineages that are segregated to the interior mountain systems.<sup>[219][220]</sup>



The architecture of the inflorescence in grasses is subject to the physical pressures of wind and shaped by the forces of natural selection facilitating wind-pollination (anemophily).<sup>[214][215]</sup>

## Fire



Plants convert carbon dioxide into biomass and emit oxygen into the atmosphere. Approximately 350 million years ago (at the end of the Devonian period), the amount of photosynthesis brought the concentration of atmospheric oxygen above 17%, which allowed combustion to occur.<sup>[221]</sup> Fire releases CO<sub>2</sub> and converts fuel into ash and tar. Fire is a significant ecological parameter that raises many issues pertaining to its control and suppression.<sup>[222]</sup> While the issue of fire in relation to ecology and plants has been recognized for a long time,<sup>[223]</sup> Charles Cooper brought attention to the issue of forest fires in relation to the ecology of forest fire suppression and management in the 1960s.<sup>[224][225]</sup>

Native North Americans were among the first to influence fire regimes by controlling their spread near their homes or by lighting fires to stimulate the production of herbaceous foods and basketry materials.<sup>[226]</sup> Fire creates a heterogenous ecosystem age and canopy structure, and the altered soil nutrient supply and cleared canopy structure opens new ecological niches for seedling establishment.<sup>[227][228]</sup> Most ecosystems are adapted to natural fire cycles. Plants, for example, are equipped with a variety of adaptations to deal with forest fires. Some species (e.g., *Pinus halepensis*) cannot germinate until after their seeds have lived through a fire or been exposed to certain compounds from smoke. Environmentally triggered germination of seeds is called serotiny.<sup>[229][230]</sup> Fire plays a major role in the persistence and resilience of ecosystems.<sup>[197]</sup>

## Soils

Soil is the living top layer of mineral and organic dirt that covers the surface of the planet. It is the chief organizing centre of most ecosystem functions, and it is of critical importance in agricultural science and ecology. The decomposition of dead organic matter (for example, leaves on the forest floor), results in soils containing minerals and nutrients that feed into plant production. The whole of the planet's soil ecosystems is called the pedosphere where a large biomass of the Earth's biodiversity organizes into trophic levels. Invertebrates that feed and shred larger leaves, for example, create smaller bits for smaller organisms in the feeding chain. Collectively, these organisms are the detritivores that regulate soil formation.<sup>[231][232]</sup> Tree roots, fungi, bacteria, worms, ants, beetles, centipedes, spiders, mammals, birds, reptiles, amphibians and other less familiar creatures all work to create the trophic web of life in soil ecosystems. Soils form composite phenotypes where inorganic matter is enveloped into the physiology of a whole community. As organisms feed and migrate through soils they physically displace materials, an ecological process called bioturbation. This aerates soils and stimulates heterotrophic growth and production. Soil microorganisms are influenced by and feed back into the trophic dynamics of the ecosystem. No single axis of causality can be discerned to segregate the biological from geomorphological systems in soils.<sup>[233][234]</sup> Paleoecological studies of soils places the origin for bioturbation to a time before the Cambrian period. Other events, such as the evolution of trees and the colonization of land in the Devonian period played a significant role in the early development of ecological trophism in soils.<sup>[99][232][235]</sup>

## Biogeochemistry and climate

Ecologists study and measure nutrient budgets to understand how these materials are regulated, flow, and recycled through the environment.<sup>[141][142][194]</sup> This research has led to an understanding that there is global feedback between ecosystems and the physical parameters of this planet, including minerals, soil, pH, ions, water and atmospheric gases. Six major elements (hydrogen, carbon, nitrogen, oxygen, sulfur, and phosphorus; H, C, N, O, S, and P) form the constitution of all biological macromolecules and feed into the Earth's geochemical processes. From the smallest scale of biology, the combined effect of billions upon billions of ecological processes amplify and ultimately regulate the biogeochemical cycles of the Earth. Understanding the relations and cycles mediated between these elements and their ecological pathways has significant bearing toward understanding global biogeochemistry.<sup>[236]</sup>

The ecology of global carbon budgets gives one example of the linkage between biodiversity and biogeochemistry. It is estimated that the Earth's oceans hold 40,000 gigatonnes (Gt) of carbon, that vegetation and soil hold 2070 Gt, and that fossil fuel emissions are 6.3 Gt carbon per year.<sup>[237]</sup> There have been major restructurings in these global carbon budgets during the Earth's history, regulated to a large extent by the ecology of the land. For example, through the early-mid Eocene volcanic outgassing, the oxidation of methane stored in wetlands, and seafloor gases increased atmospheric CO<sub>2</sub> (carbon dioxide) concentrations to levels as high as 3500 ppm.<sup>[238]</sup>

In the Oligocene, from 25 to 32 million years ago, there was another significant restructuring of the global carbon cycle as grasses evolved C4 photosynthesis and expanded their ranges. This new photosynthetic pathway evolved in response to the drop in atmospheric CO<sub>2</sub> concentrations below 550 ppm.<sup>[239]</sup> The relative abundance and distribution of biodiversity alters the dynamics between organisms and their environment such that ecosystems can be both cause and effect in relation to climate change. Human-driven modifications to the planet's ecosystems (e.g., disturbance, biodiversity loss, agriculture) contributes to rising atmospheric greenhouse gas levels. Transformation of the global carbon cycle in the next century is projected to raise planetary temperatures, lead to more extreme fluctuations in weather, alter species distributions, and increase extinction rates. The effect of global warming is already being registered in melting glaciers, melting mountain ice caps, and rising sea levels. Consequently, species distributions are changing along waterfronts and in continental areas where migration patterns and breeding grounds are tracking the prevailing shifts in climate. Large sections of permafrost are also melting to create a new mosaic of flooded areas having increased rates of soil decomposition activity that raises methane (CH<sub>4</sub>) emissions. There is concern over increases in atmospheric methane in the context of the global carbon cycle, because methane is a greenhouse gas that

is 23 times more effective at absorbing long-wave radiation than CO<sub>2</sub> on a 100-year time scale.<sup>[240]</sup> Hence, there is a relationship between global warming, decomposition and respiration in soils and wetlands producing significant climate feedbacks and globally altered biogeochemical cycles.<sup>[137][241][242][243][244][245]</sup>

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## Notes

- A. ^ In Ernst Haeckel's (1866) footnote where the term ecology originates, he also gives attribute to Ancient Greek: *χώρας khōrā* "χώρα", meaning "dwelling place, distributional area" - quoted from Stauffer (1957)
- B. ^ This is a copy of Haeckel's original definition (Original: Haeckel, E. (1866) *Generelle Morphologie der Organismen. Allgemeine Grundzüge der organischen Formen- Wissenschaft, mechanisch begründet durch die von Charles Darwin reformirte Descendenz-Theorie*. 2 vols. Reimer, Berlin.) translated and quoted from Stauffer (1957).
- C. ^ Foster & Clark (2008) note how Smut's holism contrasts starkly against his racial political views as the father of apartheid.
- D. ^ First introduced in MacArthur & Wilson's (1967) book of notable mention in the history and theoretical science of ecology, *The Theory of Island Biogeography*
- E. ^ Aristotle wrote about this concept in *Metaphysics* ([<http://classics.mit.edu/Aristotle/metaphysics.mb.txt> Quoted from the Internet Classics Archive. Book VIII, Part 6): "To return to the difficulty which has been stated with respect both to definitions and to numbers, what is the cause of their unity? In the case of all things which have several parts and in which the totality is not, as it were, a mere heap, but the whole is something beside the parts, there is a cause; for even in bodies contact is the cause of unity in some cases, and in others viscosity or some other such quality."

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## External links

- Ecology (Stanford Encyclopedia of Philosophy) (<http://plato.stanford.edu/entries/ecology/>)
- The Nature Education Knowledge Project: Ecology (<http://www.nature.com/scitable/knowledge/ecology-102>)
- Ecology Journals List of ecological scientific journals (<http://ekolojinet.com/journals.html>)
- Ecology Dictionary - Explanation of Ecological Terms (<http://ecologydictionary.org/>)
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- The Encyclopedia of Earth – Wilderness: Biology & Ecology (<http://www.eoearth.org/topics/view/49660/>)
- Ecology and Society - A journal of integrative science for resilience and sustainability (<http://www.ecologyandsociety.org/>)
- Science Aid: Ecology, U.K. High School (GCSE, Alevel) Ecology (<http://scienceaid.co.uk/biology/ecology/index.html>)

## Ecosystem

An **ecosystem** is a community of living organisms (plants, animals and microbes) in conjunction with the nonliving components of their environment (things like air, water and mineral soil), interacting as a system.<sup>[2]</sup> These components are regarded as linked together through nutrient cycles and energy flows.<sup>[3]</sup> As ecosystems are defined by the network of interactions among organisms, and between organisms and their environment,<sup>[4]</sup> they can come in any size but usually encompass specific, limited spaces<sup>[5]</sup> (although it is sometimes said that the entire planet is an ecosystem).<sup>[6]</sup>

Energy, water, nitrogen and soil minerals are other essential abiotic components of an ecosystem. The energy that flows through ecosystems is obtained primarily from the sun. It generally enters the system through photosynthesis, a process that also captures carbon from the atmosphere. By feeding on plants and on one-another, animals play an important role in the movement of matter and energy through the system. They also influence the quantity of plant and microbial biomass present. By breaking down dead organic matter, decomposers release carbon back to the atmosphere and facilitate nutrient cycling by converting nutrients stored in dead biomass back to a form that can be readily used by plants and other microbes.<sup>[7]</sup>



Coral reefs are a highly productive marine ecosystem.<sup>[1]</sup>

Ecosystems are controlled both by external and internal factors. External factors such as climate, the parent material which forms the soil and topography, control the overall structure an ecosystem and the way things work within it, but are not themselves influenced by the ecosystem.<sup>[8]</sup> Other external factors include time and potential biota. Ecosystems are dynamic entities—invariably, they are subject to periodic disturbances and are in the process of recovering from some past disturbance.<sup>[9]</sup> Ecosystems in similar environments that are located in different parts of the world can end up doing things very differently simply because they have different pools of species present.<sup>[8]</sup> The introduction of non-native species can cause substantial shifts in ecosystem function. Internal factors not only control ecosystem processes but are also controlled by them and are often subject to feedback loops.<sup>[8]</sup> While the resource inputs are generally controlled by external processes like climate and parent material, the availability of these resources within the ecosystem is controlled by internal factors like decomposition, root competition or shading.<sup>[8]</sup> Other internal factors include disturbance, succession and the types of species present. Although humans exist and operate within ecosystems, their cumulative effects are large enough to influence external factors like climate.<sup>[8]</sup>



Rainforest ecosystems are rich in biodiversity.  
This is the Gambia River in Senegal's  
Niokolo-Koba National Park.

Biodiversity affects ecosystem functioning, as do the processes of disturbance and succession. Ecosystems provide a variety of goods and services upon which people depend; the principles of ecosystem management suggest that rather than managing individual species, natural resources should be managed at the level of the ecosystem itself. Classifying ecosystems into ecologically homogeneous units is an important step towards effective ecosystem management, but there is no single, agreed-upon way to do this.

## History and development

Arthur Tansley, a British ecologist, was the first person to use the term "ecosystem" in a published work.<sup>[10][11]</sup> Tansley devised the concept to draw attention to the importance of transfers of materials between organisms and their environment.<sup>[12]</sup> He later refined the term, describing it as "The whole system, ... including not only the organism-complex, but also the whole complex of physical factors forming what we call the environment".<sup>[13]</sup> Tansley regarded ecosystems not simply as natural units, but as mental isolates.<sup>[13]</sup> Tansley later<sup>[14]</sup> defined the spatial extent of ecosystems using the term ecotope.

G. Evelyn Hutchinson, a pioneering limnologist who was a contemporary of Tansley's, combined Charles Elton's ideas about trophic ecology with those of Russian geochemist Vladimir Vernadsky to suggest that mineral nutrient availability in a lake limited algal production which would, in turn, limit the abundance of animals that feed on algae. Raymond Lindeman took these ideas one step further to suggest that the flow of energy through a lake was the primary driver of the ecosystem. Hutchinson's students, brothers Howard T. Odum and Eugene P. Odum, further developed a "systems approach" to the study of ecosystems, allowing them to study the flow of energy and material through ecological systems.<sup>[12]</sup>



## Ecosystem processes

Energy and carbon enter ecosystems through photosynthesis, are incorporated into living tissue, transferred to other organisms that feed on the living and dead plant matter, and eventually released through respiration.<sup>[15]</sup> Most mineral nutrients, on the other hand, are recycled within ecosystems.<sup>[16]</sup>

Ecosystems are controlled both by external and internal factors. External factors, also called state factors, control the overall structure an ecosystem and the way things work within it, but are not themselves influenced by the ecosystem. The most important of these is climate.<sup>[8]</sup> Climate determines the biome in which the ecosystem is embedded. Rainfall patterns and temperature seasonality determine the amount of water available to the ecosystem and the supply of energy available (by influencing photosynthesis).<sup>[8]</sup> Parent material, the underlying geological material that gives rise to soils, determines the nature of the soils present, and influences the supply of mineral nutrients. Topography also controls ecosystem processes by affecting things like microclimate, soil development and the movement of water through a system. This may be the difference between the ecosystem present in wetland situated in a small depression on the landscape, and one present on an adjacent steep hillside.<sup>[8]</sup>

Other external factors that play an important role in ecosystem functioning include time and potential biota. Ecosystems are dynamic entities—invariably, they are subject to periodic disturbances and are in the process of recovering from some past disturbance.<sup>[9]</sup> Time plays a role in the development of soil from bare rock and the recovery of a community from disturbance.<sup>[8]</sup> Similarly, the set of organisms that can potentially be present in an area can also have a major impact on ecosystems. Ecosystems in similar environments that are located in different parts of the world can end up doing things very differently simply because they have different pools of species present.<sup>[8]</sup> The introduction of non-native species can cause substantial shifts in ecosystem function.

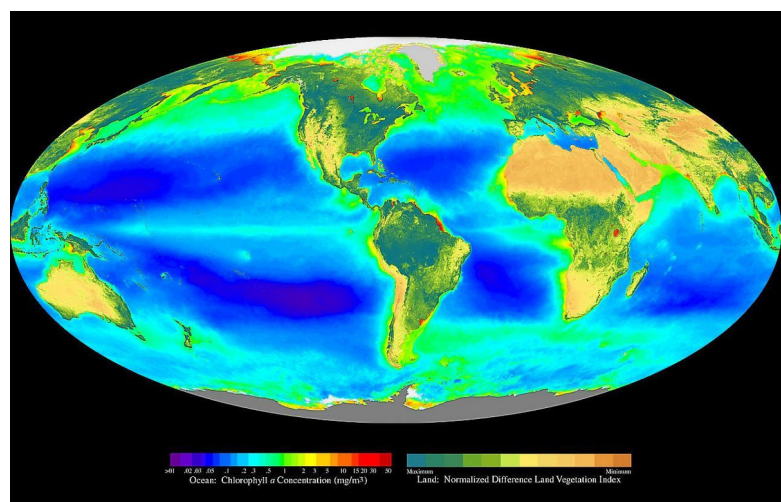
Unlike external factors, internal factors in ecosystems not only control ecosystem processes, but are also controlled by them. Consequently, they are often subject to feedback loops.<sup>[8]</sup> While the resource inputs are generally controlled by external processes like climate and parent material, the availability of these resources within the ecosystem is controlled by internal factors like decomposition, root competition or shading.<sup>[8]</sup> Other factors like disturbance, succession or the types of species present are also internal factors. Human activities are important in almost all ecosystems. Although humans exist and operate within ecosystems, their cumulative effects are large enough to influence external factors like climate.<sup>[8]</sup>

## Primary production

Primary production is the production of organic matter from inorganic carbon sources. Overwhelmingly, this occurs through photosynthesis. The energy incorporated through this process supports life on earth, while the carbon makes up much of the organic matter in living and dead biomass, soil carbon and fossil fuels. It also drives the carbon cycle, which influences global climate via the greenhouse effect.

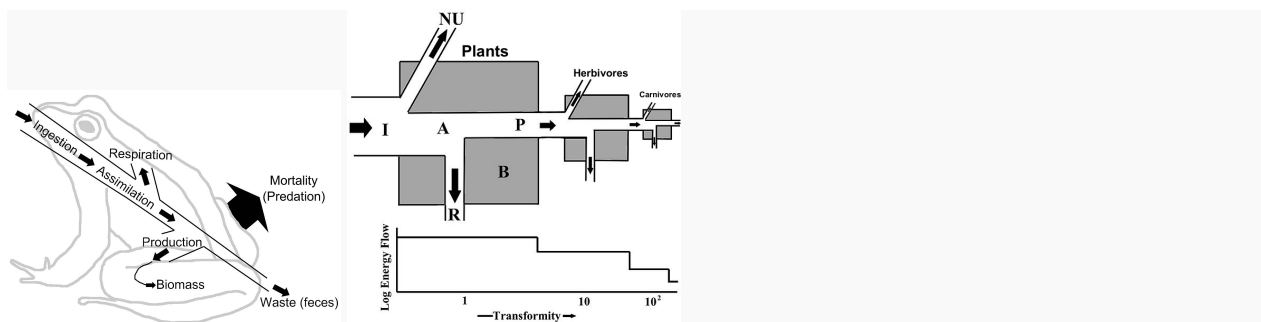
Through the process of photosynthesis, plants capture energy from light and use it to combine carbon dioxide and water to produce carbohydrates and oxygen. The photosynthesis carried out

by all the plants in an ecosystem is called the gross primary production (GPP).<sup>[17]</sup> About 48–60% of the GPP is consumed in plant respiration. The remainder, that portion of GPP that is not used up by respiration, is known as the net primary production (NPP).<sup>[15]</sup> Total photosynthesis is limited by a range of environmental factors. These include the amount of light available, the amount of leaf area a plant has to capture light (shading by other plants is a major limitation of photosynthesis), rate at which carbon dioxide can be supplied to the chloroplasts to support photosynthesis, the availability of water, and the availability of suitable temperatures for carrying out photosynthesis.<sup>[17]</sup>



Global oceanic and terrestrial phototroph abundance, from September 1997 to August 2000. As an estimate of autotroph biomass, it is only a rough indicator of primary production potential, and not an actual estimate of it. Provided by the SeaWiFS Project, NASA/Goddard Space Flight Center and ORBIMAGE.

## Energy flow



**Left:** Energy flow diagram of a frog. The frog represents a node in an extended food web. The energy ingested is utilized for metabolic processes and transformed into biomass. The energy flow continues on its path if the frog is ingested by predators, parasites, or as a decaying carcass in soil. This energy flow diagram illustrates how energy is lost as it fuels the metabolic process that transform the energy and nutrients into biomass.

**Right:** An expanded three link energy food chain (1. plants, 2. herbivores, 3. carnivores) illustrating the relationship between food flow diagrams and energy transformity. The transformity of energy becomes degraded, dispersed, and diminished from higher quality to lesser quantity as the energy within a food chain flows from one trophic species into another. Abbreviations: I=input, A=assimilation, R=respiration, NU=not utilized, P=production, B=biomass.<sup>[18]</sup>

The carbon and energy incorporated into plant tissues (net primary production) is either consumed by animals while the plant is alive, or it remains uneaten when the plant tissue dies and becomes detritus. In terrestrial ecosystems,

roughly 90% of the NPP ends up being broken down by decomposers. The remainder is either consumed by animals while still alive and enters the plant-based trophic system, or it is consumed after it has died, and enters the detritus-based trophic system. In aquatic systems, the proportion of plant biomass that gets consumed by herbivores is much higher.<sup>[19]</sup> In trophic systems photosynthetic organisms are the primary producers. The organisms that consume their tissues are called primary consumers or secondary producers—herbivores. Organisms which feed on microbes (bacteria and fungi) are termed microbivores. Animals that feed on primary consumers—carnivores—are secondary consumers. Each of these constitutes a trophic level.<sup>[19]</sup> The sequence of consumption—from plant to herbivore, to carnivore—forms a food chain. Real systems are much more complex than this—organisms will generally feed on more than one form of food, and may feed at more than one trophic level. Carnivores may capture some prey which are part of a plant-based trophic system and others that are part of a detritus-based trophic system (a bird that feeds both on herbivorous grasshoppers and earthworms, which consume detritus). Real systems, with all these complexities, form food webs rather than food chains.<sup>[19]</sup>

## Decomposition

The carbon and nutrients in dead organic matter are broken down by a group of processes known as decomposition. This releases nutrients that can then be re-used for plant and microbial production, and returns carbon dioxide to the atmosphere (or water) where it can be used for photosynthesis. In the absence of decomposition, dead organic matter would accumulate in an ecosystem and nutrients and atmospheric carbon dioxide would be depleted.<sup>[20]</sup> Approximately 90% of terrestrial NPP goes directly from plant to decomposer.<sup>[19]</sup>

Decomposition processes can be separated into three categories—leaching, fragmentation and chemical alteration of dead material. As water moves through dead organic matter, it dissolves and carries with it the water-soluble components. These are then taken up by organisms in the soil, react with mineral soil, or are transported beyond the confines of the ecosystem (and are considered "lost" to it).<sup>[20]</sup> Newly shed leaves and newly dead animals have high concentrations of water-soluble components, and include sugars, amino acids and mineral nutrients. Leaching is more important in wet environments, and much less important in dry ones.<sup>[20]</sup>

Fragmentation processes break organic material into smaller pieces, exposing new surfaces for colonization by microbes. Freshly shed leaf litter may be inaccessible due to an outer layer of cuticle or bark, and cell contents are protected by a cell wall. Newly dead animals may be covered by an exoskeleton. Fragmentation processes, which break through these protective layers, accelerate the rate of microbial decomposition.<sup>[20]</sup> Animals fragment detritus as they hunt for food, as does passage through the gut. Freeze-thaw cycles and cycles of wetting and drying also fragment dead material.<sup>[20]</sup>

The chemical alteration of dead organic matter is primarily achieved through bacterial and fungal action. Fungal hyphae produce enzymes which can break through the tough outer structures surrounding dead plant material. They also produce enzymes which break down lignin, which allows to them access to both cell contents and to the nitrogen in the lignin. Fungi can transfer carbon and nitrogen through their hyphal networks and thus, unlike bacteria, are not dependent solely on locally available resources.<sup>[20]</sup>

Decomposition rates vary among ecosystems. The rate of decomposition is governed by three sets of factors—the physical environment (temperature, moisture and soil properties), the quantity and quality of the dead material available to decomposers, and the nature of the microbial community itself.<sup>[21]</sup> Temperature controls the rate of microbial respiration; the higher the temperature, the faster microbial decomposition occurs. It also affects soil moisture, which slows microbial growth and reduces leaching. Freeze-thaw cycles also affect decomposition—freezing temperatures kill soil microorganisms, which allows leaching to play a more important role in moving nutrients around. This can be especially important as the soil thaws in the Spring, creating a pulse of nutrients which become available.<sup>[21]</sup>

Decomposition rates are low under very wet or very dry conditions. Decomposition rates are highest in wet, moist conditions with adequate levels of oxygen. Wet soils tend to become deficient in oxygen (this is especially true in

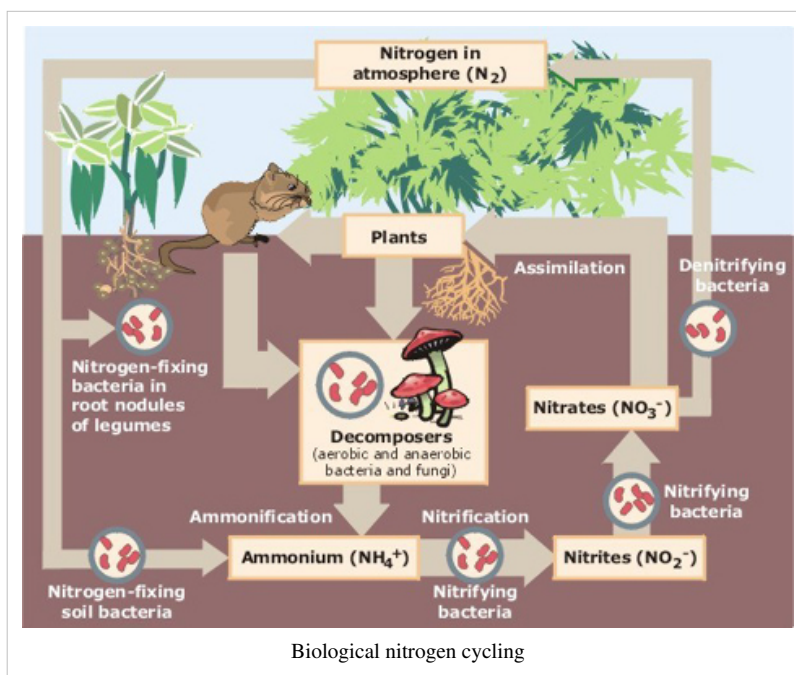
wetlands), which slows microbial growth. In dry soils, decomposition slows as well, but bacteria continue to grow (albeit at a slower rate) even after soils become too dry to support plant growth. When the rains return and soils become wet, the osmotic gradient between the bacterial cells and the soil water causes the cells to gain water quickly. Under these conditions, many bacterial cells burst, releasing a pulse of nutrients.<sup>[21]</sup> Decomposition rates also tend to be slower in acidic soils.<sup>[21]</sup> Soils which are rich in clay minerals tend to have lower decomposition rates, and thus, higher levels of organic matter.<sup>[21]</sup> The smaller particles of clay result in a larger surface area that can hold water. The higher the water content of a soil, the lower the oxygen content<sup>[22]</sup> and consequently, the lower the rate of decomposition. Clay minerals also bind particles of organic material to their surface, making them less accessible to microbes.<sup>[21]</sup> Soil disturbance like tilling increase decomposition by increasing the amount of oxygen in the soil and by exposing new organic matter to soil microbes.<sup>[21]</sup>

The quality and quantity of the material available to decomposers is another major factor that influences the rate of decomposition. Substances like sugars and amino acids decompose readily and are considered "labile". Cellulose and hemicellulose, which are broken down more slowly, are "moderately labile". Compounds which are more resistant to decay, like lignin or cutin, are considered "recalcitrant".<sup>[21]</sup> Litter with a higher proportion of labile compounds decomposes much more rapidly than does litter with a higher proportion of recalcitrant material. Consequently, dead animals decompose more rapidly than dead leaves, which themselves decompose more rapidly than fallen branches.<sup>[21]</sup> As organic material in the soil ages, its quality decreases. The more labile compounds decompose quickly, leaving and increasing proportion of recalcitrant material. Microbial cell walls also contain a recalcitrant materials like chitin, and these also accumulate as the microbes die, further reducing the quality of older soil organic matter.<sup>[21]</sup>

## Nutrient cycling

Ecosystems continually exchange energy and carbon with the wider environment; mineral nutrients, on the other hand, are mostly cycled back and forth between plants, animals, microbes and the soil. Most nitrogen enters ecosystems through biological nitrogen fixation, is deposited through precipitation, dust, gases or is applied as fertilizer.<sup>[16]</sup> Since most terrestrial ecosystems are nitrogen-limited, nitrogen cycling is an important control on ecosystem production.<sup>[16]</sup>

Until modern times, nitrogen fixation was the major source of nitrogen for ecosystems. Nitrogen fixing bacteria either live symbiotically with plants, or live freely in the soil. The energetic cost is high for plants which support nitrogen-fixing symbionts—as much as 25% of GPP when measured in controlled conditions. Many members of the legume plant family support nitrogen-fixing symbionts. Some cyanobacteria are also capable of nitrogen fixation. These are phototrophs, which carry out photosynthesis. Like other nitrogen-fixing bacteria, they can either be free-living or have symbiotic relationships with plants.<sup>[16]</sup> Other sources of nitrogen include acid deposition produced through the combustion of fossil fuels, ammonia gas which evaporates from agricultural fields which have had fertilizers applied to them, and dust.<sup>[16]</sup> Anthropogenic nitrogen inputs account for about 80% of all nitrogen fluxes in ecosystems.<sup>[16]</sup>



When plant tissues are shed or are eaten, the nitrogen in those tissues becomes available to animals and microbes. Microbial decomposition releases nitrogen compounds from dead organic matter in the soil, where plants, fungi and bacteria compete for it. Some soil bacteria use organic nitrogen-containing compounds as a source of carbon, and release ammonium ions into the soil. This process is known as nitrogen mineralization. Others convert ammonium to nitrite and nitrate ions, a process known as nitrification. Nitric oxide and nitrous oxide are also produced during nitrification.<sup>[16]</sup> Under nitrogen-rich and oxygen-poor conditions, nitrates and nitrites are converted to nitrogen gas, a process known as denitrification.<sup>[16]</sup>

Other important nutrients include phosphorus, sulfur, calcium, potassium, magnesium and manganese.<sup>[23]</sup> Phosphorus enters ecosystems through weathering. As ecosystems age this supply diminishes, making phosphorus-limitation more common in older landscapes (especially in the tropics).<sup>[23]</sup> Calcium and sulfur are also produced by weathering, but acid deposition is an important source of sulfur in many ecosystems. Although magnesium and manganese are produced by weathering, exchanges between soil organic matter and living cells account for a significant portion of ecosystem fluxes. Potassium is primarily cycled between living cells and soil organic matter.<sup>[23]</sup>

## Function and biodiversity

Ecosystem processes are broad generalizations that actually take place through the actions of individual organisms. The nature of the organisms—the species, functional groups and trophic levels to which they belong—dictates the sorts of actions these individuals are capable of carrying out, and the relative efficiency with which they do so. Thus, ecosystem processes are driven by the number of species in an ecosystem, the exact nature of each individual species, and the relative abundance organisms within these species.<sup>[25]</sup> Biodiversity plays an important role in ecosystem functioning.<sup>[26]</sup>

Ecological theory suggests that in order to coexist, species must have some level of limiting similarity—they must be different from one another in some fundamental way, otherwise one species would competitively exclude the other.<sup>[27]</sup> Despite this, the cumulative effect of additional species in an ecosystem is not linear—additional species may enhance nitrogen retention, for example, but beyond some level of species richness, additional species may have little additive effect.<sup>[25]</sup> The addition (or loss) of species which are ecologically similar to those already present in an ecosystem tends to only have a small effect on ecosystem function. Ecologically distinct species, on the other hand, have a much larger effect. Similarly, dominant species have a large impact on ecosystem function, while rare species tend to have a small effect. Keystone species tend to have a effect on ecosystem function that is disproportionate to their abundance in an ecosystem.<sup>[25]</sup>

## Ecosystem goods and services

Ecosystems provide a variety of goods and services upon which people depend.<sup>[28]</sup> Ecosystem goods include the "tangible, material products"<sup>[29]</sup> of ecosystem processes—food, construction material, medicinal plants—in addition to less tangible items like tourism and recreation, and genes from wild plants and animals that can be used to improve domestic species.<sup>[28]</sup> Ecosystem services, on the other hand, are generally "improvements in the condition



Loch Lomond in Scotland forms a relatively isolated ecosystem. The fish community of this lake has remained stable over a long period until a number of introductions in the 1970s restructured its food web.<sup>[24]</sup>



Spiny forest at Ifaty, Madagascar, featuring various *Adansonia* (baobab) species, *Alluaudia procera* (Madagascar ocotillo) and other vegetation.



or location of things of value".<sup>[29]</sup> These include things like the maintenance of hydrological cycles, cleaning air and water, the maintenance of oxygen in the atmosphere, crop pollination and even things like beauty, inspiration and opportunities for research.<sup>[28]</sup> While ecosystem goods have traditionally been recognized as being the basis for things of economic value, ecosystem services tend to be taken for granted.<sup>[29]</sup> While Gretchen Daily's original definition distinguished between ecosystem goods and ecosystem services, Robert Costanza and colleagues' later work and that of the Millennium Ecosystem Assessment lumped all of these together as ecosystem services.<sup>[29]</sup>

## Ecosystem management

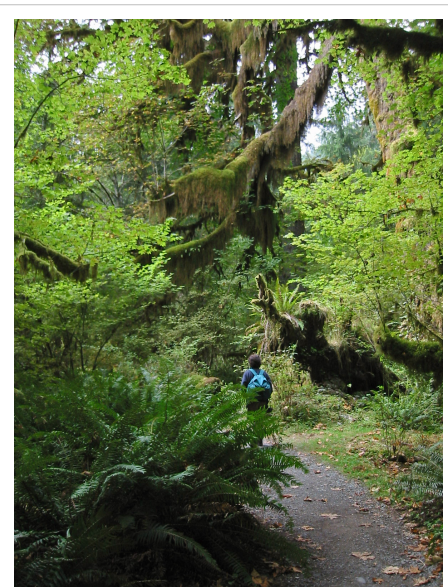
When natural resource management is applied to whole ecosystems, rather than single species, it is termed ecosystem management.<sup>[30]</sup> A variety of definitions exist: F. Stuart Chapin and coauthors define it as "the application of ecological science to resource management to promote long-term sustainability of ecosystems and the delivery of essential ecosystem goods and services",<sup>[31]</sup> while Norman Christensen and coauthors defined it as "management driven by explicit goals, executed by policies, protocols, and practices, and made adaptable by monitoring and research based on our best understanding of the ecological interactions and processes necessary to sustain ecosystem structure and function"<sup>[28]</sup> and Peter Brussard and colleagues defined it as "managing areas at various scales in such a way that ecosystem services and biological resources are preserved while appropriate human use and options for livelihood are sustained".<sup>[32]</sup>

Although definitions of ecosystem management abound, there is a common set of principles which underlie these definitions.<sup>[31]</sup> A fundamental principle is the long-term sustainability of the production of goods and services by the ecosystem;<sup>[31]</sup> "intergenerational sustainability [is] a precondition for management, not an afterthought".<sup>[28]</sup> It also requires clear goals with respect to future trajectories and behaviors of the system being managed. Other important requirements include a sound ecological understanding of the system, including connectedness, ecological dynamics and the context in which the system is embedded. Other important principles include an understanding of the role of humans as components of the ecosystems and the use of adaptive management.<sup>[28]</sup> While ecosystem management can be used as part of a plan for wilderness conservation, it can also be used in intensively managed ecosystems<sup>[28]</sup> (see, for example, agroecosystem and close to nature forestry).

## Ecosystem dynamics

Ecosystems are dynamic entities—invariably, they are subject to periodic disturbances and are in the process of recovering from some past disturbance.<sup>[9]</sup> When an ecosystem is subject to some sort of perturbation, it responds by moving away from its initial state. The tendency of a system to remain close to its equilibrium state, despite that disturbance, is termed its resistance. On the other hand, the speed with which it returns to its initial state after disturbance is called its resilience.<sup>[9]</sup>

From one year to another, ecosystems experience variation in their biotic and abiotic environments. A drought, an especially cold winter and a pest outbreak all constitute short-term variability in environmental conditions. Animal populations vary from year to year, building up during resource-rich periods and crashing as they overshoot their food supply. These changes play out in changes in NPP, decomposition rates, and other ecosystem processes.<sup>[9]</sup> Longer-term changes also shape ecosystem processes—the forests of



Temperate rainforest on the Olympic Peninsula in Washington state.



eastern North America still show legacies of cultivation which ceased 200 years ago, while methane production in eastern Siberian lakes is controlled by organic matter which accumulated during the Pleistocene.<sup>[9]</sup>

Disturbance also plays an important role in ecological processes. F. Stuart Chapin and coauthors define disturbance as "a relatively discrete event in time and space that alters the structure of populations, communities and ecosystems and causes changes in resources availability or the physical environment".<sup>[33]</sup> This can range from tree falls and insect outbreaks to hurricanes and wildfires to volcanic eruptions and can cause large changes in plant, animal and microbe populations, as well soil organic matter content.<sup>[9]</sup> Disturbance is followed by succession, a "directional change in ecosystem structure and functioning resulting from biotically driven changes in resources supply."<sup>[33]</sup>



The High Peaks Wilderness Area in the 6000000-acre (**unknown operator: u'strong'** ha) Adirondack Park is an example of a diverse ecosystem.

The frequency and severity of disturbance determines the way it impacts ecosystem function. Major disturbance like a volcanic eruption or glacial advance and retreat leave behind soils that lack plants, animals or organic matter. Ecosystems that experience disturbances that sever undergo primary succession. Less severe disturbance like forest fires, hurricanes or cultivation result in secondary succession.<sup>[9]</sup> More severe disturbance and more frequent disturbance result in longer recovery times. Ecosystems recover more quickly from less severe disturbance events.<sup>[9]</sup>

The early stages of primary succession are dominated by species with small propagules (seed and spores) which can be dispersed long distances. The early colonizers—often algae, cyanobacteria and lichens—stabilize the substrate. Nitrogen supplies are limited in new soils, and nitrogen-fixing species tend to play an important role early in primary succession. Unlike in primary succession, the species that dominate secondary succession, are usually present from the start of the process, often in the soil seed bank. In some systems the successional pathways are fairly consistent, and thus, are easy to predict. In others, there are many possible pathways—for example, the introduced nitrogen-fixing legume, *Myrica faya*, alter successional trajectories in Hawai'ian forests.<sup>[9]</sup>

The theoretical ecologist Robert Ulanowicz has used information theory tools to describe the structure of ecosystems, emphasizing mutual information (correlations) in studied systems. Drawing on this methodology and prior observations of complex ecosystems, Ulanowicz depicts approaches to determining the stress levels on ecosystems and predicting system reactions to defined types of alteration in their settings (such as increased or reduced energy flow, and eutrophication).<sup>[34]</sup>

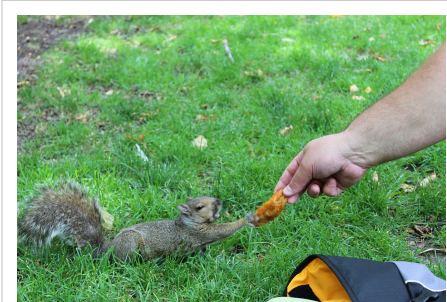
## Ecosystem ecology

Ecosystem ecology studies "the flow of energy and materials through organisms and the physical environment". It seeks to understand the processes which govern the stocks of material and energy in ecosystems, and the flow of matter and energy through them. The study of ecosystems can cover 10 orders of magnitude, from the surface layers of rocks to the surface of the planet.<sup>[35]</sup>

There is no single definition of what constitutes an ecosystem.<sup>[36]</sup> German ecologist Ernst-Detlef Schulze and coauthors defined an ecosystem as an area which is "uniform regarding the biological turnover, and contains all the fluxes above and below the ground area under consideration." They explicitly reject Gene Likens' use of entire river catchments as "too wide a demarcation" to be a single ecosystem, given the level of heterogeneity within such an area.<sup>[37]</sup> Other authors have suggested that an ecosystem can encompass a much larger area, even the whole planet.<sup>[6]</sup> Schulze and coauthors also rejected the idea that a single rotting log could be studied as an ecosystem because the size of the flows between the log and its surroundings are too large, relative to the proportion cycles within the log.<sup>[37]</sup> Philosopher of science Mark Sagoff considers the failure to define "the kind of object it studies" to be an

obstacle to the development of theory in ecosystem ecology.<sup>[36]</sup>

Ecosystems can be studied through a variety of approaches—theoretical studies, studies monitoring specific ecosystems over long periods of time, those that look at differences between ecosystems to elucidate how they work and direct manipulative experimentation.<sup>[38]</sup> Studies can be carried out at a variety of scales, from microcosms and mesocosms which serve as simplified representations of ecosystems, through whole-ecosystem studies.<sup>[39]</sup> American ecologist Stephen R. Carpenter has argued that microcosm experiments can be "irrelevant and diversionary" if they are not carried out in conjunction with field studies carried out at the ecosystem scale, because microcosm experiments often fail to accurately predict ecosystem-level dynamics.<sup>[40]</sup>

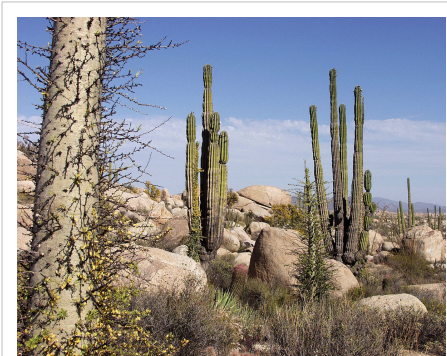


The ecosystem of public parks often includes humans feeding the wildlife.

The Hubbard Brook Ecosystem Study, established in the White Mountains, New Hampshire in 1963, was the first successful attempt to study an entire watershed as an ecosystem. The study used stream chemistry as a means of monitoring ecosystem properties, and developed a detailed biogeochemical model of the ecosystem.<sup>[41]</sup> Long-term research at the site led to the discovery of acid rain in North America in 1972, and was able to document the consequent depletion of soil cations (especially calcium) over the next several decades.<sup>[42]</sup>

## Classification

Classifying ecosystems into ecologically homogeneous units is an important step towards effective ecosystem management.<sup>[43]</sup> A variety of systems exist, based on vegetation cover, remote sensing, and bioclimatic classification systems.<sup>[43]</sup> American geographer Robert Bailey defines a hierarchy of ecosystem units ranging from microecosystems (individual homogeneous sites, on the order of 10 square kilometres (**unknown operator: u'strong'** sq mi) in area), through mesoecosystems (landscape mosaics, on the order of 1000 square kilometres (**unknown operator: u'strong'** sq mi)) to macroecosystems (ecoregions, on the order of 100000 square kilometres (**unknown operator: u'strong'** sq mi)).<sup>[44]</sup>



Flora of Baja California Desert, Cataviña region, Mexico.

Bailey outlined five different methods for identifying ecosystems:

*gestalt* ("a whole that is not derived through considerable of its parts"), in which regions are recognized and boundaries drawn intuitively; a map overlay system where different layers like geology, landforms and soil types are overlain to identify ecosystems; multivariate clustering of site attributes; digital image processing of remotely sensed data grouping areas based on their appearance or other spectral properties; or by a "controlling factors method" where a subset of factors (like soils, climate, vegetation physiognomy or the distribution of plant or animal species) are selected from a large array of possible ones are used to delineate ecosystems.<sup>[45]</sup> In contrast with Bailey's methodology, Puerto Rican ecologist Ariel Lugo and coauthors identified ten characteristics of an effective classification system: that it be based on georeferenced, quantitative data; that it should minimize subjectivity and explicitly identify criteria and assumptions; that it should be structured around the factors that drive ecosystem processes; that it should reflect the hierarchical nature of ecosystems; that it should be flexible enough to conform to the various scales at which ecosystem management operates; that it should be tied to reliable measures of climate so that it can "anticipat[e] global climate change; that it be applicable worldwide; that it should be validated against independent data; that it take into account the sometimes complex relationship between climate, vegetation and

ecosystem functioning; and that it should be able to adapt and improve as new data become available".<sup>[43]</sup>

## Examples of ecosystems

- Aquatic ecosystem
- Forest
- Greater Yellowstone Ecosystem
- Human ecosystem
- Large marine ecosystem
- Littoral zone
- Marine ecosystem
- Riparian zone
- River ecosystem
- Subsurface Lithoautotrophic Microbial Ecosystem
- Urban ecosystem
- Mobile Cave



A freshwater ecosystem in Gran Canaria, an island of the Canary Islands.

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## External links

- Millennium Ecosystem Assessment (<http://www.millenniumassessment.org/en/index.aspx>) (2005)
  - The State of the Nation's Ecosystems (U.S.) (<http://www.heinzctr.org/ecosystems>)
  - Bering Sea Climate and Ecosystem (Current status) (<http://www.beringclimate.noaa.gov/>)
  - Arctic Climate and Ecosystem (Current status) (<http://www.arctic.noaa.gov/detect/>)
  - Teaching about Ecosystems (<http://www.ericdigests.org/2004-1/ecosystems.htm>)
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# Human interrelationship

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## Wilderness

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**Wilderness** or **wildland** is a natural environment on Earth that has not been significantly modified by human activity. It may also be defined as: "The most intact, undisturbed wild natural areas left on our planet—those last truly wild places that humans do not control and have not developed with roads, pipelines or other industrial infrastructure."<sup>[1]</sup> Wilderness areas can be found in preserves, estates, farms, conservation preserves, ranches, National Forests, National Parks and even in urban areas along rivers, gulches or otherwise undeveloped areas. These areas are considered important for the survival of certain species, biodiversity, ecological studies, conservation, solitude, and recreation. Wilderness is deeply valued for cultural, spiritual, moral, and aesthetic reasons. Some nature writers believe wilderness areas are vital for the human spirit and creativity.<sup>[2]</sup> They may also preserve historic genetic traits and provide habitat for wild flora and fauna that may be difficult to recreate in zoos, arboretums or laboratories.

The word *wilderness* derives from the notion of "wildness"—in other words, that which is not controllable by humans. The word's etymology is from the Old English *wildeornes*, which in turn derives from *wildeor* meaning wild beast (wild + deor = beast, deer) (The Collins English Dictionary, 2000). From this point of view, it is the wildness of a place that makes it a wilderness. The mere presence or activity of people does not disqualify an area from



Old growth European Beech forest in Biogradska Gora National Park, Montenegro



Forrester Island Wilderness in the U.S. state of Alaska



being "wilderness." Many ecosystems that are, or have been, inhabited or influenced by activities of people may still be considered "wild." This way of looking at wilderness includes areas within which natural processes operate without human interference.

The WILD Foundation states that wilderness areas have two dimensions: they must be biologically intact and legally protected.<sup>[3][4]</sup> The World Conservation Union (IUCN) classifies wilderness at two levels, Ia (Strict Nature Reserves) and Ib (Wilderness Areas). Most scientists and conservationists agree that no place on earth is completely untouched by humanity, either due to past occupation by indigenous people, or through global processes such as climate change. Activities on the margins of specific wilderness areas, such as fire suppression and the interruption of animal migration also affect the interior of wildernesses.<sup>[5]</sup>

Especially in wealthier, industrialized nations, it has a specific legal meaning as well: as land where development is prohibited by law. Many nations have designated wilderness, including Australia, Canada, New Zealand, South Africa and the

United States. Many new parks are currently being planned and legally passed by various Parliaments and Legislatures at the urging of dedicated individuals around the globe who believe that "in the end, dedicated, inspired people empowered by effective legislation will ensure that the spirit and services of wilderness will thrive and permeate our society, preserving a world that we are proud to hand over to those who come after us."<sup>[6]</sup>



Perkuć Reserve in Puszcza Augustowska



This is a redwood tree in northern California redwood forest, where many redwood trees are managed for preservation and longevity.

## History

Looked at through the lens of the visual arts, nature and wildness have been important subjects in various epochs of world history. An early tradition of landscape art occurred in the Tang Dynasty (618-907). The tradition of representing nature *as it is* became one of the aims of Chinese painting and was a significant influence in Asian art. Artists in the tradition of Shan shui (lit. *mountain-water-picture*), learned to depict mountains and rivers "from the perspective of nature as a whole and on the basis of their understanding of the laws of nature... as if seen through the eyes of a bird." In the 13th century, Shih Erh Chi recommended avoiding painting "scenes lacking any places made inaccessible by nature."<sup>[7]</sup>



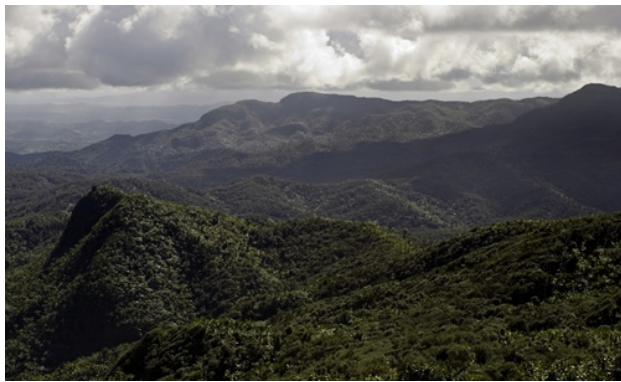
Innoko Wilderness, Alaska USA in the summer.

The idea of wilderness having intrinsic value emerged in the Western world in the 19th century. British artists John Constable and JMW Turner turned their attention to capturing the beauty of the natural world in their paintings. Prior to that, paintings had been primarily of religious scenes or of human beings. William Wordsworth's poetry described the wonder of the natural world, which had formerly been viewed as a threatening place. Increasingly the valuing of nature became an aspect of Western culture.<sup>[8]</sup>

Over the course of the 19th century wilderness became viewed not as a place to fear but a place to enjoy and protect, hence came the conservation movement in the latter half of the 19th century. Rivers were rafted and mountains were climbed solely for the sake of recreation, not to determine their geographical context. Early conservationists advocated the creation of a legal mechanism by which boundaries could be set on capitalistic ventures in order to preserve natural and unique lands for the enjoyment and use of future generations. This profound shift in wilderness thought reached a pinnacle in the US with the passage of the Wilderness Act of 1964, which allowed for parts of U.S. National Forests to be designated as "wilderness preserves". Similar acts, such as the 1975 Eastern Wilderness Act, followed.

The 21st century has seen another slight shift in wilderness thought and theory. It is now understood that simply drawing lines around a piece of land and declaring it a wilderness does not necessarily make it a wilderness. All landscapes are intricately connected and what happens outside a wilderness certainly affects what happens inside it. For example, air pollution from Los Angeles and the California Central Valley affects Kern Canyon and Sequoia National Park. The national park has miles of "wilderness" but the air is filled with pollution from the valley. This gives rise to the paradox of what a wilderness really is; a key issue in 21st century wilderness thought.

## Preservation



El Toro Wilderness within the Caribbean National Forest in Puerto Rico

For most of human history, the greater part of the Earth's terrain was wilderness, and human attention was concentrated in settled areas. The first known laws to protect parts of nature date back to the Babylonian Empire and Chinese Empire. Ashoka, the Great Mauryan King, defined the first laws in the world to protect flora and fauna in Edicts of Ashoka around 3rd Century B.C. In the Middle Ages, the Kings of England initiated one of the world's first conscious efforts to protect natural areas. They were motivated by a desire to be able to hunt wild animals in private hunting preserves rather than a desire to

protect wilderness. Nevertheless, in order to have animals to hunt they would have to protect wildlife from subsistence hunting and the land from villagers gathering firewood.<sup>[8]</sup> Similar measures were introduced in other European countries.

Early in the 19th century, Wordsworth and other romanticists in the U.K., concerned about "the excesses of industrialization and urbanization," called for a return to natural environments. This movement achieved some gains in protecting sensitive ecosystems, but a more successful form of environmentalism emerged in Germany by the mid-19th century. "Scientific Conservation," as it was called, advocated "the efficient utilization of natural resources through the application of science and technology." Concepts of forest management based on the German approach were applied in other parts of the world, but with varying degrees of success.<sup>[9]</sup>

By the later 19th century it had become clear that in many countries wild areas had either disappeared or were in danger of disappearing. This realization gave rise to the conservation movement in the USA, partly through the efforts of writers and activists such as John Burroughs, Aldo Leopold, and John Muir, and politicians such as U.S. President Teddy Roosevelt.

## National parks

The creation of National Parks, beginning in the 19th century, preserved some especially attractive and notable areas, but the pursuits of commerce, lifestyle, and recreation combined with increases in human population have continued to result in human modification of relatively untouched areas. Such human activity often negatively impacts native flora and fauna. As such, to better protect critical habitats and preserve low-impact recreational opportunities, legal concepts of "wilderness" were established in many countries, beginning with the United States (see below).



Cook Lake in the Bridger Wilderness, Bridger-Teton National Forest, Wyoming, U.S.

The first National Park was Yellowstone, which was signed into law by U.S. President Ulysses S. Grant on March 1, 1872.<sup>[10]</sup> The Act of Dedication declared Yellowstone a land "hereby reserved and withdrawn from settlement, occupancy, or sale under the laws of the United States, and dedicated and set apart as a public park or pleasuring ground for the benefit and enjoyment of the people."<sup>[11]</sup>



The world's second national park, the Royal National Park, located just 32 km to the south of Sydney, Australia, was established in 1879.<sup>[12]</sup>

The U.S. concept of national parks soon caught on in Canada, which created Banff National Park in the 1880s, at the same time as the transcontinental Canadian Pacific Railway was being built. The creation of this and other parks showed a growing appreciation of wild nature, but also an economic reality. The railways wanted to entice people to travel west. Parks such as Banff and Yellowstone gained favor as the railroads advertised travel to "the great wild spaces" of North America. When outdoorsman Teddy Roosevelt became president of the United States, he began to enlarge the U.S. National Parks system, and established the National Forest system.<sup>[8]</sup>

By the 1920s, travel across North America by train to experience the "wilderness" (often viewing it only through windows) had become wildly popular. This led to the commercialization of some of Canada's National Parks with the building of great hotels such as the Banff Springs Hotel and Chateau Lake Louise.

## Conservation vs. preservation



Bachalpsee in the Swiss Alps; generally mountainous areas are less affected by human activity

Two opposing factions had emerged within the environmental movement by the early 20th century: the conservationists and the preservationists. The conservationists (such as Gifford Pinchot) focused on the *proper use of nature*, whereas the preservationists sought the *protection of nature from use*.<sup>[9]</sup> Put another way, conservation sought to regulate human use while preservation sought to eliminate human impact altogether. Wilderness preservation is central to deep ecology; a philosophy that recognizes an inherent worth of all living beings, regardless of their instrumental utility to human needs.<sup>[13]</sup>

The idea of protecting nature for nature's sake began to gain more recognition in the 1930s with American writers like Aldo Leopold, calling for a "land ethic" and urging wilderness protection. It had become increasingly clear that wild spaces were disappearing rapidly and that decisive action was needed to save them.

Global conservation became an issue at the time of the dissolution of the British Empire in Africa in the late 1940s. The British established great wildlife preserves there. As before, this interest in conservation had an economic motive: in this case, big game hunting. Nevertheless, this led to growing recognition in the 1950s and the early 1960s of the need to protect large spaces for wildlife conservation worldwide. The World Wildlife Fund (WWF), founded in 1961, grew to be one of the largest conservation organizations in the world.<sup>[8]</sup>

Preservation again came to the fore in the 1960s with the publication of Rachel Carson's *Silent Spring* in 1962, which was the genesis of the modern environmental movement. Major environmental groups such as the Sierra Club shifted from protesting to working with politicians to influence environmental policy.<sup>[9]</sup>

Nevertheless, initiatives for wilderness conservation continue to increase. There are a growing number of projects to protect tropical rainforests through conservation initiatives. There are also large-scale projects to conserve wilderness regions, such as Canada's Boreal Forest Conservation Framework. The Framework calls for conservation of 50 percent of the 6,000,000 square kilometres of boreal forest in Canada's north.<sup>[14]</sup> In addition to the World Wildlife Fund, organizations such as The WILD Foundation, The Nature Conservancy, Conservation International, The Wilderness Society (United States) and many others are active in such conservation efforts.

## Formal Wilderness Designation

### New Zealand

There are seven wilderness areas in New Zealand as defined by the National Parks Act 1980 and the Conservation Act 1987 that fall well within the IUCN definition. Wilderness areas cannot have any human intervention and can only have indigenous species re-introduced into the area if it is compatible with conservation management strategies.

In New Zealand wilderness areas are remote blocks of land that have high natural character. The Conservation Act 1987 prevents any access by vehicles and livestock, the construction of tracks and buildings, and all indigenous natural resources are protected.<sup>[15]</sup> They are generally over 40,000 ha in size.<sup>[16]</sup>

### United States

*Main article:* National Wilderness Preservation System



**The Great Swamp** of New Jersey, donated for federal protection by concerned residents, was designated as the **first wilderness refuge in the United States** - winter scene photographed in March, 2008

In the United States, a Wilderness Area is an area of federal land set aside by an act of Congress. Human activities in wilderness areas are restricted to scientific study and non-mechanized recreation; horses are permitted but mechanized vehicles and equipment, such as cars and bicycles, are not.

The United States was the first country to officially designate land as "wilderness" through the Wilderness Act of 1964. The Wilderness Act was—and is still—an important part of wilderness designation because it created the legal definition of wilderness and founded the National Wilderness Preservation System. The Wilderness Act defines wilderness as “an area where the earth and its community of life are untrammelled by man, where man himself is a visitor who does not

remain.”<sup>[17]</sup>

Wilderness designation helps preserve the natural state of the land and protects flora and fauna by prohibiting development and providing for non-mechanized recreation only.

The first administratively protected wilderness area in the United States was the Gila National Forest. In 1922, Aldo Leopold, then a ranking member of the U.S. Forest Service, proposed a new management strategy for the Gila National Forest. His proposal was adopted in 1924, and 750 thousand acres of the Gila National Forest became the Gila Wilderness.<sup>[18]</sup>

**The Great Swamp in New Jersey** was the first formally designated wilderness refuge in the United States. It was declared a wildlife refuge on November 3, 1960. In 1966 it was declared a National Natural Landmark and, in 1968, it was given wilderness status. Properties in the swamp had been acquired by a small group of residents of the area, who donated the assembled properties to the federal government as a park for perpetual protection. Today the refuge amounts to 7600 acres (**unknown operator: u'strong'** km<sup>2</sup>) that are within thirty miles of Manhattan.<sup>[19]</sup>

While wilderness designations were originally granted by an Act of Congress for Federal land that retained a "primeval character", meaning that it had not suffered from human habitation or development, the Eastern Wilderness Act of 1975 extended the protection of the NWPS to areas in the eastern States that were not initially considered for inclusion in the Wilderness Act. This act allowed lands that did not meet the constraints of size, roadlessness, or human impact to be designated as wilderness areas under the belief that they could be returned to a "primeval" state through preservation.<sup>[20]</sup>

Approximately **unknown operator: u',** acres (**unknown operator: u'strong'unknown operator: u',**km<sup>2</sup>) are designated as wilderness in the United States. This accounts for 4.82% of the country's total land area; however, 54% of that amount is found in Alaska (recreation and development in Alaskan wilderness is often less restrictive), while only 2.58% of the lower continental United States is designated as wilderness. Following the Omnibus Public Land Management Act of 2009 there are 756 separate wilderness designations in the United States ranging in size from Florida's Pelican Island at 5 acres (**unknown operator: u'strong' m<sup>2</sup>**) to Alaska's Wrangell-Saint Elias at **unknown operator: u',** acres (**unknown operator: u'strong'unknown operator: u',**km<sup>2</sup>).

## Greece

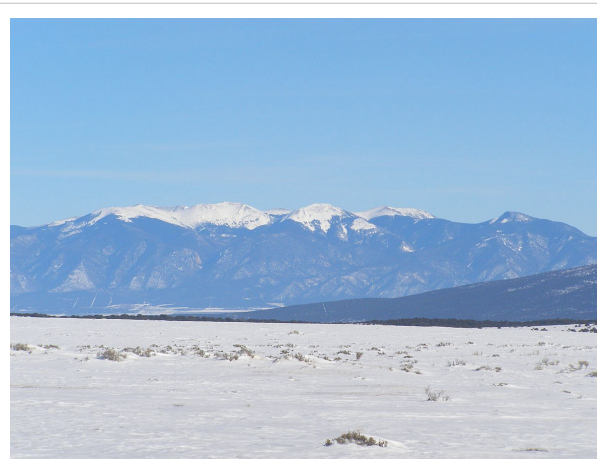
In Greece there are some parks called "ethniki drimoi" (εθνικοί δρυμοί, national forests) that are under protection of the Greek government. Such parks include: Olympus, Parnassos and Parnitha National Parks.

## International movement

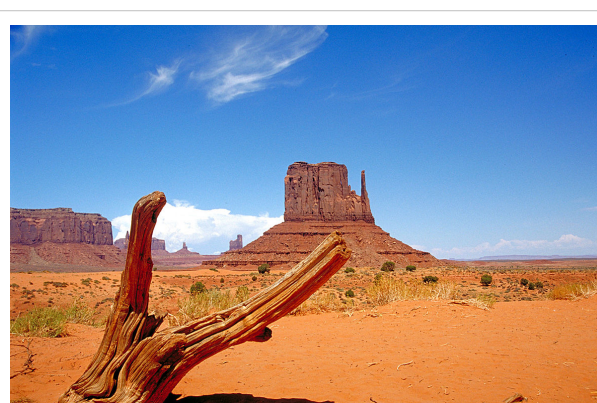
At the forefront of the international wilderness movement has been The WILD Foundation, its founder Ian Player and its network of sister and partner organizations around the globe. The pioneer World Wilderness Congress in 1977 introduced the wilderness concept as an issue of international importance, and began the process of defining the term in biological and social contexts. Today, this work is continued by many international groups who still look to the World Wilderness Congress as the international venue for wilderness and to The WILD Foundation network for wilderness tools and action. The WILD Foundation

also publishes the standard references for wilderness professionals and others involved in the issues: *Wilderness Management: Stewardship and Protection of Resources and Values*, the *International Journal of Wilderness*, *A Handbook on International Wilderness Law and Policy* and *Protecting Wild Nature on Native Lands* are the backbone of information and management tools for international wilderness issues.

The Wilderness Task Force within the World Commission on Protected Areas (WTF/WCPA) of the International Union for the Conservation of Nature (IUCN) plays a critical role in defining legal and management guidelines for wilderness at the international level and is also a clearing-house for information on wilderness issues. The IUCN



Latir Peak Wilderness, taken from milepost 394 along US-285, ten miles north of Tres Piedras and 14 miles south of the New Mexico and Colorado border.



Monument Valley in Utah, USA.



Protected Areas Classification System defines wilderness as “A large area of unmodified or slightly modified land, and/or sea retaining its natural character and influence, without permanent or significant habitation, which is protected and managed so as to preserve its natural condition (Category 1b).” The WILD Foundation founded the WTF/WCPA in 2002 and remains co-chair.

## Extent



The Ahklun Mountains and the Togiak Wilderness within the Togiak National Wildlife Refuge in the U.S. state of Alaska

According to a major study, *Wilderness: Earth's Last Wild Places*, carried out by Conservation International, 46% of the world's land mass is wilderness. For purposes of this report, "wilderness" was defined as an area that "has 70% or more of its original vegetation intact, covers at least **unknown operator: u','** square kilometers (**unknown operator: u'strong'unknown operator: u','**sq mi) and must have fewer than five people per square kilometer."<sup>[21]</sup> However, an IUCN/UNEP report published in 2003, found that only 10.9% of the world's land mass is currently a Category 1 Protected Area, that is, either a *strict nature reserve* (5.5%) or *protected wilderness* (5.4%).<sup>[22]</sup> Such areas

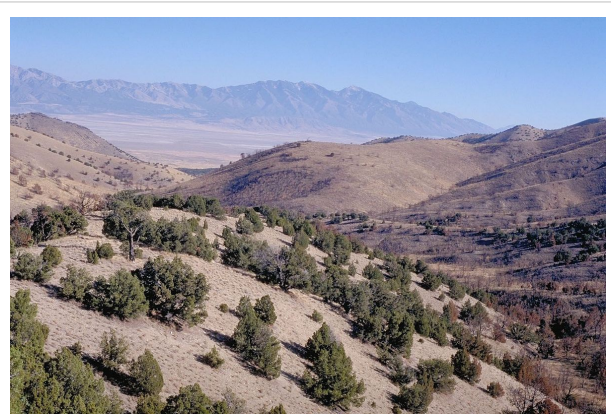
remain relatively untouched by humans. Of course, there are large tracts of lands in National Parks and other protected areas that would also qualify as wilderness. However, many protected areas have some degree of human modification or activity, so a definitive estimate of true wilderness is difficult.

The Wildlife Conservation Society generated a *human footprint* using a number of indicators, the absence of which indicate wildness: human population density, human access via roads and rivers, human infrastructure for agriculture and settlements and the presence of industrial power (lights visible from space). The society estimates that 26% of the Earth's land mass falls into the category of "Last of the wild." The wildest regions of the world include the tundra, the taiga, the Amazonian rain forest, the Tibetan Plateau, the Australian outback and deserts such as the Sahara, and the Gobi.<sup>[23]</sup> However from the 1970s, numerous geoglyphs have been discovered on deforested land in the Amazon rainforest, leading to claims about Pre-Columbian civilizations.<sup>[24][25]</sup> The BBC's *Unnatural Histories* claimed that the Amazon rainforest, rather than being a pristine wilderness, has been shaped by man for at least 11,000 years through practices such as forest gardening and terra preta.<sup>[26]</sup>

It should be noted that the percentage of land area designated "wilderness" does not reflect "quality" of remaining wilderness, part of which is barren areas with low biodiversity. Of the last natural wilderness areas, the taiga—which is mostly wilderness—represents 11% of the total land mass in the Northern Hemisphere.<sup>[27]</sup> Tropical rainforest represent a further 7% of the world's land base.<sup>[28]</sup> Estimates of the Earth's remaining wilderness underscore the rate at which these lands are being developed, with dramatic declines in biodiversity as a consequence.

## Critique

The American concept of wilderness has been criticized by some nature writers. For example, William Cronon writes that what he calls a wilderness ethic or cult may "teach us to be dismissive or even contemptuous of such humble places and experiences", and that "wilderness tends to privilege some parts of nature at the expense of others", using as an example "the mighty canyon more inspiring than the humble marsh."<sup>[29]</sup> This is most clearly visible with the fact that nearly all U.S. National Parks preserve spectacular canyons and mountains, and it was not until the 1940s that a swamp became a national park—the Everglades. In the mid-20th century national parks started to protect biodiversity, not simply attractive scenery.



Cedar Mountain Wilderness in northern Utah, USA.

Cronon also believes the passion to save wilderness "poses a serious threat to responsible environmentalism" and writes that it allows people to "give ourselves permission to evade responsibility for the lives we actually lead....to the extent that we live in an urban-industrial civilization" but at the same time pretend to ourselves that our real home is in the wilderness."<sup>[29]</sup>

Michael Pollan has argued that the wilderness ethic leads people to dismiss areas whose wildness is less than absolute. In his book *Second Nature*, Pollan writes that "once a landscape is no longer 'virgin' it is typically written off as fallen, lost to nature, irredeemable."<sup>[30]</sup> Another challenge to the conventional notion of wilderness comes from Robert Winkler in his book, *Going Wild: Adventures with Birds in the Suburban Wilderness*. "On walks in the unpeopled parts of the suburbs," Winkler writes, "I've witnessed the same wild creatures, struggles for survival, and natural beauty that we associate with true wilderness."<sup>[31]</sup>

Another source of criticism is that the criteria for wilderness designation is vague and open to interpretation. For example, the Wilderness Act states that wilderness must be roadless. The definition given for roadless is "the absences of roads which have been improved and maintained by mechanical means to insure relatively regular and continuous use."<sup>[32]</sup> However, there have been added sub-definitions that have, in essence, made this standard unclear and open to interpretation.

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This spiny forest at Ifaty, Madagascar features various *Adansonia* (baobab) species, *Alluaudia procera* (Madagascar ocotillo) and other vegetation.

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## Further reading

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- Gutkind, L (Ed). (2002). *On Nature: Great Writers on the Great Outdoors*. ISBN 1-58542-173-1

## External links

- Wilderness Information Network (<http://www.wilderness.net>)
- Wilderness Articles, Survival Techniques, Edible Plants (<http://www.wildsurvive.com>)
- CFACT ([http://www.cfact.org/site/view\\_article.asp?idCategory=11&idarticle=231](http://www.cfact.org/site/view_article.asp?idCategory=11&idarticle=231)) - news about wilderness areas
- Community of Outdoor Enthusiasts Provides Real Life Experiences and Educational Articles ([http://www.theskipper.ca/index.php?utm\\_campaign=Dictionaries&utm\\_content=Text-Link&utm\\_medium=Link&utm\\_source=Wikipedia](http://www.theskipper.ca/index.php?utm_campaign=Dictionaries&utm_content=Text-Link&utm_medium=Link&utm_source=Wikipedia))
- Aldo Leopold Wilderness Research Institute (<http://leopold.wilderness.net/>)
- Wilderness Task Force/World Commission on Protect Areas (<http://www.wildernesstaskforce.org/>)
- Campaign for America's Wilderness (<http://www.leaveitwild.org/>)

## Definitions

- Detailed maps of United States wilderness designations (<http://www.wilderness.net/index.cfm?fuse=NWPS>)
  - What is Wilderness? (<http://www.wilderdom.com/wilderness/WildernessDefinition.html>) - Definition & discussion of wilderness as a human construction
  - Wilderness and the American Mind (<http://www.erraticimpact.com/~ecologic/html/nash.htm>) - by Roderick Nash
  - The Trouble with Wilderness; or, Getting Back to the Wrong Nature (<http://www.utne.com/1996-05-01/GettingBacktotheWrongNature.aspx>) by William Cronon.
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# Natural environment

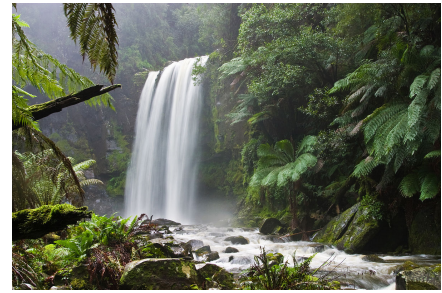
The **natural environment** encompasses all living and non-living things occurring naturally on Earth or some region thereof. It is an environment that encompasses the interaction of all living species.<sup>[1]</sup> The concept of the *natural environment* can be distinguished by components:

- Complete ecological units that function as natural systems without massive human intervention, including all vegetation, microorganisms, soil, rocks, atmosphere, and natural phenomena that occur within their boundaries.
- Universal natural resources and physical phenomena that lack clear-cut boundaries, such as air, water, and climate, as well as energy, radiation, electric charge, and magnetism, not originating from human activity.

The natural environment is contrasted with the built environment, which comprises the areas and components that are strongly influenced by humans. A geographical area is regarded as a natural environment.

It is difficult to find *absolutely natural* environments, and it is common that the naturalness varies in a continuum, from ideally 100% natural in one extreme to 0% natural in the other. More precisely, we can consider the different aspects or components of an environment, and see that their degree of naturalness is not uniform.<sup>[2]</sup> If, for instance, we take an agricultural field, and consider the mineralogic composition and the structure of its soil, we will find that whereas the first is quite similar to that of an undisturbed forest soil, the structure is quite different.

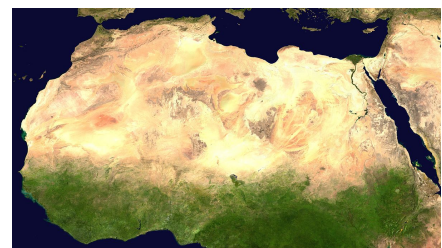
*Natural environment* is often used as a synonym for habitat. For instance, when we say that the natural environment of giraffes is the savanna.



Land management policies have been developed to preserve the natural characteristics of Hoptoun Falls, Australia while allowing ample access for visitors.

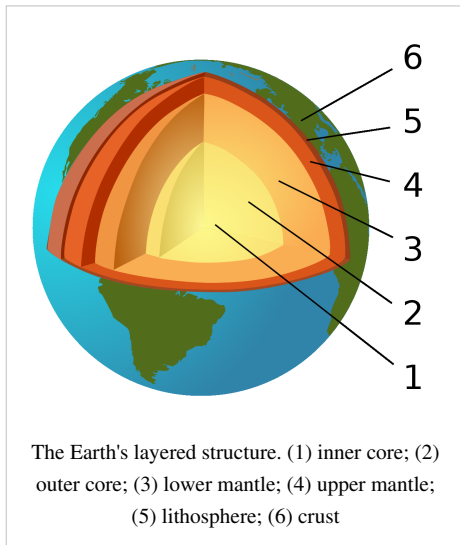


Bachalpsee in the Swiss Alps; generally mountainous areas are less affected by human activity.



A satellite image of the Sahara desert; the world's largest hot desert and third largest desert after Antarctica and the Arctic.

## Composition



Earth science generally recognizes 4 spheres, the lithosphere, the hydrosphere, the atmosphere, and the biosphere<sup>[3]</sup> as correspondent to rocks, water, air, and life. Some scientists include, as part of the spheres of the Earth, the cryosphere (corresponding to ice) as a distinct portion of the hydrosphere, as well as the pedosphere (corresponding to soil) as an active and intermixed sphere. Earth science (also known as geoscience, the geosciences or the Earth Sciences), is an all-embracing term for the sciences related to the planet Earth.<sup>[4]</sup> There are four major disciplines in earth sciences, namely geography, geology, geophysics and geodesy. These major disciplines use physics, chemistry, biology, chronology and mathematics to build a qualitative and quantitative understanding of the principal areas or *spheres* of the Earth system.

## Geological activity



A volcanic fissure and lava channel.

The Earth's crust, or lithosphere, is the outermost solid surface of the planet and is chemically and mechanically different from underlying mantle. It has been generated largely by igneous processes in which magma (molten rock) cools and solidifies to form solid rock. Beneath the lithosphere lies the mantle which is heated by the decay of radioactive elements. The mantle though solid is in a state of rheic convection. This convection process causes the lithospheric plates to move, albeit slowly. The resulting process is known as plate tectonics.<sup>[5][6][7]</sup> Volcanoes result primarily from the melting of subducted crust material or of rising mantle at mid-ocean ridges and

mantle plumes.

## Water on Earth

### Oceans

An ocean is a major body of saline water, and a component of the hydrosphere. Approximately 71% of the Earth's surface (an area of some 362 million square kilometers) is covered by ocean, a continuous body of water that is customarily divided into several principal oceans and smaller seas. More than half of this area is over 3,000 meters (9,800 ft) deep. Average oceanic salinity is around 35 parts per thousand (ppt) (3.5%), and nearly all seawater has a salinity in the range of 30 to 38 ppt. Though generally recognized as several 'separate' oceans, these waters comprise one global, interconnected body of salt water often referred to as the World Ocean or global ocean.<sup>[8][9]</sup> This concept of a global ocean as a continuous body of water with relatively free interchange among its parts is of fundamental importance to oceanography.<sup>[10]</sup> The major oceanic divisions are defined in part by the continents, various archipelagos, and other criteria: these divisions are (in descending order of size) the Pacific Ocean, the Atlantic Ocean, the Indian Ocean, the Southern Ocean and the Arctic Ocean.



Coral reefs have significant marine biodiversity.



## Rivers



The Columbia River, along the border of the United States U.S. states of Oregon and Washington (U.S. state) Washington.

A river is a natural watercourse,<sup>[11]</sup> usually freshwater, flowing toward an ocean, a lake, a sea or another river. In a few cases, a river simply flows into the ground or dries up completely before reaching another body of water. Small rivers may also be termed by several other names, including stream, creek and brook. In the United States a river is generally classified as a watercourse more than 60 feet (18 metres) wide. The water in a river is usually in a channel, made up of a stream bed between banks. In larger rivers there is also a wider floodplain shaped by waters over-topping the channel. Flood plains may be very wide in relation to the size of the river channel. Rivers are a part of the hydrological cycle. Water within a river is generally collected from precipitation through surface runoff, groundwater recharge, springs, and the release of water stored in glaciers and snowpacks.



A rocky stream in the U.S. state of Hawaii.

### Streams

A stream is a flowing body of water with a current, confined within a bed and stream banks. Streams play an important corridor role in connecting fragmented habitats and thus in conserving biodiversity. The study of streams and waterways in general is known as *surface hydrology*.<sup>[12]</sup> Types of streams include creeks, tributaries, which do not reach an ocean and connect with another stream or river, brooks, which are typically small streams and sometimes sourced from a spring or seep and tidal inlets.

## Lakes

A lake (from Latin *lacus*) is a terrain feature, a body of water that is localized to the bottom of basin. A body of water is considered a lake when it is inland, is not part of an ocean, is larger and deeper than a pond, and is fed by a river.<sup>[13][14]</sup>

Natural lakes on Earth are generally found in mountainous areas, rift zones, and areas with ongoing or recent glaciation. Other lakes are found in endorheic basins or along the courses of mature rivers. In some parts of the world, there are many lakes because of chaotic drainage patterns left over from the last Ice Age. All lakes are temporary over geologic time scales, as they will slowly fill in with sediments or spill out of the basin containing them.



The Lácar Lake is a lake of glacial origin in the province of Neuquén, Argentina.

### Ponds

A pond is a body of standing water, either natural or man-made, that is usually smaller than a lake. A wide variety of man-made bodies of water are classified as ponds, including water gardens designed for aesthetic ornamentation, fish ponds designed for commercial fish



A swamp area in Everglades National Park, Florida, USA.

breeding, and solar ponds designed to store thermal energy. Ponds and lakes are distinguished from streams via current speed. While currents in streams are easily observed, ponds and lakes possess thermally driven micro-currents and moderate wind driven currents. These features distinguish a pond from many other aquatic terrain features, such as stream pools and tide pools.

## Atmosphere, climate and weather

The atmosphere of the Earth serves as a key factor in sustaining the planetary ecosystem. The thin layer of gases that envelops the Earth is held in place by the planet's gravity. Dry air consists of 78% nitrogen, 21% oxygen, 1% argon and other inert gases, such as carbon dioxide. The remaining gases are often referred to as trace gases,<sup>[16]</sup> among which are the greenhouse gases such as water vapor, carbon dioxide, methane, nitrous oxide, and ozone. Filtered air includes trace amounts of many other chemical compounds. Air also contains a variable amount of water vapor and suspensions of water droplets and ice crystals seen as clouds. Many natural substances may be present in tiny amounts in an unfiltered air sample, including dust, pollen and spores, sea spray, volcanic ash, and meteoroids. Various industrial pollutants also may be present, such as chlorine (elementary or in compounds), fluorine compounds, elemental mercury, and sulphur compounds such as sulphur dioxide [SO<sub>2</sub>].

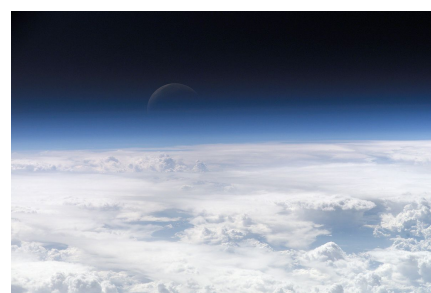
The ozone layer of the Earth's atmosphere plays an important role in depleting the amount of ultraviolet (UV) radiation that reaches the surface. As DNA is readily damaged by UV light, this serves to protect life at the surface. The atmosphere also retains heat during the night, thereby reducing the daily temperature extremes.

### Atmospheric layers

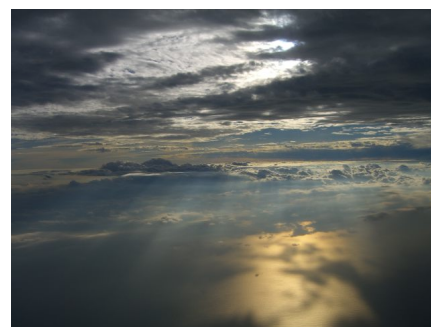
#### Principal layers

Earth's atmosphere can be divided into five main layers. These layers are mainly determined by whether temperature increases or decreases with altitude. From highest to lowest, these layers are:

- Exosphere: The outermost layer of Earth's atmosphere extends from the exobase upward, mainly composed of hydrogen and helium.

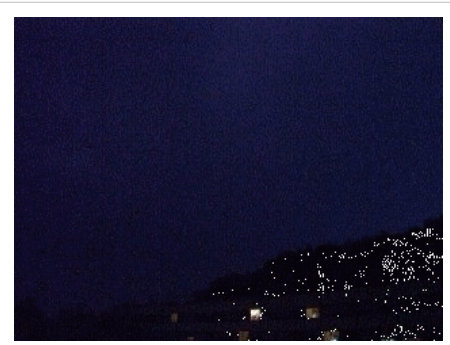


Atmospheric gases scatter blue light more than other wavelengths, creating a blue halo when seen from space.



A view of Earth's troposphere from an airplane.

- Thermosphere: The top of the thermosphere is the bottom of the exosphere, called the exobase. Its height varies with solar activity and ranges from about 350–800 km (**unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'; unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'**). The International Space Station orbits in this layer, between 320 and 380 km (**unknown operator: u'strong' and unknown operator: u'strong' mi**).
- Mesosphere: The mesosphere extends from the stratopause to 80–85 km (**unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'; unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'**). It is the layer where most meteors burn up upon entering the atmosphere.
- Stratosphere: The stratosphere extends from the tropopause to about 51 km (**unknown operator: u'strong' mi; unknown operator: u'strong' ft**). The stratopause, which is the boundary between the stratosphere and mesosphere, typically is at 50 to 55 km (**unknown operator: u'strong' to unknown operator: u'strong' mi; unknown operator: u'strong' to unknown operator: u'strong' ft**).
- Troposphere: The troposphere begins at the surface and extends to between 7 km (**unknown operator: u'strong' ft**) at the poles and 17 km (**unknown operator: u'strong' ft**) at the equator, with some variation due to weather. The troposphere is mostly heated by transfer of energy from the surface, so on average the lowest part of the troposphere is warmest and temperature decreases with altitude. The tropopause is the boundary between the troposphere and stratosphere.



Lightning is an atmospheric discharge of electricity accompanied by thunder, which typically occurs during thunderstorms, and sometimes during volcanic eruptions or dust storms.<sup>[15]</sup>

#### Other layers

Within the five principal layers determined by temperature are several layers determined by other properties.

- The ozone layer is contained within the stratosphere. It is mainly located in the lower portion of the stratosphere from about 15–35 km (**unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'; unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong'**), though the thickness varies seasonally and geographically. About 90% of the ozone in our atmosphere is contained in the stratosphere.
- The ionosphere, the part of the atmosphere that is ionized by solar radiation, stretches from 50 to 1000 km (**unknown operator: u'strong' to unknown operator: u'strong' mi; unknown operator: u'strong' to unknown operator: u'strong' ft**) and typically overlaps both the exosphere and the thermosphere. It forms the inner edge of the magnetosphere.
- The homosphere and heterosphere: The homosphere includes the troposphere, stratosphere, and mesosphere. The upper part of the heterosphere is composed almost completely of hydrogen, the lightest element.
- The planetary boundary layer is the part of the troposphere that is nearest the Earth's surface and is directly affected by it, mainly through turbulent diffusion.



## Effects of global warming



The Retreat of glaciers since 1850 of Aletsch Glacier in the Swiss Alps (situation in 1979, 1991 and 2002), due to global warming.

The potential dangers of global warming are being increasingly studied by a wide global consortium of scientists. These scientists are increasingly concerned about the potential long-term effects of global warming on our natural environment and on the planet. Of particular concern is how climate change and global warming caused by anthropogenic, or human-made releases of greenhouse gases, most notably carbon dioxide, can act interactively, and have adverse effects upon the planet, its natural environment and humans' existence. Efforts have been increasingly focused on the mitigation of greenhouse gases that are causing climatic changes, on developing adaptative strategies

to global warming, to assist humans, animal and plant species, ecosystems, regions and nations in adjusting to the effects of global warming. Some examples of recent collaboration to address climate change and global warming include:

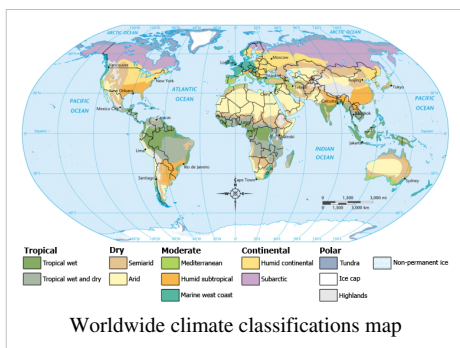
- The United Nations Framework Convention Treaty and convention on Climate Change, to stabilize greenhouse gas concentrations in the atmosphere at a level that would prevent dangerous anthropogenic interference with the climate system.<sup>[17]</sup>
- The Kyoto Protocol, which is the protocol to the international Framework Convention on Climate Change treaty, again with the objective of reducing greenhouse gases in an effort to prevent anthropogenic climate change.<sup>[18]</sup>
- The Western Climate Initiative, to identify, evaluate, and implement collective and cooperative ways to reduce greenhouse gases in the region, focusing on a market-based cap-and-trade system.<sup>[19]</sup>



Another view of the Aletsch Glacier in the Swiss Alps and because of global warming it has been decreasing

A significantly profound challenge is to identify the natural environmental dynamics in contrast to environmental changes not within natural variances. A common solution is to adapt a static view neglecting natural variances to exist. Methodologically, this view could be defended when looking at processes which change slowly and short time series, while the problem arrives when fast processes turns essential in the object of the study.

## Climate



Climate encompasses the statistics of temperature, humidity, atmospheric pressure, wind, rainfall, atmospheric particle count and numerous other meteorological elements in a given region over long periods of time. Climate can be contrasted to weather, which is the present condition of these same elements over periods up to two weeks.

Climates can be classified according to the average and typical ranges of different variables, most commonly temperature and precipitation.

The most commonly used classification scheme is the one originally developed by Wladimir Köppen. The Thornthwaite system,<sup>[20]</sup> in use since 1948, incorporates evapotranspiration in addition to temperature and precipitation information and is used in studying animal species diversity and potential impacts of climate changes.

## Weather

Weather is a set of all the phenomena occurring in a given atmospheric area at a given time.<sup>[21]</sup> Most weather phenomena occur in the troposphere,<sup>[22][23]</sup> just below the stratosphere. Weather refers, generally, to day-to-day temperature and precipitation activity, whereas climate is the term for the average atmospheric conditions over longer periods of time.<sup>[24]</sup> When used without qualification, "weather" is understood to be the weather of Earth.

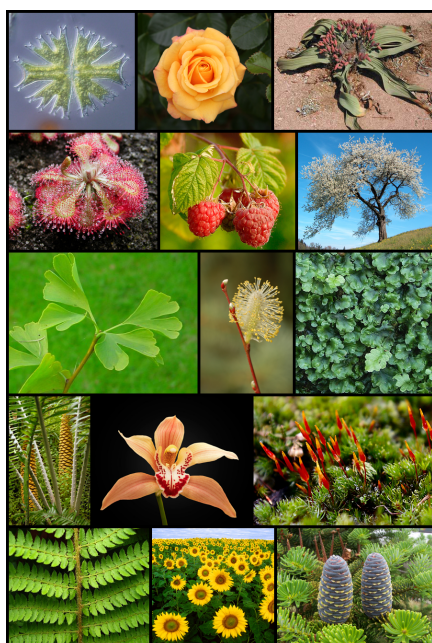
Weather occurs due to density (temperature and moisture) differences between one place and another. These differences can occur due to the sun angle at any particular spot, which varies by latitude from the tropics. The strong temperature contrast between polar and tropical air gives rise to the jet stream. Weather systems in the mid-latitudes, such as extratropical cyclones, are caused by instabilities of the jet stream flow. Because the Earth's axis is tilted relative to its orbital plane, sunlight is incident at different angles at different times of the year. On the Earth's surface, temperatures usually range  $\pm 40^{\circ}\text{C}$  ( $100^{\circ}\text{F}$  to  $-40^{\circ}\text{F}$ ) annually. Over thousands of years, changes in the Earth's orbit have affected the amount and distribution of solar energy received by the Earth and influence long-term climate

Surface temperature differences in turn cause pressure differences. Higher altitudes are cooler than lower altitudes due to differences in compressional heating. Weather forecasting is the application of science and technology to predict the state of the atmosphere for a future time and a given location. The atmosphere is a chaotic system, and small changes to one part of the system can grow to have large effects on the system as a whole. Human attempts to control the weather have occurred throughout human history, and there is evidence that human activity such as agriculture and industry has inadvertently modified weather patterns.



Rainbows are optical and meteorological phenomenon that causes a spectrum of light to appear in the sky when the Sun shines onto droplets of moisture in the Earth's atmosphere.

## Life



There are many plant species on the planet.

Evidence suggest that life on Earth has existed for about 3.7 billion years.<sup>[25]</sup> All known life forms share fundamental molecular mechanisms, and based on these observations, theories on the origin of life attempt to find a mechanism explaining the formation of a primordial single cell organism from which all life originates. There are many different hypotheses regarding the path that might have been taken from simple organic molecules via pre-cellular life to protocells and metabolism.

Although there is no universal agreement on the definition of life, scientists generally accept that the biological manifestation of life is characterized by organization, metabolism, growth, adaptation, response to stimuli and reproduction.<sup>[26]</sup> Life may also be said to be simply the characteristic state of organisms. In biology, the science of living organisms, "life" is the condition which distinguishes active organisms from inorganic matter, including the capacity for growth, functional activity and the continual change preceding death.<sup>[27][28]</sup>

A diverse array of living organisms (life forms) can be found in the biosphere on Earth, and properties common to these organisms—plants, animals, fungi, protists, archaea, and bacteria—are a carbon- and water-based cellular form with complex organization and heritable genetic information. Living organisms undergo metabolism, maintain homeostasis, possess a capacity to grow, respond to stimuli, reproduce and, through natural selection, adapt to their environment in successive generations. More complex living organisms can communicate through various means.



An example of the many animal species on the Earth.

## Ecosystems



Rainforests often have a great deal of biodiversity with many plant and animal species. This is the Gambia River in Senegal's Niokolo-Koba National Park.

An ecosystem (also called as environment) is a natural unit consisting of all plants, animals and micro-organisms (biotic factors) in an area functioning together with all of the non-living physical (abiotic) factors of the environment.<sup>[29]</sup>

Central to the ecosystem concept is the idea that living organisms are continually engaged in a highly interrelated set of relationships with every other element constituting the environment in which they exist. Eugene Odum, one of the founders of the science of ecology, stated: "Any unit that includes all of the organisms (ie: the "community") in a given area interacting with the physical environment so that a flow of energy leads to clearly defined trophic structure, biotic diversity, and material cycles (i.e.: exchange of materials between living and

nonliving parts) within the system is an ecosystem."<sup>[30]</sup>



The human ecosystem concept is then grounded in the deconstruction of the human/nature dichotomy, and the emergent premise that all species are ecologically integrated with each other, as well as with the abiotic constituents of their biotope.

A greater number or variety of species or biological diversity of an ecosystem may contribute to greater resilience of an ecosystem, because there are more species present at a location to respond to change and thus "absorb" or reduce its effects. This reduces the effect before the ecosystem's structure is fundamentally changed to a different state. This is not universally the case and there is no proven relationship between the species diversity of an ecosystem and its ability to provide goods and services on a sustainable level.

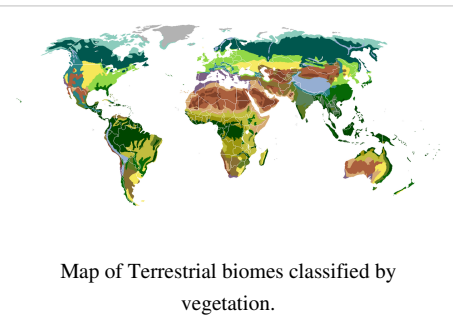
The term ecosystem can also pertain to human-made environments, such as human ecosystems and human-influenced ecosystems, and can describe any situation where there is relationship between living organisms and their environment. Fewer areas on the surface of the earth today exist free from human contact, although some genuine wilderness areas continue to exist without any forms of human intervention.



Old-growth forest and a creek on Larch Mountain, in the U.S. state of Oregon.

## Biomes

Biomes are terminologically similar to the concept of ecosystems, and are climatically and geographically defined areas of ecologically similar climatic conditions on the Earth, such as communities of plants, animals, and soil organisms, often referred to as ecosystems. Biomes are defined on the basis of factors such as plant structures (such as trees, shrubs, and grasses), leaf types (such as broadleaf and needleleaf), plant spacing (forest, woodland, savanna), and climate. Unlike ecozones, biomes are not defined by genetic, taxonomic, or historical similarities. Biomes are often identified with particular patterns of ecological succession and climax vegetation.

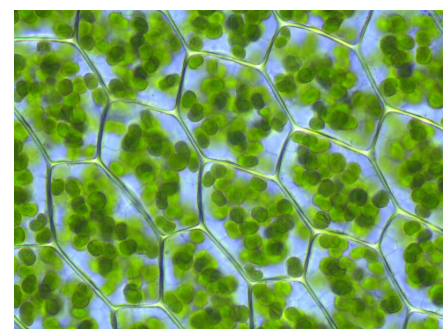


Map of Terrestrial biomes classified by vegetation.

## Biogeochemical cycles

Global biogeochemical cycles are critical to life, most notably those of water, oxygen, carbon, nitrogen and phosphorus.<sup>[31]</sup>

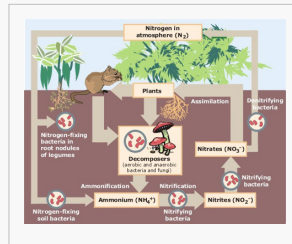
- The nitrogen cycle is the transformation of nitrogen and nitrogen-containing compounds in nature. It is a cycle which includes gaseous components.
- The water cycle, is the continuous movement of water on, above, and below the surface of the Earth. Water can change states among liquid, vapor, and ice at various places in the water cycle. Although the balance of water on Earth remains fairly constant over time, individual water molecules can come and go.
- The carbon cycle is the biogeochemical cycle by which carbon is exchanged among the biosphere, pedosphere, geosphere, hydrosphere, and atmosphere of the Earth.



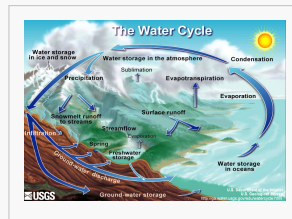
Chloroplasts conduct photosynthesis and are found in plant cells and other eukaryotic organisms. These are Chloroplasts visible in the cells of *Plagiomnium affine* — Many-fruited Thyme-moss.

- The oxygen cycle is the movement of oxygen within and between its three main reservoirs: the atmosphere, the biosphere, and the lithosphere. The main driving factor of the oxygen cycle is photosynthesis, which is responsible for the modern Earth's atmospheric composition and life.
- The phosphorus cycle is the movement of phosphorus through the lithosphere, hydrosphere, and biosphere. The atmosphere does not play a significant role in the movements of phosphorus, because phosphorus and phosphorus compounds are usually solids at the typical ranges of temperature and pressure found on Earth.

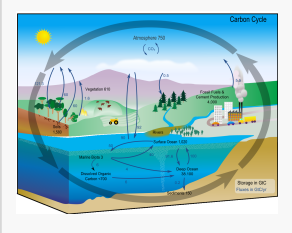
### Biogeochemical cycles



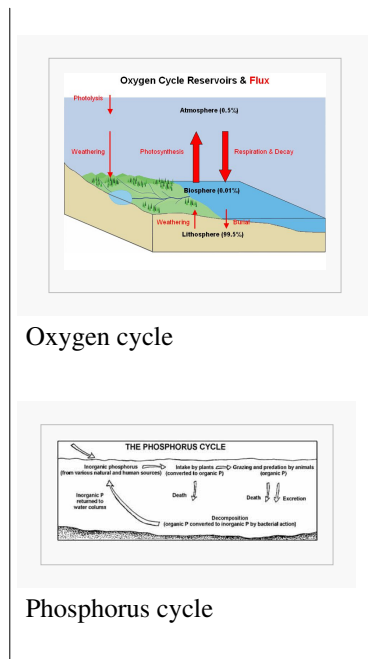
### Nitrogen cycle



### Water cycle



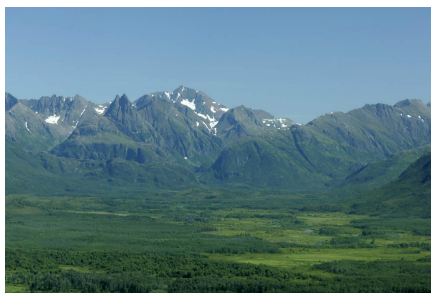
### Carbon cycle



## Wilderness



A conifer forest in the Swiss Alps (National Park).



The Ahklun Mountains and the Togiak Wilderness within the Togiak National Wildlife Refuge in the U.S. state of Alaska.

Wilderness is generally defined as a natural environment on Earth that has not been significantly modified by human activity. The WILD Foundation goes into more detail, defining wilderness as: "The most intact, undisturbed wild natural areas left on our planet - those last truly wild places that humans do not control and have not developed with roads, pipelines or other industrial infrastructure."<sup>[32]</sup> Wilderness areas and protected parks are considered important for the survival of certain species, ecological studies, conservation, solitude, and recreation. Wilderness is deeply valued for cultural, spiritual, moral, and aesthetic reasons. Some nature writers believe wilderness areas are vital for the human spirit and creativity.<sup>[33]</sup>

The word, "wilderness", derives from the notion of wildness; in other words that which is not controllable by humans. The word's etymology is from the Old English *wildeornes*, which in turn derives from *wildeor* meaning *wild beast* (wild + deor = beast, deer).<sup>[34]</sup> From this point of view, it is the wildness of a place that makes it a wilderness. The mere presence or activity of people does not disqualify an area from being "wilderness." Many ecosystems that are, or have been, inhabited or influenced by activities of people may still be considered "wild." This way of looking at wilderness includes areas within which natural processes operate without very noticeable human interference.

Wildlife includes all non-domesticated plants, animals and other organisms. Domesticating wild plant and animal species for human benefit has occurred many times all over the planet, and has a major impact on the environment, both positive and negative. Wildlife can be found in all ecosystems. Deserts, rain forests, plains, and other areas—including the most

developed urban sites—all have distinct forms of wildlife. While the term in popular culture usually refers to animals that are untouched by human factors, most scientists agree that wildlife around the world is impacted by human activities.

## Challenges



Before flue-gas desulfurization was installed, the air-polluting emissions from this power plant in New Mexico contained excessive amounts of sulfur dioxide



Amazon Rainforest in Brazil. The tropical rainforests of South America contain the largest diversity of species on Earth, including some that have evolved within the past few hundred thousand years.<sup>[35][36]</sup>

It is the common understanding of *natural environment* that underlies environmentalism — a broad political, social, and philosophical movement that advocates various actions and policies in the interest of protecting what nature remains in the natural environment, or restoring or expanding the role of nature in this environment. While true wilderness is increasingly rare, *wild* nature (e.g., unmanaged forests, uncultivated grasslands, wildlife, wildflowers) can be found in many locations previously inhabited by humans.

Goals commonly expressed by environmental scientists include:

- Reduction and clean up of pollution, with future goals of zero pollution;
- Cleanly converting non-recyclable materials into energy through direct combustion or after conversion into secondary fuels;
- Reducing societal consumption of non-renewable fuels;
- Development of alternative, green, low-carbon or renewable energy sources;
- Conservation and sustainable use of scarce resources such as water, land, and air;
- Protection of representative or unique or pristine ecosystems;
- Preservation of threatened and endangered species extinction;
- The establishment of nature and biosphere reserves under various types of protection; and, most generally, the protection of biodiversity and ecosystems upon which all human and other life on earth depends.

Very large development projects - megaprojects - pose special instructions and risks to the natural environments. Major dams and power plants are cases in point. The challenge to the environment from such projects is growing because more and bigger megaprojects are being built, in developed and developing nations alike.<sup>[37]</sup>

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## External links

- UNEP - United Nations Environment Programme (<http://www.unep.org/>)
- BBC - Science and Nature (<http://www.bbc.co.uk/sn/>).
- Science.gov - Environment & Environmental Quality ([http://www.science.gov/browse/w\\_123.htm](http://www.science.gov/browse/w_123.htm))



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# Matter and energy

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## Matter

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**Matter** is generally considered to be anything that has mass and volume. The volume is determined by the space in three dimensions that it occupies. The mass is determined by its rest mass (or invariant mass), which is measured by the acceleration a body has when a force is applied. The greater the mass, the slower the acceleration for the same force<sup>[1]</sup>. Matter is thus a general term for the substance of which all observable physical objects consist.<sup>[2][3]</sup> Typically, matter includes atoms and other particles that have mass, but this definition confuses mass and matter, which are not the same.<sup>[4]</sup> Different fields use the term in different and sometimes incompatible ways; there is no single agreed scientific meaning of the word "matter," even though the term "mass" is better-defined.

For much of the history of the natural sciences people have contemplated the exact nature of matter. The idea that matter was built of discrete building blocks, the so-called *particulate theory of matter*, was first put forward by the Greek philosophers Leucippus (~490 BC) and Democritus (~470–380 BC).<sup>[5]</sup> All objects we see with the naked eye are composed of molecules, and molecules consist of atoms, atomic matter, made up of interacting subatomic particles, usually a nucleus of protons, neutrons and a cloud of orbiting electrons.<sup>[6][7]</sup> Einstein showed<sup>[8]</sup> that ultimately all matter is capable of being converted to energy, by the formula

$$E = mc^2$$

Where E is the energy of a piece of matter of mass m, times the  $c^2$  the speed of light squared (multiplied by itself). As the speed of light is exactly 299,792,458 metres per second or 186,272 miles per second, just a little matter is equivalent to a lot of energy.

When atomic matter is hot enough it ionises (or loses its electrons, usually above about 5,000 degrees C) and this causes it to emit the energy of light. The light we see by in this way usually comes from the sun (or other stars at night), either by daylight or by fossil fuels, stored from sunlight captured many millions of years ago. Atomic matter at lower temperatures also can reflect light, absorbing some at specific wavelengths, which determines the colours of the objects we see<sup>[9]</sup>

Matter is commonly said to exist in four *states* (or *phases*): solid, liquid, gas and plasma. However, advances in experimental techniques have realized other phases, previously only theoretical constructs, such as Bose–Einstein condensates and fermionic condensates. A focus on an elementary-particle view of matter also leads to new phases of matter, such as the quark–gluon plasma.<sup>[10]</sup>

In physics and chemistry, matter exhibits both wave-like and particle-like properties, the so-called wave–particle duality.<sup>[11][12][13]</sup>

In the realm of cosmology, extensions of the term *matter* are invoked to include dark matter and dark energy, concepts introduced to explain some odd phenomena of the observable universe, such as the galactic rotation curve. These exotic forms of "matter" do not refer to matter as "building blocks", but rather to currently poorly understood forms of mass and energy.<sup>[14]</sup>

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## Definitions

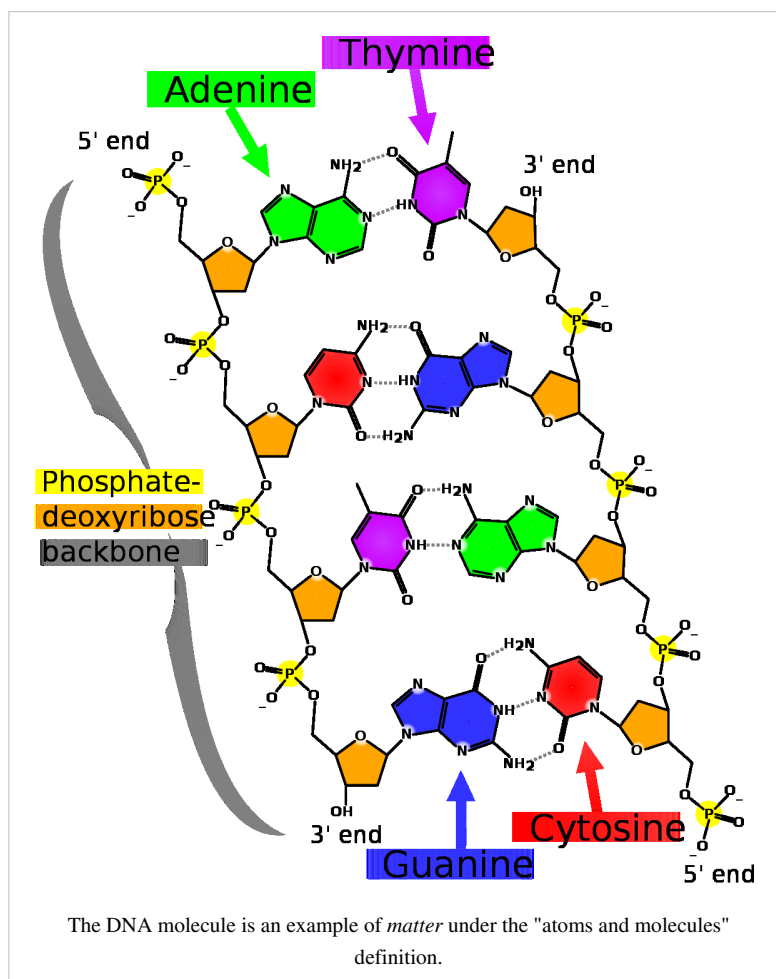
### Common definition

The common definition of matter is *anything that has both mass and volume (occupies space)*.<sup>[15][16]</sup> For example, a car would be said to be made of matter, as it occupies space, and has mass.

The observation that matter occupies space goes back to antiquity. However, an explanation for why matter occupies space is recent, and is argued to be a result of the Pauli exclusion principle.<sup>[17][18]</sup> Two particular examples where the exclusion principle clearly relates matter to the occupation of space are white dwarf stars and neutron stars, discussed further below.

### Relativity

In the context of relativity, mass is not an additive quantity.<sup>[2]</sup> Thus, in relativity usually a more general view is taken that it is not mass, but the energy–momentum tensor that quantifies the amount of matter. Matter therefore is anything that contributes to the energy–momentum of a system, that is, anything that is not purely gravity.<sup>[19][20]</sup> This view is commonly held in fields that deal with general relativity such as cosmology.



### Atoms and molecules definition

A definition of "matter" that is based upon its physical and chemical structure is: *matter is made up of atoms and molecules*.<sup>[21]</sup> As an example, deoxyribonucleic acid molecules (DNA) are matter under this definition because they are made of atoms. This definition can be extended to include charged atoms and molecules, so as to include plasmas (gases of ions) and electrolytes (ionic solutions), which are not obviously included in the atoms and molecules definition. Alternatively, one can adopt the *protons, neutrons and electrons* definition.

### Protons, neutrons and electrons definition

A definition of "matter" more fine-scale than the atoms and molecules definition is: *matter is made up of what atoms and molecules are made of*, meaning anything made of positively charged protons, neutral neutrons, and negatively charged electrons.<sup>[22]</sup> This definition goes beyond atoms and molecules, however, to include substances made from these building blocks that are *not* simply atoms or molecules, for example white dwarf matter — typically, carbon and oxygen nuclei in a sea of degenerate electrons. At a microscopic level, the constituent "particles" of matter such as protons, neutrons and electrons obey the laws of quantum mechanics and exhibit wave–particle duality. At an even deeper level, protons and neutrons are made up of quarks and the force fields (gluons) that bind them together

(see Quarks and leptons definition below).

## Quarks and leptons definition

As may be seen from the above discussion, many early definitions of what can be called *ordinary matter* were based upon its structure or "building blocks". On the scale of elementary particles, a definition that follows this tradition can be stated as: *ordinary matter is everything that is composed of elementary fermions, namely quarks and leptons.*<sup>[23][24]</sup> The connection between these formulations follows.

Leptons (the most famous being the electron), and quarks (of which baryons, such as protons and neutrons, are made) combine to form atoms, which in turn form molecules. Because atoms and molecules are said to be matter, it is natural to phrase the definition as: *ordinary matter is anything that is made of the same things that atoms and molecules are made of.* (However, notice that one also can make from these building blocks matter that is *not* atoms or molecules.) Then, because electrons

are leptons, and protons and neutrons are made of quarks, this definition in turn leads to the definition of matter as being "quarks and leptons", which are the two types of elementary fermions. Carithers and Grannis state: *Ordinary matter is composed entirely of first-generation particles, namely the [up] and [down] quarks, plus the electron and its neutrino.*<sup>[25]</sup> (Higher generations particles quickly decay into first-generation particles, and thus are not commonly encountered.<sup>[26]</sup>)

This definition of ordinary matter is more subtle than it first appears. All the particles that make up ordinary matter (leptons and quarks) are elementary fermions, while all the force carriers are elementary bosons.<sup>[27]</sup> The W and Z bosons that mediate the weak force are not made of quarks or leptons, and so are not ordinary matter, even if they have mass.<sup>[28]</sup> In other words, mass is not something that is exclusive to ordinary matter.

The quark-lepton definition of ordinary matter, however, identifies not only the elementary building blocks of matter, but also includes composites made from the constituents (atoms and molecules, for example). Such composites contain an interaction energy that holds the constituents together, and may constitute the bulk of the mass of the composite. As an example, to a great extent, the mass of an atom is simply the sum of the masses of its constituent protons, neutrons and electrons. However, digging deeper, the protons and neutrons are made up of quarks bound together by gluon fields (see dynamics of quantum chromodynamics) and these gluons fields contribute significantly to the mass of hadrons.<sup>[29]</sup> In other words, most of what composes the "mass" of ordinary

Three generations of matter (fermions)				
	I	II	III	
mass →	2.4 MeV/c <sup>2</sup>	1.27 GeV/c <sup>2</sup>	171.2 GeV/c <sup>2</sup>	0
charge →	$\frac{2}{3}$	$\frac{2}{3}$	$\frac{2}{3}$	0
spin →	$\frac{1}{2}$	$\frac{1}{2}$	$\frac{1}{2}$	1
name →	u up	c charm	t top	γ photon
Quarks	4.8 MeV/c <sup>2</sup>	104 MeV/c <sup>2</sup>	4.2 GeV/c <sup>2</sup>	0
	$-\frac{1}{3}$	$-\frac{1}{3}$	$-\frac{1}{3}$	0
	$\frac{1}{2}$	$\frac{1}{2}$	$\frac{1}{2}$	1
	d down	s strange	b bottom	g gluon
Leptons	<2.2 eV/c <sup>2</sup>	<0.17 MeV/c <sup>2</sup>	<15.5 MeV/c <sup>2</sup>	91.2 GeV/c <sup>2</sup>
	0	0	0	0
	$\frac{1}{2}$	$\frac{1}{2}$	$\frac{1}{2}$	1
	ν <sub>e</sub> electron neutrino	ν <sub>μ</sub> muon neutrino	ν <sub>τ</sub> tau neutrino	Z <sup>0</sup> Z boson
Gauge bosons	0.511 MeV/c <sup>2</sup>	105.7 MeV/c <sup>2</sup>	1.777 GeV/c <sup>2</sup>	80.4 GeV/c <sup>2</sup>
	-1	-1	-1	±1
	$\frac{1}{2}$	$\frac{1}{2}$	$\frac{1}{2}$	1
	e electron	μ muon	τ tau	W <sup>±</sup> W boson

Under the "quarks and leptons" definition, the elementary and composite particles made of the quarks (in purple) and leptons (in green) would be "matter"; while the gauge bosons (in red) would not be "matter". However, interaction energy inherent to composite particles (for example, gluons involved in neutrons and protons) contribute to the mass of ordinary matter.

matter is due to the binding energy of quarks within protons and neutrons.<sup>[30]</sup> For example, the sum of the mass of the three quarks in a nucleon is approximately  $12.5 \text{ MeV}/c^2$ , which is low compared to the mass of a nucleon (approximately  $938 \text{ MeV}/c^2$ ).<sup>[26][31]</sup> The bottom line is that most of the mass of everyday objects comes from the interaction energy of its elementary components.

Smaller building blocks?

The Standard Model groups matter particles into three generations, where each generation consists of two quarks and two leptons. The first generation is the *up* and *down* quarks, the *electron* and the *electron neutrino*; the second includes the *charm* and *strange* quarks, the *muon* and the *muon neutrino*; the third generation consists of the *top* and *bottom* quarks and the *tau* and *tau neutrino*.<sup>[32]</sup> The most natural explanation for this would be that quarks and leptons of higher generations are excited states of the first generations. If this turns out to be the case, it would imply that quarks and leptons are composite particles, rather than elementary particles.<sup>[33]</sup>

Structure

In particle physics, fermions are particles which obey Fermi–Dirac statistics. Fermions can be elementary, like the electron, or composite, like the proton and the neutron. In the Standard Model there are two types of elementary fermions: quarks and leptons, which are discussed next.

Quarks

Quarks are particles of spin- $\frac{1}{2}$ , implying that they are fermions. They carry an electric charge of  $-\frac{1}{3} e$  (down-type quarks) or  $+\frac{2}{3} e$  (up-type quarks). For comparison, an electron has a charge of  $-1 e$ . They also carry colour charge, which is the equivalent of the electric charge for the strong interaction. Quarks also undergo radioactive decay, meaning that they are subject to the weak interaction. Quarks are massive particles, and therefore are also subject to gravity.

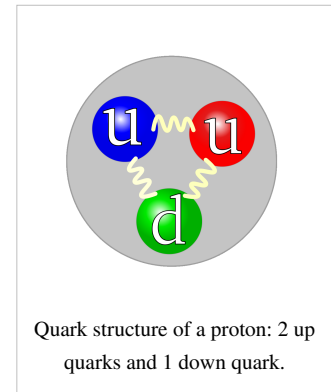
Quark properties<sup>[34]</sup>

name	symbol	spin	electric charge (e)	mass (MeV/c <sup>2</sup> )	mass comparable to	antiparticle	antiparticle symbol
up-type quarks							
up	u	$\frac{1}{2}$	$+\frac{2}{3}$	1.5 to 3.3	~ 5 electrons	antiup	u
charm	c	$\frac{1}{2}$	$+\frac{2}{3}$	1160 to 1340	~ 1 proton	anticharm	c
top	t	$\frac{1}{2}$	$+\frac{2}{3}$	169,100 to 173,300	~ 180 protons or ~ 1 tungsten atom	antitop	t
down-type quarks							
down	d	$\frac{1}{2}$	$-\frac{1}{3}$	3.5 to 6.0	~ 10 electrons	antidown	d
strange	s	$\frac{1}{2}$	$-\frac{1}{3}$	70 to 130	~ 200 electrons	antistrange	s
bottom	b	$\frac{1}{2}$	$-\frac{1}{3}$	4130 to 4370	~ 5 protons	antibottom	b

### Baryonic matter

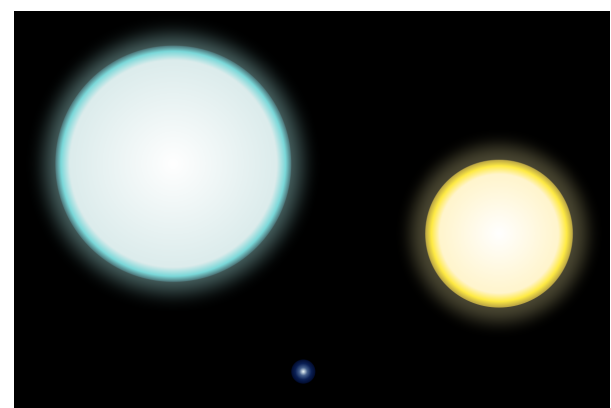
Baryons are strongly interacting fermions, and so are subject to Fermi-Dirac statistics. Amongst the baryons are the protons and neutrons, which occur in atomic nuclei, but many other unstable baryons exist as well. The term baryon is usually used to refer to triquarks — particles made of three quarks. "Exotic" baryons made of four quarks and one antiquark are known as the pentaquarks, but their existence is not generally accepted.

Baryonic matter is the part of the universe that is made of baryons (including all atoms). This part of the universe does not include dark energy, dark matter, black holes or various forms of degenerate matter, such as compose white dwarf stars and neutron stars. Microwave light seen by Wilkinson Microwave Anisotropy Probe (WMAP), suggests that only about 4.6% of that part of the universe within range of the best telescopes (that is, matter that may be visible because light could reach us from it), is made of baryonic matter. About 23% is dark matter, and about 72% is dark energy.<sup>[35]</sup>



### Degenerate matter

In physics, **degenerate matter** refers to the ground state of a gas of fermions at a temperature near absolute zero.<sup>[36]</sup> The Pauli exclusion principle requires that only two fermions can occupy a quantum state, one spin-up and the other spin-down. Hence, at zero temperature, the fermions fill up sufficient levels to accommodate all the available fermions, and for the case of many fermions the maximum kinetic energy called the Fermi energy and the pressure of the gas becomes very large and dependent upon the number of fermions rather than the temperature, unlike normal states of matter.



A comparison between the white dwarf IK Pegasi B (center), its A-class companion IK Pegasi A (left) and the Sun (right). This white dwarf has a surface temperature of 35,500 K.

Degenerate matter is thought to occur during the evolution of heavy stars.<sup>[37]</sup> The demonstration by Subrahmanyan Chandrasekhar that white dwarf stars have a maximum allowed mass because of the exclusion principle caused a revolution in the theory of star evolution.<sup>[38]</sup>

Degenerate matter includes the part of the universe that is made up of neutron stars and white dwarfs.

### Strange matter

**Strange matter** is a particular form of quark matter, usually thought of as a 'liquid' of up, down, and strange quarks. It is to be contrasted with nuclear matter, which is a liquid of neutrons and protons (which themselves are built out of up and down quarks), and with non-strange quark matter, which is a quark liquid containing only up and down quarks. At high enough density, strange matter is expected to be color superconducting. Strange matter is hypothesized to occur in the core of neutron stars, or, more speculatively, as isolated droplets that may vary in size from femtometers (strangelets) to kilometers (quark stars).

### Two meanings of the term "strange matter"

In particle physics and astrophysics, the term is used in two ways, one broader and the other more specific.

1. The broader meaning is just quark matter that contains three flavors of quarks: up, down, and strange. In this definition, there is a critical pressure and an associated critical density, and when nuclear matter (made of protons and neutrons) is compressed beyond this density, the protons and neutrons dissociate into quarks, yielding quark matter (probably strange matter).
2. The narrower meaning is quark matter that is *more stable than nuclear matter*. The idea that this could happen is the "strange matter hypothesis" of Bodmer<sup>[39]</sup> and Witten.<sup>[40]</sup> In this definition, the critical pressure is zero: the true ground state of matter is *always* quark matter. The nuclei that we see in the matter around us, which are droplets of nuclear matter, are actually metastable, and given enough time (or the right external stimulus) would decay into droplets of strange matter, i.e. strangelets.

### Leptons

Leptons are particles of spin- $\frac{1}{2}$ , meaning that they are fermions. They carry an electric charge of  $-1\text{ e}$  (charged leptons) or  $0\text{ e}$  (neutrinos). Unlike quarks, leptons do not carry colour charge, meaning that they do not experience the strong interaction. Leptons also undergo radioactive decay, meaning that they are subject to the weak interaction. Leptons are massive particles, therefore are subject to gravity.

#### Lepton properties

name	symbol	spin	electric charge (e)	mass (MeV/c <sup>2</sup> )	mass comparable to	antiparticle	antiparticle symbol
charged leptons <sup>[41]</sup>							
electron	e <sup>−</sup>	$\frac{1}{2}$	−1	0.5110	1 electron	antielectron	e <sup>+</sup>
muon	μ <sup>−</sup>	$\frac{1}{2}$	−1	105.7	~ 200 electrons	antimuon	μ <sup>+</sup>
tau	τ <sup>−</sup>	$\frac{1}{2}$	−1	1,777	~ 2 protons	antitau	τ <sup>+</sup>
neutrinos <sup>[42]</sup>							
electron neutrino	$\nu_e$	$\frac{1}{2}$	0	< 0.000460	< $\frac{1}{1000}$ electron	electron antineutrino	$\bar{\nu}_e$
muon neutrino	$\nu_\mu$	$\frac{1}{2}$	0	< 0.19	< $\frac{1}{2}$ electron	muon antineutrino	$\bar{\nu}_\mu$
tau neutrino	$\nu_\tau$	$\frac{1}{2}$	0	< 18.2	< 40 electrons	tau antineutrino	$\bar{\nu}_\tau$



## Phases

In bulk, matter can exist in several different forms, or states of aggregation, known as *phases*,<sup>[44]</sup> depending on ambient pressure, temperature and volume.<sup>[45]</sup> A phase is a form of matter that has a relatively uniform chemical composition and physical properties (such as density, specific heat, refractive index, and so forth). These phases include the three familiar ones (solids, liquids, and gases), as well as more exotic states of matter (such as plasmas, superfluids, supersolids, Bose–Einstein condensates, ...). A *fluid* may be a liquid, gas or plasma. There are also paramagnetic and ferromagnetic phases of magnetic materials. As conditions change, matter may change from one phase into another. These phenomena are called phase transitions, and are studied in the field of thermodynamics. In nanomaterials, the vastly increased ratio of surface area to volume results in matter that can exhibit properties entirely different from those of bulk material, and not well described by any bulk phase (see nanomaterials for more details).

Phases are sometimes called *states of matter*, but this term can lead to confusion with thermodynamic states.

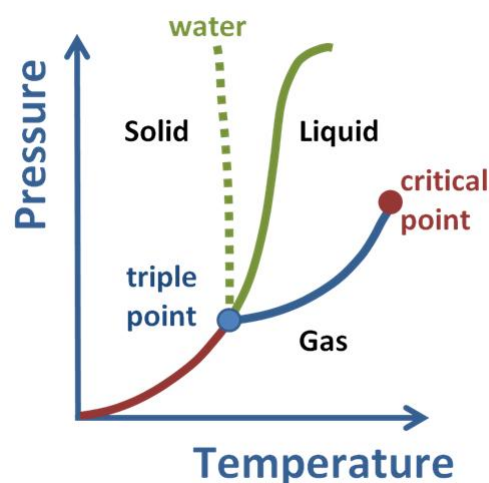
For example, two gases maintained at different pressures are in different *thermodynamic states* (different pressures), but in the same *phase* (both are gases).

## Antimatter

In particle physics and quantum chemistry, **antimatter** is matter that is composed of the antiparticles of those that constitute ordinary matter. If a particle and its antiparticle come into contact with each other, the two annihilate; that is, they may both be converted into other particles with equal energy in accordance with Einstein's equation  $E = mc^2$ . These new particles may be high-energy photons (gamma rays) or other particle–antiparticle pairs. The resulting particles are endowed with an amount of kinetic energy equal to the difference between the rest mass of the products of the annihilation and the rest mass of the original particle–antiparticle pair, which is often quite large.

Antimatter is not found naturally on Earth, except very briefly and in vanishingly small quantities (as the result of radioactive decay, lightning or cosmic rays). This is because antimatter which came to exist on Earth outside the confines of a suitable physics laboratory would almost instantly meet the ordinary matter that Earth is made of, and be annihilated. Antiparticles and some stable antimatter (such as antihydrogen) can be made in tiny amounts, but not in enough quantity to do more than test a few of its theoretical properties.

There is considerable speculation both in science and science fiction as to why the observable universe is apparently almost entirely matter, and whether other places are almost entirely antimatter instead. In the early universe, it is thought that matter and antimatter were equally represented, and the disappearance of antimatter requires an asymmetry in physical laws called the charge parity (or CP symmetry) violation. CP symmetry violation can be obtained from the Standard Model,<sup>[46]</sup> but at this time the apparent asymmetry of matter and antimatter in the visible universe is one of the great unsolved problems in physics. Possible processes by which it came about are explored in

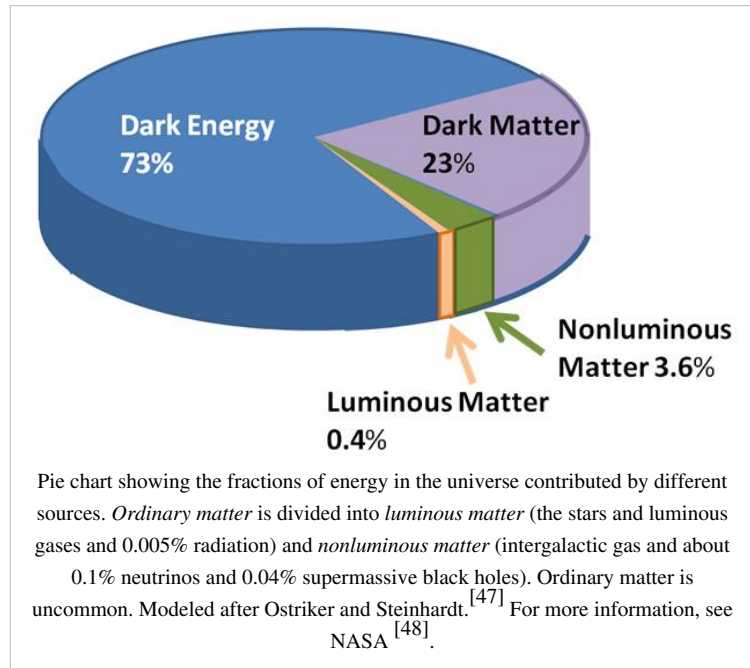


Phase diagram for a typical substance at a fixed volume. Vertical axis is Pressure, horizontal axis is Temperature. The green line marks the freezing point (above the green line is *solid*, below it is *liquid*) and the blue line the boiling point (above it is *liquid* and below it is *gas*). So, for example, at higher  $T$ , a higher  $P$  is necessary to maintain the substance in liquid phase. At the triple point the three phases; liquid, gas and solid; can coexist. Above the critical point there is no detectable difference between the phases. The dotted line shows the anomalous behavior of water: ice melts at constant temperature with increasing pressure.<sup>[43]</sup>

more detail under baryogenesis.

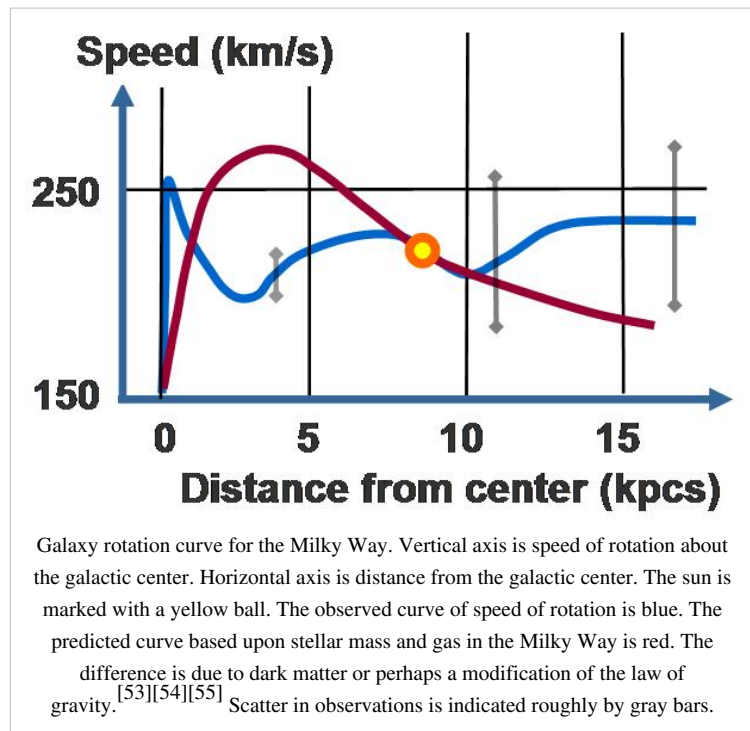
## Other types of matter

Ordinary matter, in the quarks and leptons definition, constitutes about 4% of the energy of the observable universe. The remaining energy is theorized to be due to exotic forms, of which 23% is dark matter<sup>[49][50]</sup> and 73% is dark energy.<sup>[51][52]</sup>



## Dark matter

In astrophysics and cosmology, **dark matter** is matter of unknown composition that does not emit or reflect enough electromagnetic radiation to be observed directly, but whose presence can be inferred from gravitational effects on visible matter.<sup>[14][56]</sup> Observational evidence of the early universe and the big bang theory require that this matter have energy and mass, but is not composed of either elementary fermions (as above) OR gauge bosons. The commonly accepted view is that most of the dark-matter is non-baryonic in nature.<sup>[14]</sup> As such, it is composed of particles as yet unobserved in the laboratory. Perhaps they are supersymmetric particles,<sup>[57]</sup> which are not Standard Model particles, but relics formed at very high energies in the early phase of the universe and still floating about.<sup>[14]</sup>



## Dark energy

In cosmology, **dark energy** is the name given to the antigravitating influence that is accelerating the rate of expansion of the universe. It is known not to be composed of known particles like protons, neutrons or electrons, nor of the particles of dark matter, because these all gravitate.<sup>[58][59]</sup>

Fully 70% of the matter density in the universe appears to be in the form of dark energy. Twenty-six percent is dark matter. Only 4% is ordinary matter. So less than 1 part in 20 is made out of matter we have observed experimentally or described in the standard model of particle physics. Of the other 96%, apart from the properties just mentioned, we know absolutely nothing.

— Lee Smolin: *The Trouble with Physics*, p. 16

## Exotic matter

Exotic matter is a hypothetical concept of particle physics. It covers any material which violates one or more classical conditions or is not made of known baryonic particles. Such materials would possess qualities like negative mass or being repelled rather than attracted by gravity.

## Historical development

### Origins

The pre-Socratics were among the first recorded speculators about the underlying nature of the visible world. Thales (c. 624 BC–c. 546 BC) regarded water as the fundamental material of the world. Anaximander (c. 610 BC–c. 546 BC) posited that the basic material was wholly characterless or limitless: the Infinite (*apeiron*). Anaximenes (flourished 585 BC, d. 528 BC) posited that the basic stuff was *pneuma* or air. Heraclitus (c. 535–c. 475 BC) seems to say the basic element is fire, though perhaps he means that all is change. Empedocles (c. 490–430 BC) spoke of four elements of which everything was made: earth, water, air, and fire.<sup>[60]</sup> Meanwhile, Parmenides argued that change does not exist, and Democritus argued that everything is composed of minuscule, inert bodies of all shapes called atoms, a philosophy called atomism. All of these notions had deep philosophical problems.<sup>[61]</sup>

Aristotle (384 BC – 322 BC) was the first to put the conception on a sound philosophical basis, which he did in his natural philosophy, especially in *Physics* book I.<sup>[62]</sup> He adopted as reasonable suppositions the four Empedoclean elements, but added a fifth, aether. Nevertheless these elements are not basic in Aristotle's mind. Rather they, like everything else in the visible world, are composed of the basic *principles* matter and form.

The word Aristotle uses for matter, ὕλη (*hyle* or *hule*), can be literally translated as wood or timber, that is, "raw material" for building.<sup>[63]</sup> Indeed, Aristotle's conception of matter is intrinsically linked to something being made or composed. In other words, in contrast to the early modern conception of matter as simply occupying space, matter for Aristotle is definitionally linked to process or change: matter is what underlies a change of substance.

For example, a horse eats grass: the horse changes the grass into itself; the grass as such does not persist in the horse, but some aspect of it—its matter—does. The matter is not specifically described (e.g., as atoms), but consists of whatever persists in the change of substance from grass to horse. Matter in this understanding does not exist independently (i.e., as a substance), but exists interdependently (i.e., as a "principle") with form and only insofar as it underlies change. It can be helpful to conceive of the relationship of matter and form as very similar to that between parts and whole. For Aristotle, matter as such can only *receive* actuality from form; it has no activity or actuality in itself, similar to the way that parts as such only have their existence *in* a whole (otherwise they would be independent wholes).

## Early modernity

René Descartes (1596–1650) originated the modern conception of matter. He was primarily a geometer. Instead of, like Aristotle, deducing the existence of matter from the physical reality of change, Descartes arbitrarily postulated matter to be an abstract, mathematical substance that occupies space:

So, extension in length, breadth, and depth, constitutes the nature of bodily substance; and thought constitutes the nature of thinking substance. And everything else which can be attributed to body presupposes extension, and is only a mode of that which is extended

— René Descartes, *Principles of Philosophy*<sup>[64]</sup>

For Descartes, matter has only the property of extension, so its only activity aside from locomotion is to exclude other bodies<sup>[65]</sup>; this is the mechanical philosophy. Descartes makes an absolute distinction between mind, which he defines as unextended, thinking substance, and matter, which he defines as unthinking, extended substance.<sup>[66]</sup> They are independent things. In contrast, Aristotle defines matter and the formal/forming principle as complementary *principles* which together compose one independent thing (substance). In short, Aristotle defines matter (roughly speaking) as what things are actually made of (with a *potential* independent existence), but Descartes elevates matter to an actual independent thing in itself.

The continuity and difference between Descartes' and Aristotle's conceptions is noteworthy. In both conceptions, matter is passive or inert. In the respective conceptions matter has different relationships to intelligence. For Aristotle, matter and intelligence (form) exist together in an interdependent relationship, whereas for Descartes, matter and intelligence (mind) are definitionally opposed, independent substances.<sup>[67]</sup>

Descartes' justification for restricting the inherent qualities of matter to extension is its permanence, but his real criterion is not permanence (which equally applied to color and resistance), but his desire to use geometry to explain all material properties.<sup>[68]</sup> Like Descartes, Hobbes, Boyle, and Locke argued that the inherent properties of bodies were limited to extension, and that so-called secondary qualities, like color, were only products of human perception.<sup>[69]</sup>

Isaac Newton (1643–1727) inherited Descartes' mechanical conception of matter. In the third of his "Rules of Reasoning in Philosophy," Newton lists the universal qualities of matter as "extension, hardness, impenetrability, mobility, and inertia."<sup>[70]</sup> Similarly in *Optics* he conjectures that God created matter as "solid, massy, hard, impenetrable, movable particles", which were "even so very hard as never to wear or break in pieces."<sup>[71]</sup> The "primary" properties of matter were amenable to mathematical description, unlike "secondary" qualities such as color or taste. Like Descartes, Newton rejected the essential nature of secondary qualities.<sup>[72]</sup>

Newton developed Descartes' notion of matter by restoring to matter intrinsic properties in addition to extension (at least on a limited basis), such as mass. Newton's use of gravitational force, which worked "at a distance," effectively repudiated Descartes' mechanics, in which interactions happened exclusively by contact.<sup>[73]</sup>

Though Newton's gravity would seem to be a *power* of bodies, Newton himself did not admit it to be an *essential* property of matter. Carrying the logic forward more consistently, Joseph Priestley argued that corporeal properties transcend contact mechanics: chemical properties require the *capacity* for attraction.<sup>[73]</sup> He argued matter has other inherent powers besides the so-called primary qualities of Descartes, et al.<sup>[74]</sup>

Since Priestley's time, there has been a massive expansion in knowledge of the constituents of the material world (viz., molecules, atoms, subatomic particles), but there has been no further development in the *definition* of matter. Rather the question has been set aside. Noam Chomsky summarizes the situation that has prevailed since that time:

What is the concept of body that finally emerged?[...] The answer is that there is no clear and definite conception of body.[...] Rather, the material world is whatever we discover it to be, with whatever properties it must be assumed to have for the purposes of explanatory theory. Any intelligible theory that offers genuine explanations and that can be assimilated to the core notions of physics becomes part of the theory of the material world, part of our account of body. If we have such a theory in some domain, we seek to assimilate it

to the core notions of physics, perhaps modifying these notions as we carry out this enterprise.

— Noam Chomsky, '**Language and problems of knowledge: the Managua lectures**, p. 144<sup>[73]</sup>

So matter is whatever physics studies and the object of study of physics is matter: there is no independent general definition of matter, apart from its fitting into the methodology of measurement and controlled experimentation. In sum, the boundaries between what constitutes matter and everything else remains as vague as the demarcation problem of delimiting science from everything else.<sup>[75]</sup>

## Late nineteenth and early twentieth centuries

In the 19th century, following the development of the periodic table, and of atomic theory, atoms were seen as being the fundamental constituents of matter; atoms formed molecules and compounds.<sup>[76]</sup>

The common definition in terms of occupying space and having mass is in contrast with most physical and chemical definitions of matter, which rely instead upon its structure and upon attributes not necessarily related to volume and mass. At the turn of the nineteenth century, the knowledge of matter began a rapid evolution.

Aspects of the Newtonian view still held sway. James Clerk Maxwell discussed matter in his work *Matter and Motion*.<sup>[77]</sup> He carefully separates "matter" from space and time, and defines it in terms of the object referred to in Newton's first law of motion.

However, the Newtonian picture was not the whole story. In the 19th century, the term "matter" was actively discussed by a host of scientists and philosophers, and a brief outline can be found in Levere.<sup>[78]</sup> A textbook discussion from 1870 suggests matter is what is made up of atoms:<sup>[79]</sup>

Three divisions of matter are recognized in science: masses, molecules and atoms.

A Mass of matter is any portion of matter appreciable by the senses.

A Molecule is the smallest particle of matter into which a body can be divided without losing its identity.

An Atom is a still smaller particle produced by division of a molecule.

Rather than simply having the attributes of mass and occupying space, matter was held to have chemical and electrical properties. The famous physicist J. J. Thomson wrote about the "constitution of matter" and was concerned with the possible connection between matter and electrical charge.<sup>[80]</sup>

## Later developments

There is an entire literature concerning the "structure of matter", ranging from the "electrical structure" in the early 20th century,<sup>[81]</sup> to the more recent "quark structure of matter", introduced today with the remark: *Understanding the quark structure of matter has been one of the most important advances in contemporary physics*.<sup>[82]</sup> In this connection, physicists speak of *matter fields*, and speak of particles as "quantum excitations of a mode of the matter field".<sup>[11][12]</sup> And here is a quote from de Sabbata and Gasperini: "With the word "matter" we denote, in this context, the sources of the interactions, that is spinor fields (like quarks and leptons), which are believed to be the fundamental components of matter, or scalar fields, like the Higgs particles, which are used to introduced mass in a gauge theory (and which, however, could be composed of more fundamental fermion fields)."<sup>[83]</sup>

The modern conception of matter has been refined many times in history, in light of the improvement in knowledge of just *what* the basic building blocks are, and in how they interact.

In the late 19th century with the discovery of the electron, and in the early 20th century, with the discovery of the atomic nucleus, and the birth of particle physics, matter was seen as made up of electrons, protons and neutrons interacting to form atoms. Today, we know that even protons and neutrons are not indivisible, they can be divided into quarks, while electrons are part of a particle family called leptons. Both quarks and leptons are elementary particles, and are currently seen as being the fundamental constituents of matter.<sup>[84]</sup>

These quarks and leptons interact through four fundamental forces: gravity, electromagnetism, weak interactions, and strong interactions. The Standard Model of particle physics is currently the best explanation for all of physics, but despite decades of efforts, gravity cannot yet be accounted for at the quantum-level; it is only described by classical physics (see quantum gravity and graviton).<sup>[85]</sup> Interactions between quarks and leptons are the result of an exchange of force-carrying particles (such as photons) between quarks and leptons.<sup>[86]</sup> The force-carrying particles are not themselves building blocks. As one consequence, mass and energy (which cannot be created or destroyed) cannot always be related to matter (which can be created out of non-matter particles such as photons, or even out of pure energy, such as kinetic energy). Force carriers are usually not considered matter: the carriers of the electric force (photons) possess energy (see Planck relation) and the carriers of the weak force (W and Z bosons) are massive, but neither are considered matter either.<sup>[87]</sup> However, while these particles are not considered matter, they do contribute to the total mass of atoms, subatomic particles, and all systems which contain them.<sup>[88][89]</sup>

## Summary

The term "matter" is used throughout physics in a bewildering variety of contexts: for example, one refers to "condensed matter physics",<sup>[90]</sup> "elementary matter",<sup>[91]</sup> "partonic" matter, "dark" matter, "anti"-matter, "strange" matter, and "nuclear" matter. In discussions of matter and antimatter, normal matter has been referred to by Alfvén as *koinomatter*.<sup>[92]</sup> It is fair to say that in physics, there is no broad consensus as to a general definition of matter, and the term "matter" usually is used in conjunction with a specifying modifier.

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## External links

- Visionlearning Module on Matter ([http://www.visionlearning.com/library/module\\_viewer.php?mid=49&l=&c3=](http://www.visionlearning.com/library/module_viewer.php?mid=49&l=&c3=))
- Matter in the universe (<http://www.newuniverse.co.uk/Matter.html>) How much Matter is in the Universe?
- NASA on superfluid core of neutron star ([http://imagine.gsfc.nasa.gov/docs/ask\\_astro/answers/970213.html](http://imagine.gsfc.nasa.gov/docs/ask_astro/answers/970213.html))

## Energy

In physics, **energy** (Ancient Greek: ἐνέργεια *energeia* "activity, operation"<sup>[1]</sup>) is an indirectly observed quantity that is often understood as the ability of a physical system to do work on other physical systems.<sup>[2][3]</sup> Since work is defined as a force acting through a distance (a length of space), energy is always equivalent to the ability to exert pulls or pushes against the basic forces of nature, along a path of a certain length.

The total energy contained in an object is identified with its mass, and energy cannot be created or destroyed. When matter (ordinary material particles) is changed into energy (such as energy of motion, or into

radiation), the **mass** of the system does not change through the transformation process. However, there may be mechanistic limits as to how much of the matter in an object may be changed into other types of energy and thus into work, on other systems. Energy, like mass, is a scalar physical quantity. In the International System of Units (SI), energy is measured in joules, but in many fields other units, such as kilowatt-hours and kilocalories, are customary. All of these units translate to units of work, which is always defined in terms of forces and the distances that the forces act through.

A system can transfer energy to another system by simply transferring matter to it (since matter is equivalent to energy, in accordance with its mass). However, when energy is transferred by means other than matter-transfer, the transfer produces changes in the second system, as a result of work done on it. This work manifests itself as the effect of force(s) applied through distances within the target system. For example, a system can emit energy to another by transferring (radiating) electromagnetic energy, but this creates forces upon the particles that absorb the radiation. Similarly, a system may transfer energy to another by physically impacting it, but in that case the energy of motion in an object, called kinetic energy, results in forces acting over distances (new energy) to appear in another object that is struck. Transfer of thermal energy by heat occurs by both of these mechanisms: heat can be transferred by electromagnetic radiation, or by physical contact in which direct particle-particle impacts transfer kinetic energy.

Energy may be stored in systems without being present as matter, or as kinetic or electromagnetic energy. Stored energy is created whenever a particle has been moved through a field it interacts with (requiring a force to do so), but the energy to accomplish this is stored as a new position of the particles in the field—a configuration that must be "held" or fixed by a different type of force (otherwise, the new configuration would resolve itself by the field pushing or pulling the particle back toward its previous position). This type of energy "stored" by force-fields and



In a typical lightning strike, 500 megajoules of electric potential energy are converted into 500 megajoules (total) of light energy, sound energy, thermal energy, and so on.

particles that have been forced into a new physical configuration in the field by doing work on them by another system, is referred to as potential energy. A simple example of potential energy is the work needed to lift an object in a gravity field, up to a support. Each of the basic forces of nature is associated with a different type of potential energy, and all types of potential energy (like all other types of energy) appears as system mass, whenever present. For example, a compressed spring will be slightly more massive than before it was compressed. Likewise, whenever energy is transferred between systems by any mechanism, an associated mass is transferred with it.

Any form of energy may be transformed into another form. For example, all types of potential energy are converted into kinetic energy when the objects are given freedom to move to different position (as for example, when an object falls off a support). When energy is in a form other than thermal energy, it may be transformed with good or even perfect efficiency, to any other type of energy, including electricity or production of new particles of matter. With thermal energy, however, there are often limits to the efficiency of the conversion to other forms of energy, as described by the second law of thermodynamics.

In all such energy transformation processes, the total energy remains the same, and a transfer of energy from one system to another, results in a loss to compensate for any gain. This principle, the conservation of energy, was first postulated in the early 19th century, and applies to any isolated system. According to Noether's theorem, the conservation of energy is a consequence of the fact that the laws of physics do not change over time.<sup>[4]</sup>

Although the total energy of a system does not change with time, its value may depend on the frame of reference. For example, a seated passenger in a moving airplane has zero kinetic energy relative to the airplane, but non-zero kinetic energy (and higher total energy) relative to the Earth.

## History

The word *energy* derives from the Greek ἐνέργεια *energeia*, which possibly appears for the first time in the work of Aristotle in the 4th century BCE.

The concept of energy emerged out of the idea of *vis viva* (living force), which Gottfried Leibniz defined as the product of the mass of an object and its velocity squared; he believed that total *vis viva* was conserved. To account for slowing due to friction, Leibniz theorized that thermal energy consisted of the random motion of the constituent parts of matter, a view shared by Isaac Newton, although it would be more than a century until this was generally accepted. In 1807, Thomas Young was possibly the first to use the term "energy" instead of *vis viva*, in its modern sense.<sup>[5]</sup> Gustave-Gaspard Coriolis described "kinetic energy" in 1829 in its modern sense, and in 1853, William Rankine coined the term "potential energy". It was argued for some years whether energy was a substance (the caloric) or merely a physical quantity, such as momentum.

William Thomson (Lord Kelvin) amalgamated all of these laws into the laws of thermodynamics, which aided in the rapid development of explanations of chemical processes by Rudolf Clausius, Josiah Willard Gibbs, and Walther Nernst. It also led to a mathematical formulation of the concept of entropy by Clausius and to the introduction of laws of radiant energy by Jožef Stefan.

During a 1961 lecture<sup>[6]</sup> for undergraduate students at the California Institute of Technology, Richard Feynman, a celebrated physics teacher and Nobel Laureate, said this about the concept of energy:

There is a fact, or if you wish, a *law*, governing all natural phenomena that are known to date. There is no known exception to this law—it is exact so far as we know. The law is called the *conservation of energy*. It



Thomas Young – the first to use the term "energy" in the modern sense.

states that there is a certain quantity, which we call energy, that does not change in manifold changes which nature undergoes. That is a most abstract idea, because it is a mathematical principle; it says that there is a numerical quantity which does not change when something happens. It is not a description of a mechanism, or anything concrete; it is just a strange fact that we can calculate some number and when we finish watching nature go through her tricks and calculate the number again, it is the same.

—*The Feynman Lectures on Physics*

Since 1918 it has been known that the law of conservation of energy is the direct mathematical consequence of the translational symmetry of the quantity conjugate to energy, namely time. That is, energy is conserved because the laws of physics do not distinguish between different instants of time (see Noether's theorem).

## Energy in various contexts

The concept of energy and its transformations is useful in explaining and predicting most natural phenomena. The *direction* of transformations in energy (what kind of energy is transformed to what other kind) is often described by entropy (equal energy spread among all available degrees of freedom) considerations, as in practice all energy transformations are permitted on a small scale, but certain larger transformations are not permitted because it is statistically unlikely that energy or matter will randomly move into more concentrated forms or smaller spaces.

The concept of energy is widespread in all sciences.

- In the context of chemistry, energy is an attribute of a substance as a consequence of its atomic, molecular or aggregate structure. Since a chemical transformation is accompanied by a change in one or more of these kinds of structure, it is invariably accompanied by an increase or decrease of energy of the substances involved. Some energy is transferred between the surroundings and the reactants of the reaction in the form of heat or light; thus the products of a reaction may have more or less energy than the reactants. A reaction is said to be exergonic if the final state is lower on the energy scale than the initial state; in the case of endergonic reactions the situation is the reverse. Chemical reactions are invariably not possible unless the reactants surmount an energy barrier known as the activation energy. The *speed* of a chemical reaction (at given temperature  $T$ ) is related to the activation energy  $E$ , by the Boltzmann's population factor  $e^{-E/kT}$  – that is the probability of molecule to have energy greater than or equal to  $E$  at the given temperature  $T$ . This exponential dependence of a reaction rate on temperature is known as the Arrhenius equation. The activation energy necessary for a chemical reaction can be in the form of thermal energy.
- In biology, energy is an attribute of all biological systems from the biosphere to the smallest living organism. Within an organism it is responsible for growth and development of a biological cell or an organelle of a biological organism. Energy is thus often said to be stored by cells in the structures of molecules of substances such as carbohydrates (including sugars), lipids, and proteins, which release energy when reacted with oxygen in respiration. In human terms, the human equivalent (H-e) (Human energy conversion) indicates, for a given amount of energy expenditure, the relative quantity of energy needed for human metabolism, assuming an average human energy expenditure of 12,500 kJ per day and a basal metabolic rate of 80 watts. For example, if our bodies run (on average) at 80 watts, then a light bulb running at 100 watts is running at 1.25 human equivalents ( $100 \div 80$ ) i.e. 1.25 H-e. For a difficult task of only a few seconds' duration, a person can put out thousands of watts, many times the 746 watts in one official horsepower. For tasks lasting a few minutes, a fit human can generate perhaps 1,000 watts. For an activity that must be sustained for an hour, output drops to around 300; for an activity kept up all day, 150 watts is about the maximum.<sup>[7]</sup> The human equivalent assists understanding of energy flows in physical and biological systems by expressing energy units in human terms: it provides a “feel” for the use of a given amount of energy<sup>[8]</sup>
- In geology, continental drift, mountain ranges, volcanoes, and earthquakes are phenomena that can be explained in terms of energy transformations in the Earth's interior.,<sup>[9]</sup> while meteorological phenomena like wind, rain, hail, snow, lightning, tornadoes and hurricanes, are all a result of energy transformations brought about by solar energy



on the atmosphere of the planet Earth.

- In cosmology and astronomy the phenomena of stars, nova, supernova, quasars and gamma ray bursts are the universe's highest-output energy transformations of matter. All stellar phenomena (including solar activity) are driven by various kinds of energy transformations. Energy in such transformations is either from gravitational collapse of matter (usually molecular hydrogen) into various classes of astronomical objects (stars, black holes, etc.), or from nuclear fusion (of lighter elements, primarily hydrogen).

Energy transformations in the universe over time are characterized by various kinds of potential energy that has been available since the Big Bang, later being "released" (transformed to more active types of energy such as kinetic or radiant energy), when a triggering mechanism is available.

Familiar examples of such processes include nuclear decay, in which energy is released that was originally "stored" in heavy isotopes (such as uranium and thorium), by nucleosynthesis, a process ultimately using the gravitational potential energy released from the gravitational collapse of supernovae, to store energy in the creation of these heavy elements before they were incorporated into the solar system and the Earth. This energy is triggered and released in nuclear fission bombs. In a slower process, **radioactive decay** of these atoms in the core of the Earth releases heat. This thermal energy drives plate tectonics and may lift mountains, via orogenesis. This slow lifting represents a kind of gravitational potential energy storage of the thermal energy, which may be later released to active kinetic energy in landslides, after a triggering event. Earthquakes also release stored elastic potential energy in rocks, a store that has been produced ultimately from the same radioactive heat sources. Thus, according to present understanding, familiar events such as landslides and earthquakes release energy that has been stored as potential energy in the Earth's gravitational field or elastic strain (mechanical potential energy) in rocks. Prior to this, they represent release of energy that has been stored in heavy atoms since the collapse of long-dead supernova stars created these atoms.

In another similar chain of transformations beginning at the dawn of the universe, **nuclear fusion** of hydrogen in the Sun also releases another store of potential energy which was created at the time of the Big Bang. At that time, according to theory, space expanded and the universe cooled too rapidly for hydrogen to completely fuse into heavier elements. This meant that hydrogen represents a store of potential energy that can be released by fusion. Such a fusion process is triggered by heat and pressure generated from gravitational collapse of hydrogen clouds when they produce stars, and some of the fusion energy is then transformed into sunlight. Such sunlight from our Sun may again be stored as gravitational potential energy after it strikes the Earth, as (for example) water evaporates from oceans and is deposited upon mountains (where, after being released at a hydroelectric dam, it can be used to drive turbines or generators to produce electricity). Sunlight also drives many weather phenomena, save those generated by volcanic events. An example of a solar-mediated weather event is a hurricane, which occurs when large unstable areas of warm ocean, heated over months, give up some of their thermal energy suddenly to power a few days of violent air movement. Sunlight is also captured by plants as *chemical potential energy* in photosynthesis, when carbon dioxide and water (two low-energy compounds) are converted into the high-energy compounds carbohydrates, lipids, and proteins. Plants also release oxygen during photosynthesis, which is utilized by living organisms as an electron acceptor, to release the energy of carbohydrates, lipids, and proteins. Release of the energy stored during photosynthesis as heat or light may be triggered suddenly by a spark, in a forest fire, or it may be made available more slowly for animal or human metabolism, when these molecules are ingested, and catabolism is triggered by enzyme action.

Through all of these transformation chains, potential energy stored at the time of the Big Bang is later released by intermediate events, sometimes being stored in a number of ways over time between releases, as more active energy. In all these events, one kind of energy is converted to other types of energy, including heat.

## Distinction between energy and power

Although in everyday usage the terms *energy* and *power* are essentially synonyms, scientists and engineers distinguish between them. In its technical sense, power is not at all the same as energy, but is the **rate** at which energy is converted (or, equivalently, at which work is performed). Thus a hydroelectric plant, by allowing the water above the dam to pass through turbines, converts the water's potential energy into kinetic energy and ultimately into electric energy, whereas the amount of electric energy that is generated **per unit of time** is the electric power generated. The same amount of energy converted through a shorter period of time is more power over that shorter time.

## Conservation of energy

Energy is subject to the **law of conservation of energy**. According to this law, energy can neither be created (produced) nor destroyed by itself. It can only be transformed.

Most kinds of energy (with gravitational energy being a notable exception)<sup>[10]</sup> are subject to strict local conservation laws as well. In this case, energy can only be exchanged between adjacent regions of space, and all observers agree as to the volumetric density of energy in any given space. There is also a global law of conservation of energy, stating that the total energy of the universe cannot change; this is a corollary of the local law, but not vice versa.<sup>[6][11]</sup> Conservation of energy is the mathematical consequence of translational symmetry of time (that is, the indistinguishability of time intervals taken at different time)<sup>[12]</sup> - see Noether's theorem.

According to Conservation of energy the total inflow of energy into a system must equal the total outflow of energy from the system, plus the change in the energy contained within the system.

This law is a fundamental principle of physics. It follows from the translational symmetry of time, a property of most phenomena below the cosmic scale that makes them independent of their locations on the time coordinate. Put differently, yesterday, today, and tomorrow are physically indistinguishable.

This is because energy is the quantity which is canonical conjugate to time. This mathematical entanglement of energy and time also results in the uncertainty principle - it is impossible to define the exact amount of energy during any definite time interval. The uncertainty principle should not be confused with energy conservation - rather it provides mathematical limits to which energy can in principle be defined and measured.

In quantum mechanics energy is expressed using the Hamiltonian operator. On any time scales, the uncertainty in the energy is by

$$\Delta E \Delta t \geq \frac{\hbar}{2}$$

which is similar in form to the Heisenberg uncertainty principle (but not really mathematically equivalent thereto, since  $H$  and  $t$  are not dynamically conjugate variables, neither in classical nor in quantum mechanics).

In particle physics, this inequality permits a qualitative understanding of virtual particles which carry momentum, exchange by which and with real particles, is responsible for the creation of all known fundamental forces (more accurately known as fundamental interactions). Virtual photons (which are simply lowest quantum mechanical energy state of photons) are also responsible for electrostatic interaction between electric charges (which results in Coulomb law), for spontaneous radiative decay of excited atomic and nuclear states, for the Casimir force, for van der Waals bond forces and some other observable phenomena.

## Applications of the concept of energy

Energy is subject to a strict global conservation law; that is, whenever one measures (or calculates) the total energy of a system of particles whose interactions do not depend explicitly on time, it is found that the total energy of the system always remains constant.<sup>[13]</sup>

- The total energy of a system can be subdivided and classified in various ways. For example, it is sometimes convenient to distinguish potential energy (which is a function of coordinates only) from kinetic energy (which is a function of coordinate time derivatives only). It may also be convenient to distinguish gravitational energy, electric energy, thermal energy, and other forms. These classifications overlap; for instance, thermal energy usually consists partly of kinetic and partly of potential energy.
- The *transfer* of energy can take various forms; familiar examples include work, heat flow, and advection, as discussed below.
- The word "energy" is also used outside of physics in many ways, which can lead to ambiguity and inconsistency. The vernacular terminology is not consistent with technical terminology. For example, while energy is always conserved (in the sense that the total energy does not change despite energy transformations), energy can be converted into a form, e.g., thermal energy, that cannot be utilized to perform work. When one talks about "conserving energy by driving less," one talks about conserving fossil fuels and preventing useful energy from being lost as heat. This usage of "conserve" differs from that of the law of conservation of energy.<sup>[11]</sup>

In classical physics energy is considered a scalar quantity, the canonical conjugate to time. In special relativity energy is also a scalar (although not a Lorentz scalar but a time component of the energy-momentum 4-vector).<sup>[14]</sup> In other words, energy is invariant with respect to rotations of space, but not invariant with respect to rotations of space-time (= boosts).

## Energy transfer

Because energy is strictly conserved and is also locally conserved (wherever it can be defined), it is important to remember that by the definition of energy the transfer of energy between the "system" and adjacent regions is work. A familiar example is *mechanical work*. In simple cases this is written as the following equation:

$$\Delta E = W \quad (1)$$

if there are no other energy-transfer processes involved. Here  $E$  is the amount of energy transferred, and  $W$  represents the work done on the system.

More generally, the energy transfer can be split into two categories:

$$\Delta E = W + Q \quad (2)$$

where  $Q$  represents the heat flow into the system.

There are other ways in which an open system can gain or lose energy. In chemical systems, energy can be added to a system by means of adding substances with different chemical potentials, which potentials are then extracted (both of these process are illustrated by fueling an auto, a system which gains in energy thereby, without addition of either work or heat). Winding a clock would be adding energy to a mechanical system. These terms may be added to the above equation, or they can generally be subsumed into a quantity called "energy addition term  $E$ " which refers to *any* type of energy carried over the surface of a control volume or system volume. Examples may be seen above, and many others can be imagined (for example, the kinetic energy of a stream of particles entering a system, or energy from a laser beam adds to system energy, without either being either work-done or heat-added, in the classic senses).

$$\Delta E = W + Q + E \quad (3)$$

Where  $E$  in this general equation represents other additional advected energy terms not covered by work done on a system, or heat added to it.

Energy is also transferred from potential energy ( $E_p$ ) to kinetic energy ( $E_k$ ) and then back to potential energy constantly. This is referred to as conservation of energy. In this closed system, energy cannot be created or destroyed; therefore, the initial energy and the final energy will be equal to each other. This can be demonstrated by the following:

$$E_{pi} + E_{ki} = E_{pF} + E_{kF} \quad (4)$$

The equation can then be simplified further since  $E_p = mgh$  (mass times acceleration due to gravity times the height) and  $E_k = \frac{1}{2}mv^2$  (half mass times velocity squared). Then the total amount of energy can be found by adding  $E_p + E_k = E_{total}$ .

## Energy and the laws of motion

In classical mechanics, energy is a conceptually and mathematically useful property, as it is a conserved quantity. Several formulations of mechanics have been developed using energy as a core concept.

### The Hamiltonian

The total energy of a system is sometimes called the Hamiltonian, after William Rowan Hamilton. The classical equations of motion can be written in terms of the Hamiltonian, even for highly complex or abstract systems. These classical equations have remarkably direct analogs in nonrelativistic quantum mechanics.<sup>[15]</sup>

### The Lagrangian

Another energy-related concept is called the Lagrangian, after Joseph Louis Lagrange. This is even more fundamental than the Hamiltonian, and can be used to derive the equations of motion. It was invented in the context of classical mechanics, but is generally useful in modern physics. The Lagrangian is defined as the kinetic energy *minus* the potential energy.

Usually, the Lagrange formalism is mathematically more convenient than the Hamiltonian for non-conservative systems (such as systems with friction).

### Noether's Theorem

Noether's (first) theorem (1918) states that any differentiable symmetry of the action of a physical system has a corresponding conservation law.

Noether's theorem has become a fundamental tool of modern theoretical physics and the calculus of variations. A generalization of the seminal formulations on constants of motion in Lagrangian and Hamiltonian mechanics (1788 and 1833, respectively), it does not apply to systems that cannot be modeled with a Lagrangian; for example, dissipative systems with continuous symmetries need not have a corresponding conservation law.

## Energy and thermodynamics

### Internal energy

Internal energy is the sum of all microscopic forms of energy of a system. It is the energy needed to create the system. It is related to the potential energy, e.g., molecular structure, crystal structure, and other geometric aspects, as well as the motion of the particles, in form of kinetic energy. Thermodynamics is chiefly concerned with changes in internal energy and not its absolute value, which is impossible to determine with thermodynamics alone.<sup>[16]</sup>

### The first law of thermodynamics

The first law of thermodynamics asserts that energy is conserved<sup>[17]</sup> and that heat flow is a form of energy transfer. For homogeneous systems, with a well-defined temperature and pressure, a commonly used corollary of the first law is that, for a system subject only to pressure forces and heat transfer (e.g., a cylinder-full of gas), the differential change in the internal energy of the system (with a *gain* in energy signified by a positive quantity) is given as

$$dE = TdS - PdV,$$

where the first term on the right is the heat transferred into the system, expressed in terms of temperature  $T$  and entropy  $S$  (in which entropy increases and the change  $dS$  is positive when the system is heated), and the last term on the right hand side is identified as work done on the system, where pressure is  $P$  and volume  $V$  (the negative sign results since compression of the system requires work to be done on it and so the volume change,  $dV$ , is negative when work is done on the system).

This equation is highly specific, ignoring all chemical, electrical, nuclear, and gravitational forces, effects such as advection of any form of energy other than heat and pV-work. The general formulation of the first law (i.e., conservation of energy) is valid even in situations in which the system is not homogeneous. For these cases the change in internal energy of a *closed* system is expressed in a general form by

$$dE = \delta Q + \delta W$$

where  $\delta Q$  is the heat supplied to the system and  $\delta W$  is the work applied to the system.

### Equipartition of energy

The energy of a mechanical harmonic oscillator (a mass on a spring) is alternatively kinetic and potential. At two points in the oscillation cycle it is entirely kinetic, and alternatively at two other points it is entirely potential. Over the whole cycle, or over many cycles, net energy is thus equally split between kinetic and potential. This is called equipartition principle; total energy of a system with many degrees of freedom is equally split among all available degrees of freedom.

This principle is vitally important to understanding the behavior of a quantity closely related to energy, called entropy. Entropy is a measure of evenness of a distribution of energy between parts of a system. When an isolated system is given more degrees of freedom (i.e., given new available energy states that are the same as existing states), then total energy spreads over **all** available degrees equally without distinction between "new" and "old" degrees. This mathematical result is called the second law of thermodynamics.

## Oscillators, phonons, and photons

In an ensemble (connected collection) of unsynchronized oscillators, the average energy is spread equally between kinetic and potential types.

In a solid, thermal energy (often referred to loosely as heat content) can be accurately described by an ensemble of thermal phonons that act as mechanical oscillators. In this model, thermal energy is equally kinetic and potential.

In an ideal gas, the interaction potential between particles is essentially the delta function which stores no energy: thus, all of the thermal energy is kinetic.

Because an electric oscillator (LC circuit) is analogous to a mechanical oscillator, its energy must be, on average, equally kinetic and potential. It is entirely arbitrary whether the magnetic energy is considered kinetic and whether the electric energy is considered potential, or vice versa. That is, either the inductor is analogous to the mass while the capacitor is analogous to the spring, or vice versa.

1. By extension of the previous line of thought, in free space the electromagnetic field can be considered an ensemble of oscillators, meaning that radiation energy can be considered equally potential and kinetic. This model is useful, for example, when the electromagnetic Lagrangian is of primary interest and is interpreted in terms of potential and kinetic energy.

2. On the other hand, in the key equation  $m^2c^4 = E^2 - p^2c^2$ , the contribution  $mc^2$  is called the rest energy, and all other contributions to the energy are called kinetic energy. For a particle that has mass, this implies that the kinetic energy is  $0.5p^2/m$  at speeds much smaller than  $c$ , as can be proved by writing  $E = mc^2 \sqrt{1 + p^2m^{-2}c^{-2}}$  and expanding the square root to lowest order. By this line of reasoning, the energy of a photon is entirely kinetic, because the photon is massless and has no rest energy. This expression is useful, for example, when the energy-versus-momentum relationship is of primary interest. The two analyses are entirely consistent. The electric and magnetic degrees of freedom in item 1 are *transverse* to the direction of motion, while the speed in item 2 is *along* the direction of motion. For non-relativistic particles these two notions of potential versus kinetic energy are numerically equal, so the ambiguity is harmless, but not so for relativistic particles.

## Work and virtual work

Work, a form of energy, is force times distance.

$$W = \int_C \mathbf{F} \cdot d\mathbf{s}$$

This says that the work ( $W$ ) is equal to the line integral of the force  $\mathbf{F}$  along a path  $C$ ; for details see the mechanical work article.

Work and thus energy is frame dependent. For example, consider a ball being hit by a bat. In the center-of-mass reference frame, the bat does no work on the ball. But, in the reference frame of the person swinging the bat, considerable work is done on the ball.

## Quantum mechanics

In quantum mechanics energy is defined in terms of the energy operator as a time derivative of the wave function. The Schrödinger equation equates the energy operator to the full energy of a particle or a system. In results can be considered as a definition of measurement of energy in quantum mechanics. The Schrödinger equation describes the space- and time-dependence of slow changing (non-relativistic) wave function of quantum systems. The solution of this equation for bound system is discrete (a set of permitted states, each characterized by an energy level) which results in the concept of quanta. In the solution of the Schrödinger equation for any oscillator (vibrator) and for electromagnetic waves in a vacuum, the resulting energy states are related to the frequency by the Planck equation  $E = h\nu$  (where  $h$  is the Planck's constant and  $\nu$  the frequency). In the case of electromagnetic wave these



energy states are called quanta of light or photons.

## Relativity

When calculating kinetic energy (work to accelerate a mass from zero speed to some finite speed) relativistically - using Lorentz transformations instead of Newtonian mechanics, Einstein discovered an unexpected by-product of these calculations to be an energy term which does not vanish at zero speed. He called it rest mass energy - energy which every mass must possess even when being at rest. The amount of energy is directly proportional to the mass of body:

$$E = mc^2,$$

where

$m$  is the mass,

$c$  is the speed of light in vacuum,

$E$  is the rest mass energy.

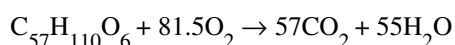
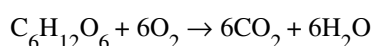
For example, consider electron-positron annihilation, in which the rest mass of individual particles is destroyed, but the inertia equivalent of the system of the two particles (its invariant mass) remains (since all energy is associated with mass), and this inertia and invariant mass is carried off by photons which individually are massless, but as a system retain their mass. This is a reversible process - the inverse process is called pair creation - in which the rest mass of particles is created from energy of two (or more) annihilating photons. In this system the matter (electrons and positrons) is destroyed and changed to non-matter energy (the photons). However, the total system mass and energy do not change during this interaction.

In general relativity, the stress-energy tensor serves as the source term for the gravitational field, in rough analogy to the way mass serves as the source term in the non-relativistic Newtonian approximation.<sup>[14]</sup>

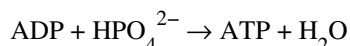
It is not uncommon to hear that energy is "equivalent" to mass. It would be more accurate to state that every energy has an inertia and gravity equivalent, and because mass is a form of energy, then mass too has inertia and gravity associated with it.

## Energy and life

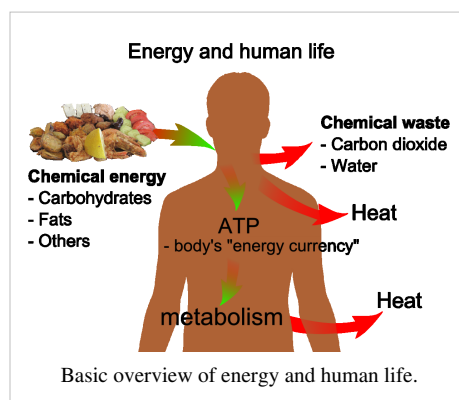
Any living organism relies on an external source of energy—radiation from the Sun in the case of green plants; chemical energy in some form in the case of animals—to be able to grow and reproduce. The daily 1500–2000 Calories (6–8 MJ) recommended for a human adult are taken as a combination of oxygen and food molecules, the latter mostly carbohydrates and fats, of which glucose ( $C_6H_{12}O_6$ ) and stearin ( $C_{57}H_{110}O_6$ ) are convenient examples. The food molecules are oxidised to carbon dioxide and water in the mitochondria



and some of the energy is used to convert ADP into ATP



The rest of the chemical energy in the carbohydrate or fat is converted into heat: the ATP is used as a sort of "energy currency", and some of the chemical energy it contains when split and reacted with water, is used for other metabolism (at each stage of a metabolic pathway, some chemical energy is converted into heat). Only a tiny fraction of the original chemical energy is used for work.<sup>[18]</sup>



gain in kinetic energy of a sprinter during a 100 m race: 4 kJ

gain in gravitational potential energy of a 150 kg weight lifted through 2 metres: 3kJ

Daily food intake of a normal adult: 6–8 MJ

It would appear that living organisms are remarkably inefficient (in the physical sense) in their use of the energy they receive (chemical energy or radiation), and it is true that most real machines manage higher efficiencies. In growing organisms the energy that is converted to heat serves a vital purpose, as it allows the organism tissue to be highly ordered with regard to the molecules it is built from. The second law of thermodynamics states that energy (and matter) tends to become more evenly spread out across the universe: to concentrate energy (or matter) in one specific place, it is necessary to spread out a greater amount of energy (as heat) across the remainder of the universe ("the surroundings").<sup>[19]</sup> Simpler organisms can achieve higher energy efficiencies than more complex ones, but the complex organisms can occupy ecological niches that are not available to their simpler brethren. The conversion of a portion of the chemical energy to heat at each step in a metabolic pathway is the physical reason behind the pyramid of biomass observed in ecology: to take just the first step in the food chain, of the estimated 124.7 Pg/a of carbon that is fixed by photosynthesis, 64.3 Pg/a (52%) are used for the metabolism of green plants,<sup>[20]</sup> i.e. reconverted into carbon dioxide and heat.

## Measurement

Because energy is defined as the ability to do work on objects, there is no absolute measure of energy. Only the transition of a system from one state into another can be defined and thus energy is measured in relative terms. The choice of a baseline or zero point is often arbitrary and can be made in whatever way is most convenient for a problem.

## Methods

The methods for the measurement of energy often deploy methods for the measurement of still more fundamental concepts of science, namely mass, distance, radiation, temperature, time, electric charge and electric current.

Conventionally the technique most often employed is calorimetry, a thermodynamic technique that relies on the measurement of temperature using a thermometer or of intensity of radiation using a bolometer.

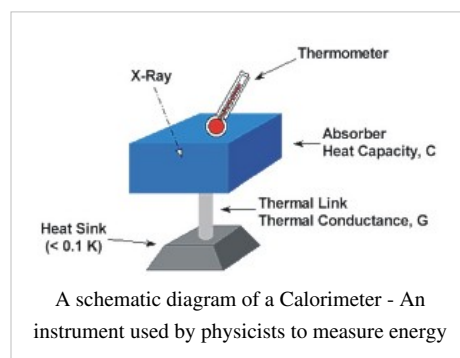
## Units

Throughout the history of science, energy has been expressed in several different units such as ergs and calories. At present, the accepted unit of measurement for energy is the SI unit of energy, the joule. In addition to the joule, other units of energy include the kilowatt hour (kWh) and the British thermal unit (Btu). These are both larger units of energy. One kWh is equivalent to exactly 3.6 million joules, and one Btu is equivalent to about 1055 joules.<sup>[21]</sup>

## Energy density

**Energy density** is a term used for the amount of useful energy stored in a given system or region of space per unit volume.

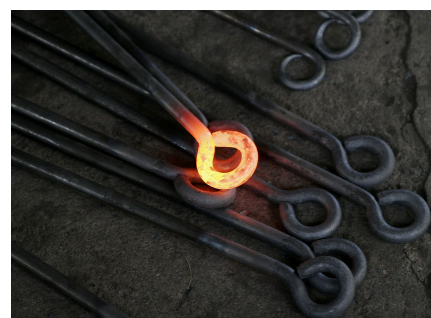
For fuels, the energy per unit volume is sometimes a useful parameter. In a few applications, comparing, for example, the effectiveness of hydrogen fuel to gasoline it turns out that hydrogen has a higher specific energy than does gasoline, but, even in liquid form, a much lower energy *density*.



## Forms of energy

In the context of physical sciences, several forms of energy have been defined. These include:

- Thermal energy, thermal energy in transit is called heat
- Chemical energy
- Electric energy
- Radiant energy, the energy of electromagnetic radiation
- Nuclear energy
- Magnetic energy
- Elastic energy
- Sound energy
- Mechanical energy
- Luminous energy
- Mass ( $E=mc^2$ )



Heat, a form of energy, is partly potential energy and partly kinetic energy.

These forms of energy may be divided into two main groups; kinetic energy and potential energy. Other familiar types of energy are a varying mix of both potential and kinetic energy.

Energy may be transformed between these forms, some with 100% energy conversion efficiency and others with less. Items that transform between these forms are called transducers.

The above list of the known possible forms of energy is not necessarily complete. Whenever physical scientists discover that a certain phenomenon appears to violate the law of energy conservation, new forms may be added, as is the case with dark energy, a hypothetical form of energy that permeates all of space and tends to increase the rate of expansion of the universe.

Classical mechanics distinguishes between potential energy, which is a function of the position of an object, and kinetic energy, which is a function of its movement. Both position and movement are relative to a frame of reference, which must be specified: this is often (and originally) an arbitrary fixed point on the surface of the Earth, the *terrestrial* frame of reference. It has been attempted to categorize *all* forms of energy as either kinetic or potential: this is not incorrect, but neither is it clear that it is a real simplification, as Feynman points out:

These notions of potential and kinetic energy depend on a notion of length scale. For example, one can speak of *macroscopic* potential and kinetic energy, which do not include thermal potential and kinetic energy. Also what is called chemical potential energy is a macroscopic notion, and closer examination shows that it is really the sum of the potential *and* kinetic energy on the atomic and subatomic scale. Similar remarks apply to nuclear "potential" energy and most other forms of energy. This dependence on length scale is non-problematic if the various length scales are decoupled, as is often the case ... but confusion can arise when different length scales are coupled, for instance when friction converts macroscopic work into microscopic thermal energy.

## Transformations of energy

One form of energy can often be readily transformed into another with the help of a device- for instance, a battery, from chemical energy to electric energy; a dam: gravitational potential energy to kinetic energy of moving water (and the blades of a turbine) and ultimately to electric energy through an electric generator. Similarly, in the case of a chemical explosion, chemical potential energy is transformed to kinetic energy and thermal energy in a very short time. Yet another example is that of a pendulum. At its highest points the kinetic energy is zero and the gravitational potential energy is at maximum. At its lowest point the kinetic energy is at maximum and is equal to the decrease of potential energy. If one (unrealistically) assumes that there is no friction, the conversion of energy between these processes is perfect, and the pendulum will continue swinging forever.

Energy gives rise to weight when it is trapped in a system with zero momentum, where it can be weighed. It is also equivalent to mass, and this mass is always associated with it. Mass is also equivalent to a certain amount of energy, and likewise always appears associated with it, as described in mass-energy equivalence. The formula  $E = mc^2$ , derived by Albert Einstein (1905) quantifies the relationship between rest-mass and rest-energy within the concept of special relativity. In different theoretical frameworks, similar formulas were derived by J. J. Thomson (1881), Henri Poincaré (1900), Friedrich Hasenöhl (1904) and others (see Mass-energy equivalence#History for further information).

Matter may be destroyed and converted to energy (and vice versa), but mass cannot ever be destroyed; rather, mass remains a constant for both the matter and the energy, during any process when they are converted into each other. However, since  $c^2$  is extremely large relative to ordinary human scales, the conversion of ordinary amount of matter (for example, 1 kg) to other forms of energy (such as heat, light, and other radiation) can liberate tremendous amounts of energy ( $\sim 9 \times 10^{16}$  joules = 21 megatons of TNT), as can be seen in nuclear reactors and nuclear weapons. Conversely, the mass equivalent of a unit of energy is minuscule, which is why a loss of energy (loss of mass) from most systems is difficult to measure by weight, unless the energy loss is very large. Examples of energy transformation into matter (i.e., kinetic energy into particles with rest mass) are found in high-energy nuclear physics.

Transformation of energy into useful work is a core topic of thermodynamics. In nature, transformations of energy can be fundamentally classed into two kinds: those that are thermodynamically reversible, and those that are thermodynamically irreversible. A reversible process in thermodynamics is one in which no energy is dissipated (spread) into empty energy states available in a volume, from which it cannot be recovered into more concentrated forms (fewer quantum states), without degradation of even more energy. A reversible process is one in which this sort of dissipation does not happen. For example, conversion of energy from one type of potential field to another, is reversible, as in the pendulum system described above. In processes where heat is generated, quantum states of lower energy, present as possible excitations in fields between atoms, act as a reservoir for part of the energy, from which it cannot be recovered, in order to be converted with 100% efficiency into other forms of energy. In this case, the energy must partly stay as heat, and cannot be completely recovered as usable energy, except at the price of an increase in some other kind of heat-like increase in disorder in quantum states, in the universe (such as an expansion of matter, or a randomization in a crystal).

As the universe evolves in time, more and more of its energy becomes trapped in irreversible states (i.e., as heat or other kinds of increases in disorder). This has been referred to as the inevitable thermodynamic heat death of the universe. In this heat death the energy of the universe does not change, but the fraction of energy which is available to do work through a heat engine, or be transformed to other usable forms of energy (through the use of generators attached to heat engines), grows less and less.

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- [18] These examples are solely for illustration, as it is not the energy available for work which limits the performance of the athlete but the power output of the sprinter and the force of the weightlifter. A worker stacking shelves in a supermarket does more work (in the physical sense) than either of the athletes, but does it more slowly.
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- Energy (<http://www.dmoz.org/Science/Technology/Energy/>) at the Open Directory Project

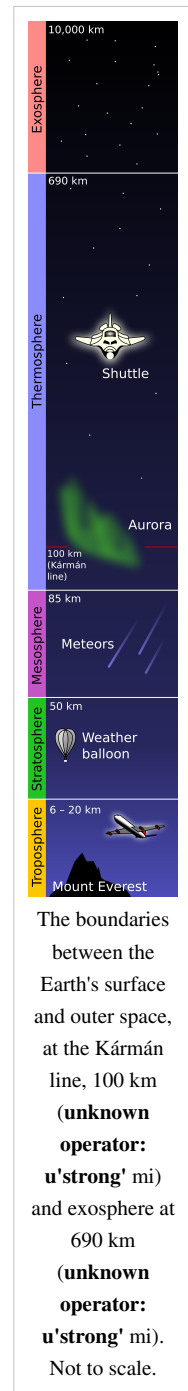
# Nature beyond Earth

## Outer space

**Outer space**, or simply *space*, is the void that exists between celestial bodies, including the Earth.<sup>[1]</sup> It is not completely empty, but consists of a hard vacuum containing a low density of particles: predominantly a plasma of hydrogen and helium, as well as electromagnetic radiation, magnetic fields, and neutrinos. Observations and theory suggest that it also contains dark matter and dark energy. The baseline temperature, as set by the background radiation left over from the Big Bang, is only 3 Kelvin (K); in contrast, temperatures in the coronae of stars can reach over a million Kelvin. Plasma with an extremely low density (less than one hydrogen atom per cubic meter) and high temperature (millions of Kelvin) in the space between galaxies accounts for most of the baryonic (ordinary) matter in outer space; local concentrations have condensed into stars and galaxies. Intergalactic space takes up most of the volume of the Universe, but even galaxies and star systems consist almost entirely of empty space.

There is no firm boundary where space begins. However the Kármán line, at an altitude of 100 km (**unknown operator: u'strong'** mi) above sea level, is conventionally used as the start of outer space for the purpose of space treaties and aerospace records keeping. The framework for international space law was established by the Outer Space Treaty, which was passed by the United Nations in 1967. This treaty precludes any claims of national sovereignty and permits all states to explore outer space freely. In 1979, the Moon Treaty made the surfaces of objects such as planets, as well as the orbital space around these bodies, the jurisdiction of the international community. Additional resolutions regarding the peaceful uses of outer space have been drafted by the United Nations, but these have not precluded the deployment of weapons into outer space, including the live testing of anti-satellite weapons.

Humans began the physical exploration of space during the twentieth century with the advent of high-altitude balloon flights, followed by the development of single and multi-stage rocket launchers. Earth orbit was achieved by Yuri Gagarin in 1961 and unmanned spacecraft have since reached all of the known planets in the Solar System. Achieving orbit requires a minimum velocity of 28400 km/h (**unknown operator: u'strong'** mph); much faster than any conventional aircraft. Outer space represents a challenging environment for human exploration because of the dual hazards of vacuum and radiation. Microgravity has a deleterious effect on human physiology, resulting in muscle atrophy and bone loss. As of yet space travel has been limited to low Earth orbit and the Moon for manned flight, and the vicinity of the Solar System for unmanned; the remainder of outer space remains inaccessible to humans other than by passive observation with telescopes.





## Discovery

In 350 BC, Greek philosopher Aristotle suggested that *nature abhors a vacuum*, a principle that became known as the *horror vacui*. This concept built upon a 5th century BCE ontological argument by the Greek philosopher Parmenides, who denied the possible existence of a void in space.<sup>[2]</sup> Based on this idea that a vacuum could not exist, in the West it was widely held for many centuries that space could not be empty.<sup>[3]</sup> As late as the 17th century, the French philosopher René Descartes argued that the entirety of space must be filled.<sup>[4]</sup>

In ancient China, there were various schools of thought concerning the nature of the heavens, some of which bear a resemblance to the modern understanding. In the 2nd century CE, astronomer Zhang Heng became convinced that space must be infinite, extending well beyond the mechanism that supported the Sun and the stars. The surviving books of the Hsüan Yeh school said that the heavens were boundless, "empty and void of substance". Likewise, the "sun, moon, and the company of stars float in the empty space, moving or standing still".<sup>[5]</sup>

The Italian scientist Galileo Galilei knew that air had mass and so was subject to gravity. In 1640, he demonstrated that an established force resisted the formation of a vacuum. However, it would remain for his pupil Evangelista Torricelli to create an apparatus that would produce a vacuum in 1643. This experiment resulted in the first mercury barometer and created a scientific sensation in Europe. The French mathematician Blaise Pascal reasoned that if the column of mercury was supported by air then the column ought to be shorter at higher altitude where the air pressure is lower.<sup>[6]</sup> In 1648, his brother-in-law, Florin Périer, repeated the experiment on the Puy-de-Dôme mountain in central France and found that the column was shorter by three inches. This decrease in pressure was further demonstrated by carrying a half-full balloon up a mountain and watching it gradually inflate, then deflate upon descent.<sup>[7]</sup>



The original Magdeburg hemispheres (lower left) used to demonstrate Otto von Guericke's vacuum pump (right)

In 1650, German scientist Otto von Guericke constructed the first vacuum pump: a device that would further refute the principle of *horror vacui*. He correctly noted that the atmosphere of the Earth surrounds the planet like a shell, with the density gradually declining with altitude. He concluded that there must be a vacuum between the Earth and the Moon.<sup>[8]</sup>

Back in the 15th century, German theologian Nicolaus Cusanus speculated that the Universe lacked a center and a circumference. He believed that the Universe, while not infinite, could not be held as finite as it lacked any bounds within which it could be contained.<sup>[9]</sup> These ideas led to speculations as to the infinite dimension of space by the Italian philosopher Giordano Bruno in the 16th century. He extended the Copernican heliocentric cosmology to the concept of an infinite Universe filled with a substance he called aether, which did not cause resistance to the motions of heavenly bodies.<sup>[10]</sup> English philosopher William Gilbert arrived at a similar conclusion, arguing that the stars are visible to us only because they are surrounded by a thin aether or a void.<sup>[11]</sup> This concept of an aether originated with

ancient Greek philosophers, including Aristotle, who conceived of it as the medium through which the heavenly bodies moved.<sup>[12]</sup>

The concept of a Universe filled with a luminiferous aether remained in vogue among some scientists until the early 20th century. This form of aether was viewed as the medium through which light could propagate.<sup>[13]</sup> In 1887, the Michelson–Morley experiment tried to detect the Earth's motion through this medium by looking for changes in the speed of light depending on the direction of the planet's motion. However, the null result indicated something was wrong with the concept. The idea of the luminiferous aether was then abandoned. It was replaced by Albert

Einstein's theory of special relativity, which holds that the speed of light in a vacuum is a fixed constant, independent of the observer's motion or frame of reference.<sup>[14][15]</sup>

The first professional astronomer to support the concept of an infinite Universe was the Englishman Thomas Digges in 1576.<sup>[16]</sup> But the scale of the Universe remained unknown until the first successful measurement of the distance to a nearby star in 1838 by the German astronomer Friedrich Bessel. He showed that the star 61 Cygni had a parallax of just 0.31 arcseconds (compared to the modern value of 0.287"). This corresponds to a distance of over 10 light years.<sup>[17]</sup> The distance to the Andromeda Galaxy was determined in 1923 by American astronomer Edwin Hubble by measuring the brightness of cepheid variables in that galaxy, a new technique discovered by Henrietta Leavitt.<sup>[18]</sup> This established that the Andromeda galaxy, and by extension all galaxies, lay well outside the Milky Way.<sup>[19]</sup>

The modern concept of outer space is based on the "Big Bang" cosmology, first proposed in 1931 by the Belgian physicist Georges Lemaître.<sup>[20]</sup> This theory holds that the observable Universe originated from a very compact form that has since undergone continuous expansion. Matter that remained following the initial expansion has since undergone gravitational collapse to create stars, galaxies and other astronomical objects, leaving behind a deep vacuum that forms what is now called outer space.<sup>[21]</sup> As light has a finite velocity, this theory also constrains the size of the directly observable Universe. This leaves open the question as to whether the Universe is finite or infinite.

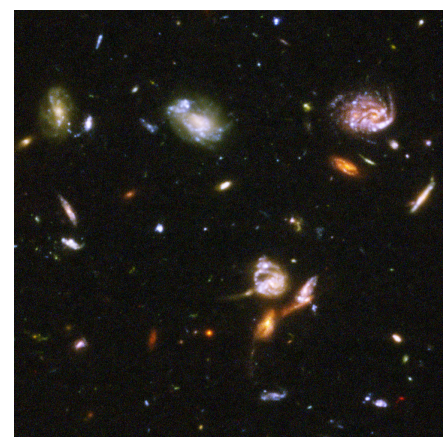
The term *outer space* was used as early as 1842 by the English poet Lady Emmeline Stuart-Wortley in her poem "The Maiden of Moscow".<sup>[22]</sup> The expression *outer space* was used as an astronomical term by Alexander von Humboldt in 1845.<sup>[23]</sup> It was later popularized in the writings of H. G. Wells in 1901.<sup>[24]</sup> The shorter term *space* is actually older, first used to mean the region beyond Earth's sky in John Milton's *Paradise Lost* in 1667.<sup>[25]</sup>

## Environment

Outer space is the closest natural approximation to a perfect vacuum. It has effectively no friction, allowing stars, planets and moons to move freely along their ideal orbits. However, even the deep vacuum of intergalactic space is not devoid of matter, as it contains a few hydrogen atoms per cubic meter.<sup>[26]</sup> By comparison, the air we breathe contains about  $10^{25}$  molecules per cubic meter.<sup>[27]</sup> The sparse density of matter in outer space means that electromagnetic radiation can travel great distances without being scattered: the mean free path of a photon in intergalactic space is about  $10^{23}$  km, or 10 billion light years.<sup>[28]</sup> In spite of this, extinction, which is the absorption and scattering of photons by dust and gas, is an important factor in galactic and intergalactic astronomy.<sup>[29]</sup>

Stars, planets and moons retain their atmospheres by gravitational attraction. Atmospheres have no clearly delineated boundary: the density of atmospheric gas gradually decreases with distance from the object until it becomes indistinguishable from the surrounding environment.<sup>[30]</sup> The Earth's atmospheric pressure drops to about  $3.2 \times 10^{-2}$  Pa at 100 kilometres (**unknown operator: u'strong'** miles) of altitude,<sup>[31]</sup> compared to 100 kPa for the International Union of Pure and Applied Chemistry (IUPAC) definition of standard pressure. Beyond this altitude, isotropic gas pressure rapidly becomes insignificant when compared to radiation pressure from the Sun and the dynamic pressure of the solar wind. The thermosphere in this range has large gradients of pressure, temperature and composition, and varies greatly due to space weather.<sup>[32]</sup>

On the Earth, temperature is defined in terms of the kinetic activity of the surrounding atmosphere. However the temperature of the vacuum cannot be measured in this way. Instead, the temperature is determined by measurement



Part of the Hubble Ultra-Deep Field image showing a typical section of space containing galaxies interspersed by deep vacuum. Given the finite speed of light, this view covers the last 13 billion years of the history of outer space.

of the radiation. All of the observable Universe is filled with photons that were created during the Big Bang, which is known as the cosmic microwave background radiation (CMB). (There is quite likely a correspondingly large number of neutrinos called the cosmic neutrino background.) The current black body temperature of the background radiation is about 3 K (–**unknown operator: u'strong'** °C; –**unknown operator: u'strong'** °F).<sup>[33]</sup> Some regions of outer space can contain highly energetic particles that have a much higher temperature than the CMB, such as the corona of the Sun where temperatures can range over 1.2–2.6 MK.<sup>[34]</sup>

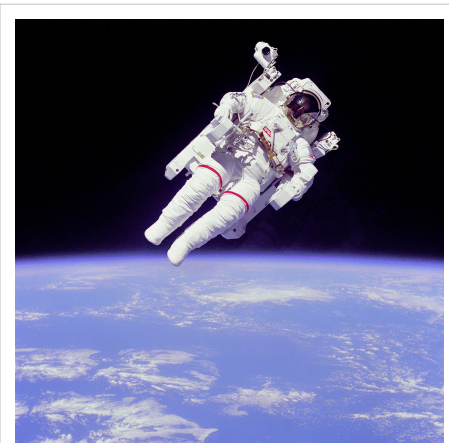
Outside of a protective atmosphere and magnetic field, there are few obstacles to the passage through space of energetic subatomic particles known as cosmic rays. These particles have energies ranging from about  $10^6$  eV up to an extreme  $10^{20}$  eV of ultra-high-energy cosmic rays.<sup>[35]</sup> The peak flux of cosmic rays occurs at energies of about  $10^9$  eV, with approximately 87% protons, 12% helium nuclei and 1% heavier nuclei. In the high energy range, the flux of electrons is only about 1% of that of protons.<sup>[36]</sup> Cosmic rays can damage electronic components and pose a health threat to space travelers.<sup>[37]</sup>

## Effect on human bodies

Contrary to popular belief,<sup>[38]</sup> a person suddenly exposed to a vacuum would not explode, freeze to death or die from boiling blood. However, sudden exposure to very low pressure, such as during a rapid decompression, could cause pulmonary barotrauma—a rupture of the lungs, due to the large pressure differential between inside and outside of the chest.<sup>[39]</sup> Even if the victim's airway is fully open, the flow of air through the windpipe may be too slow to prevent the rupture.<sup>[40]</sup> Rapid decompression can rupture eardrums and sinuses, bruising and blood seep can occur in soft tissues, and shock can cause an increase in oxygen consumption that leads to hypoxia.<sup>[39]</sup>

As a consequence of rapid decompression, any oxygen dissolved in the blood would empty into the lungs to try to equalize the partial pressure gradient. Once the deoxygenated blood arrived at the brain, humans and animals will lose consciousness after a few seconds and die of hypoxia within minutes.<sup>[41]</sup> Blood and other body fluids boil when the pressure drops below 6.3 kPa, and this condition is called ebullism.<sup>[42]</sup> The steam may bloat the body to twice its normal size and slow circulation, but tissues are elastic and porous enough to prevent rupture. Ebullism is slowed by the pressure containment of blood vessels, so some blood remains liquid.<sup>[43][44]</sup> Swelling and ebullism can be reduced by containment in a flight suit. Shuttle astronauts wear a fitted elastic garment called the Crew Altitude Protection Suit (CAPS) which prevents ebullism at pressures as low as 2 kPa.<sup>[45]</sup> Space suits are needed at 8 km (**unknown operator: u'strong'** mi) to provide enough oxygen for breathing and to prevent water loss, while above 20 km (**unknown operator: u'strong'** mi) they are essential to prevent ebullism.<sup>[46]</sup> Most space suits use around 30–39 kPa of pure oxygen, about the same as on the Earth's surface. This pressure is high enough to prevent ebullism, but evaporation of blood could still cause decompression sickness and gas embolisms if not managed.<sup>[47]</sup>

Because humans are optimized for life in Earth gravity, exposure to weightlessness has been shown to have deleterious effects on the health. Initially, more than 50% of astronauts experience space motion sickness. This can cause nausea and vomiting, vertigo, headaches, lethargy, and overall malaise. The duration of space sickness varies, but it typically lasts for 1–3 days, after which the body adjusts to the new environment. Longer term exposure to weightlessness results in muscle atrophy and deterioration of the skeleton, or spaceflight osteopenia. These effects can be minimized through a regimen of exercise.<sup>[48]</sup> Other effects include fluid redistribution, slowing of the cardiovascular system, decreased production of red blood cells, balance disorders, and a weakening of the immune system. Lesser symptoms include loss of body mass, nasal congestion, sleep disturbance, and puffiness of the



Because of the hazards of a vacuum, astronauts must wear a pressurized space suit while outside their spacecraft.

face.<sup>[49]</sup>

For long duration space travel, radiation can pose an acute health hazard. Exposure to radiation sources such as high-energy, ionizing cosmic rays can result in fatigue, nausea, vomiting, as well as damage to the immune system and changes to the white blood cell count. Over longer durations, symptoms include an increase in the risk of cancer, plus damage to the eyes, nervous system, lungs and the gastrointestinal tract.<sup>[50]</sup> On a round-trip Mars mission lasting three years, nearly the entire body would be traversed by high energy nuclei, each of which can cause ionization damage to cells. Fortunately, most such particles are significantly attenuated by the shielding provided by the aluminum walls of a spacecraft, and can be further diminished by water containers and other barriers. However, the impact of the cosmic rays upon the shielding produces additional radiation that can affect the crew. Further research will be needed to assess the radiation hazards and determine suitable countermeasures.<sup>[51]</sup>

## Boundary

There is no clear boundary between Earth's atmosphere and space, as the density of the atmosphere gradually decreases as the altitude increases. There are several standard boundary designations, namely:

- The Fédération Aéronautique Internationale has established the Kármán line at an altitude of 100 km (**unknown operator: u'strong' mi**) as a working definition for the boundary between aeronautics and astronautics. This is used because at an altitude of roughly 100 km (**unknown operator: u'strong' mi**), as Theodore von Kármán calculated, a vehicle would have to travel faster than orbital velocity in order to derive sufficient aerodynamic lift from the atmosphere to support itself.<sup>[52]</sup>
- The United States designates people who travel above an altitude of 50 miles (**unknown operator: u'strong' km**) as astronauts.<sup>[53]</sup>
- NASA's mission control uses 76 mi (**unknown operator: u'strong' km**) as their re-entry altitude (termed the Entry Interface), which roughly marks the boundary where atmospheric drag becomes noticeable (depending on the ballistic coefficient of the vehicle), thus leading shuttles to switch from steering with thrusters to maneuvering with air surfaces.<sup>[54]</sup>



SpaceShipOne completed the first manned private spaceflight in 2004, reaching an altitude of 100.124 km (**unknown operator: u'strong' mi**).

In 2009, scientists at the University of Calgary reported detailed measurements with an instrument called the Supra-Thermal Ion Imager (an instrument that measures the direction and speed of ions), which allowed them to establish a boundary at 118 km (**unknown operator: u'strong' mi**) above Earth. The boundary represents the midpoint of a gradual transition over tens of kilometers from the relatively gentle winds of the Earth's atmosphere to the more violent flows of charged particles in space, which can reach speeds well over 268 m/s (**unknown operator: u'strong' mph**).<sup>[55][56]</sup>



## Legal status

The Outer Space Treaty provides the basic framework for international space law. It covers the legal use of outer space by nation states, and includes in its definition of *outer space* the Moon and other celestial bodies. The treaty states that outer space is free for all nation states to explore and is not subject to claims of national sovereignty. It also prohibits the deployment of nuclear weapons in outer space. The treaty was passed by the United Nations General Assembly in 1963 and signed in 1967 by the USSR, the United States of America and the United Kingdom. As of January 1, 2008 the treaty has been ratified by 98 states and signed by an additional 27 states.<sup>[57]</sup>

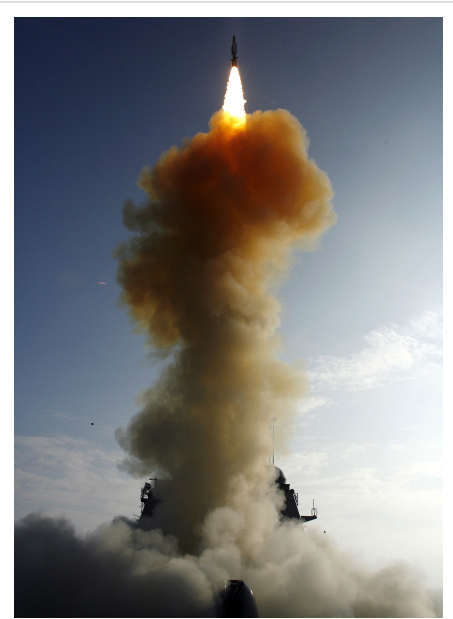
Beginning in 1958, outer space has been the subject of multiple resolutions by the United Nations General Assembly. Of these, more than 50 have been concerning the international co-operation in the peaceful uses of outer space and preventing an arms race in space.<sup>[58]</sup> Four additional space law treaties have been negotiated and drafted by the UN's Committee on the Peaceful Uses of Outer Space. Still, there remains no legal prohibition against deploying conventional weapons in space, and anti-satellite weapons have been successfully tested by the US, USSR and China.<sup>[59]</sup> The 1979 Moon Treaty turned the jurisdiction of all heavenly bodies (including the orbits around such bodies) over to the international community. However, this treaty has not been ratified by any nation that currently practices manned spaceflight.<sup>[60]</sup>

In 1976 eight equatorial states (Ecuador, Colombia, Brazil, Congo, Zaire, Uganda, Kenya, and Indonesia) met in Bogotá, Colombia. They made the "Declaration of the First Meeting of Equatorial Countries," also known as "the Bogotá Declaration", where they made a claim to control the segment of the geosynchronous orbital path corresponding to each country.<sup>[61]</sup> These claims are not internationally accepted.<sup>[62]</sup>

## Space versus orbit

A spacecraft enters orbit when it has enough horizontal velocity for its centripetal acceleration due to gravity to be less than or equal to the centrifugal acceleration due to the horizontal component of its velocity. For a low Earth orbit, this velocity is about 7900 m/s (**unknown operator: u'strong'** km/h; **unknown operator: u'strong'** mph);<sup>[63]</sup> by contrast, the fastest airplane speed ever achieved (excluding speeds achieved by deorbiting spacecraft) was 2200 m/s (**unknown operator: u'strong'** km/h; **unknown operator: u'strong'** mph) in 1967 by the North American X-15.<sup>[64]</sup>

To achieve an orbit, a spacecraft must travel faster than a sub-orbital spaceflight. The energy required to reach Earth orbital velocity at an altitude of 600 km (**unknown operator: u'strong'** mi) is about 36 MJ/kg, which is six times the energy needed merely to climb to the corresponding altitude.<sup>[65]</sup> Spacecraft with a perigee below about **unknown operator: u', 'unknown operator: u', 'unknown operator: u', (unknown operator: u'strong' unknown operator: u', 'mi)** are subject to drag from the Earth's atmospheric, which will cause the orbital altitude to decrease. The rate of orbital decay depends on the satellite's cross-sectional area and mass, as well as variations in the air density of the upper atmosphere. Below about 300 km (**unknown operator: u'strong'** mi), decay becomes more rapid with lifetimes measured in days. Once a satellite descends to 180 km (**unknown operator: u'strong'** mi), it will start to burn up in the atmosphere.<sup>[66]</sup> The escape velocity required to pull free of Earth's gravitational field altogether and move into interplanetary space is about 11200 m/s (**unknown operator: u'strong'** km/h; **unknown operator: u'strong'** mph).<sup>[67]</sup>



2008 launch of the SM-3 missile used to destroy American spy satellite USA-193

Earth's gravity reaches out far past the Van Allen radiation belt and keeps the Moon in orbit at an average distance of 384403 km (**unknown operator: u'strong'** mi). The region of space where the gravity of a planet tends to dominate the motion of objects in the presence of other perturbing bodies (such as another planet) is known as the Hill sphere. For Earth, this sphere has a radius of about 1500000 km (**unknown operator: u'strong'** mi).<sup>[68]</sup>

## Regions

Space is a partial vacuum: its different regions are defined by the various atmospheres and "winds" that dominate within them, and extend to the point at which those winds give way to those beyond. Geospace extends from Earth's atmosphere to the outer reaches of Earth's magnetic field, whereupon it gives way to the solar wind of interplanetary space. Interplanetary space extends to the heliopause, whereupon the solar wind gives way to the winds of the interstellar medium. Interstellar space then continues to the edges of the galaxy, where it fades into the intergalactic void.

## Geospace

Geospace is the region of outer space near the Earth. Geospace includes the upper region of the atmosphere, as well as the magnetosphere.<sup>[69]</sup> The outer boundary of geospace is the magnetopause, which forms an interface between the planet's magnetosphere and the solar wind. The inner boundary is the ionosphere.<sup>[70]</sup> As the physical properties and behavior of near Earth space is affected by the behavior of the Sun and space weather, the field of geospace is interlinked with heliophysics; the study of the Sun and its impact on the Solar System planets.<sup>[71]</sup>



Aurora australis observed from the Space Shuttle *Discovery*, on STS-39, May 1991  
(orbital altitude: 260 km)

The volume of geospace defined by the magnetopause is compacted in the direction of the Sun by the pressure of the solar wind, giving it a typical subsolar distance of 10 Earth radii from the center of the planet. However, the tail can extend outward to more than 100–200 Earth radii.<sup>[72]</sup> The Van Allen radiation belt lies within the geospace. The region between Earth's atmosphere and the Moon is sometimes referred to as *cis-lunar space*. The Moon passes through geospace roughly four days each month, during which time the surface is shielded from the solar wind.<sup>[73]</sup>

Geospace is populated by electrically charged particles at very low densities, the motions of which are controlled by the Earth's magnetic field. These plasmas form a medium from which storm-like disturbances powered by the solar wind can drive electrical currents into the Earth's upper atmosphere. During geomagnetic storms two regions of geospace, the radiation belts and the ionosphere, can become strongly disturbed. These storms increase fluxes of energetic electrons that can permanently damage satellite electronics, disrupting telecommunications and GPS technologies, and can also be a hazard to astronauts, even in low Earth orbit. They also create aurorae seen near the magnetic poles.<sup>[74]</sup>

Although it meets the definition of outer space, the atmospheric density within the first few hundred kilometers above the Kármán line is still sufficient to produce significant drag on satellites.<sup>[66]</sup> This region contains material left over from previous manned and unmanned launches that are a potential hazard to spacecraft. Some of this debris re-enters Earth's atmosphere periodically.<sup>[75]</sup>



## Interplanetary

Interplanetary space, the space around the Sun and planets of the Solar System, is the region dominated by the interplanetary medium, which extends out to the heliopause where the influence of the galactic environment starts to dominate over the magnetic field and particle flux from the Sun. Interplanetary space is defined by the solar wind, a continuous stream of charged particles emanating from the Sun that creates a very tenuous atmosphere (the heliosphere) for billions of miles into space. This wind has a particle density of 5–10 protons/cm<sup>3</sup> and is moving at a velocity of 350–400 km/s (**unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong'unknown operator: u'strong'**).<sup>[76]</sup> The distance and strength of the heliopause varies depending on the activity level of the solar wind.<sup>[77]</sup> The discovery since 1995 of extrasolar planets means that other stars must possess their own interplanetary media.<sup>[78]</sup>



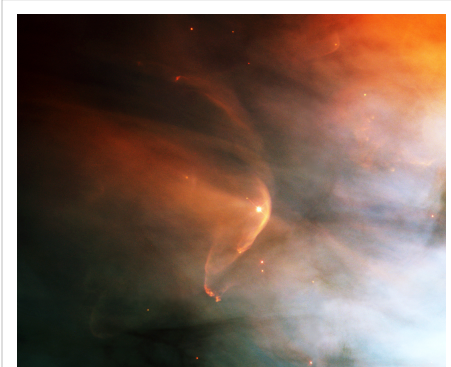
The sparse plasma (blue) and dust (white) in the tail of comet Hale–Bopp are being shaped by pressure from solar radiation and the solar wind, respectively

The volume of interplanetary space is a nearly total vacuum, with a mean free path of about one astronomical unit at the orbital distance of the Earth. However, this space is not completely empty, and is sparsely filled with cosmic rays, which include ionized atomic nuclei and various subatomic particles. There is also gas, plasma and dust, small meteors, and several dozen types of organic molecules discovered to date by microwave spectroscopy.<sup>[79]</sup>

Interplanetary space contains the magnetic field generated by the Sun.<sup>[76]</sup> There are also magnetospheres generated by planets such as Jupiter, Saturn, Mercury and the Earth that have their own magnetic fields. These are shaped by the influence of the solar wind into the approximation of a teardrop shape, with the long tail extending outward behind the planet. These magnetic fields can trap particles from the solar wind and other sources, creating belts of magnetic particles such as the Van Allen radiation belt. Planets without magnetic fields, such as Mars, have their atmospheres gradually eroded by the solar wind.<sup>[80]</sup>

## Interstellar

Interstellar space is the physical space within a galaxy not occupied by stars or their planetary systems. The interstellar medium resides—by definition—in interstellar space. The average density of matter in this region is about 10<sup>6</sup> particles per m<sup>3</sup>, but this varies from a low of about 10<sup>4</sup>–10<sup>5</sup> in regions of sparse matter up to about 10<sup>8</sup>–10<sup>10</sup> in dark nebula. Regions of star formation may reach 10<sup>12</sup>–10<sup>14</sup> particles per m<sup>3</sup>. Nearly 70% of this mass consists of lone hydrogen atoms. This is enriched with helium atoms as well as trace amounts of heavier atoms formed through stellar nucleosynthesis. These atoms can be ejected into the interstellar medium by stellar winds, or when evolved stars begin to shed their outer envelopes such as during the formation of a planetary nebula. The cataclysmic explosion of a supernova will generate an expanding shock wave consisting of ejected materials, as well as galactic cosmic rays. A number of molecules exist in interstellar space, as can tiny, 0.1 μm dust particles.<sup>[81]</sup>



Bow shock formed by the magnetosphere of LL Orionis (center) as it collides with the Orion Nebula flow

The local interstellar medium is a region of space within 100 parsecs (pc) of the Sun, which is of interest both for its proximity and for its interaction with the Solar System. This volume nearly coincides with a region of space known as the Local Bubble, which is characterized by a lack of dense, cold clouds. It forms a cavity in the Orion Arm of the Milky Way galaxy, with dense molecular clouds lying along the borders, such as those in the constellations of

Ophiuchus and Taurus. (The actual distance to the border of this cavity varies from 60 to 250 pc or more.) This volume contains about  $10^4$ – $10^5$  stars and the local interstellar gas counterbalances the astrospheres that surround these stars, with the volume of each sphere varying depending on the local density of the interstellar medium. The Local Bubble contains dozens of warm interstellar clouds with temperatures of up to 7,000 K and radii of 0.5–5 pc.<sup>[82]</sup>

When stars are moving at a sufficiently high peculiar velocity, their astrosphere can generate a bow shock as it collides with the interstellar medium. For decades it was assumed that the Sun had a bow shock. In 2012, data from IBEX and Voyagers showed that the Sun's bow shock does not exist. Instead, these authors argue that a subsonic bow wave defines the transition from the solar wind flow to the interstellar medium.<sup>[83][84]</sup> A bow shock is the third boundary of an astrosphere after the termination shock and the astropause (called the heliopause in the Solar System).<sup>[84]</sup>

## Intergalactic

Intergalactic space is the physical space between galaxies. The huge spaces between galaxy clusters are called the voids. Present estimates put the average energy density of the Universe at the equivalent of 5.9 protons per cubic meter, including dark energy, dark matter, and ordinary, baryonic matter, or atoms. The atoms account for only 4.6% of the total energy density, or a density of one proton per four cubic meters.<sup>[85]</sup> The density of the universe, however, is clearly not uniform; it ranges from relatively high density in galaxies—including very high density in structures within galaxies, such as planets, stars, and black holes—to conditions in vast voids that have much lower density, at least in terms of visible matter.<sup>[86]</sup>

Surrounding and stretching between galaxies, there is a rarefied plasma<sup>[87]</sup> that is organized in a cosmic filamentary structure.<sup>[88]</sup> This material is called the intergalactic medium (IGM). The density of the IGM is 5-200 times the average density of the Universe.<sup>[89]</sup> It consists mostly of ionized hydrogen; i.e. a plasma consisting of equal numbers of electrons and protons. As gas falls into the intergalactic medium from the voids, it heats up to temperatures of  $10^5$  K to  $10^7$  K,<sup>[90]</sup> which is high enough so that collisions between atoms have enough energy to cause the bound electrons to escape from the hydrogen nuclei; this is why the IGM is ionized. At these temperatures, it is called the warm-hot intergalactic medium (WHIM). (Although the gas is very hot by terrestrial standards,  $10^5$  K is often called "warm" in astrophysics.) Computer simulations and observations indicate that up to half of the atomic matter in the universe might exist in this warm-hot, rarefied state.<sup>[89][91][92]</sup> When gas falls from the filamentary structures of the WHIM into the galaxy clusters at the intersections of the cosmic filaments, it can heat up even more, reaching temperatures of  $10^8$  K and above in the so-called intracluster medium.<sup>[93]</sup>

## Exploration and applications

For the majority of human history, space was explored by remote observation; initially with the unaided eye and then with the telescope. Prior to the advent of reliable rocket technology, the closest that humans had come to reaching outer space was through the use of balloon flights. In 1935, the U.S. *Explorer II* manned balloon flight had reached an altitude of 22 km (**unknown operator: u'strong'** mi).<sup>[94]</sup> This was greatly exceeded in 1942 when the third launch of the German A-4 rocket climbed to an altitude of about 80 km (**unknown operator: u'strong'** mi). In 1957, the unmanned satellite *Sputnik 1* was launched by a Russian R-7 rocket, achieving Earth orbit at an altitude of 215–939 kilometres (**unknown operator:**



Soviet stamp depicting Yuri Gagarin, first human in space

**u'strong'unknown operator: u'strong'unknown operator: u'strong' unknown operator: u'strong')**<sup>[95]</sup> This was followed by the first human spaceflight in 1961, when Yuri Gagarin was sent into orbit on Vostok 1. The first humans to escape Earth orbit were Frank Borman, Jim Lovell and William Anders in 1968 on board Apollo 8, which achieved lunar orbit<sup>[96]</sup> and reached a maximum distance of 377349 km (**unknown operator: u'strong'** mi) from the Earth.<sup>[97]</sup>

In order to explore the other planets, a spacecraft must first reach escape velocity, which will allow it to travel beyond Earth orbit. The first spacecraft to accomplish this feat was the Soviet Luna 1, which performed a fly-by of the Moon in 1959.<sup>[98]</sup> In 1961, Venera 1 became the first planetary probe. It revealed the presence of the solar wind and performed the first fly-by of the planet Venus, although contact was lost before reaching Venus. The first successful planetary mission was the Mariner 2 fly-by of Venus in 1962.<sup>[99]</sup> The first spacecraft to perform a fly-by of Mars was Mariner 4, which reached the planet in 1964. Since that time, unmanned spacecraft have successfully examined each of the Solar System's planets, as well their moons and many minor planets and comets. They remain a fundamental tool for the exploration of outer space, as well as observation of the Earth.<sup>[100]</sup>

The absence of air makes outer space (and the surface of the Moon) ideal locations for astronomy at all wavelengths of the electromagnetic spectrum, as evidenced by the spectacular pictures sent back by the Hubble Space Telescope, allowing light from about 13.7 billion years ago—almost to the time of the Big Bang—to be observed. However, not every location in space is ideal for a telescope. The interplanetary zodiacal dust emits a diffuse near-infrared radiation that can mask the emission of faint sources such as extrasolar planets. Moving an infrared telescope out past the dust will increase the effectiveness of the instrument.<sup>[101]</sup> Likewise, a site like the Daedalus crater on the far side of the Moon could shield a radio telescope from the radio frequency interference that hampers Earth-based observations.<sup>[102]</sup>

The deep vacuum of space could make it an attractive environment for certain industrial processes, such as those that require ultraclean surfaces.<sup>[103]</sup>

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# Universe

The **universe** is commonly defined as the totality of everything that can exist or that we know exists.<sup>[1]</sup> This includes all matter and energy; planets, stars, and galaxies; and the contents of intergalactic space.<sup>[2][3]</sup> Definitions and usage vary and similar terms include the *cosmos*, the *world* and *nature*. Scientific observation of earlier stages in the development of the universe, which can be seen at great distances, suggests that the universe has been governed by the same physical laws and constants throughout most of its extent and history. There are various multiverse theories, in which physicists have suggested that our universe might be one among many universes that likewise exist.<sup>[4][5]</sup>

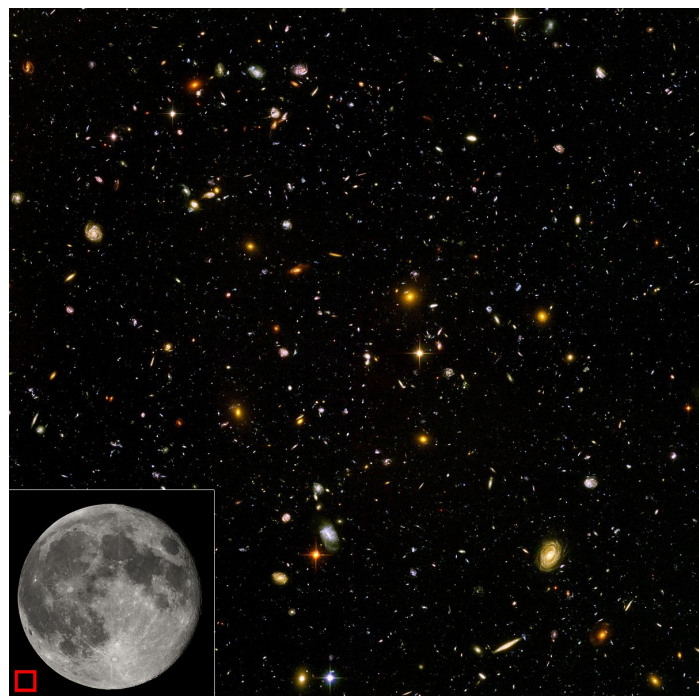
## History

### Observational history

Throughout recorded history, several cosmologies and cosmogonies have been proposed to account for observations of the universe. The earliest quantitative geocentric models were developed by the ancient Greek philosophers. Over the centuries, more precise observations and improved theories of gravity led to Copernicus's heliocentric model and the Newtonian model of the Solar System, respectively. Further improvements in astronomy led to the realization that the Solar System is embedded in a galaxy composed of billions of stars, the Milky Way, and that other galaxies exist outside it, as far as astronomical instruments can reach. Careful studies of the distribution of these galaxies and their spectral lines have led to much of modern cosmology. Discovery of the red shift and cosmic microwave background radiation revealed that the universe is expanding and apparently had a beginning.

### History of the Universe

According to the prevailing scientific model of the universe, known as the Big Bang, the universe expanded from an extremely hot, dense phase called the Planck epoch, in which all the matter and energy of the observable universe was concentrated. Since the Planck epoch, the universe has been expanding to its present form, possibly with a brief period (less than  $10^{-32}$  seconds) of cosmic inflation. Several independent experimental measurements support this theoretical expansion and, more generally, the Big Bang theory. Recent observations indicate that this expansion is accelerating because of dark energy, and that most of the matter in the universe may be in a form which cannot be detected by present instruments, called dark matter<sup>[6]</sup>. The common use of the "dark matter" and "dark energy" placeholder names for the unknown entities purported to account for about 95% of the mass-energy density of



This high-resolution image of the Hubble Ultra-Deep Field shows a diverse range of galaxies, each consisting of billions of stars. The equivalent area of sky that the picture occupies is shown as a red box in the lower left corner. The smallest, reddest galaxies, about 100, are some of the most distant galaxies to have been imaged by an optical telescope, existing at the time shortly after the Big Bang.

the universe demonstrates the present observational and conceptual shortcomings and uncertainties concerning the nature and ultimate fate of the universe.<sup>[7]</sup>

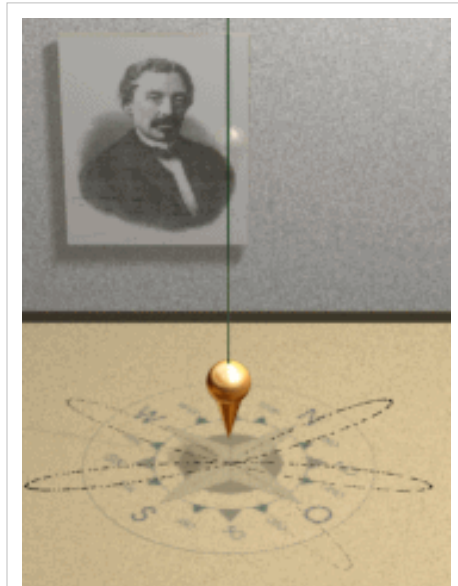
Current interpretations of astronomical observations indicate that the age of the universe is  $13.75 \pm 0.17$  billion years,<sup>[8]</sup> (whereas the decoupling of light and matter, see CMBR, happened already 380,000 years after the Big Bang), and that the diameter of the observable universe is at least 93 billion light years or  $8.80 \times 10^{26}$  metres.<sup>[9]</sup> According to general relativity, space can expand faster than the speed of light, although we can view only a small portion of the universe due to the limitation imposed by light speed. Since we cannot observe space beyond the limitations of light (or any electromagnetic radiation), it is uncertain whether the size of the universe is finite or infinite.

## Etymology, synonyms and definitions

The word *universe* derives from the Old French word *Univers*, which in turn derives from the Latin word *universum*.<sup>[10]</sup> The Latin word was used by Cicero and later Latin authors in many of the same senses as the modern English word is used.<sup>[11]</sup> The Latin word derives from the poetic contraction *Unvorsum* — first used by Lucretius in Book IV (line 262) of his *De rerum natura* (*On the Nature of Things*) — which connects *un*, *uni* (the combining form of *unus*, or "one") with *vorsum*, *versum* (a noun made from the perfect passive participle of *vertere*, meaning "something rotated, rolled, changed").<sup>[11]</sup>

An alternative interpretation of *unvorsum* is "everything rotated as one" or "everything rotated by one". In this sense, it may be considered a translation of an earlier Greek word for the universe, *περιφορά*, (*periforá*, "circumambulation"), originally used to describe a course of a meal, the food being carried around the circle of dinner guests.<sup>[12]</sup> This Greek word refers to celestial spheres, an early Greek model of the universe. Regarding Plato's Metaphor of the sun, Aristotle suggests that the rotation of the sphere of fixed stars inspired by the prime mover, motivates, in turn, terrestrial change via the Sun. Careful astronomical and physical measurements (such as the Foucault pendulum) are required to prove the Earth rotates on its axis.

A term for "universe" in ancient Greece was *τὸ πᾶν* (*tò pán*, The All, Pan (mythology)). Related terms were matter, (*τὸ ὄλον*, *tò ólon*, see also Hyle, lit. wood) and place (*τὸ κενόν*, *tò kenón*).<sup>[13][14]</sup> Other synonyms for the universe among the ancient Greek philosophers included *κόσμος* (cosmos) and *φύσις* (meaning Nature, from which we derive the word physics).<sup>[15]</sup> The same synonyms are found in Latin authors (*totum*, *mundus*, *natura*)<sup>[16]</sup> and survive in modern languages, e.g., the German words *Das All*, *Weltall*, and *Natur* for universe. The same synonyms are found in English, such as everything (as in the theory of everything), the cosmos (as in cosmology), the world (as in the many-worlds hypothesis), and Nature (as in natural laws or natural philosophy).<sup>[17]</sup>



Artistic rendition (highly exaggerated) of a Foucault pendulum showing that the Earth is not stationary, but rotates.

## Broadest definition: reality and probability

The broadest definition of the universe is found in *De divisione naturae* by the medieval philosopher and theologian Johannes Scotus Eriugena, who defined it as simply everything: everything that is created and everything that is not created.

## Definition as reality

More customarily, the universe is defined as everything that exists, (has existed, and will exist) . According to our current understanding, the universe consists of three principles: spacetime, forms of energy, including momentum and matter, and the physical laws that relate them.

## Definition as connected space-time

It is possible to conceive of disconnected space-times, each existing but unable to interact with one another. An easily visualized metaphor is a group of separate soap bubbles, in which observers living on one soap bubble cannot interact with those on other soap bubbles, even in principle. According to one common terminology, each "soap bubble" of space-time is denoted as a universe, whereas our particular space-time is denoted as *the universe*, just as we call our moon *the Moon*. The entire collection of these separate space-times is denoted as the multiverse.<sup>[18]</sup> In principle, the other unconnected universes may have different dimensionalities and topologies of space-time, different forms of matter and energy, and different physical laws and physical constants, although such possibilities are currently speculative.

## Definition as observable reality

According to a still-more-restrictive definition, the universe is everything within our connected space-time that could have a chance to interact with us and vice versa. According to the general theory of relativity, some regions of space may never interact with ours even in the lifetime of the universe, due to the finite speed of light and the ongoing expansion of space. For example, radio messages sent from Earth may never reach some regions of space, even if the universe would live forever; space may expand faster than light can traverse it. It is worth emphasizing that those distant regions of space are taken to exist and be part of reality as much as we are; yet we can never interact with them. The spatial region within which we can affect and be affected is denoted as the observable universe. Strictly speaking, the observable universe depends on the location of the observer. By traveling, an observer can come into contact with a greater region of space-time than an observer who remains still, so that the observable universe for the former is larger than for the latter. Nevertheless, even the most rapid traveler will not be able to interact with all of space. Typically, the observable universe is taken to mean the universe observable from our vantage point in the Milky Way Galaxy.

## Size, age, contents, structure, and laws

The universe is immensely large and possibly infinite in volume. The region visible from Earth (the observable universe) is a sphere with a radius of about 46 billion light years,<sup>[19]</sup> based on where the expansion of space has taken the most distant objects observed. For comparison, the diameter of a typical galaxy is only 30,000 light-years, and the typical distance between two neighboring galaxies is only 3 million light-years.<sup>[20]</sup> As an example, our Milky Way Galaxy is roughly 100,000 light years in diameter,<sup>[21]</sup> and our nearest sister galaxy, the Andromeda Galaxy, is located roughly 2.5 million light years away.<sup>[22]</sup> There are probably more than 100 billion ( $10^{11}$ ) galaxies in the observable universe.<sup>[23]</sup> Typical galaxies range from dwarfs with as few as ten million<sup>[24]</sup> ( $10^7$ ) stars up to giants with one trillion<sup>[25]</sup> ( $10^{12}$ ) stars, all orbiting the galaxy's center of mass. A 2010 study by astronomers estimated that the observable universe contains 300 sextillion ( $3 \times 10^{23}$ ) stars.<sup>[26]</sup>

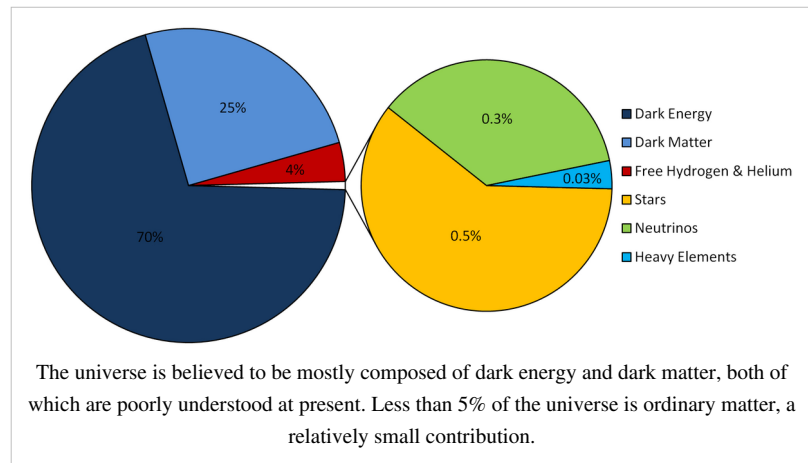
The observable matter is spread homogeneously (*uniformly*) throughout the universe, when averaged over distances longer than 300 million light-years.<sup>[27]</sup> However, on smaller length-scales, matter is observed to form "clumps", i.e., to cluster hierarchically; many atoms are condensed into stars, most stars into galaxies, most galaxies into clusters, superclusters and, finally, the largest-scale structures such as the Great Wall of galaxies. The observable

matter of the universe is also spread *isotropically*, meaning that no direction of observation seems different from any other; each region of the sky has roughly the same content.<sup>[28]</sup> The universe is also bathed in a highly isotropic microwave radiation that corresponds to a thermal equilibrium blackbody spectrum of roughly 2.725 kelvin.<sup>[29]</sup> The hypothesis that the large-scale universe is homogeneous and isotropic is known as the cosmological principle,<sup>[30]</sup> which is supported by astronomical observations.

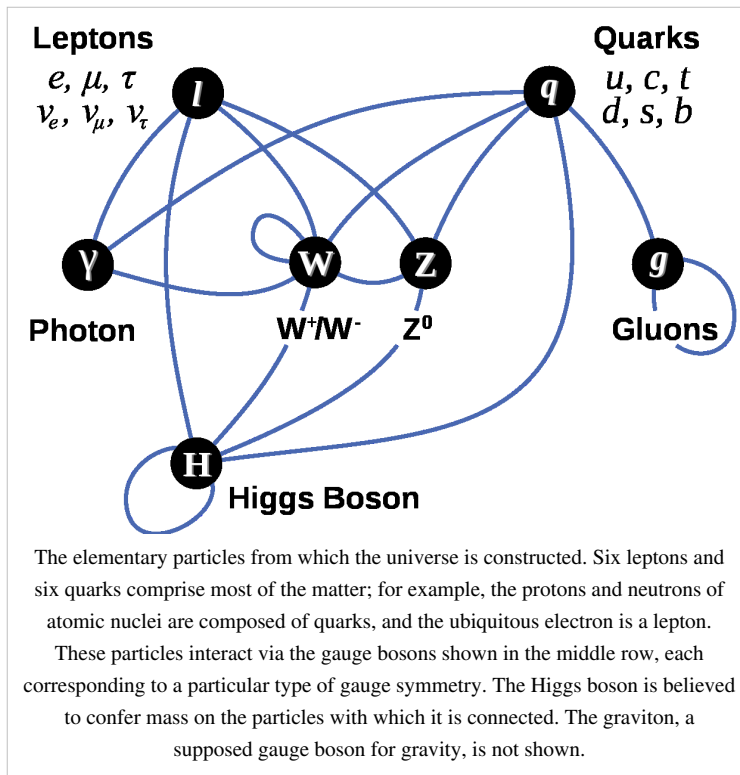
The present overall density of the universe is very low, roughly  $9.9 \times 10^{-30}$  grams per cubic centimetre. This mass-energy appears to consist of 73% dark energy, 23% cold dark matter and 4% ordinary matter. Thus the density of atoms is on the order of a single hydrogen atom for every four cubic meters of volume.<sup>[31]</sup> The properties of dark energy and dark matter are largely unknown. Dark matter gravitates as ordinary matter, and thus works to slow the expansion of the universe; by contrast, dark energy accelerates its expansion.

The most precise estimate of the universe's age is  $13.72 \pm 0.12$  billion years old, based on observations of the cosmic microwave background radiation.<sup>[32]</sup> Independent estimates (based on measurements such as radioactive dating) agree at 13–15 billion years.<sup>[33]</sup> The universe has not been the same at all times in its history; for example, the relative populations of quasars and galaxies have changed and space itself appears to have expanded. This expansion accounts for how Earth-bound scientists can observe the light from a galaxy 30 billion light years away, even if that light has traveled for only 13 billion years; the very space between them has expanded. This expansion is consistent with the observation that the light from distant galaxies has been redshifted; the photons emitted have been stretched to longer wavelengths and lower frequency during their journey. The rate of this spatial expansion is accelerating, based on studies of Type Ia supernovae and corroborated by other data.

The relative fractions of different chemical elements — particularly the lightest atoms such as hydrogen, deuterium and helium — seem to be identical throughout the universe and throughout its observable history.<sup>[34]</sup> The universe seems to have much more matter than antimatter, an asymmetry possibly related to the observations of CP violation.<sup>[35]</sup> The universe appears to have no net electric charge, and therefore gravity appears to be the dominant interaction on cosmological length scales. The universe also appears to have neither net momentum nor angular momentum. The absence of net charge and momentum would follow from accepted physical laws (Gauss's law and the non-divergence of the stress-energy-momentum pseudotensor, respectively), if the universe were finite.<sup>[36]</sup>







The universe appears to have a smooth space-time continuum consisting of three spatial dimensions and one temporal (time) dimension. On the average, space is observed to be very nearly flat (close to zero curvature), meaning that Euclidean geometry is experimentally true with high accuracy throughout most of the Universe.<sup>[37]</sup> Spacetime also appears to have a simply connected topology, at least on the length-scale of the observable universe. However, present observations cannot exclude the possibilities that the universe has more dimensions and that its spacetime may have a multiply connected global topology, in analogy with the cylindrical or toroidal topologies of two-dimensional spaces.<sup>[38]</sup>

The universe appears to behave in a manner that regularly follows a set of physical laws

and physical constants.<sup>[39]</sup> According to the prevailing Standard Model of physics, all matter is composed of three generations of leptons and quarks, both of which are fermions. These elementary particles interact via at most three fundamental interactions: the electroweak interaction which includes electromagnetism and the weak nuclear force; the strong nuclear force described by quantum chromodynamics; and gravity, which is best described at present by general relativity. The first two interactions can be described by renormalized quantum field theory, and are mediated by gauge bosons that correspond to a particular type of gauge symmetry. A renormalized quantum field theory of general relativity has not yet been achieved, although various forms of string theory seem promising. The theory of special relativity is believed to hold throughout the universe, provided that the spatial and temporal length scales are sufficiently short; otherwise, the more general theory of general relativity must be applied. There is no explanation for the particular values that physical constants appear to have throughout our universe, such as Planck's constant  $h$  or the gravitational constant  $G$ . Several conservation laws have been identified, such as the conservation of charge, momentum, angular momentum and energy; in many cases, these conservation laws can be related to symmetries or mathematical identities.

## Fine tuning

It appears that many of the properties of the universe have special values in the sense that a universe where these properties only differ slightly would not be able to support intelligent life.<sup>[40][41]</sup> Not all scientists agree that this fine-tuning exists.<sup>[42][43]</sup> In particular, it is not known under what conditions intelligent life could form and what form or shape that would take. A relevant observation in this discussion is that for an observer to exist to observe fine-tuning, the universe must be able to support intelligent life. As such the conditional probability of observing a universe that is fine-tuned to support intelligent life is 1. This observation is known as the anthropic principle and is particularly relevant if the creation of the universe was probabilistic or if multiple universes with a variety of properties exist (see below).

## Historical models

Many models of the cosmos (cosmologies) and its origin (cosmogonies) have been proposed, based on the then-available data and conceptions of the universe. Historically, cosmologies and cosmogonies were based on narratives of gods acting in various ways. Theories of an impersonal universe governed by physical laws were first proposed by the Greeks and Indians. Over the centuries, improvements in astronomical observations and theories of motion and gravitation led to ever more accurate descriptions of the universe. The modern era of cosmology began with Albert Einstein's 1915 general theory of relativity, which made it possible to quantitatively predict the origin, evolution, and conclusion of the universe as a whole. Most modern, accepted theories of cosmology are based on general relativity and, more specifically, the predicted Big Bang; however, still more careful measurements are required to determine which theory is correct.

## Creation

Many cultures have stories describing the origin of the world, which may be roughly grouped into common types. In one type of story, the world is born from a world egg; such stories include the Finnish epic poem *Kalevala*, the Chinese story of Pangu or the Indian Brahmanda Purana. In related stories, the creation idea is caused by a single entity emanating or producing something by him- or herself, as in the Tibetan Buddhism concept of Adi-Buddha, the ancient Greek story of Gaia (Mother Earth), the Aztec goddess Coatlicue myth, the ancient Egyptian god Atum story, or the Genesis creation narrative. In another type of story, the world is created from the union of male and female deities, as in the Maori story of Rangi and Papa. In other stories, the universe is created by crafting it from pre-existing materials, such as the corpse of a dead god — as from Tiamat in the Babylonian epic *Enuma Elish* or from the giant Ymir in Norse mythology — or from chaotic materials, as in Izanagi and Izanami in Japanese mythology. In other stories, the universe emanates from fundamental principles, such as Brahman and Prakrti, the creation myth of the Serers<sup>[44]</sup>, or the yin and yang of the Tao.

## Philosophical models

Further information: Cosmology

From the 6th century BCE, the pre-Socratic Greek philosophers developed the earliest known philosophical models of the universe. The earliest Greek philosophers noted that appearances can be deceiving, and sought to understand the underlying reality behind the appearances. In particular, they noted the ability of matter to change forms (e.g., ice to water to steam) and several philosophers proposed that all the apparently different materials of the world are different forms of a single primordial material, or arche. The first to do so was Thales, who proposed this material is Water. Thales' student, Anaximander, proposed that everything came from the limitless apeiron. Anaximenes proposed Air on account of its perceived attractive and repulsive qualities that cause the arche to condense or dissociate into different forms. Anaxagoras, proposed the principle of Nous (Mind). Heraclitus proposed fire (and spoke of logos). Empedocles proposed the elements: earth, water, air and fire. His four element theory became very popular. Like Pythagoras, Plato believed that all things were composed of number, with the Empedocles' elements taking the form of the Platonic solids. Democritus, and later philosophers—most notably Leucippus—proposed that the universe was composed of indivisible atoms moving through void (vacuum). Aristotle did not believe that was feasible because air, like water, offers resistance to motion. Air will immediately rush in to fill a void, and moreover, without resistance, it would do so indefinitely fast.

Although Heraclitus argued for eternal change, his quasi-contemporary Parmenides made the radical suggestion that all change is an illusion, that the true underlying reality is eternally unchanging and of a single nature. Parmenides denoted this reality as τὸ ἐν (The One). Parmenides' theory seemed implausible to many Greeks, but his student Zeno of Elea challenged them with several famous paradoxes. Aristotle responded to these paradoxes by developing the notion of a potential countable infinity, as well as the infinitely divisible continuum. Unlike the eternal and unchanging cycles of time, he believed the world was bounded by the celestial spheres, and thus magnitude was only

finitely multiplicative.

The Indian philosopher Kanada, founder of the Vaisheshika school, developed a theory of atomism and proposed that light and heat were varieties of the same substance.<sup>[45]</sup> In the 5th century AD, the Buddhist atomist philosopher Dignāga proposed atoms to be point-sized, durationless, and made of energy. They denied the existence of substantial matter and proposed that movement consisted of momentary flashes of a stream of energy.<sup>[46]</sup>

The theory of temporal finitism was inspired by the doctrine of Creation shared by the three Abrahamic religions: Judaism, Christianity and Islam. The Christian philosopher, John Philoponus, presented the philosophical arguments against the ancient Greek notion of an infinite past and future. Philoponus' arguments against an infinite past were used by the early Muslim philosopher, Al-Kindi (Alkindus); the Jewish philosopher, Saadia Gaon (Saadia ben Joseph); and the Muslim theologian, Al-Ghazali (Algazel). Borrowing from Aristotle's *Physics* and *Metaphysics*, they employed two logical arguments against an infinite past, the first being the "argument from the impossibility of the existence of an actual infinite", which states:<sup>[47]</sup>

"An actual infinite cannot exist."

"An infinite temporal regress of events is an actual infinite."

"∴ An infinite temporal regress of events cannot exist."

The second argument, the "argument from the impossibility of completing an actual infinite by successive addition", states:<sup>[47]</sup>

"An actual infinite cannot be completed by successive addition."

"The temporal series of past events has been completed by successive addition."

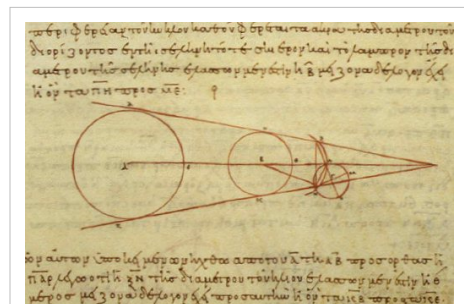
"∴ The temporal series of past events cannot be an actual infinite."

Both arguments were adopted by Christian philosophers and theologians, and the second argument in particular became more famous after it was adopted by Immanuel Kant in his thesis of the first antinomy concerning time.<sup>[47]</sup>

## Astronomical models

Astronomical models of the universe were proposed soon after astronomy began with the Babylonian astronomers, who viewed the universe as a flat disk floating in the ocean, and this forms the premise for early Greek maps like those of Anaximander and Hecataeus of Miletus.

Later Greek philosophers, observing the motions of the heavenly bodies, were concerned with developing models of the universe based more profoundly on empirical evidence. The first coherent model was proposed by Eudoxus of Cnidos. According to Aristotle's physical interpretation of the model, celestial spheres eternally rotate with uniform motion around a stationary Earth. Normal matter, is entirely contained within the terrestrial sphere. This model was also refined by Callippus and after concentric spheres were abandoned, it was brought into nearly perfect agreement with astronomical observations by Ptolemy. The success of such a model is largely due to the mathematical fact that any function (such as the position of a planet) can be decomposed into a set of circular functions (the Fourier modes). Other Greek scientists, such as the Pythagorean philosopher Philolaus postulated that at the center of the universe was a "central fire" around which the Earth, Sun, Moon and Planets revolved in uniform circular motion.<sup>[48]</sup> The Greek astronomer Aristarchus of Samos was the first known individual to propose a heliocentric model of the universe. Though the original text has been lost, a reference in Archimedes' book *The Sand Reckoner* describes Aristarchus' heliocentric theory. Archimedes wrote: (translated into English)



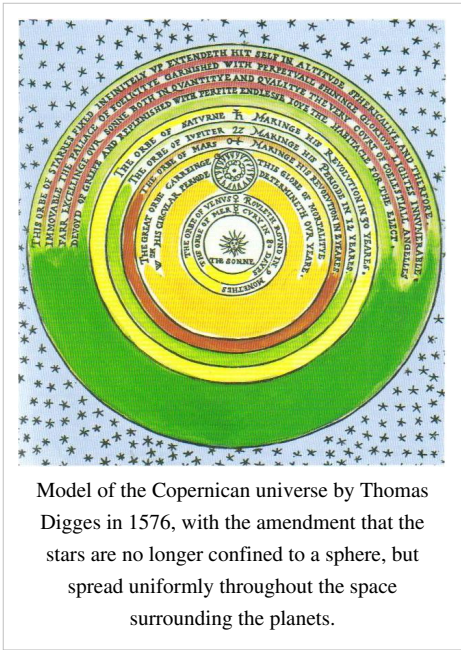
Aristarchus's 3rd century BCE calculations on the relative sizes of from left the Sun, Earth and Moon, from a 10th century AD Greek copy

You King Gelon are aware the 'universe' is the name given by most astronomers to the sphere the center of which is the center of the Earth, while its radius is equal to the straight line between the center of the Sun and the center of the Earth. This is the common account as you have heard from astronomers. But Aristarchus has brought out a book consisting of certain hypotheses, wherein it appears, as a consequence of the assumptions made, that the universe is many times greater than the 'universe' just mentioned. His hypotheses are that the fixed stars and the Sun remain unmoved, that the Earth revolves about the Sun on the circumference of a circle, the Sun lying in the middle of the orbit, and that the sphere of fixed stars, situated about the same center as the Sun, is so great that the circle in which he supposes the Earth to revolve bears such a proportion to the distance of the fixed stars as the center of the sphere bears to its surface.

Aristarchus thus believed the stars to be very far away, and saw this as the reason why there was no visible parallax, that is, an observed movement of the stars relative to each other as the Earth moved around the Sun. The stars are in fact much farther away than the distance that was generally assumed in ancient times, which is why stellar parallax is only detectable with telescopes. The geocentric model, consistent with planetary parallax, was assumed to be an explanation for the unobservability of the parallel phenomenon, stellar parallax. The rejection of the heliocentric view was apparently quite strong, as the following passage from Plutarch suggests (*On the Apparent Face in the Orb of the Moon*):

Cleanthes [a contemporary of Aristarchus and head of the Stoics] thought it was the duty of the Greeks to indict Aristarchus of Samos on the charge of impiety for putting in motion the Hearth of the universe [i.e. the earth], . . . supposing the heaven to remain at rest and the earth to revolve in an oblique circle, while it rotates, at the same time, about its own axis. [1]

The only other astronomer from antiquity known by name who supported Aristarchus' heliocentric model was Seleucus of Seleucia, a Hellenistic astronomer who lived a century after Aristarchus.<sup>[49][50][51]</sup> According to Plutarch, Seleucus was the first to prove the heliocentric system through reasoning, but it is not known what arguments he used. Seleucus' arguments for a heliocentric theory were probably related to the phenomenon of tides.<sup>[52]</sup> According to Strabo (1.1.9), Seleucus was the first to state that the tides are due to the attraction of the Moon, and that the height of the tides depends on the Moon's position relative to the Sun.<sup>[53]</sup> Alternatively, he may have proved the heliocentric theory by determining the constants of a geometric model for the heliocentric theory and by developing methods to compute planetary positions using this model, like what Nicolaus Copernicus later did in the 16th century.<sup>[54]</sup> During the Middle Ages, heliocentric models may have also been proposed by the Indian astronomer, Aryabhata,<sup>[55]</sup> and by the Persian astronomers, Albumasar<sup>[56]</sup> and Al-Sijzi.<sup>[57]</sup>



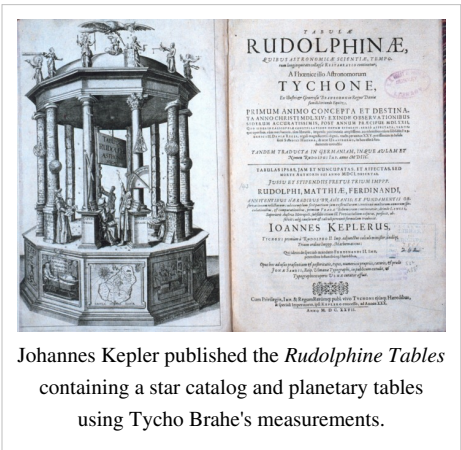
The Aristotelian model was accepted in the Western world for roughly two millennia, until Copernicus revived Aristarchus' theory that the astronomical data could be explained more plausibly if the earth rotated on its axis and if the sun were placed at the center of the universe.

In the center rests the sun. For who would place this lamp of a very beautiful temple in another or better place than this wherefrom it can illuminate everything at the same time?

—Nicolaus Copernicus, in Chapter 10, Book 1 of *De Revolutionibus Orbium Coelestrum* (1543)

As noted by Copernicus himself, the suggestion that the Earth rotates was very old, dating at least to Philolaus (c. 450 BC), Heraclides Ponticus (c. 350 BC) and Ecphantus the Pythagorean. Roughly a century before Copernicus, Christian scholar Nicholas of Cusa also proposed that the Earth rotates on its axis in his book, *On Learned Ignorance* (1440).<sup>[58]</sup> Aryabhata (476–550), Brahmagupta (598–668), Albumasar and Al-Sijzi, also proposed that the Earth rotates on its axis. The first empirical evidence for the Earth's rotation on its axis, using the phenomenon of comets, was given by Tusi (1201–1274) and Ali Qushji (1403–1474).

This cosmology was accepted by Isaac Newton, Christiaan Huygens and later scientists.<sup>[59]</sup> Edmund Halley (1720)<sup>[60]</sup> and Jean-Philippe de Cheseaux (1744)<sup>[61]</sup> noted independently that the assumption of an infinite space filled uniformly with stars would lead to the prediction that the nighttime sky would be as bright as the sun itself; this became known as Olbers' paradox in the 19th century.<sup>[62]</sup> Newton believed that an infinite space uniformly filled with matter would cause infinite forces and instabilities causing the matter to be crushed inwards under its own gravity.<sup>[59]</sup> This instability was clarified in 1902 by the Jeans instability criterion.<sup>[63]</sup> One solution to these paradoxes is the Charlier universe, in which the matter is arranged hierarchically (systems of orbiting bodies that are themselves orbiting in a larger system, *ad infinitum*) in a fractal way such that the universe has a negligibly small overall density; such a cosmological model had also been proposed earlier in 1761 by Johann Heinrich Lambert.<sup>[64]</sup> A significant astronomical advance of the 18th century was the realization by Thomas Wright, Immanuel Kant and others of nebulae.<sup>[65]</sup>



The modern era of physical cosmology began in 1917, when Albert Einstein first applied his general theory of relativity to model the structure and dynamics of the universe.<sup>[66]</sup>

## Theoretical models

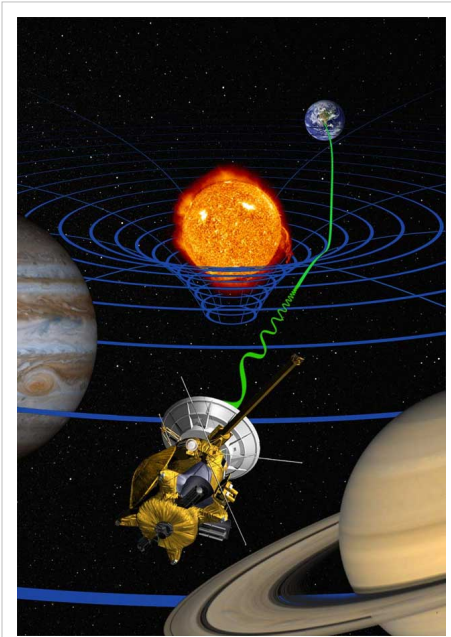
Of the four fundamental interactions, gravitation is dominant at cosmological length scales; that is, the other three forces play a negligible role in determining structures at the level of planetary systems, galaxies and larger-scale structures. Since all matter and energy gravitate, gravity's effects are cumulative; by contrast, the effects of positive and negative charges tend to cancel one another, making electromagnetism relatively insignificant on cosmological length scales. The remaining two interactions, the weak and strong nuclear forces, decline very rapidly with distance; their effects are confined mainly to sub-atomic length scales.

## General theory of relativity

Given gravitation's predominance in shaping cosmological structures, accurate predictions of the universe's past and future require an accurate theory of gravitation. The best theory available is Albert Einstein's general theory of relativity, which has passed all experimental tests hitherto. However, since rigorous experiments have not been carried out on cosmological length scales, general relativity could conceivably be inaccurate. Nevertheless, its cosmological predictions appear to be consistent with observations, so there is no compelling reason to adopt another theory.

General relativity provides a set of ten nonlinear partial differential equations for the spacetime metric (Einstein's field equations) that must be solved from the distribution of mass-energy and momentum throughout the universe. Since these are unknown in exact detail, cosmological models have been based on the cosmological principle, which states that the universe is homogeneous and isotropic. In effect, this principle asserts that the gravitational effects of the various galaxies making up the universe are equivalent to those of a fine dust distributed uniformly throughout the universe with the same average density. The assumption of a uniform dust makes it easy to solve Einstein's field equations and predict the past and future of the universe on cosmological time scales.

Einstein's field equations include a cosmological constant ( $\Lambda$ ),<sup>[66][67]</sup> that corresponds to an energy density of empty space.<sup>[68]</sup> Depending on its sign, the cosmological constant can either slow (negative  $\Lambda$ ) or accelerate (positive  $\Lambda$ ) the expansion of the universe. Although many scientists, including Einstein, had speculated that  $\Lambda$  was zero,<sup>[69]</sup> recent astronomical observations of type Ia supernovae have detected a large amount of "dark energy" that is accelerating the universe's expansion.<sup>[70]</sup> Preliminary studies suggest that this dark energy corresponds to a positive  $\Lambda$ , although alternative theories cannot be ruled out as yet.<sup>[71]</sup> Russian physicist Zel'dovich suggested that  $\Lambda$  is a measure of the zero-point energy associated with virtual particles of quantum field theory, a pervasive vacuum energy that exists everywhere, even in empty space.<sup>[72]</sup> Evidence for such zero-point energy is observed in the Casimir effect.



High-precision test of general relativity by the Cassini space probe (artist's impression): radio signals sent between the Earth and the probe (green wave) are delayed by the warping of space and time (blue lines) due to the Sun's mass.



## Special relativity and space-time

The universe has at least three spatial and one temporal (time) dimension. It was long thought that the spatial and temporal dimensions were different in nature and independent of one another. However, according to the special theory of relativity, spatial and temporal separations are interconvertible (within limits) by changing one's motion.

To understand this interconversion, it is helpful to consider the analogous interconversion of spatial separations along the three spatial dimensions. Consider the two endpoints of a rod of length  $L$ . The length can be determined from the differences in the three coordinates  $\Delta x$ ,  $\Delta y$  and  $\Delta z$  of the two endpoints in a given reference frame

$$L^2 = \Delta x^2 + \Delta y^2 + \Delta z^2$$

using the Pythagorean theorem. In a rotated reference frame, the coordinate differences differ, but they give the same length

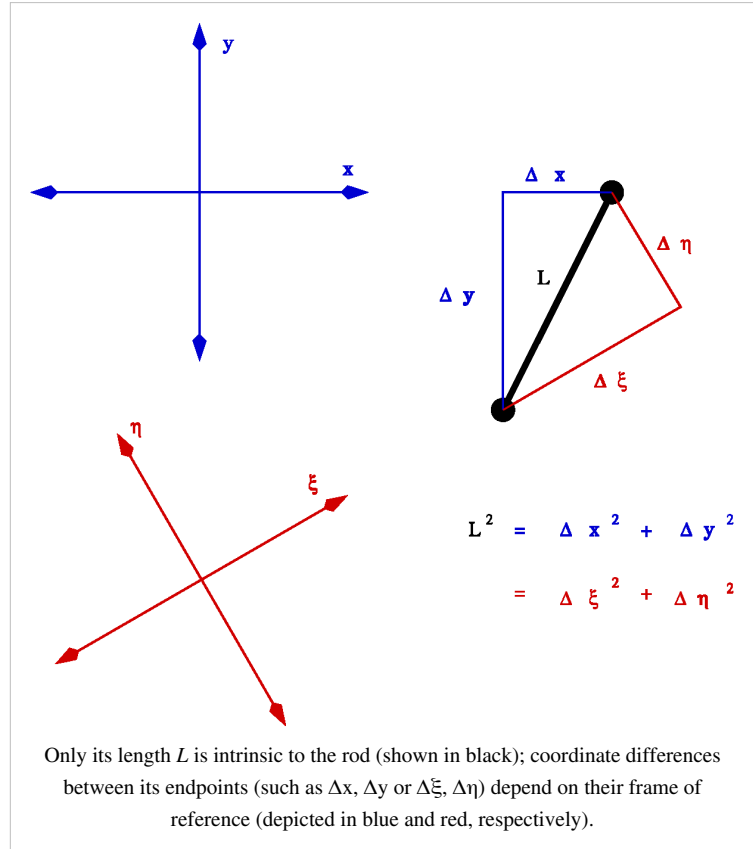
$$L^2 = \Delta \xi^2 + \Delta \eta^2 + \Delta \zeta^2.$$

Thus, the coordinates differences  $(\Delta x, \Delta y, \Delta z)$  and  $(\Delta \xi, \Delta \eta, \Delta \zeta)$  are not intrinsic to the rod, but merely reflect the reference frame used to describe it; by contrast, the length  $L$  is an intrinsic property of the rod. The coordinate differences can be changed without affecting the rod, by rotating one's reference frame.

The analogy in spacetime is called the interval between two events; an event is defined as a point in spacetime, a specific position in space and a specific moment in time. The spacetime interval between two events is given by

$$s^2 = L_1^2 - c^2 \Delta t_1^2 = L_2^2 - c^2 \Delta t_2^2$$

where  $c$  is the speed of light. According to special relativity, one can change a spatial and time separation  $(L_1, \Delta t_1)$  into another  $(L_2, \Delta t_2)$  by changing one's reference frame, as long as the change maintains the spacetime interval  $s$ . Such a change in reference frame corresponds to changing one's motion; in a moving frame, lengths and times are different from their counterparts in a stationary reference frame. The precise manner in which the coordinate and time differences change with motion is described by the Lorentz transformation.



## Solving Einstein's field equations

The distances between the spinning galaxies increase with time, but the distances between the stars within each galaxy stay roughly the same, due to their gravitational interactions. This animation illustrates a closed Friedmann universe with zero cosmological constant  $\Lambda$ ; such a universe oscillates between a Big Bang and a Big Crunch.

In non-Cartesian (non-square) or curved coordinate systems, the Pythagorean theorem holds only on infinitesimal length scales and must be augmented with a more general metric tensor  $g_{\mu\nu}$ , which can vary from place to place and which describes the local geometry in the particular coordinate system. However, assuming the cosmological principle that the universe is homogeneous and isotropic everywhere, every point in space is like every other point; hence, the metric tensor must be the same everywhere. That leads to a single form for the metric tensor, called the Friedmann–Lemaître–Robertson–Walker metric

$$ds^2 = -c^2 dt^2 + R(t)^2 \left( \frac{dr^2}{1 - kr^2} + r^2 d\theta^2 + r^2 \sin^2 \theta d\phi^2 \right)$$

where  $(r, \theta, \phi)$  correspond to a spherical coordinate system. This metric has only two undetermined parameters: an overall length scale  $R$  that can vary with time, and a curvature index  $k$  that can be only 0, 1 or  $-1$ , corresponding to flat Euclidean geometry, or spaces of positive or negative curvature. In cosmology, solving for the history of the universe is done by calculating  $R$  as a function of time, given  $k$  and the value of the cosmological constant  $\Lambda$ , which is a (small) parameter in Einstein's field equations. The equation describing how  $R$  varies with time is known as the Friedmann equation, after its inventor, Alexander Friedmann.<sup>[73]</sup>

The solutions for  $R(t)$  depend on  $k$  and  $\Lambda$ , but some qualitative features of such solutions are general. First and most importantly, the length scale  $R$  of the universe can remain constant *only* if the universe is perfectly isotropic with positive curvature ( $k=1$ ) and has one precise value of density everywhere, as first noted by Albert Einstein. However, this equilibrium is unstable and since the universe is known to be inhomogeneous on smaller scales,  $R$  must change, according to general relativity. When  $R$  changes, all the spatial distances in the universe change in tandem; there is an overall expansion or contraction of space itself. This accounts for the observation that galaxies appear to be flying apart; the space between them is stretching. The stretching of space also accounts for the apparent paradox that two galaxies can be 40 billion light years apart, although they started from the same point 13.7 billion years ago and never moved faster than the speed of light.

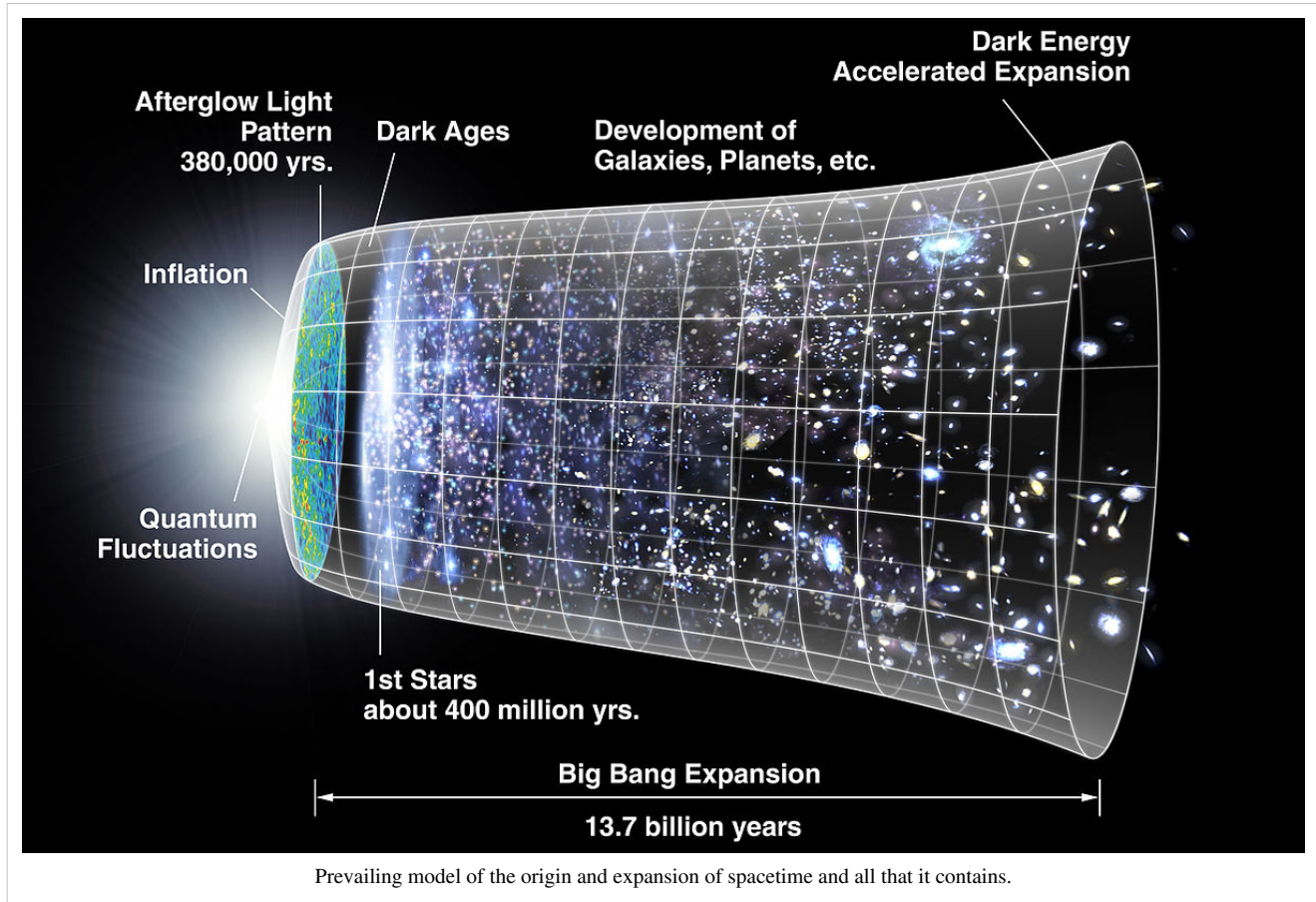
Second, all solutions suggest that there was a gravitational singularity in the past, when  $R$  goes to zero and matter and energy became infinitely dense. It may seem that this conclusion is uncertain since it is based on the questionable assumptions of perfect homogeneity and isotropy (the cosmological principle) and that only the gravitational interaction is significant. However, the Penrose–Hawking singularity theorems show that a singularity should exist for very general conditions. Hence, according to Einstein's field equations,  $R$  grew rapidly from an unimaginably hot, dense state that existed immediately following this singularity (when  $R$  had a small, finite value); this is the essence of the Big Bang model of the universe. A common misconception is that the Big Bang model predicts that matter and energy exploded from a single point in space and time; that is false. Rather, space itself was created in the Big Bang and imbued with a fixed amount of energy and matter distributed uniformly throughout; as space expands (i.e., as  $R(t)$  increases), the density of that matter and energy decreases.

Space has no boundary – that is empirically more certain than any external observation. However, that does not imply that space is infinite...(translated, original German)

Bernhard Riemann (Habilitationsvortrag, 1854)

Third, the curvature index  $k$  determines the sign of the mean spatial curvature of spacetime averaged over length scales greater than a billion light years. If  $k=1$ , the curvature is positive and the universe has a finite volume. Such universes are often visualized as a three-dimensional sphere  $S^3$  embedded in a four-dimensional space. Conversely, if  $k$  is zero or negative, the universe *may* have infinite volume, depending on its overall topology. It may seem

counter-intuitive that an infinite and yet infinitely dense universe could be created in a single instant at the Big Bang when  $R=0$ , but exactly that is predicted mathematically when  $k$  does not equal 1. For comparison, an infinite plane has zero curvature but infinite area, whereas an infinite cylinder is finite in one direction and a torus is finite in both. A toroidal universe could behave like a normal universe with periodic boundary conditions, as seen in "wrap-around" video games such as *Asteroids*; a traveler crossing an outer "boundary" of space going *outwards* would reappear instantly at another point on the boundary moving *inwards*.



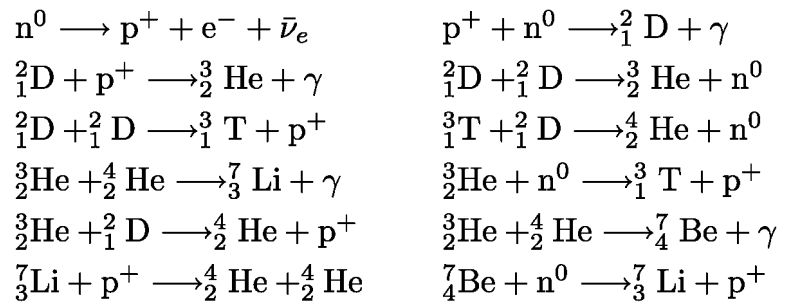
The ultimate fate of the universe is still unknown, because it depends critically on the curvature index  $k$  and the cosmological constant  $\Lambda$ . If the universe is sufficiently dense,  $k$  equals  $+1$ , meaning that its average curvature throughout is positive and the universe will eventually recollapse in a Big Crunch, possibly starting a new universe in a Big Bounce. Conversely, if the universe is insufficiently dense,  $k$  equals 0 or  $-1$  and the universe will expand forever, cooling off and eventually becoming inhospitable for all life, as the stars die and all matter coalesces into black holes (the Big Freeze and the heat death of the universe). As noted above, recent data suggests that the expansion speed of the universe is not decreasing as originally expected, but increasing; if this continues indefinitely, the universe will eventually rip itself to shreds (the Big Rip). Experimentally, the universe has an overall density that is very close to the critical value between recollapse and eternal expansion; more careful astronomical observations are needed to resolve the question.

## Big Bang model

The prevailing Big Bang model accounts for many of the experimental observations described above, such as the correlation of distance and redshift of galaxies, the universal ratio of hydrogen:helium atoms, and the ubiquitous, isotropic microwave radiation background. As noted above, the redshift arises from the metric expansion of space; as the space itself expands, the wavelength of a photon traveling through space likewise increases, decreasing its energy. The longer a photon has been traveling, the more expansion it has undergone; hence, older photons from

more distant galaxies are the most red-shifted. Determining the correlation between distance and redshift is an important problem in experimental physical cosmology.

Other experimental observations can be explained by combining the overall expansion of space with nuclear and atomic physics. As the universe expands, the energy density of the electromagnetic radiation decreases more quickly than does that of matter, since the energy of a photon decreases with its wavelength. Thus, although the energy density of the universe is now dominated by matter, it was once



Chief nuclear reactions responsible for the relative abundances of light atomic nuclei observed throughout the universe.

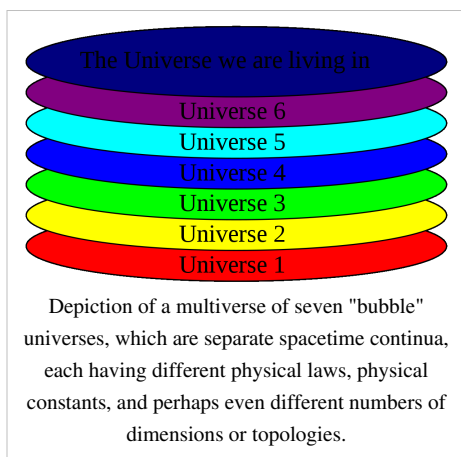
dominated by radiation; poetically speaking, all was light. As the universe expanded, its energy density decreased and it became cooler; as it did so, the elementary particles of matter could associate stably into ever larger combinations. Thus, in the early part of the matter-dominated era, stable protons and neutrons formed, which then associated into atomic nuclei. At this stage, the matter in the universe was mainly a hot, dense plasma of negative electrons, neutral neutrinos and positive nuclei. Nuclear reactions among the nuclei led to the present abundances of the lighter nuclei, particularly hydrogen, deuterium, and helium. Eventually, the electrons and nuclei combined to form stable atoms, which are transparent to most wavelengths of radiation; at this point, the radiation decoupled from the matter, forming the ubiquitous, isotropic background of microwave radiation observed today.

Other observations are not answered definitively by known physics. According to the prevailing theory, a slight imbalance of matter over antimatter was present in the universe's creation, or developed very shortly thereafter, possibly due to the CP violation that has been observed by particle physicists. Although the matter and antimatter mostly annihilated one another, producing photons, a small residue of matter survived, giving the present matter-dominated universe. Several lines of evidence also suggest that a rapid cosmic inflation of the universe occurred very early in its history (roughly  $10^{-35}$  seconds after its creation). Recent observations also suggest that the cosmological constant ( $\Lambda$ ) is not zero and that the net mass-energy content of the universe is dominated by a dark energy and dark matter that have not been characterized scientifically. They differ in their gravitational effects. Dark matter gravitates as ordinary matter does, and thus slows the expansion of the universe; by contrast, dark energy serves to accelerate the universe's expansion.

## Multiverse theory

Some speculative theories have proposed that this universe is but one of a set of disconnected universes, collectively denoted as the multiverse, challenging or enhancing more limited definitions of the universe.<sup>[18][74]</sup> Scientific multiverse theories are distinct from concepts such as alternate planes of consciousness and simulated reality, although the idea of a larger universe is not new; for example, Bishop Étienne Tempier of Paris ruled in 1277 that God could create as many universes as he saw fit, a question that was being hotly debated by the French theologians.<sup>[75]</sup>

Max Tegmark developed a four part classification scheme for the different types of multiverses that scientists have suggested in various



problem domains. An example of such a theory is the chaotic inflation model of the early universe.<sup>[76]</sup> Another is the many-worlds interpretation of quantum mechanics. Parallel worlds are generated in a manner similar to quantum superposition and decoherence, with all states of the wave function being realized in separate worlds. Effectively, the multiverse evolves as a universal wavefunction. If the big bang that created our multiverse created an ensemble of multiverses, the wave function of the ensemble would be entangled in this sense.

The least controversial category of multiverse in Tegmark's scheme is Level I, which describes distant space-time events "in our own universe". If space is infinite, or sufficiently large and uniform, identical instances of the history of Earth's entire Hubble volume occur every so often, simply by chance. Tegmark calculated our nearest so-called doppelgänger, is  $10^{10^{115}}$  meters away from us (a double exponential function larger than a googolplex).<sup>[77][78]</sup> In principle, it would be impossible to scientifically verify an identical Hubble volume. However, it does follow as a fairly straightforward consequence from otherwise unrelated scientific observations and theories. Tegmark suggests that statistical analysis exploiting the anthropic principle provides an opportunity to test multiverse theories in some cases. Generally, science would consider a multiverse theory that posits neither a common point of causation, nor the possibility of interaction between universes, to be an idle speculation.

## Shape of the universe

The shape or geometry of the universe includes both local geometry in the observable universe and global geometry, which we may or may not be able to measure. Shape can refer to curvature and topology. More formally, the subject in practice investigates which 3-manifold corresponds to the spatial section in comoving coordinates of the four-dimensional space-time of the universe. Cosmologists normally work with a given space-like slice of spacetime called the comoving coordinates. In terms of observation, the section of spacetime that can be observed is the backward light cone (points within the cosmic light horizon, given time to reach a given observer). If the observable universe is smaller than the entire universe (in some models it is many orders of magnitude smaller), one cannot determine the global structure by observation: one is limited to a small patch.

Among the Friedmann–Lemaître–Robertson–Walker (FLRW) models, the presently most popular shape of the Universe found to fit observational data according to cosmologists is the infinite flat model,<sup>[79]</sup> while other FLRW models include the Poincaré dodecahedral space<sup>[80][81]</sup> and the Picard horn.<sup>[82]</sup> The data fit by these FLRW models of space especially include the Wilkinson Microwave Anisotropy Probe (WMAP) maps of cosmic background radiation. NASA released the first WMAP cosmic background radiation data in February 2003. In 2009 the Planck observatory was launched to observe the microwave background at higher resolution than WMAP, possibly providing more information on the shape of the Universe. The data should be released in late 2012.

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## External links

- Is there a hole in the universe? (<http://www.howstuffworks.com/hole-in-universe.htm>) at HowStuffWorks
- *Stephen Hawking's Universe* (<http://www.pbs.org/wnet/hawking/html/home.html>) – Why is the universe the way it is?
- Cosmology FAQ ([http://www.astro.ucla.edu/~wright/cosmology\\_faq.html](http://www.astro.ucla.edu/~wright/cosmology_faq.html))
- Cosmos – An "illustrated dimensional journey from microcosmos to macrocosmos" (<http://www.shekpvar.net/~dna/Publications/Cosmos/cosmos.html>)
- Illustration comparing the sizes of the planets, the sun, and other stars (<http://www.co-intelligence.org/newsletter/comparisons.html>)
- Logarithmic Maps of the Universe (<http://www.astro.princeton.edu/~mjuric/universe/>)
- My So-Called Universe (<http://web.archive.org/web/20101225211703/http://www.slate.com/id/2087206/nav/navoa/>) – Arguments for and against an infinite and parallel universes
- The Dark Side and the Bright Side of the Universe ([http://cosmology.lbl.gov/talks/Ho\\_07.pdf](http://cosmology.lbl.gov/talks/Ho_07.pdf)) Princeton University, Shirley Ho
- Richard Powell: *An Atlas of the Universe* (<http://www.atlasoftheuniverse.com/>) – Images at various scales, with explanations
- Multiple Big Bangs (<http://www.npr.org/templates/story/story.php?storyId=1142346>)
- Universe – Space Information Centre (<http://www.exploreuniverse.com/ic/>)
- Exploring the Universe (<http://www.nasa.gov/topics/universe/index.html>) at Nasa.gov

## Videos

- The Known Universe (<http://www.youtube.com/embed/17jymDn0W6U>) created by the American Museum of Natural History
  - Understand The Size Of The Universe (<http://www.youtube.com/embed/0fKBhvDjuy0>) – by Powers of Ten
  - 3-D Video (01:46) - Over a Million Galaxies of Billions of Stars each - BerkeleyLab/animated (<http://www.youtube.com/watch?v=08LBltePDZw>)
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